Report of Induced Polarization Surveys and

Total Field Magnetic Surveys

On the Turtlepond Project

Turtle Pond, Emmons Lake, Prig, North Glatz, Glatz, and Double E Grids

Turtlepond Lake Area, Ontario

Claim Nos
. 4219025, 1247471, 1247472, 4219034, 4219035,
4219032, 3012581, and 4219039

Kenora Mining Division

For

Canadian Arrow Mines Ltd.

December 15, 2008
Timmins, Ontario

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Table of Contents

Appendices

List of Maps

1.0 Introduction

The Turtle Pond, Emmons Lake, Prig, Glatz, North Glatz and Double E grids are located on the Turtlepond property of Canadian Arrow Mines Ltd., located in Turtlepond Area, Kenora Mining Division. The grids, in Turtlepond Area, cover portions or all of mining claim numbers 4219025, 1247471, 1247472, 4219034, 4219035, 4219032, 3012581, and 4219039

During July and August 2008, a geophysical survey program consisting of induced polarization and resistivity surveys and total field magnetic surveys was conducted over a portions of these claim groups. Ray Meikle and Associates of North Bay, Ontario, carried out the geophysical surveys. The geophysical surveys were performed in order to evaluate and map the presence of disseminated to massive sulphides with respect to their location, width, and concentrations.

2.0 Location And Access

The Turtlepond Property is located 30 km southeast of Dryden Ontario, within Turtlepond Lake Township (G-2595), in the Kenora Mining Division. The grids are located in the west central portion of Turtlepond township. The project is centered at UTM 520609 E 5487918 N Zone 15, NTS 52F/10SW.

The grids are accessed by traveling along Hwy 502 approximately 30 km south from Dryden. A gravel road going east at this point for approximately 1.0 km is used to access all of the grids and crosses the central portions of claim group (see figures 1 to 2).

3.0 Summary of 2008 Geophysical Program

The geophysical program consisted of induced polarization and resistivity surveying (I.P.) and total field magnetic surveys. These surveys were carried out on a grids of previously established lines generally spaced every 100 meters and chained and marked every 25 meters. The grid are summarized in Table 1.0.

