

Report of an Induced Polarization Survey

On the

2.34618

Byers Property

Byers Township, Ontario

Claim Nos. 4207658 4201796

Porcupine Mining Division

For

Amador Gold Corp.

March 31, 2007 Timmins, Ontario Matthew Johnston Consulting Geophysicist 1226 Gatineau Blvd. Timmins, Ont. P4R 1E3

Table of Contents

		Page No.
1.0	Introduction	2
2.0	Location and Access	2
3.0	Summary of 2007 Geophysical and Gridding Program	2
4.0	Discussion of Results	5
5.0	Conclusions and Recommendations	6

Statement of Qualifications

Appendices

Appendix A Geophysical Instruments and Survey Methods

List of Maps

Мар	Scale
7 I.P./Resistivity Pseudo-Sections Lines 2000E to 2700E	1:2500
Filtered Resistivity Contours with I.P. Anomalies Plan Map	1:5000

1.0 Introduction

The Byers property of Amador Gold Corp. consists of two unpatented mining claims, 4207658 and 4201796, located in Byers Township, Porcupine Mining Division. During March 2007, a geophysical survey program consisting of induced polarization and resistivity surveys was conducted over a portion of this claim group. Ray Meikle and Associates of North Bay, Ontario carried out the geophysical surveys. The I.P. surveys were performed in order to evaluate and map the presence of disseminated to massive sulphides with respect to their location, width, and concentrations.

