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Work Assessment Report

on the

JUNIOR LAKE PROPERTY

2014 Ground Geophysics Program
3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) Survey
(VW Deposit, VW West, B4-7 East areas)

Falcon Lake Area
Thunder Bay North Mines and Minerals Division
Ontario

NTS 52I/08 and 42L/05

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January 5, 2016
Thunder Bay, Ontario

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1 SUMMARY

The Junior Lake property is located approximately 230 kilometres north-northeast of the city of Thunder Bay, Ontario, within the central portion of the Caribou-O'Sullivan Greenstone Belt. The property is host to two NI 43-101 compliant mineral resources – the B4-7 Ni-Cu-Co-PGE deposit and the VW Ni deposit, located 3 kilometres apart. Other occurrences of Ni-Cu-PGE, Cu, Cu-Zn, Cr, Li and Au are known on the property.

This report covers the 2014 ground geophysics program conducted on the VW Deposit, VW West and B4-7 East areas, located in the central portion of the Junior Lake property. During January and February 2014, an Orion 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) survey was conducted by Quantec Geoscience Ltd. (Quantec), Toronto, for Landore Resources.

The survey area covered 739.02 hectares, extending from cut grid line 300E in the B4-7 Nickel-Copper-Cobalt-PGEs deposit, through the VW Nickel deposit to line 3700E, and from 700N to 1500S. This survey is located directly adjacent to the east of Landore Resources' 2012 Orion 3D DCIP + MT survey block covering the B4-7 West zone. The attached report by Quantec describes the survey logistics, data processing, interpretation, and provides the specifications of the survey.

This DCIP + MT ground geophysics survey acquired three sets of data in multi-directions; DC, IP and MT, and is a true three dimensional survey. Sophisticated digital signal processing was utilized to obtain high resolution imaging at depths up to 1000+ metres below surface. DC resistivity was utilized to identify prospective nickel mineralization, and IP chargeability used to investigate potential copper and PGE targets.

The 2014 DCIP-MT ground geophysics survey followed up on an identical survey conducted in late 2012 covering from line 1400W to 400E in the B4-7 deposit. The 2012 DC resistivity results supported an Exploration Target identified along strike and down dip from the B4-7 resource containing a potential 1.5 Mt to 2.0 Mt of sulphide mineralisation of similar grade range to that which has been outlined to-date (Tuomi, 2013).

The B4-7 combined open pit and underground NI 43-101 compliant resource is 2,695,000 tonnes at 1.24% Nickel equivalent (NiEq) for 33,248 tonnes of contained metal, all within the Indicated category (2013 Mineral Resource estimate and a NI 43-101 Technical Report on the B4-7 Deposit, RPA, Toronto, Canada).

The VW NI 43-101 compliant resource is 3.73 million tonnes at 0.49% NiEq in the Indicated category at a cut-off grade of 0.25 per cent. nickel. The resource holds a further 0.72 million tonnes at 0.49% NiEq in the Inferred category at the same cut-off grade giving a combined total of 21,760 tonnes of contained NiEq metal. The deposit remains open down plunge at depth and along strike to the east and to the west. (2010 Mineral Resource estimate and a NI 43-101 Technical Report on the VW Deposit, RPA, Toronto, Canada).

Results from the 2014 survey have been highly encouraging, delineating nine significant new zones ranging from approximately 400m to 1,200m in length of potential nickel sulphide mineralization along strike and adjacent to the existing B4-7 Nickel-Copper-Cobalt-PGEs resource and the VW Nickel resource. Numerous targets are at depth and below existing Landore exploration drilling.

The DCIP + MT geophysics survey on the VW Deposit, VW West and B4-7 East areas was successful in identifying potential massive sulphide mineralization as well as disseminated sulphide mineralization containing economic grades of nickel, copper, cobalt, PGEs and gold. Further drilling is warranted to follow up on promising geophysics results.

The 2014 DCIP + MT ground geophysics program included program preparation, survey implementation, geophysical and geological analysis of results. The total amount from this exploration program claimed for assessment credit is $1,014,222.58.

2 INTRODUCTION

This report and accompanying documentation presents the results of the 2014 Orion 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) ground geophysics program on Landore Resources Canada Inc.'s Junior Lake property. The Junior Lake property is located approximately 230 kilometres north-northeast of the city of Thunder Bay, Ontario, within the central portion of the Caribou-O'Sullivan Greenstone Belt. It is host to several PGE-Cu-Ni, Cu, Cu-Zn, Li, Au, and Ag occurrences. In the vicinity of the 2014 ground geophysics program, the property hosts two NI 43-101 compliant nickel deposits – the B4-7 Ni-Cu-Co-PGE deposit and the VW Ni deposit, located three kilometres apart.

The 2014 DCIP + MT ground geophysics program was conducted from cut grid line 300E in the B4-7 Nickel-Copper-Cobalt-PGEs deposit, through the VW Nickel deposit to line 3700E, and from 700N to 1500S, and is located directly adjacent to the east of Landore Resources' 2012 ORION 3D DCIP+MT survey block covering the B4-7 West zone. The survey was conducted by Quantec Geoscience Ltd., Toronto, for Landore Resources.

The ground geophysical survey results indicate that the geophysical horizons hosting the B4-7 deposit and the VW deposit extends to the east and west, and is highly prospective for further sulphide mineralization. The survey results have also identified several other areas of prospective sulphide mineralization the VW West area which warrant follow-up drilling.

This report is submitted to the Ontario Ministry of Northern Development and Mines Geoscience Assessment Office to claim assessment credit.

3 PROPERTY DESCRIPTION AND LOCATION

The Junior Lake property is located approximately 230 km north-northeast of Thunder Bay, Ontario, and approximately 75 km east-northeast of the village of Armstrong, Ontario (Figure 2-1). The centre of the property is located at 87º59’4”W longitude and 50º23’9”N latitude; NAD83 UTM coordinates Zone 16, 430,000E and 5,580,000N. The property area is within the NTS 1:50,000 Jackfish Lake and Toronto Lake topographic map sheets NTS 52I/08 and 44L/05, respectively. The Junior Lake property claims and leases are located on the Falcon Lake, Junior Lake, Toronto Lake, Kapikotongwa River, Summit Lake, and Willet Lake claim maps (Thunder Bay Mining Division areas NTS 52I/08NE and SE, 42L/05NW, SE and SW).

LAND TENURE

Landore’s Junior Lake property consists of 95 mineral claims (1,145 units) and four mining leases totaling 3,793 hectares (Tables 3-1 and 3-2, Figure 3-2).

Landore held a 100% interest in claims TB1077140 to TB1077142, TB1217179 to TB1217181, and TB1233556 and TB1233557, subject to a 2% net smelter return (NSR) royalty held by Wing Resources Inc. The above claims, except TB1077140, have been taken to lease. The B4-7 Deposit lies on patented claims PA39127, PA39128 and lease CLM460, whereas the VW Deposit lies on lease CLM461.

The exploration work undertaken by Landore prior to 28th August, 2008 was on mining leases in which Landore held a 100% interest: mining claims TB1077142, TB1217179. These claims were taken to lease (CLM 461) on 28th August, 2008.

Figure 3-1: Junior Lake Property Location

Figure 3-2: Junior Lake Property Leases and Claims

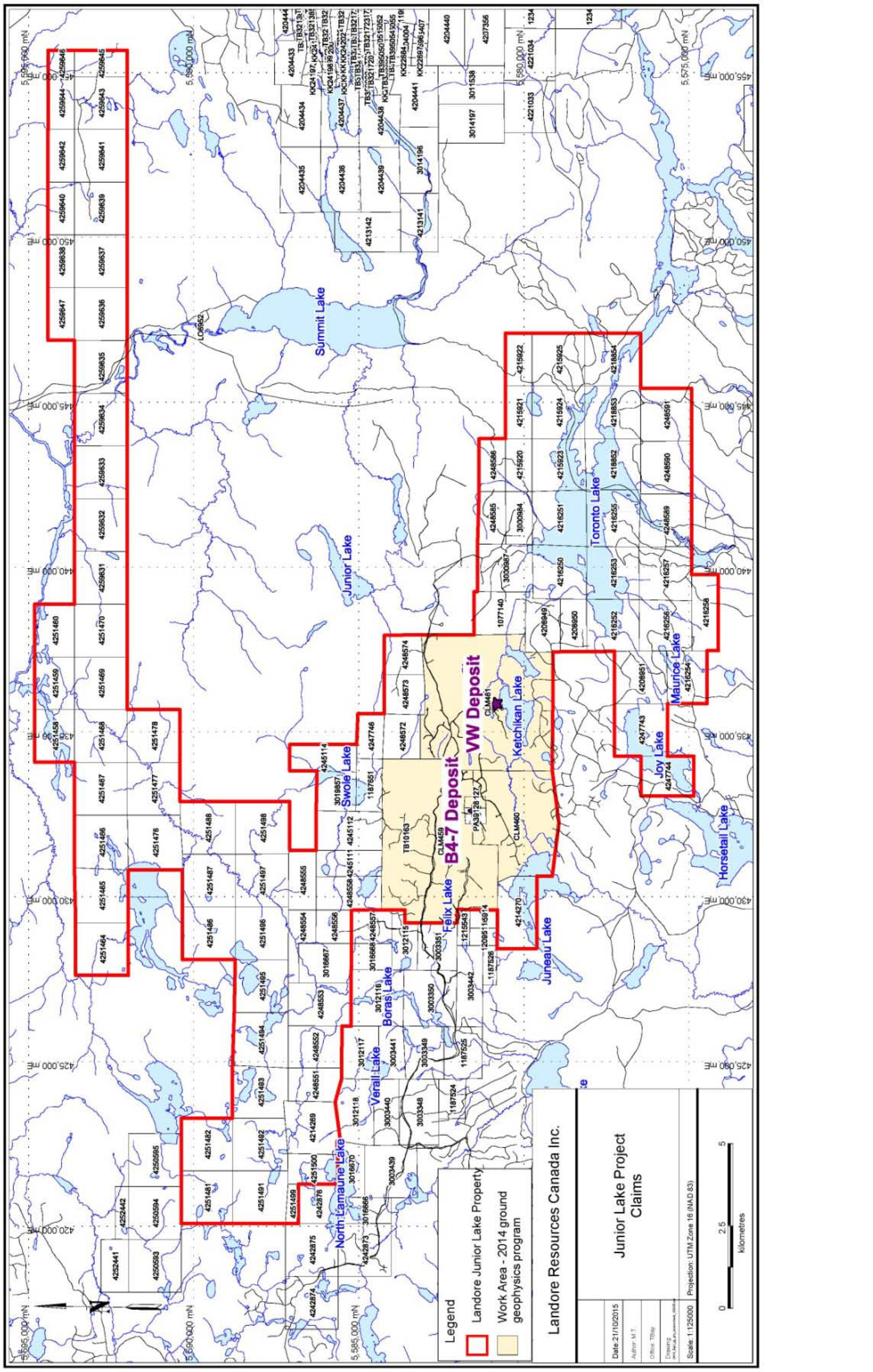


Table 3-1: Landore Mineral Claims (100% Interest)