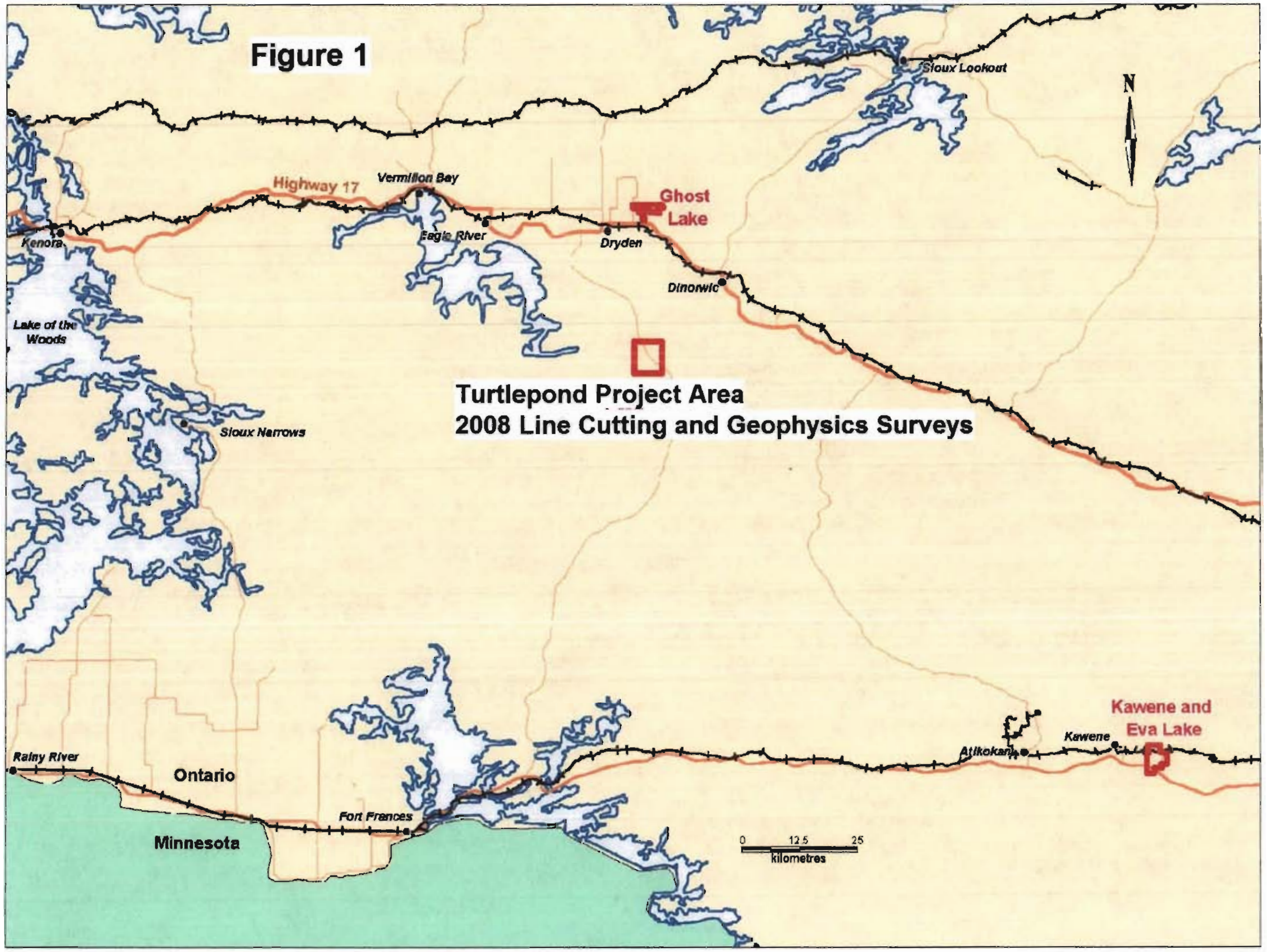
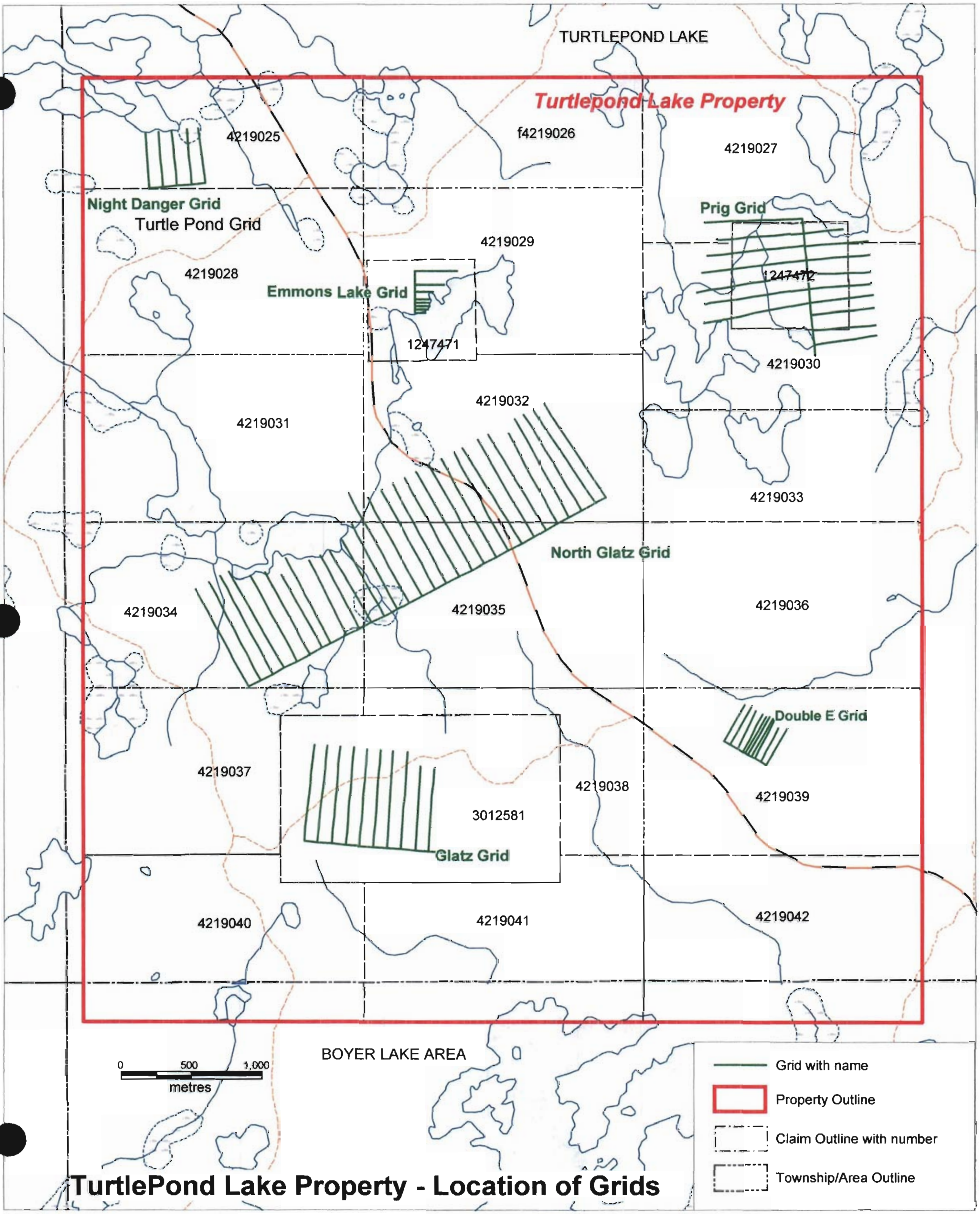