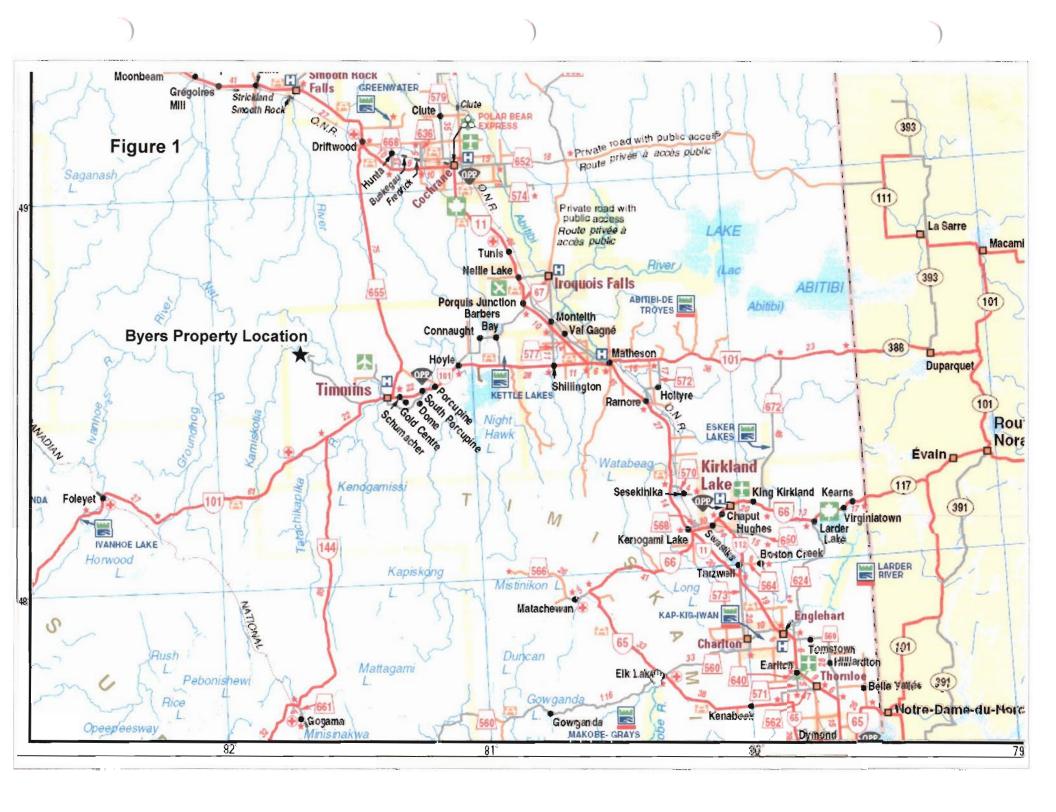
2.0 Location And Access

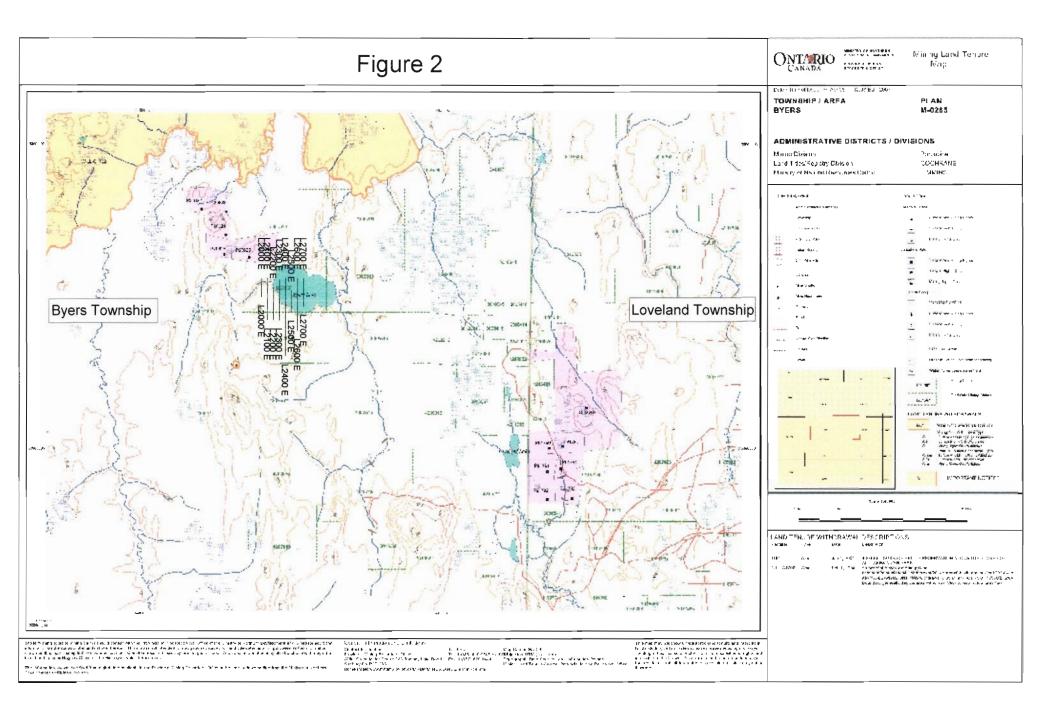
The property is located approximately 30 km northeast of the city of Timmins and northwest Kamiscotia Lake, in the north central portion of Byers Township, Porcupine Mining Division. Access to the property is via highway 101 west from Timmins to the Kamiscotia highway. Travel north on the Kamiscotia highway for approximately 20 kilometres. From this point local logging roads and trails provide access to the grid area via snowmobile, four wheel drive vehicles and walking (see figures 1 and 2).

3.0 Summary of 2007 Geophysical Program

The geophysical program consisted of induced polarization and resistivity surveying. This survey was carried out on a grid of previously cut lines oriented at 0° and spaced every 100 meters and chained and marked every 25 meters. The I.P. survey was performed using a pole-dipole electrode configuration. The dipole a spacing was 25 metres and increasing separations of n=1, n=2, n=3, n=4, n=5 and n=6 times the dipole spacing was measured in order to map the response at depth. A total of approximately **7.2-km** of I.P. data was measured and recorded by Ray Meikle and Associates of North Bay, Ontario.

The I.P. equipment used for the survey consisted of a Phoenix IPT-1 transmitter operating in the time domain powered by a 2 kilowatt MG-2 motor generator. The





chargeability (measured in mV/V) between the transmitted current and the received voltage is recorded by an Iris Elrec IP-6 I.P. receiver which records the chargeability and the apparent resistivity for each set of dipoles. The chargeability measured in this survey is an equivalent measure of the polarization of the underlying lithology.