Claim	Calculated Area (ha)	Units	Area	Claim	Calculated Area (ha)	Units	Area
1077140	201.533	9	Junior Lake	4251487	255.274	16	Junior Lake
1187651	126.417	8	Junior Lake	4251488	262.535	16	Junior Lake
3000984	129.049	8	Toronto Lake	4251491	263.875	16	Falcon Lake
3000987	241.242	14	Toronto Lake	4251492	268.399	16	Falcon Lake
3016667	191.05	12	Falcon Lake	4251493	262.351	16	Falcon Lake
3019857	143.257	9	Junior Lake	4251494	271.898	16	Falcon Lake
4208949	174.532	10	Toronto Lake	4251495	258.004	16	Falcon Lake
4208950	127.147	8	Toronto Lake	4251496	255.947	16	Junior Lake
4208951	252.102	16	Toronto Lake	4251497	257.811	16	Junior Lake
4215920	128.463	8	Toronto Lake	4251458	191.468	12	Kapikotongwa R.
4215921	128.45	8	Toronto Lake	4251459	191.47	12	Kapikotongwa R.
4215922	128.443	8	Willet Lake	4251460	191.467	12	Kapikotongwa R.
4215923	255.325	16	Toronto Lake	4251464	255.269	16	Junior Lake
4215924	255.568	16	Toronto Lake	4251465	262.916	16	Junior Lake
4215925	255.604	16	Willet Lake	4251466	262.901	16	Junior Lake
4216250	262.832	16	Toronto Lake	4251467	255.274	16	Junior Lake
4216251	272.55	16	Toronto Lake	4251468	255.272	16	Junior Lake
4216252	250.524	16	Toronto Lake	4251469	255.284	16	Junior Lake
4216253	252.657	16	Toronto Lake	4251470	255.289	16	Junior Lake
4216254	194.379	12	Toronto Lake	4251476	269.928	16	Junior Lake
4216255	277.21	16	Toronto Lake	4251477	255.274	16	Junior Lake
4216256	244.297	16	Toronto Lake	4251478	255.285	16	Junior Lake
4248585	132.61	8	Junior Lake	4251481	247.747	16	Falcon Lake
4248586	127.648	8	Junior Lake	4251498	262.503	16	Junior Lake
4248589	251.612	16	Toronto Lake	4251499	42.8057	3	Falcon Lake
4248590	246.935	16	Toronto Lake	4251500	135.24	8	Falcon Lake
4248591	246.981	16	Toronto Lake	4259631	255.3975	16	Junior Lake
4216257	259.137	16	Toronto Lake	4259632	255.2966	16	Junior Lake
4216258	184.68	12	Toronto Lake	4259633	255.4375	16	Junior Lake
4218852	257.424	16	Toronto Lake	4259634	254.9172	16	Junior Lake
4218853	269.098	16	Toronto Lake	4259635	255.9382	16	Summit Lake
4218854	269.053	16	Willet Lake	4259636	255.4302	16	Summit Lake
4245111	157.295	10	Junior Lake	4259637	255.5525	16	Summit Lake
4245112	236.437	15	Junior Lake	4259638	127.6468	8	Summit Lake
4245114	164.202	10	Junior Lake	4259639	255.2938	16	Summit Lake
4247743	254.86	16	Toronto Lake	4259640	127.657	8	Summit Lake
4247744	191.684	12	Toronto Lake	4259641	255.3059	16	Summit Lake
4247746	108.215	6	Junior Lake	4259642	127.6557	8	Summit Lake
4248551	160.761	10	Falcon Lake	4259643	255.3034	16	Summit Lake
4248552	206.752	12	Falcon Lake	4259644	127.1522	8	Summit Lake
4248553	248.178	15	Falcon Lake	4259645	127.654	8	Summit Lake
4248554	123.24	8	Junior Lake	4259646	63.16516	4	Summit Lake
4248555	151.516	9	Junior Lake	4259647	127.6481	8	Summit Lake
4248556	107.904	8	Junior Lake	4214269	249.927	16	Falcon Lake
4248558	200.808	10	Junior Lake	4214270	258.207	16	Toronto Lake
4248572	151.125	9	Junior Lake	**95**	**18,495.98**	**1,145**	
4248573	147.231	9	Junior Lake				
4248574	137.856	9	Junior Lake				
4251482	247.088	16	Falcon Lake				
4251486	255.278	16	Junior Lake				

Table 3-2: Landore Leases (100% Interest)

Lease #	Description	G-Number	Anniversary Date	Area (ha)	Annual Rent ($)	Expiry Date	Total Work in Reserve ($)
107421	PA 39127, 39128	4000476	98-Jan-01	52.969	158.91	2019-Jan-01	1,096,271
108257	CLM459[1]	4040218	08-Aug-01	1,460.795	4,382.39	2029-Aug-01	17,284
108258	CLM461[1]	4040217	08-Aug-01	1527.388	4,582.16	2029-Aug-01	2,468,109
108259	CLM460[1]	N/A[2]	08-Aug-01	687.794	2,063.38	2029-Aug-01	0
Totals	**4 Leases**			**3,728.946**	**11,186.84**		**3,581,664**

Notes:

1. Wing Resources holds a 2% NSR on 3 claims within CLM459, 1 claim within 460 and 3 claims within 461.
2. G-number is generated when work reports are filed.

Landore has been granted four mining leases, which include mining and surface rights, over an area encompassing the B4-7 and VW Deposits. The leases cover 23 mineral claims and two patents for a total area of 3,729 ha and have been granted for 21 years renewable for further terms of 21 years (Table 3-2).

Within the mining leases, Landore has the right, subject to provisions of certain Acts and reservations, to:

- sink shafts, excavations, etc., for mining purposes;
- construct dams, reservoirs, railways, etc., as needed; and
- erect buildings, machinery, furnaces, etc., as required, and treat ores.

There are no known environmental liabilities on the property.

4 ACCESSIBILITY

Access to the Junior Lake property from Thunder Bay is via paved provincial highways No. 17 (15 km) and No. 527 to Armstrong, with an overall distance of approximately 235 km. From Armstrong, the Buchanan Forest Products Inc. gravel haulage road (BHR) is taken east to kilometre 105, where a skidder haulage road leads approximately one kilometre to the Landore Junior Lake camp. Skidder and drill roads provide access on the property. The site of the 2014 DCIP + MT ground geophysics program is located in the central portion of the Junior Lake property, in the VW deposit and VW West areas.

There are no power lines or railway lines on the property; however, the main CNR line is approximately 13 kilometres to the south.

During the summer, most drill sites are accessible by 4-wheel-drive vehicles.

5 HISTORY

Routledge (2010) has summarized the exploration and development history of the Junior Lake property as:

Geological mapping and exploration in the vicinity of the Junior Lake property is recorded as early as 1917. In 1968, Canadian Dyno Mines Limited staked 333 claims in 15 groups to cover conductors picked from an airborne electromagnetic (EM) and magnetic (MAG) survey. Two groups, B3 and B4, included the Junior Lake property. The company merged with Mogul Mines Limited, and the successor, International Mogul Mines Limited, in joint venture with Coldstream Mines Limited, carried out prospecting, mapping, ground MAG and EM surveys, soil sampling, and trenching on the B3 and B4 claim groups. Eight diamond drill holes totaling 674.8 m (2,213.9 ft.) were drilled to test conductors in January 1969, resulting in the discovery of the B4-7 zone. The discovery hole, No. 69-5, intersected 8.26 m (27.1 ft.) of massive pyrrhotite-pyrite-chalcopyrite mineralization grading 0.80% Ni and 0.53% Cu. The B4-7 deposit was delineated by an additional 30 holes (6,850 m, or 22,479 ft.) in 1969. In the same campaign, eight holes for 628.2 m (2,061 ft.) explored other conductors on the property. A detailed MAG and EM survey was also completed over the deposit and petrographic work done on core at that time.

In late 1969, 136.1 kg (300 lbs) of drill core was composited from 71 assay rejects in 11 drill holes, split to 56.7 kg (125 lbs), and submitted to SGS for flotation recovery (metallurgical) testing, which included semi-quantitative spectrographic analysis for 30 elements. A manual tonnage/grade estimate for the B4-7 deposit was carried out, to total 2,282,520 tons (2,070,689 tonnes) averaging 0.87% Ni and 0.59% Cu (Zurowski, 1970). This historical estimate is not NI 43-101 compliant.

Coldstream Mines Limited acquired 100% of the property in 1970 and took two claims to lease in 1976.

In 1983-1986, Québec Cobalt and Exploration Limited staked part of the south portion of the Junior Lake property and carried out mapping, geophysics, and soil and rock sampling. Conwest Exploration Co. Ltd., the successor to Coldstream Mines Limited, optioned the leases covering the B4-7 deposit to Menacorp Limited in 1990, which resampled B4-7 core, and then to Minatco Exploration Ltd. in 1993.

In addition to the B4-7 deposit, exploration in the Junior Lake-Lamaune area prior to Landore work also revealed two low-grade Cu-Ni zones and occurrences of copper, iron, lithium, chrome, asbestos, zinc, and gold-molybdenite. Most of the occurrences are within two kilometres of the VW and B4-7 deposits.

Landore optioned part of the property from North Coldstream Mines Limited in 1998 and additional claims from Brancote Canada in 2000.

6 GEOLOGICAL SETTING

The regional, local and property geology has been for the most part summarized from Routledge, (2010), Lester (2009b), MacTavish (2004, 2004a), and Routledge (2006). Additional contributions are from various others, including Cooper (2009), Mungall (2009), and Pressacco (2013).

6.1 Regional Geology

The Junior Lake property is located within the Wabigoon Subprovince of the Superior Province of the Precambrian Shield and within the east-west trending Caribou-O'Sullivan greenstone belt. The belt is flanked to the south by the Robinson Lake Batholith of the Lamaune Batholithic Complex and to the north by a major, east-west-striking shear zone / terrain boundary that marks the southern limit of the English River Subprovince. Northeast of the property the belt is intruded by the elliptical, tonalitic to quartz dioritic Summit Lake Batholith. The western portion of the greenstone belt has been intruded by undulating, flat-lying, NeoProterozoic-age Nipigon diabase sills and localized dykes. These sills are the discontinuous, erosional remnants of laterally extensive sills comprising the Nipigon Plate which is centred on Lake Nipigon, approximately 30 kilometres to the south (MacTavish, 2004, 2004a). The regional geology of the Junior Lake property area is shown in Figure 6-1.