Figure 1
Sioux Lookout
N
Highway 17
Vermilion Bay
Ghost
Lake
Kenora
Eagle River
Dryden
Dinorwic
Lake of the
Woods
Turtlepond Project Area
2008 Line Cutting and Geophysics Surveys
Sioux Narrows
Kawene and
Eva Lake
Kawene
Atikokan
Rainy River
Ontario
Fort Frances
Minnesota
0
12.5
25
kilometres



TurtlePond Lake Property - Location of Grids

Grid Name	IP Kilometers and n Level	Magnetic Km's.	Baseline Azimuth (degrees)	Line Spacing
Turtle Pond	1.3 n=4	2	88	100 m
Emmons Lake	0.5 n=4	1	0	100 m
Prig	6.0 n=4	16.33	355	100 m
Glatz	8.0 n=6	7.8	95	100 m
Glatz North	11 n=6	8	64	100 m
Double E	2 n=4	0	120	100 m

The I.P. surveys were performed using a pole-dipole electrode configuration. The dipole 'a' spacing was 25 meters and increasing separations of n=1, n=2, n=3, n=4, n=5 and n=6 times the dipole spacing was measured in order to map the response to depth.

The I.P. equipment used for the surveys consisted of a Phoenix IPT1 3000 watt transmitter operating in the time domain powered by a 2.5 kilowatt motor generator. The chargeability (measured in mV/V) between the transmitted current and the received voltage is recorded by a Iris Elrec Pro I.P. receiver which records the chargeability and the apparent resistivity for each set of dipoles. The chargeability measured in this survey is a measure of the polarization of the underlying lithology.

The total field magnetic surveys, using a GEM GSM-19 magnetometer with readings collected every 12.5 meters along all lines.

A description of the survey method and equipment used can be found in Appendix A.

4.0 **Discussion of Results**

The results of the I.P. surveys are presented as contoured and posted pseudo-sections of the apparent resistivity and recorded chargeability's at a scale of 1:2500 In addition plan maps at a scale of 1:5,000 showing the contours of the N=1 apparent resistivity with the interpretation and location of the I.P. anomalies is also presented. All maps accompany this report in the pocket at the back of this report.

The magnetic data has been presented on plan maps at a scale of 1:5000, showing the contours as well the interpretations (see maps in pocket).

The higher resistivity areas of the grid areas may possibly be mapping areas of bedrock ridges and sub-cropping bedrock areas. These areas are quite evident on the plan map. It is also possible the high resistivity zones may be outlining more resistive felsic lithology or silica altered horizons as well. The depths of all of the identified I.P. anomalies are interpreted to be shallow; within the range of 4 to 65 meters below surface.

Turtle Pond Grid

One anomalous I.P. Anomaly trend labeled as TP1 was mapped on the Turtle Pond grid. This is a well defined I.P. Response which flanks a moderate strength linear magnetic anomaly.

Prig Lake Grid

The magnetic survey on the Prig Lake grid indicates a relatively active magnetic background with magnetic values ranging between 56189 and 58868 nT. The background magnetic field strength is 57593 nT. The overall magnetic pattern is disrupted by several complex, linear, and slightly discontinuous moderate anomalous magnetic highs trending in a northerly strike direction. The isomagnetic contour pattern suggests an underlying lithology striking in an north-south direction. Several fault zones have been interpreted within the grid area from the magnetic survey. These anomalies may represent major lithological contacts or structural anomalies which may be significant in this area. These anomaly locations are indicated and shown on the contour map. All of the anomalies are easily identified and are labeled on the plan maps.