A description of the survey method and equipment used can be found in Appendix A.

4.0 Discussion of Results

The results of the I.P. survey are presented as contoured and posted pseudosections of the apparent resistivity and chargeabilities at a scale of 1:2500. In addition plan maps at a scale of 1:5,000 showing the contours of the filtered apparent resistivity and phase angles with the interpretation and location of the I.P. anomalies is also presented. All maps accompany this report in the pocket at the back of this report.

The resistivity data as displayed by the contoured resistivity plan map shows a wide variation of measured resistivities in the range of 35 to 11169 ohm-m with a background resistivity of approximately 821 ohm-m. The higher resistivity areas of the grid may likely be mapping areas of bedrock ridges and sub-cropping bedrock areas. These areas are quite evident on the plan map. It is also possible the high resistivity zones may be outlining more resistive felsic lithology or silica altered horizons as well. A prominent northwest-southeast trending linear resistivity high can be observed between lines 2000E and 2400E. This area is also observed to have several generally northwest-southeast trending I.P. anomaly trend; labeled as anomaly A.

The I.P. anomalies have been interpreted and are displayed on the plan map of the filtered resistivity and chargeability responses as well. Emphasis was placed on identifying I.P. anomalies, which were thought to originate within the bedrock as opposed to cultural sources (identified on the pseudo-sections); and those I.P. anomalies that, may be associated with bedrock relief. One anomaly trend was identified and labeled on the plan map (A), as well as several other isolated I.P. anomalies that are not

readily grouped into trends. The responses are interpreted to occur at depths of between 15 and 25 metres below surface.

5.0 Conclusions and Recommendations

The induced polarization surveys completed over the Byers grid were successful in mapping several zones of anomalous I.P. effects as well as mapping the bedrock resistivity. Many of the interpreted I.P. anomalies are strong and well defined and will likely require further investigation in order to determine their causes. The most promising I.P. anomalies, which are thought to arise from bedrock sources, have been interpreted and identified.

It is always difficult to quantitatively rate all of the I.P. anomalies in terms of their economic potential when searching for exploitable mineral deposits, but it is possible that some of the I.P. anomalies mapped by this survey are caused by disseminated to semi-massive metallic mineralization. This type of mineralization is often associated which valuable deposits of massive sulphides, gold and platinum group minerals.

All of the responses should be investigated further in order to determine the priority of follow-up needed. The anomalies should be further screened utilizing any other different types of geophysical surveys that may have been undertaken on the Byers grid. This would aid greatly in further refining the interpretation of the I.P.

survey. Any existing geological or geochemical information for the surveyed grid will further aid in further assessing any geophysical anomalies.

Respectively Submitted,

Matthew Johnston

Consulting Geophysicist

Statement of Qualifications

This is to certify that: MATTHEW JOHNSTON

I am a resident of Timmins; province of Ontario since June 1, 1995.