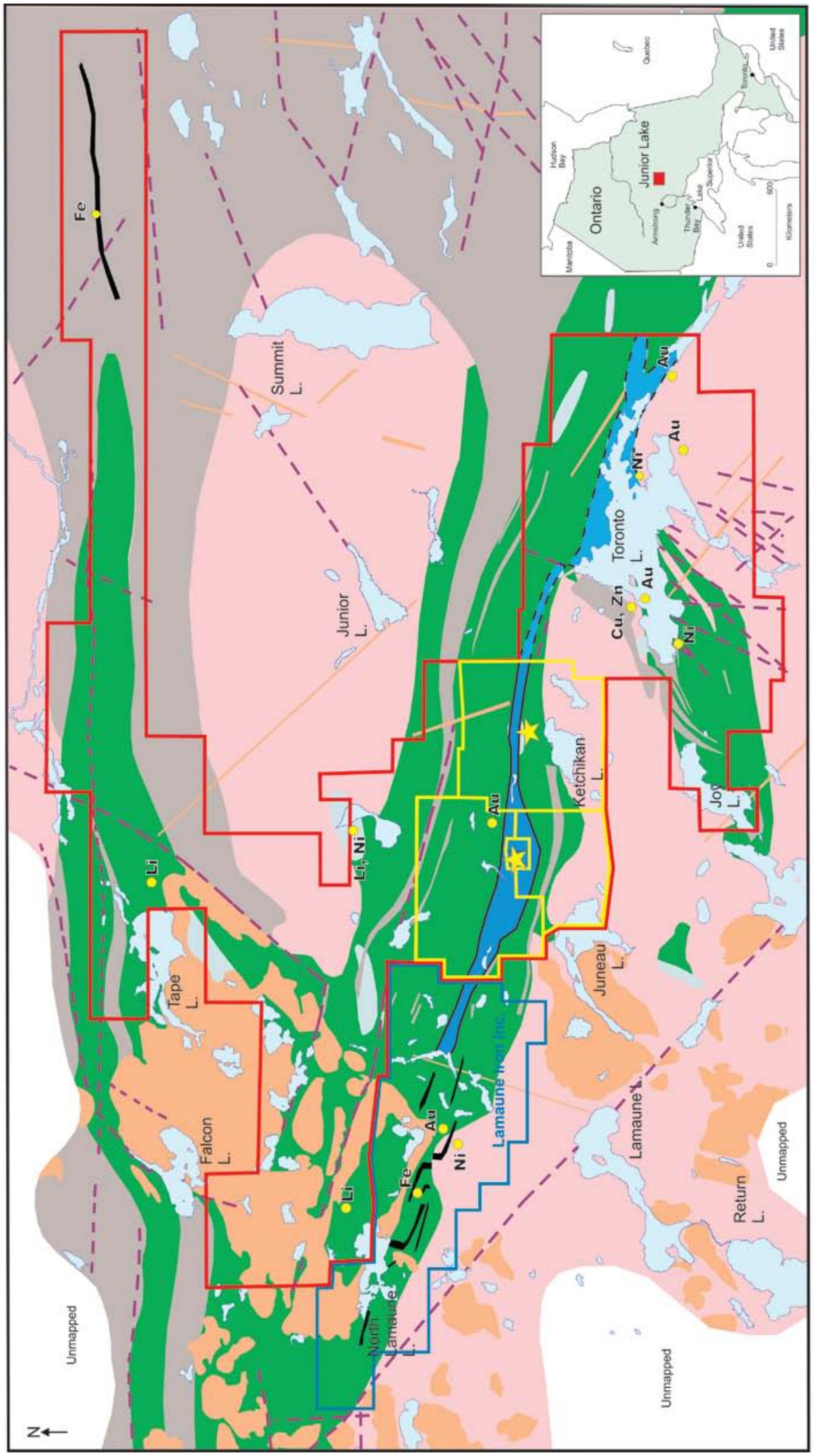


Figure 6-1: Junior Lake Regional Geology

6.2 Local and Property Geology

The supracrustal rocks and associated mafic to ultramafic intrusions of the Caribou-O'Sullivan greenstone belt are subdivided by Berger (1992) into the Archean-age Toronto and Marshall Lake groups. The two lithostratigraphic groups are similar in many respects; however, the Marshall Lake Group (MLG) contains a higher proportion of clastic metasedimentary rocks and apparently lesser amounts of mafic intrusive rocks.

The Toronto Lake Group (TLG) underlies the southern third of the Junior Lake property and consists of a bimodal assemblage of tholeiitic mafic flows and calc-alkaline rhyolitic to dacitic tuff, tuff breccias, and subordinate flows. The assemblage has been intruded by numerous mafic to ultramafic sills, dykes, and small stocks.

Four lithostratigraphic sequences defined within the TLG are as follows:

- The laterally extensive Carrot Top sequence trends west-northwest within the southern portions of the TLG and is comprised of magnetic talc-carbonatechlorite+/-tremolite schists derived from deformed and altered ultramafic rocks and clastic and chemical metasedimentary rocks. This sequence is 300 to more than 600 metres thick and hosts the D-Z iron occurrence, and several Ni-PGE (including Carrot Top and Zap Zone), Cu, Zn-Cu and Ag occurrences. Strong centimetre to metre scale folding is evident in the iron formation, and as such likely exists on a larger scale, possibly causing thickening and thinning along the main trends.
- The west-northwest trending Grassy Pond Sill intrudes the top of the TLG at its contact with the Marshall Lake Group (MLG) through the centre of the Junior Lake property. The Grassy Pond sill is a thick (100m to 500 metre wide), deformed, laterally continuous, gabbroic to locally anorthositic intrusive. The sill's most identifying characteristic is the presence of large (up to 10 cm in diameter) subhedral to euhedral plagioclase phenocrysts that often collect to form leucogabbro and anorthositic intervals of highly variable thicknesses. The Grassy Pond Sill hosts PGE, Cu and Ni occurrences, and is interpreted as being on the same geophysical structure as the B4-7 zone to the east.
- The B4-7 Sequence is a composite sequence, 1.9 kilometres long and up to 400 metres thick, of primarily mafic metavolcanic flows, intrusives and clastic and chemical metasediments that host the B4-7 Ni-Cu-Co-PGE deposit including the B4-7, Alpha and Beta Zones. The B4-7 sequence lies between the Carrot Top Sequence and the Grassy Pond Sill.
- The BAM Sequence is a composite sequence composed of mafic metavolcanic flows, mafic dykes and sills, and intermediate dykes. The BAM sequence is estimated to be 1.65 kilometres long and up to 160 m thick, possibly associated with an oblique structure. Archean Lamprophyre Dykes cut the TLG rocks.

In the north portions of the Junior Lake property, the Marshall Lake Group (MLG) includes tholeiitic, amphibolite mafic flows and calc-alkalic dacitic tuff, minor tuff breccias, and intercalated greywacke, chert and sulphide iron formation. Thin, discontinuous intermediate to felsic metavolcanic rock units also occur in the MLG. A higher portion of metasedimentary rocks and fewer mafic intrusives occur in the MLG compared to the TLG. Most of the rocks observed on the property are finely amphibolites, pillowed, mafic metavolcanic flows with well-defined pillow selvedge and a greater occurrence of plagioclase phenocrysts than observed within mafic flows south of the Grassy Pond Sill. Some outcrops exhibit an irregular, pervasive alteration, characterized by large, acicular actinolite porphyroblasts contained within a fine-grained matrix of chlorite, sericite, actinolite/tremolite, and epidote. This alteration is very similar to localized alteration observed within the Toronto Lake Group.

Pye (1968) interprets the presence of a large-scale fold on the western portion of the Junior Lake property southeast of Lamaune Lake and east-northeast-trending syncline in the vicinity of Toronto Lake to the east. The east-southeast trending, north-dipping North Lamaune Lake anticline is interpreted from magnetometer surveys tracing Iron Formation.

Structural Geology

Regional deformation rotated the supracrustal packages into near vertical orientation and developed a large west-northwest trending deformation zone (local portion referred to as the Junior Lake Shear Zone) north and west of Toronto Lake. This zone is the most prominent structural feature in the area and is characterized by narrow discrete zones of intensely sheared rock displaying dextral rotation separated by relative undeformed rock packages (Larouche, 1999). The deformation zone is evident as an aeromagnetic lineament which extends east and west of the Junior Lake property and appears to join the regional 450 km long Sydney Lake-Lake St. Joseph (SL-LSJ) Fault zone to the north, which also coincides with the boundary of the English River (ERT) and East Wabigoon subprovinces (EWT). The brittle-ductile fault zone of the SL-LSJ is steeply dipping, one to four kilometres wide, and is estimated to have accommodated about 30 km of right-lateral transcurrent displacement and 2.5 km of north vergent thrust movement (Percival, 2007).

A second, more local deformation in the east part of the property is confined to the supracrustal rocks around the periphery of the Robinson Lake Batholith, with deformation expressed as crenulation cleavage, northeast trending faults, and lineations which clearly post-date the regional deformation (Larouche, 1999).

Junior Lake Shear Zone and Associated Geology

Narrow, discrete zones of intense shearing (Junior Lake Shear Zone) form a corridor up to 800 m wide along the contact between the TLG and MLG. This shearing roughly follows the north contact of the Grassy Pond Sill. The evidence for the shear zone at Junior Lake is based on known geology and textures in drill holes and from limited exposures with deformation textures found from the micro to the macro level encompassing mylonites, cataclasites, sharp thin failure planes, and pressure-solution features such as stylolites. The widespread occurrence of pseudotachylite veinlets and infill demonstrates localized melting on failure planes.

Within the shear zone, the TLG is dominated by a large gabbro intrusive centred in the Grassy Pond Sill to VW area. It is a long linear intrusive and possibly split into several individual units. It is intruded into a mafic volcanic pile consisting of submarine pillow lavas and volcaniclastics. Cooper (2009) speculates that the gabbro has been the feeder for the volcanism and has then intruded its own lava pile.

Although the shear zone is slightly sinuous through Junior Lake, three of the mineral occurrences, Carrot Top, B4-7 and VW, fall on a straight line and Grassy Pond is only slightly to the north of this line. The length of the shear zone is uncertain, however, a length of at least 10 km has been defined. Along this length, there are variations in intensity with local domains of low deformation surrounded by high deformation zones as a result of competency contrast, general heterogeneity through the zone and lithology types. The rock succession in Junior Lake was deformed within a mobile greenstone belt and all geology became subvertical and with continued deformation within a deep ductile-regime, shear zones developed. During and post to shearing, gabbroic intrusive episodes occurred with a final pulse of very extensive vertical gabbro dikes. Major hydrothermal mineralizing events post-dated the gabbro dike swarm possibly as the result of heat from the post-tectonic sanukitoid style granites, such as high-Mg granitoid found in convergent margin settings (Cooper, 2009).

Less obvious at surface but no less voluminous are ultramafic lithologies such as peridotite, dunite, serpentinite, and their derivatives as talc dominated schistose metamorphic rocks. The ultramafic lava and/or intrusive suite was probably coeval with the basic suite but has suffered much more degradation of original texture and mineralogy within the mobile belt and shear environment. Variably textured granite and quartz diorite to tonalite gneiss and migmatite mapped along the south property boundary are part of the Robinson Lake Batholith.

Metamorphism

Metamorphism on the property is characterized by staurolite-cordierite-garnet, and rare sillimanite, in clastic metasediments; garnet-aluminosilicates-amphibole and rarely staurolite in the felsic and intermediate metavolcanic rocks; and garnet and amphibole in mafic meta-volcanic rocks. Most of the supracrustal rocks attained lower amphibolite grade metamorphic conditions, and greenschist grade metamorphism is only locally present (Larouche, 1999).

B4-7 DEPOSIT

The B4-7 Deposit is located in the south central area of the Junior Lake property. The B4-7 Deposit consists of polymetallic Ni-Cu-Co+PGE+Au mineralization hosted in massive sulphide (vein) and disseminated sulphides in a gabbro-basic volcanic setting coinciding with the Junior Lake shear zone. Strike length attains at least 600 m. Widths are up to approximately 18 m but are usually less than five metres. The B4-7 massive sulphide vein system appears to be a fairly simple dilational structure with marked pinch and swell in the vertical plane with an apparent plunge to the west. The mineralization was possibly introduced rapidly along pre-disposed failure planes under conditions of shearing. B4-7 consists of a continuous tabular body of semi-massive pyrrhotite-rich sulphides hosted in an assemblage of mafic volcanics and mafic intrusive. The contacts of the massive sulphide are typically very sharp and linear with minor wall rock contamination. Host rocks include leucogabbro, melanogabbro, gabbro, as well as mafic metavolcanics at the east end and may play an important role in hosting mineralization. Proximal to the sulphide mineralized zone are mafic schists, shear zones, metasedimentary rocks, locally iron formation, amphibolite, as well as pyroxenite, particularly at the east end.

The Alpha Zone occurs parallel to and approximately 50 m up dip of the B4-7 hanging wall in the gabbro and amphibolitized mafic volcanics. The zone, as described by MacTavish (2004a), is a broad envelope of narrow, moderate to high grade erratic, sulphide-rich carbonate and quartz carbonate veinlets and mineralized shears that generally trend oblique to stratigraphy. This hanging wall mineralization is localized in small fractures, foliations, and gashes and was probably introduced over a much longer time frame than the massive sulphides as the deformation progressed from foliation, through shear failure and into cataclasis. Disseminated sulphide mineralization consists of pyrrhotite, pyrite, and chalcopyrite, occurring as blebs and clots at the margin of gabbro intrusives against the host basic volcanics. Such mineralization is widespread but the better and more contiguous contact style mineralization occurs on the major gabbro contact. Thickness varies from <5 m to >15 m (Pressacco, 2013).