Numerous well defined I.P. Anomalies were mapped on the Prig Lake Grid and have been grouped into four trends labeled as P1 to P4. Anomalies P2, P3, and P4 also show sporadic correlation with moderate magnetic anomalies along their strike length which may suggest pyrrhotite mineralization within the underlying lithology.

Emmons Lake Grid

Only anomalous I.P. Trend was observed within the Emmons Lake grid and is labeled as E1. The anomaly on line 1200N is not fully defined and the line would have to be extended further east in order to fully define this anomaly. Both I.P. anomalies occur within a magnetic low area and may represent mineralization within an altered lithologic setting.

Glatz North Grid

The magnetic survey on the North Glatz grid indicates a relatively active magnetic background with magnetic values ranging between 56932 and 59244 nT. The background magnetic field strength is 57670 nT. The overall magnetic pattern is disrupted by several complex, linear, and slightly discontinuous moderate anomalous magnetic highs trending at an azimuth of approximately 67 degrees. The anomalous magnetic area occurs primarily between lines 1200E and 2000E from 1200N to 1400N. The isomagnetic contour pattern suggests an underlying lithology striking in an north east-southwest direction. Several fault zones have been interpreted within the grid area from the magnetic survey. These anomalies may represent major lithological contacts or structural anomalies which may be significant in this area. These anomaly locations are indicated and shown on the contour map. All of the anomalies are easily identified and are labeled on the plan maps.

Many well defined and strong I.P. anomalies were identified with the North Glatz grid and have been grouped into seven trends labeled at NG1 to NG7. Anomalies NG1, NG2, and NG3; in particular exhibit strong correlation to the magnetic anomalies in the south central part of the grid area previously discussed. No magnetic data was surveyed within the eastern part of the grid area; however the I.P. anomalies located are strong and well defined.

Glatz Grid

The magnetic survey on the Glatz grid indicates an active magnetic background with magnetic values ranging between 56486 and 59271 nT. The background magnetic field strength is 57625 nT. The overall magnetic pattern is disrupted by several complex, linear, and slightly discontinuous moderate anomalous magnetic highs trending at an azimuth of approximately 93 degrees. The anomalous magnetic area occurs primarily throughout the grid are between 4200N and 4500N. This anomalous magnetic corridor suggest an area of structural deformation and it is noted that many I.P. anomalies also occur within this area. The isomagnetic contour pattern suggests an underlying lithology striking in an east-west direction. Several fault zones have been interpreted within the grid area from the magnetic survey. These anomalies may represent major lithological contacts or structural anomalies which may be significant in this area. These anomaly locations are indicated and shown on the contour map. All of the anomalies are easily identified and are labeled on the plan maps.

Many I.P. anomalies were mapped and have been grouped into six trends labeled as G1 to G6. Anomalies G1, G2, G3, G4, and G5 all display close correlation with moderate strength magnetic anomalies possibly indicating the presence pyrrhotite with the bedrock at these locations.

Double E Grid

Several well defined I.P. anomalies were mapped on the Double E grid. One trend was identified which has been labeled as D1 as well several other well defined isolated I.P. anomalies. Anomaly D1 occurs within a resistivity low suggesting conductive sulphides or graphite as a possible source.

5.0 Conclusions and Recommendations

The induced polarization and magnetic surveys and completed over the Turtle Pond, Prig Lake, Emmons, Double E, Glatz, and North Glatz grids were successful in mapping many zones of anomalous I.P. effects, magnetic anomalies, as well as mapping the bedrock resistivity. Many of the interpreted I.P. anomalies are strong and well defined and will likely require further investigation in order to determine their causes. The most promising I.P. anomalies, which are thought to arise from bedrock sources, have been interpreted and identified.

It is always difficult to quantitatively rate all of the I.P. anomalies in terms of their economic potential when searching for exploitable mineral deposits, but it is possible that some of the I.P. anomalies mapped by this survey are caused by disseminated to semi-massive metallic mineralization. This type of mineralization is often associated which valuable deposits of massive sulphides, gold and platinum group minerals.

The I.P. anomalies; T1, G1, G2, G3, G4, G5, E1, D1, P1, P2, P3, P4, NG1, NG2, NG3, and NG5, should all be considered as priority anomalies warranting further investigation. The other well defined I.P. will also have to be considered as prospective zones requiring follow-up exploration in order to determine their source mineralization.