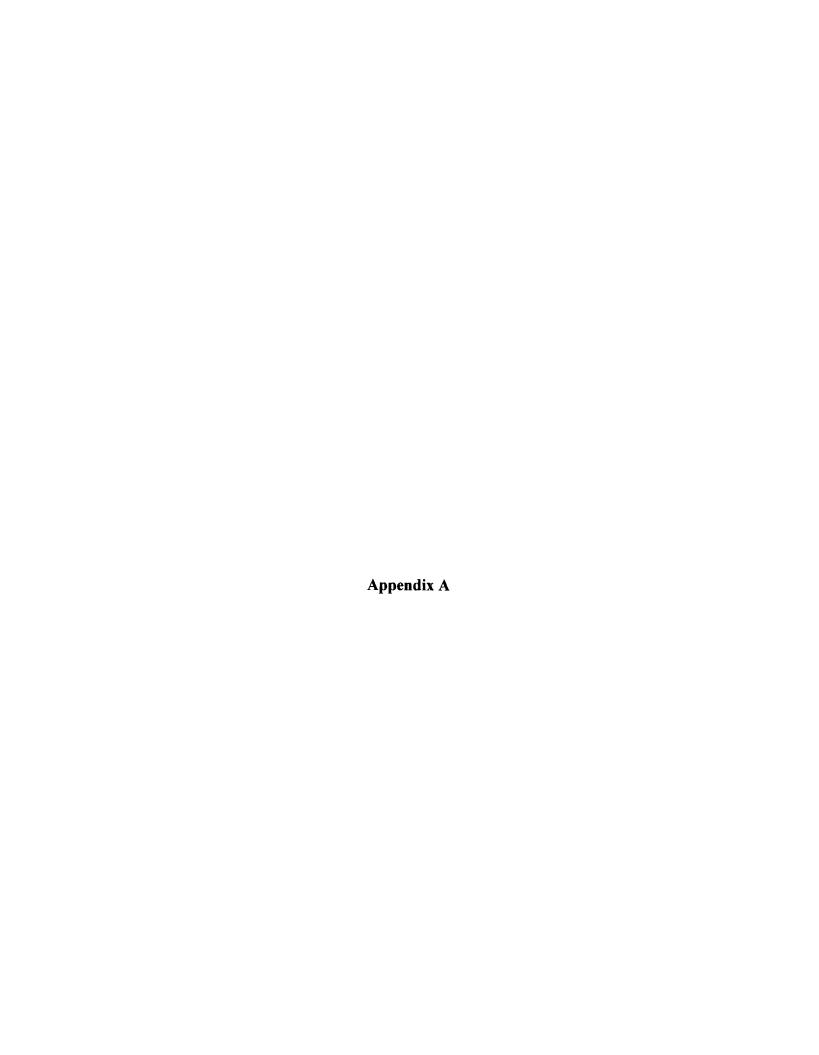
I am self-employed as a Consulting Geophysicist, based in Timmins, Ontario.

I have received a B.Sc. in geophysics from the University of Saskatchewan; Saskatoon, Saskatchewan in 1986.

I have been employed as a professional geophysicist in mining exploration, environmental and other consulting geophysical techniques since 1986.

I am registered as professional geophysicist (P.Geoph.) with the Association of Professional Engineers, Geologists and Geophysicists of the N.W.T and Nunavut (L1438).

Signed in Timmins, Ontario, this March 31, 2007



Induced Polarization Surveys

Time domain IP surveys involve measurement of the magnitude of the polarisation voltage (Vp) that results from the injection of pulsed current into the ground.

Two main mechanisms are known to be responsible for the IP effect although the exact causes are still poorly understood. The main mechanism in rocks containing metallic conductors is electrode polarisation (overvoltage effect). This results from the build up of charge on either side of conductive grains within the rock matrix as they block the flow of current. On removal of this current the ions responsible for the charge slowly diffuse back into the electrolyte (groundwater) and the potential difference across each grain slowly decays to zero. The second mechanism, membrane polarisation, results from a constriction of the flow of ions around narrow pore channels. It may also result from the excessive build up of positive ions around clay particles. This cloud of positive ions similarly blocks the passage of negative ions through pore spaces within the rock. On removal of the applied voltage the concentration of ions slowly returns to its original state resulting in the observed IP response. In TD-IP the current is usually applied in the form of a square waveform, with the polarisation voltage being measured over a series of short time intervals after each current cut-off, following a short delay of approximately 0.5s. These readings are integrated to give the area under the decay curve, which is used to define Vp. The integral voltage is divided by the observed steady voltage (the voltage due to the applied current plus the polarisation voltage) to give the apparent chargeability (Ma) measured in milliseconds. For a given charging period and integration time the measured apparent chargeability provides qualitative information on the subsurface geology.

The polarisation voltage is measured using a pair of non-polarising electrodes similar to those used in spontaneous potential measurements and other IP techniques.