MacTavish (2004a) reported the ancillary Beta Zone, identified during 2001 drilling, as a net textured extension of the massive B4-7 zone and concluded that it was the probable magmatic source for the massive sulphides injected/emplaced into the fault/breccia zone.

VW DEPOSIT

The VW Deposit lies 50 m to 150 m north of Ketchikan Lake near the southeastern end of the Junior Lake claim group on claim TB1077142. The deposit consists of a series of five mineralized subzones hosted by deformed assemblage of sheared mafic and ultramafic metavolcanics, gabbros, and chemical and clastic metasedimentary rocks. Host rocks are mafic volcanics and, to a lesser extent, mafic intrusives and metasedimentary rocks. The subzones are contained in a 125 m to 200 m wide shear (Junior Lake Shear) that dips steeply north. The VW deposit itself has been drilled over a strike length of 620 m, and dips subvertically or steeply to the north, with the deepest intersection in the most southerly subzone at 320 m (22 m elevation).

Stratigraphy of the VW deposit consists of a mixed sequence of mafic volcanics (2A, 2A F2), ultramafic volcanics (1A, 1C), mineralized volcanics (MZ 2A and 2A MZ), and mineralized gabbro (9C MZ) and gabbro (9C). This sequence youngs from south to north, with rock units described by McKay (2006) as follows:

- **Mafic-pillowed volcanics (2A F2)** are fine- to coarse-grained, grey, greenish grey, with locally well-developed pillow selvages up to three centimetres wide. Elongated, pale green to beige porphyroblasts occur locally. Chlorite alteration is moderate. Biotite alteration occurs locally. Quartz veins up to 50 cm wide are scattered throughout. The unit is locally sheared. Red-brown garnets occur but are rare. Sulphides consist of minor disseminated pyrrhotite.

- **Ultramafic volcanic (1A)** is fine- to medium-grained, steel grey, soapy and massive to well foliated. Serpentinization is pervasive. Chlorite, fuchsite and epidote alterations are moderate. White, locally beige, quartz veins up to 10 cm across are scattered throughout the unit.

- **Mafic volcanics (2A)** are fine- to coarse-grained, grey, greenish grey, foliated to massive. Chlorite and biotite alterations are moderate. Silicification is weak and patchy. Quartz veins up to seven centimetres across are scattered throughout. Sulphides, up to 1% pyrite and 0.5% chalcopyrite, occur predominantly in shear zones (?) or possibly poorly defined pillow selvages.

- **Mineralized mafic volcanics (2A MZ and MZ 2A)** are fine- to coarse-grained, grey, greenish grey, locally foliated to massive. Chlorite and biotite alterations are moderate to strong and epidote alteration is locally weak. Quartz veins are scattered throughout. The unit is sheared and fractured throughout. A 30 m intersection in hole 0405-36 contains poorly defined, very fine grained, cherty magnetite iron formation. The magnetite banding is up to 30 cm thick and rarely has sharp contacts with the hosting volcanics. The sulphide content of this magnetite iron formation is up to 15% pyrrhotite, 3% chalcopyrite, and 1% pyrite. The overall sulphide content of the unit is up to 5% pyrrhotite, 4% pyrite, and 2% chalcopyrite. In addition to the banding in the iron formation, the sulphides also occur as fracture filling, scattered blebs associated with quartz veins, and as disseminations.

- **Mineralized gabbro (9C MZ)** at the collar of hole 0405-35 is fine- to coarse-grained, grey, greenish grey and well foliated to massive. Biotite alteration and silicification are moderate to strong. Amphibole patches occur throughout. Moderate chlorite alteration occurs locally. Creamy white and beige quartz veining occurs throughout the unit. The unit is locally sheared, fractured and faulted, with minor gouge. Disseminated and blebby sulphides, up to 3% pyrite and 2% pyrrhotite, locally enhance the foliation.

VW WEST

The VW West area is located immediately adjacent to the VW nickel deposit, and extends for approximately 1.5 kilometres towards the B4-7 Ni-Cu-Co-PGE deposit. Drilling activities in this area during 2010 and 2011 focused on a prominent east-west trending geophysical anomaly and several coincident EM conductors. Five fence lines were drilled at 200 metre intervals across a 1.5 km portion of the VW West area to test the potential for a westward strike extension of the VW deposit.

The area lithology resembles that of the VW deposit itself, dominated by sequences of ultramafic rocks with ultramafic-mafic flows. Additional units include metasediments and iron formation.

Lithology:

Cross section interpretations for VW West indicate a similar lithological and laterally correlative succession as the VW nickel deposit; with a series of upturned (near vertical) interbedded and fault bounded sequences of north facing ultramafic (peridotite) (as slices and larger lithon-like masses), juxtaposed(?) or as a stratigraphic succession against a suite of thinly interbedded ultramafic-mafic flows (locally biotitic or amphibolized) hosting pelitic and metavolcanic horizons and local sulphidic iron formation chemical sediments (chert). The stratigraphy is cut by gabbro and leucogabbro intrusions (locally porphyritic to anorthositic), and the latest gabbro intruded as very fine grained (mafic) dikes often emplaced as swarms. The stratigraphic sequence appears to repeat itself to the south of the magnetic high; but limited drilling tested this area. In the entire lithologic suite of VW West one particularly important unit of the VW deposit the 2Af1 “grey rock’ (Cooper, 2011) was not noted in any appreciable abundance in the logging of drill core at VW West.

Structure and Alteration:

Numerous shear zones transect the drilled area sub parallel to the strike of the stratigraphy and in essence the entire width of much of the stratigraphy is highly strained to ductile deformed slices of varying but laterally correlative lithologic units. The well-bedded and hydrous nature of both the metavolcanic/pelite/BIF and olivine-rich peridotite act as slip planes in the stratigraphic stack and the deformation appears to be focused into these units deflected by the competency contrast of ridged lithons of either massive peridotite or massive gabbroic intrusions or dike swarms. Most strained zones are frequently accompanied by quartz-carbonate veining and alteration. The peridotite when thin or sandwiched between mafic and or pelitic or BIF is often intensely altered to carb-rich talc-schist.

Ophicarbonate is common in the upper part of the stratigraphic sequences (roughly coinciding with the northern limit of the magnetic high) and is hosted in highly strained talc-schist unit. The ophicarbonate horizon is of ultramafic aspect and is the result of carbonatisation of ultrabasic rocks and is directly associated with listvenites, [listwaenites]. These rocks are frequently associated with terrain boundaries and are a common feature in shear zones along terrain boundaries (Cooper, 2009). Pseudotachylite was observed in the upper portion of drill hole 0410-285 indicating fault movements involving mylonitization and/or partial melting. An approximate 30m true thickness of the structural zone hosting the ophicarbonates is defined in an east-west linear trend across the drill area.

Overall the stratigraphy is fairly linear along strike however a discontinuity is apparent in the magnetic data at line 1700E on the local grid. This discontinuity is apparent in the drilling as a pinching out of the massive peridotite and greater abundance of mafic flow along this section. Perhaps there is a NE-SW regional fault crossing the area here as well.

7 MINERALIZATION

7.1 B4-7 Deposit – Nickel, Copper, Cobalt, PGE, Gold

B4-7 deposit mineralization has been summarized from Pressacco (2013), with contributing information from MacTavish (2004).

Mineralization in the B4-7 zone is composed of semi-massive to massive sulphides primarily composed of pyrrhotite (25%-75%), with minor pentlandite and chalcopyrite (1%-5%). Other sulphides include pyrite, violarite, sphalerite, and covellite. Magnetite concentrations are minor (1%-5%) to moderate (5%-25%) and generally occur where oxide iron formation is proximal to the zone walls.

The pyrrhotite is fine grained and carries pentlandite and violarite exsolved as very fine grains (20 microns to 40 microns). Pentlandite occurs as relatively coarse grains as well as fine pentlandite flames. Nickel contents of up to 3% occur in the B4-7 lenses, whereas in the disseminated sulphides the nickel content is usually less than 1%.

The disseminated style of blebby and network style mineralization is, however, generally much richer in PGEs, with up to 10 g/t Pd and up to 1 g/t Pt. Both Cu and Co occur as sulphides in both styles of mineralization. PGE mineral species are represented by kotulskite and borovskite, both palladium tellurides.

MacTavish (2004) described the B4-7 zone mineralization as composed of massive, fine-grained pyrrhotite, highly variable disseminated, stringer, locally semi-massive, fine-grained chalcopyrite, and localized coarse-grained secondary pyrite. Most of the chalcopyrite is concentrated as fine to medium grains near contacts, within fragment-rich portions of the zone, or as veins and stringers within the adjacent hanging wall and footwall rocks. Pyrite occurs as coarse cubes, up to three centimetres in diameter, that usually concentrate near the upper and lower contacts of the massive sulphide zone. Cobalt content correlates directly with the amount of coarse-grained pyrite. The B4-7 zone is locally rich in disseminated magnetite, which, in most cases is correlatable to the presence of oxide facies iron formation in the adjacent hanging wall or footwall rocks.

7.2 VW Deposit – Nickel, Copper

VW deposit mineralization has been summarized by Routledge, R.E. (2010b):

There are three styles of sulphide mineralization in the VW deposit. The most important is thin lamina and veinlets following the foliation of the volcaniclastic rocks. The hydrothermal fluids appear to have been constrained by the impermeable gabbro dikes in that the highest concentrations of sulphides and grades are found immediately adjacent to dike contacts with a gradual diminution away from them.

The pyrrhotite, pyrite and chalcopyrite mineralized volcaniclastic host unit itself is usually thick, up to 25 m, but of lower overall grade than the mineralization ponded and channelled along the dikes. Mineralization in the centre of the Katrina zone may be of magmatic origin.

The third style is low grade mineralization up to 0.4% Ni occurring as fine blebs of pyrrhotite and pentlandite within ultramafic schist, peridotite and serpentinite.

The subzones within the VW deposit are composed generally of 1% to 5% sulphides consisting of pyrrhotite-pyrite-pentlandite-chalcopyrite-magnetite ± sphalerite. Locally sulphides can reach 40-50%, however, no massive mineralization has been noted. In addition to blebs/clots, lamina on foliation planes and veinlets, sulphides also occur to a lesser extent as breccia matrix, as replacement style net texture and as fracture filling. Weak to moderate sulphide mineralization at less than 2% sulphides and stronger mineralization containing more than 2% sulphides was distinguished in logging the mineralized zones.

Pyrrhotite is fine grained and carries minor pentlandite exsolved as very fine flames as well as occluded pentlandite as discrete fine grains. Free pentlandite appears to be rare. Chalcopyrite and pyrite occur as fine to medium grains.

7.3 VW West – Nickel, Copper, PGE, Chromium

Sulphide mineralization consisting of pyrite, pyrrhotite and minor chalcopyrite was observed in the gabbro, mafic volcanic and ultramafic units. These sulphides are disseminated through the host rock in relatively low abundances; trace to 2%, but in various intervals reached as high as 25% as stringers, bands, and veinlet infill semi massive to massive and network with po the dominant sulphide typically. There were few instances of centimeter-scale high abundances (>50%) of pyrite and pyrrhotite (massive sulphide) but the dominant style of mineralization is as stringers and wisps. The highest potential for Ni- bearing sulphide mineralization is that hosted in the mafic volcanic and peridotite units.

Up to 25% magnetite is disseminated through the mafic volcanic and ultramafic sequences, a contributing factor to the magnitude of the east-west trending geophysical anomaly through the area. Locally the magnetite has been shown to replace the chromium spinel and is indicative of progressive metamorphic intensity (Barnes, 1998). The degree of replacement of chromite by magnetite can be used as indicator of metamorphic grade. Locally the greenschist facies appears conducive to low grade Ni-Cr mineralization hosted in massive peridotite as reported from 0411-285, 0411-289, 0411-294, 0411-297 indicating continuity along strike.