Prior to drill testing any of the anomalies it is recommended that a program of geological mapping, prospecting and possibly trenching be undertaken in order to attempt to identify the sources of the conductive and magnetic anomalies

Any existing geological, diamond drilling or geochemical information that may exist in the mining recorder assessment files should be investigated and compiled prior to further exploration of the Turtlepond property in order to accurately assess the area of the current geophysical surveys and to determine the most effective follow-up exploration method for these anomalies.

Respectively Submitted,

Matthew Johnston

Consulting Geophysicist

Statement of Qualifications

This is to certify that: MATTHEW JOHNSTON

I am a resident of Timmins; province of Ontario since June 1, 1995.

I am self-employed as a Consulting Geophysicist, based in Timmins, Ontario.

I have received a B.Sc. in geophysics from the University of Saskatchewan; Saskatoon, Saskatchewan in 1986.

I have been employed as a professional geophysicist in mining exploration, environmental and other consulting geophysical techniques since 1986.

Signed in Timmins, Ontario, this December 15, 2008

Matthew Johnston

Appendix A

Induced Polarization Surveys

Time domain IP surveys involve measurement of the magnitude of the polarisation voltage (Vp) that results from the injection of pulsed current into the ground.

Two main mechanisms are known to be responsible for the IP effect although the exact causes are still poorly understood. The main mechanism in rocks containing metallic conductors is electrode polarisation (overvoltage effect). This results from the build up of charge on either side of conductive grains within the rock matrix as they block the flow of current. On removal of this current the ions responsible for the charge slowly diffuse back into the electrolyte (groundwater) and the potential difference across each grain slowly decays to zero. The second mechanism, membrane polarisation, results from a constriction of the flow of ions around narrow pore channels. It may also result from the excessive build up of positive ions around clay particles. This cloud of positive ions similarly blocks the passage of negative ions through pore spaces within the rock. On removal of the applied voltage the concentration of ions slowly returns to its original state resulting in the observed IP response. In TD-IP the current is usually applied in the form of a square waveform, with the polarisation voltage being measured over a series of short time intervals after each current cut-off, following a short delay of approximately 0.5s. These readings are integrated to give the area under the decay curve, which is used to define Vp. The integral voltage is divided by the observed steady voltage (the voltage due to the applied current plus the polarisation voltage) to give the apparent chargeability (Ma) measured in milliseconds or mV/V. For a given charging period and integration time the measured apparent chargeability provides qualitative information on the subsurface geology.

The polarisation voltage is measured using a pair of non-polarising electrodes similar to those used in spontaneous potential measurements and other IP techniques.

Survey Theory - Total Field Magnetics

Magnetic Survey

Theory:

The magnetic method is based on measuring alteration in the shape and magnitude of the earth's naturally occurring magnetic field caused by changes in the magnetization of the rocks in the earth. These changes in magnetization are due mainly to the presence of the magnetic minerals, of which the most common is magnetite, and to a lesser extent illuminate, pyrrhotite, and some less common minerals. Magnetic anomalies in the earth's filed are caused by changes in two types of magnetization: (1) Induced, caused by the magnetic field being altered and enhanced by increases in the magnetic susceptibility of the rocks, which is a function of the concentration of the magnetic minerals. (2) Remanent magnetism is independent of the earth's magnetic field, and is the permanent magnetization of the magnetic particles (magnetite, etc.) in the rocks. This is created when these particles orient themselves parallel to the ambient field when cooling. This magnetization may not be in the same direction as the present earth's field, due to changes in the orientation of the rock or the field. The **unit** of measurement (variations in intensity) is commonly known as the Gamma which is equivalent to the nanotesla (nT).

Method:

The magnetometer, a GEM Systems **GSM-19** with an Overhauser sensor measures the **Total Magnetic Field** (TFM) perpendicular to the earth's field (horizontal position in the polar region). The unit has no moving parts, produces an absolute and relatively high resolution measurement of the field and displays the measurement on a digital lighted display and is recorded (to memory). Initially, the tuning of the instrument should agree with the nominal value of the magnetic field for each particular area. The Overhauser procession magnetometer collected the data with a **0.2 nanoTesla accuracy.** The operator read each and every line at a 12.5 **m** intervals with the sensor attached to the top of four (56cm), aluminum tubing sections. The readings were corrected for changes in the earth's magnetic field (diurnal drift) with a similar GSM-19 magnetometer, acting as a stationary base station which automatically read and stored the readings at every 15 seconds. The data from both units was then downloaded to PC and base corrected values were computed.