Petrography of the mineralized peridotite confirms the presence of antigorite and chrysotile (fibrous asbestos) as up to 30% of the sample composition in the mineralized zones. Important associated sulphides include millerite, chromite, and magnetite after chromite (Kjarsgaard, 2011). A few rare species of supergene Ni-sulphides were noted and serve as an indication of late Ni enrichment. Results from drilling indicate Ni-Cr mineralization in the lower and least altered portions of the peridotite unit with up to 0.499% Cr over 1m (0411-289); and 0.18% Cr over 31.5m reporting 147.50m in 0411-297.

7.4 Mineralization Elsewhere on the Property

Prior to Landore ownership, exploration in the Junior Lake–Lamaune Lake area that located the B4-7 deposit in 1969 also revealed two low-grade Cu-Ni zones and occurrences of copper, iron, lithium, chromite, asbestos, zinc, and gold-molybdenite. Most of these are within two kilometres of the VW Zone.

From 1990 to 2003, Landore found nine PGE-Cu-Ni occurrences, one Cu-Pd zone, one gold zone, and Zn-Au-Ag and Zn-Co occurrences in old trenches and boulders bearing base and precious metal or arsenic mineralization. The VW deposit was discovered in 2005.

Four lithostratigraphic sequences favourable for nickel mineralization on the Junior Lake property have been identified by MacTavish (2004b) as follows:

- VW Sequence: a 1.9 km long, up to 400 m thick package of mafic metavolcanic flows, mafic intrusive dikes and sills, and clastic and chemical metasedimentary rocks that host the VW Zone.

- B4-7 Sequence: 1.9 km long and up to 400 m thick, is composed of primarily mafic metavolcanic flows (2AF1), gabbroic intrusive (9A,B,C), and clastic and chemical metasediments (6P) that lies between the Carrot Top Sequence and the Grassy Pond Sill. This sequence hosts the B4-7 Ni-Cu-Co-PGE deposit including the B4-7 massive sulphide zone and the Alpha and Beta zones.

- Grassy Pond Sill, a laterally extensive 100 m to >500 m thick gabbroic sill that hosts Cu-Ni-PGE mineralization near its base.

- Carrot Top Sequence: a complex laterally extensive 300 m to >600 m thick sequence of mafic metavolcanic flows, ultramafic schists, and clastic and chemical metasedimentary rocks that host several Ni-Cu-PGE occurrences. This sequence is located in the west portion of the Junior Lake property.

- BAM Sequence: a 1.65 km long, up to 165 m wide assemblage composed of mafic metavolcanic flows, mafic dikes and sills, and intermediate dikes that host the BAM gold occurrence. The BAM sequence is located northwest of the VW deposit in the north central portion of the Junior Lake property.

8 EXPLORATION

Cheatle (2010a) outlined the exploration history of the Junior Lake property:

Landore optioned part of the property from North Coldstream Mines Limited in 1998 and additional claims from Brancote Canada in 2000. Since then, Landore exploration has found nine PGE-Cu-Ni occurrences, one Cu-Pd zone, one gold zone, and Zn-Au-Ag and Zn-Co occurrences in old trenches and boulders bearing base and precious metals or arsenic mineralization. Landore has successfully delineated several deposits and other potential areas of significant mineralization throughout the Junior Lake property including two Ni+PGE deposits (B4-7 and VW).

Landore initial work in 2000 involved data compilation, Landsat image interpretation, prospecting, mapping, and resampling of the 1969 core, and followed up an Ontario Geological Survey (OGS) airborne EM and MAG survey flown over the area.

Ground magnetometer MaxMin II EM surveys, in addition to drilling, were completed in 2001. In 2003, Landore conducted drilling, stripping, trenching and channel sampling. All drilling data were digitized and reinterpreted, 856 core samples were assayed to fill in unsampled runs in the B4-7 deposit, in its hanging wall mineralization known as the Alpha Zone as well as in mineralization in the east extension of the B4-7 zone known as the Beta Zone.

A low level helicopter AeroTEM time-domain electromagnetic and magnetometer survey was flown in 2004. Principal geophysical sensors utilized in this survey included AeroQuest's AeroTEM© time domain helicopter electromagnetic system and a high sensitivity cesium vapour magnetometer. Bedrock EM anomalies were interpreted and graded according to the conductance.

The VW deposit was discovered in 2005 by follow-up prospecting of an AeroTEM conductor where 0.45% Ni was returned in a surface grab sample. Landore subsequently drilled the new VW deposit, as well as the Whale, NO and BAM zones, and other areas on the Junior Lake and Lamaune projects.

In 2006, Landore drilled the VW deposit, B4-7 zone, and other exploration targets including the Junior Lake, Pichette, and Lamaune claims. The 2006 campaign at the VW deposit included two surface trenches which were excavated and channel sampled. Metallurgical work included preliminary flotation and work indexes were carried out at Lakefield in September–October. Scott Wilson RPA also prepared a technical report (NI 43-101) on the B4-7 zone in 2006.

During 2007, diamond drilling of the VW and B4-7 deposits was the main focus of exploration activity. The following work was completed on the Landore property:

- Relogging of pre-2007 VW deposit drill core was initiated.
- Drill collars of the VW and B4-7 deposits and topographic control areas of the Junior Lake property were surveyed by an Ontario Land Surveyor.

- Minor line cutting was completed near Ketchikan Lake and the B4-7 deposit area to support the drilling operations.

- Baseline environmental studies were initiated and conducted by or under the guidance of Golder Associates Ltd. (Golder), of Sudbury, Ontario:
  - These studies were started in March 2007 and include quarterly sampling and analysis of lake and stream waters
  - Lake and stream sediment sampling was completed during the summer.
  - A benthic study, bathymetric study, and a fisheries study of Ketchikan Lake were completed.

- A weather station was installed at the Landore Junior Lake camp to record wind speed and direction, temperatures and three seasons of precipitation data.

- Sampling of the VW deposit drill core (quarter-cut core) was completed for metallurgical purposes.

- Claim lines were rehabilitated and the claim boundary surrounding an area to be leased was cut and surveyed in advance of filing the application to the Mining Recorder to lease the claims. Four leases have subsequently been granted.

- The land package was expanded to the southeast by staking an additional 24 claims totalling 5,056 ha.

- Aerial photography (stereo) was completed over the lease area by KBM Forestry Consulting in late 2007 to produce an air photo mosaic for exploration and infrastructure planning. The photographic data were processed to establish a detailed digital terrain topographic model (DTM).

- Golder commenced baseline aquatic studies in February 2007 on lakes and drainage tributaries in the vicinity of Junior Lake. These studies, repeated three monthly, are proceeding well and will continue through to economic studies. In addition, Golder completed a "Fish community and Fish habitat" survey of Ketchikan Lake, immediately south of the VW deposit, in addition to a bedrock resistivity survey on the northern side of the lake to determine depth of silt and evaluate bedrock competence.

- The camp was expanded and core storage was improved to hold the Junior Lake drill core on site.

- Core from previous Landore drilling in the VW deposit was relogged with a view to better understanding the controls on mineralization and identifying the disposition of mafic intrusives (dikes and sills) in the zone. In addition, further petrographic investigation was carried out on the VW deposit (Mungall, 2007). The drill hole collars were resurveyed to the Ontario base.

- In early 2007, a resource estimate was carried out by Scott Wilson RPA on the VW deposit.

In May 2008, Scott Wilson RPA prepared an updated resource estimate and NI 43-101 compliant technical report for the VW deposit. Scott Wilson RPA updated the VW deposit estimated resources to reflect 2008 to 2009 drilling and prepared a separate NI 43-101 compliant technical report (Routledge and Scott, 2009).

A non-NI 43-101 compliant mineral resource estimate to JORC standards was carried out by the Snowden Group (Snowden) on the B4-7 deposit in 2008. Scott Wilson RPA prepared resource estimates for the B4-7 deposit in 2006 and 2009.

Exploration efforts in 2009 included drilling, mapping and prospecting throughout the contiguous claims covering approximately 10 km^2, with work concentrated in the Lamaune Iron, BAM and VW areas. Additional exploration completed included prospecting and mapping at Swole Lake and Toronto Lake as well as east and west of the VW deposit.

To 2009, the VW deposit has been delineated and tested by 141 drill holes with 2,766 analyzed intervals over 2,838.36 m completed in the deposit subzones. Scott Wilson RPA has updated the VW deposit estimated resources to reflect 2008 to 2009 drilling and has prepared a separate NI 43-101 compliant technical report (Routledge and Scott, 2009).

Other exploration efforts in 2009 included mapping and prospecting throughout the contiguous claims covering approximately 10 km^2, with work concentrated in the Lamaune Iron, BAM and VW areas. Additional exploration completed included prospecting and mapping at Swole Lake and Toronto Lake as well as east and west of the VW deposit.

Overview of Recent Exploration

Recent exploration activity at Junior Lake from 2006 to 2013 has seen drilling focused on several areas including additional resource drilling at VW and B4-7 deposits, Lamaune area exploration drilling, the Whale Zone, Felix Lake, Swole Lake and B4-7 West) exploration drilling.

Other recent work, in 2007-2013, included detailed geologic mapping (B4-7, VW, BAM, Lamaune), 55 trenches over approximately 13km (Lamaune Iron, Grassy Pond, Felix Lake, Juno Lake, BAM Zone, Toronto Lake), additional geophysical work (impulse EM survey, ground magnetic, and reinterpretation and integration with historic magnetic data), as well as approximately 70 km of line cutting. Regional scale prospecting, regional reconnaissance and geologic mapping, including an airborne geophysical coverage (AeroTEM electromagnetic and magnetic) of the Toronto Lake area (various Ni, Au, PGE potential), and Swole Lake (pegmatite lithium) prospecting were also undertaken. Numerous consultant reviews and studies have been completed, including detailed Scanning Electron Microscope (SEM) and petrography studies of the VW and B4-7 deposits; relogging, resampling and reinterpretation of geology for the VW, B47, and BAM sites; as well as reviewing of regional exploration potential. Surveying of drill collars, claim lines, additional claim staking, initiation of environmental baseline study, aerial photography, and metallurgical testing were also undertaken.

In June 2011, the Lamaune block, comprised of 23 claims, for 4,096 hectares, containing the Lamaune Iron deposit as well as the Lamaune Gold prospect, was transferred into a separate private company ('Lamaune Iron Inc.').

In October 2012, a deep penetrating ORION 3D 'Direct Current Induced Polarization' (DCIP) and Magnetotellurics (MT) survey was performed over the Scorpion zone of the Junior Lake property by Quantec Geoscience Ltd. This survey encompassed the western portion of the Scorpion zone, from line 1400W eastwards to line 400E in the B4-7 deposit.

Tuomi (2013) describes the 2012 DCIP+MT survey:

This survey acquired three sets of data in multi-directions; DC (direct current), IP (induced polarization) and MT (magnetotellurics), and is a true three dimensional survey. Sophisticated digital signal processing was utilized to obtain high resolution imaging at depths up to 1000+ metres below surface. This survey utilized DC resistivity to identify prospective nickel mineralization, and used IP chargeability to investigate potential copper and PGE targets.

The survey identified three areas of interest, located in the central, eastern and northern parts of the survey area, which appear to be interconnected and geologically controlled by fault lines. A portion of the eastern survey area is drill tested and hosts the B4-7 deposit.

The DCIP + MT survey results indicate that the conductive horizon which harbours the B4-7 massive sulphide mineralization extends to the west through the Exploration Target, an area identified west along strike and down dip from the B4-7 resource containing a potential 1.5 Mt to 2.0 Mt of sulphide mineralization of similar grade range to that which has been outlined to-date (Pressacco, 2013).

Additionally, DC resistivity results have identified potential sulphide mineralization along a 1.53 km wide corridor encompassing the western Scorpion zone. IP chargeability and MT results further support the potential of the Scorpion zone area for massive and disseminated sulphide mineralization.

Subsequent drilling in winter 2013 has tested the DC resistivity and IP chargeability results at various localities along the western portion of the Scorpion zone. Drilling in the Exploration Target area between lines 175W and 300W successfully intersected B4-7 massive sulphide mineralization as well as Alpha zone disseminated sulphide mineralization.

8.1 2014 3-Dimensional (3D) Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) survey

From January 17 to February 14 2014, a 3-Dimensional (3D) Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) ground geophysics program was completed over the VW deposit, VW West, and B4-7 East areas, from line 300E to line 3700E and from 700N to 1500S. The survey, covering 739.02 hectares, was conducted by Quantec Geoscience Ltd., Toronto, for Landore Resources. This survey is located directly adjacent to the east of Landore Resources' 2012 Orion 3D DCIP + MT survey block covering the B4-7 West zone.

The DC resistivity displays a good correlation with the B4-7 deposit while the IP chargeability model is best correlated with the VW deposit. Generally, across the survey grid the areas of the highest chargeability values offset the locations of the highest conductivities. This observation is consistent throughout the entire IP chargeability model when compared with the DC resistivity model. Considering the fact that the B4-7 mineralized zone is characterized as conductive target and the VW mineralized zone is characterized as chargeable target, it manifests that these two deposits are of different sources and depositional processes and should be interpreted accordingly. Therefore, for follow-up exploration program it is important to incorporate geological conceptual model for a given deposit type and use the corresponding DC or IP models as guideline for drill targeting (Gharibi/McGill, 2014).

Results from the 2014 survey have been highly encouraging, delineating nine significant new zones ranging from approximately 400m to 1,200m in length of potential nickel sulphide mineralisation along strike and adjacent to the existing B4-7 Nickel-Copper-Cobalt-PGEs resource and the VW Nickel resource. Numerous targets are at depth and below existing Landore exploration drilling.

Table 8-1 describes priority targets as identified by Quantec.

Table 8-1: Priority Zones of Interest, 2014 Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) Survey

ID	Priority	Description
DC1/IP2	High	Eastward extension of B4-7 deposit. It correlates with anomaly DC1 south of the baseline (L50S and L100S) at depth. The extension shows a drop of more than 100m in depth in this area.
IP1	High	Westward extension of VW deposit. The VW deposit is correlated with anomaly IP1. The westward extension of the chargeability anomaly IP1 could indicate extension of the VW deposit.
IP7/DC7	High	Chargeable anomaly IP7 and conductive anomaly DC7 in the south centre of the survey grid. This area displays similar characteristics to the VW deposit with conductive and chargeable zones slightly offset.
DC1/IP4	Mid	Zone DC1 is a part of a large elongated conductive body. It corresponds with the location of chargeable anomaly IP4. Understanding the source of these anomalies could enhance understanding of the structures and mineralization in this area.
MT1/MT4	Low	Large and deep MT1/MT4 conductive area dipping northward in the northwest part of the grid. This zone determines the source of large elongated conductive body (DC1-DC4 and MT1-MT2) that is resolved across the survey grid.

Source: Gharibi/McGill (2014)

In the survey report, Gharibi/McGill (2014) notes:

The resistivity varies sharply over short lateral distance, indicating a complex subsurface dominated by faults and dike-like structures (Figure 8-1). The conceptual fault system drawn using the resistivity map at 50 m elevation indicates that these deformations are likely the result of a number of tectonic episodes, dominated by an E-W stress field system.

A number of high conductivity areas are resolved across the survey grid; anomalies DC1 through DC9 (Figure 8-2). The conductive anomalies display a general E-W elongated trend. Main conductivity feature, which is observed across the survey grid, is a large conductive body comprising anomalies DC1, DC2, DC3, and DC4 (Figure 8-2). The large conductive body appears to construct a large E-W elongated conductive structure and shows a number of lateral displacements in N-S direction. Displacements are evident along grid line L1300E between DC1 and DC2 and grid line L2000E between DC2 and DC3 and grid line L2800E between DC3 and DC4.

Additionally, resistivity cross-sections indicate that the B4-7 mineralized zone partially correlates with the anomaly DC1 to the south of the baseline (L0N). Existing drill holes in this area do not probe the possible deep extension of B4-7 to the east, south of the baseline L0N. The anomaly DC1 shows a drop of more than 100 m in depth in this area. This behavior is similar to a more than 100 m drop of the conductive anomaly associated with the B4-7 in the west, as observed in the 2012 Orion 3D survey over the same deposit. This suggests that the B4-7 deposit may be an uplifted conductive feature along a very long conductor with an E-W orientation, which is observed across the previous and present Orion survey areas. Therefore, the deep extension of anomaly DC1 south of gridline L0N resolved in the present survey is a potential target of interest to extend the existing B4-7 deposit eastward.

Conductive anomaly DC6 is resolved south of the VW mineralized zone. The relative locations of the VW zone and anomaly DC6 suggest that the VW deposit is a contact mineralization zone (Figure 8-2). The VW deposit, however, is well correlated with elevated chargeability zone IP1 in the near surface (Figure 8-3). The chargeable anomaly IP1 appears to dip northward with depth. Also, anomaly IP1 extends westward at shallow depths and merges with chargeable anomaly IP3. The extension of chargeability anomaly IP1 could indicate extension of the VW deposit westward, which makes this area a potential target.

A number of large chargeability anomalies are resolved in the west and northwest of the survey grid (Figure 8-3, anomalies IP4, IP5, and IP6). At greater depth these three anomalies merge together and construct a large chargeable body with chargeability as high as 85 milliradians. The source of this large chargeable body is unknown but the scale and geometry of the anomaly suggest that it could be structural and formational IP responses.

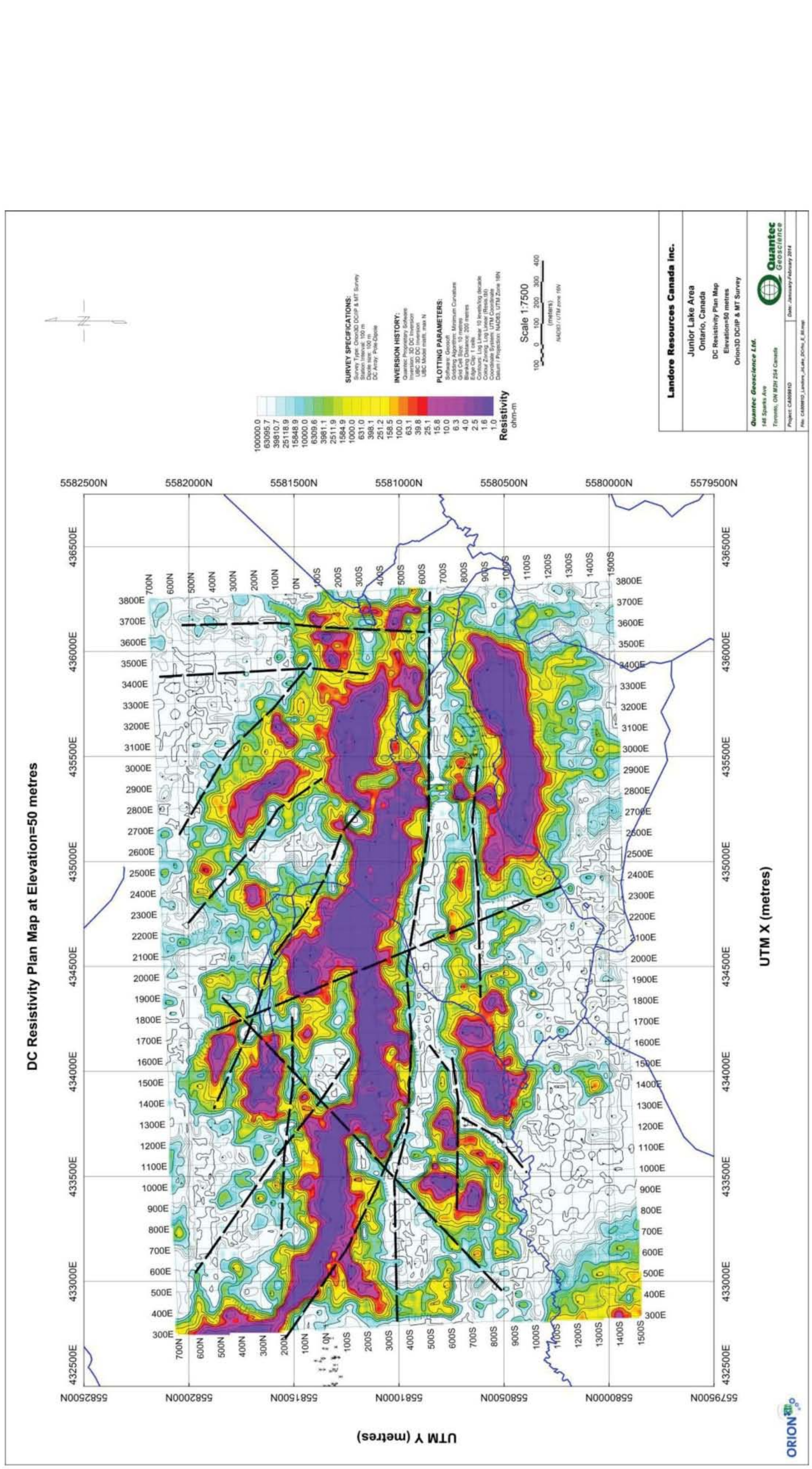
Anomaly IP2 appears to correspond with the B4-7 deposit (Figure 8-3). This anomalous zone displays connectivity with large chargeability anomalies in the east.

In the southwestern part of the survey grid complex chargeable anomalous zone is observed (Figure 8-3, IP7). This chargeable anomaly partially corresponds with and offset elevated conductive area DC7 and DC8.

In MT resistivity model, three large and connected anomalous conductivity zones are observed across the survey grid. These conductive anomalies construct a very large elongated conductive body in a nearly E-W direction with a resistivity as low as 1 Ω·m. The characteristic of this conductive body is in agreement with the conductive body resolved in the DC model. The large conductive body extends to great depths with decrease in conductivity level and lateral extents. At elevation below -400 m the conductive body is limited to the northwest part of the survey grid with a resistivity of a few hundred Ω·m. Additionally, the large conductive body dips slightly northward with depth and its lateral extents reduce to the northwest part of the survey area (Figure 8-5). This characteristic was also observed in previous Orion 3D survey in 2012 where a large conductive body resolved in MT model shows a north dipping behaviour.

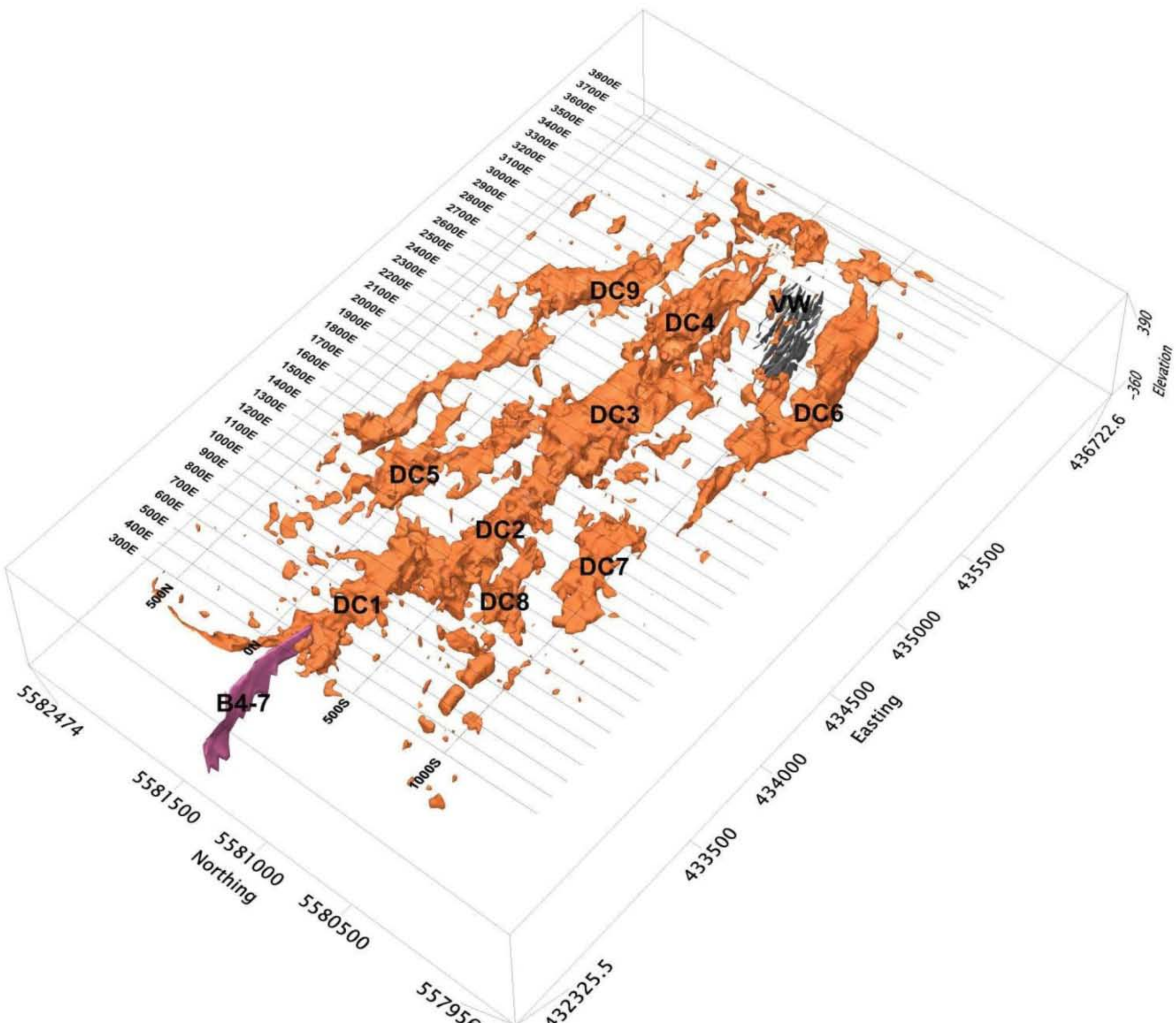
Further exploration work is required to ascertain scope and continuity of this mineralization and test prospective geophysical targets.

Figure 8-1: DC resistivity plan at 50m elevation



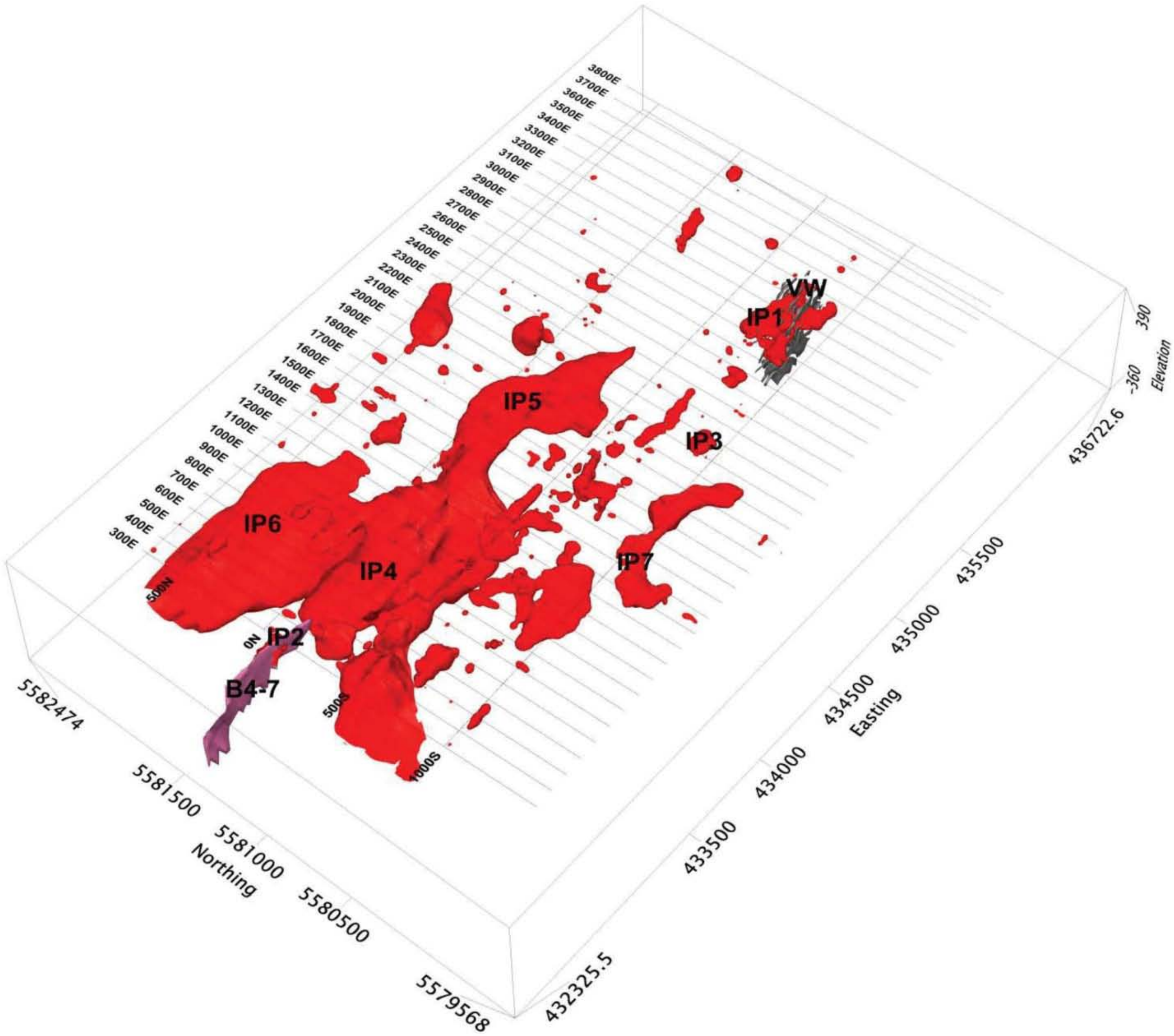
DC resistivity plan map at 50 m elevation with conseptual faults and structures.

Figure 8-2: 2 ohm-m DC Resistivity Isosurface



2 Ω·m DC Resistivity isosurface.

Figure 8-3: 70 milliradian Chargeability Isosurface

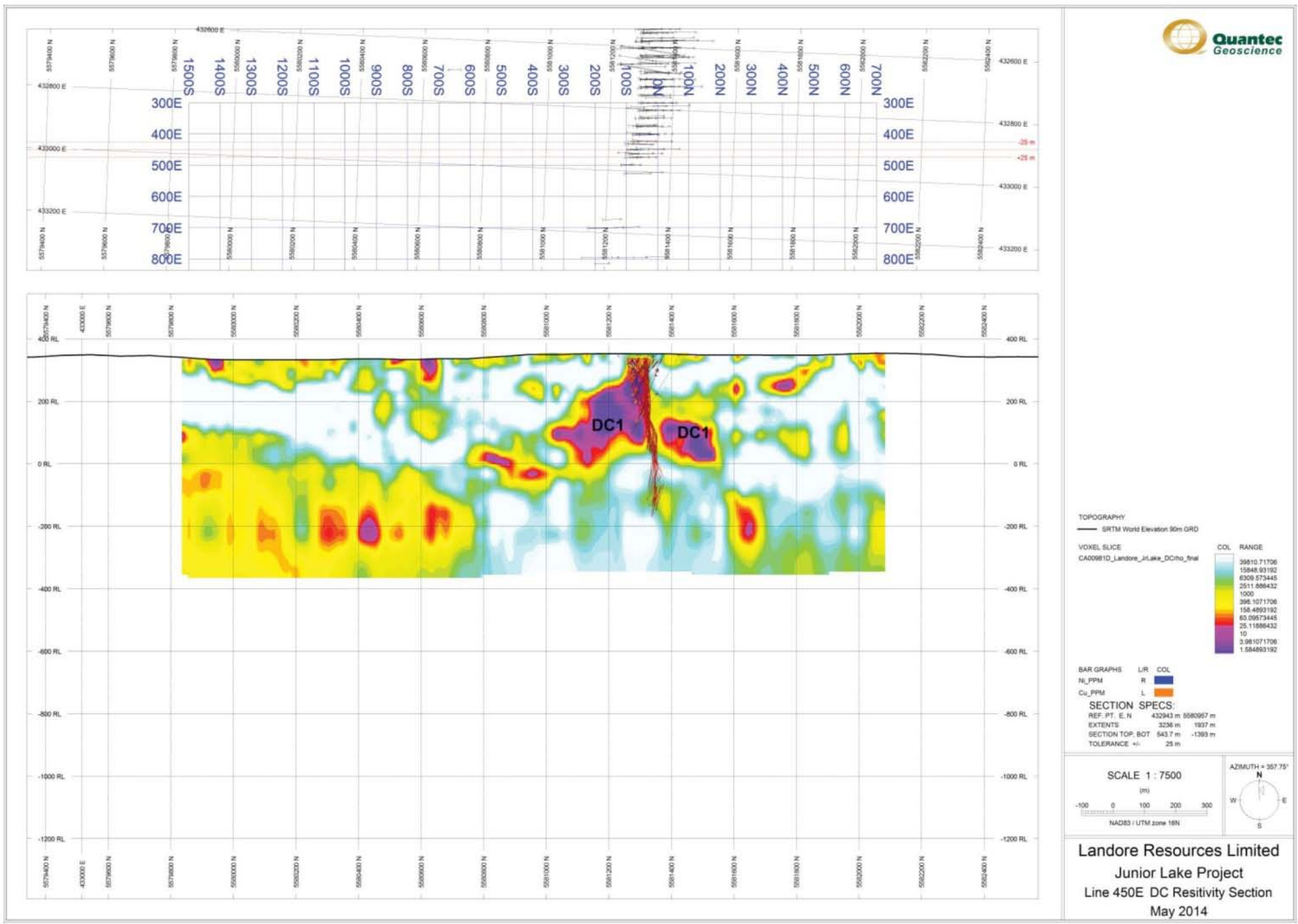


70 milliradian chargeability isosurface.

Figure 8-4: 700 ohm-m MT Resistivity Isosurface

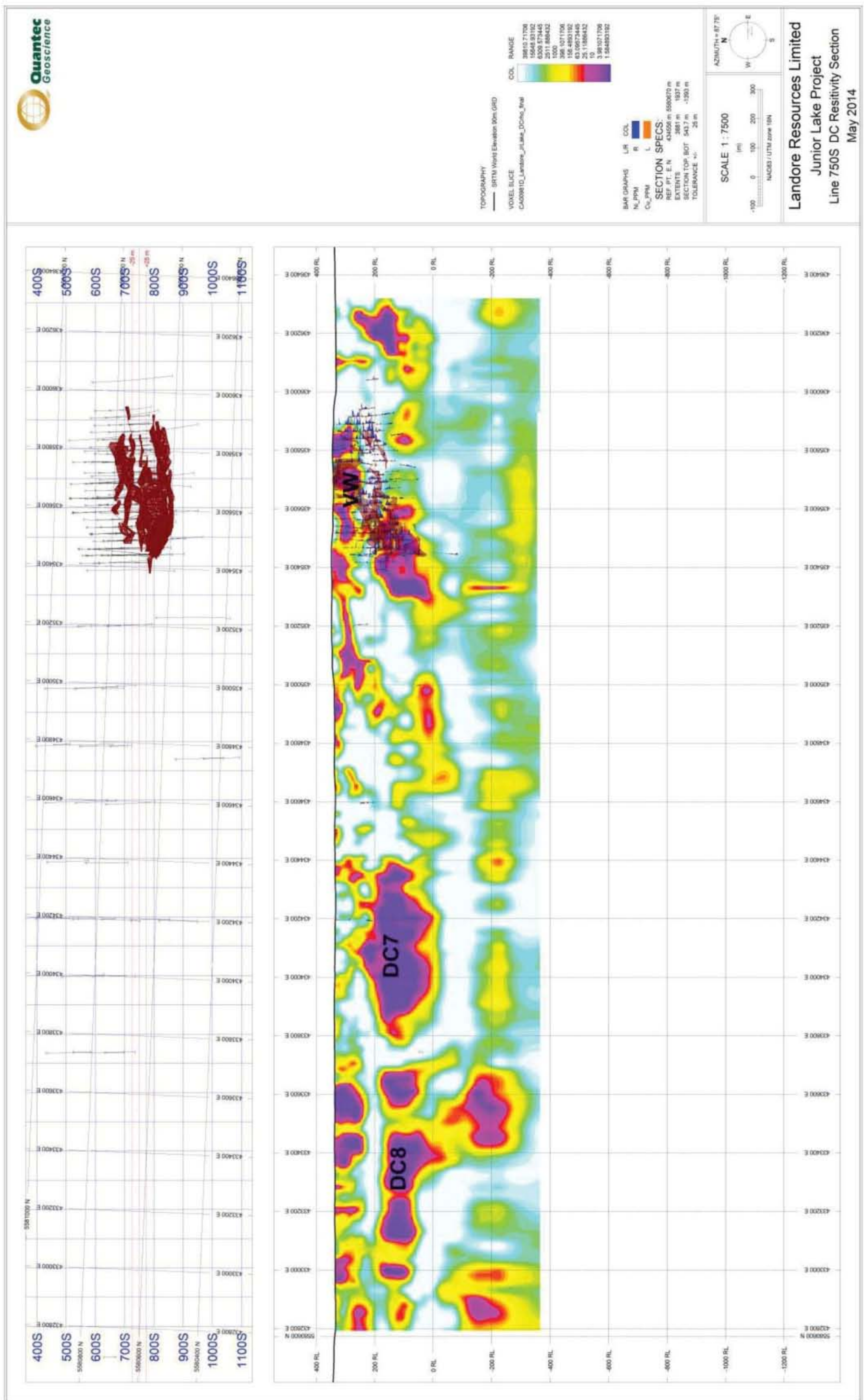
700 Ω·m MT resistivity isosurface.

Figure 8-5: N-S DC Resistivity cross-section Line 450E



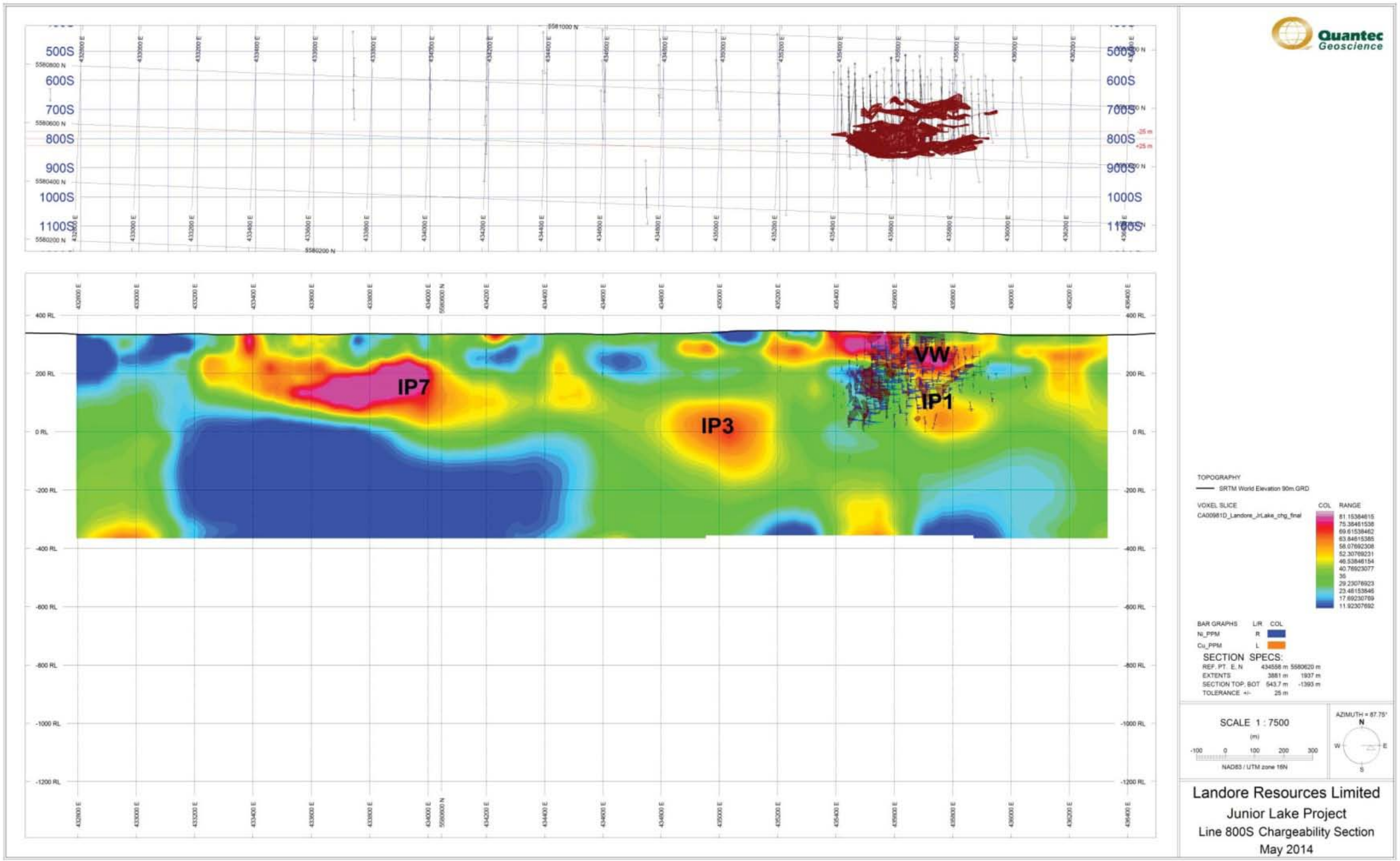
N-S DC resistivity cross-section along Landore grid line 450E (looking west) with conductive zone DC1 and B4-7 mineralized zone (wireframe).

Figure 8-6: E-W DC Resistivity Cross-Section Line 750S



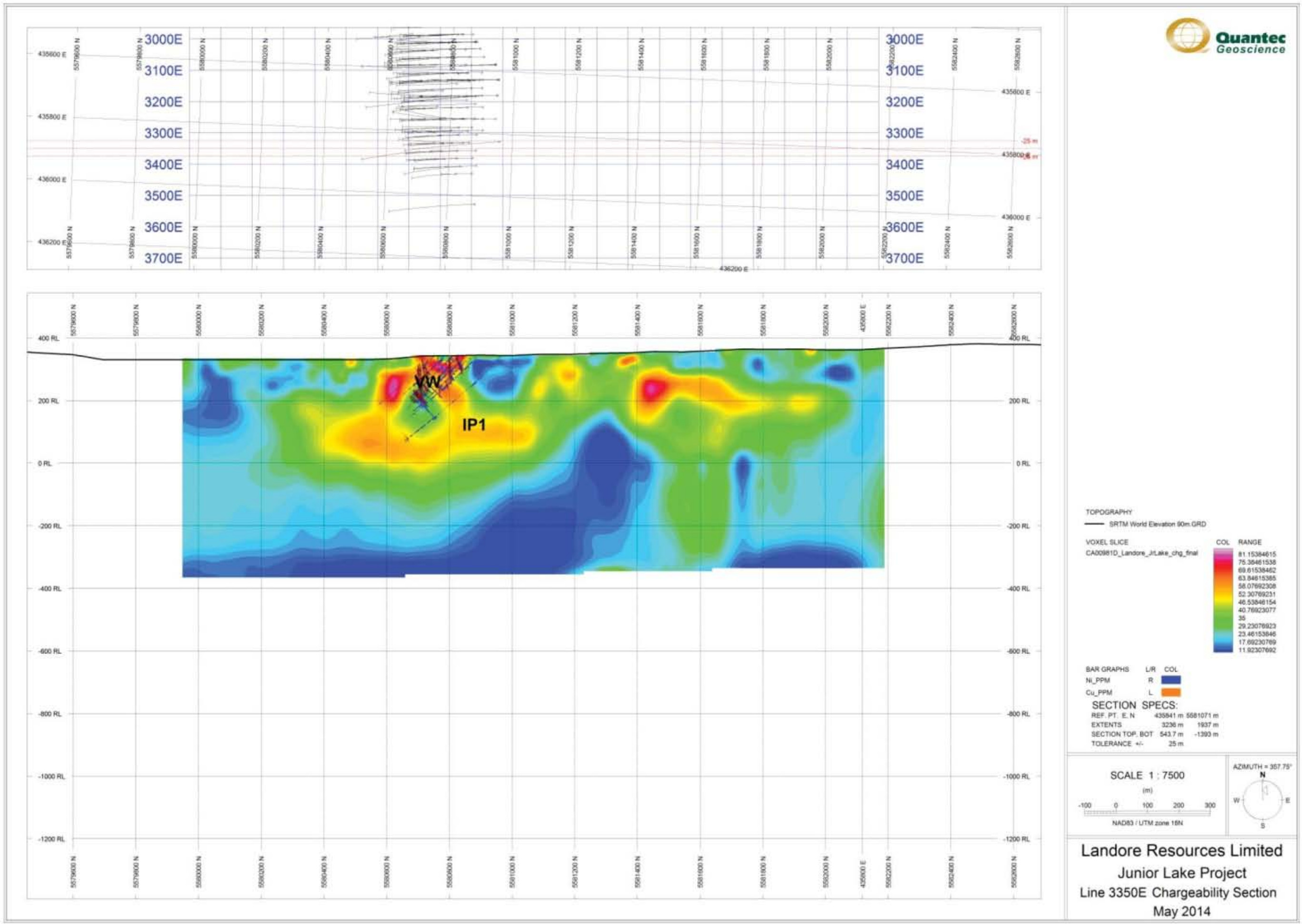
E-W DC resistivity cross-section along Landore grid line 750S (looking north) with conductive zones DC7 and DC8 and VW mineralized zone (transparent shell).

Figure 8-7: E-W IP Chargeability Cross-Section Line 800S



E-W IP chargeability cross-section along Landore grid line 800S (looking north) with chargeable zones IP1, IP3, and IP7 and VW mineralized zone (transparent shell).

Figure 8-8: N-S Chargeability Cross-Section Line 3350E



N-S IP chargeability cross-section along Landore grid line 3350E (looking west) with chargeable zone IP1 and VW mineralized zone (transparent shell).

Figure 8-9: N-S DC Resistivity Cross-Section 1700E with Drilling

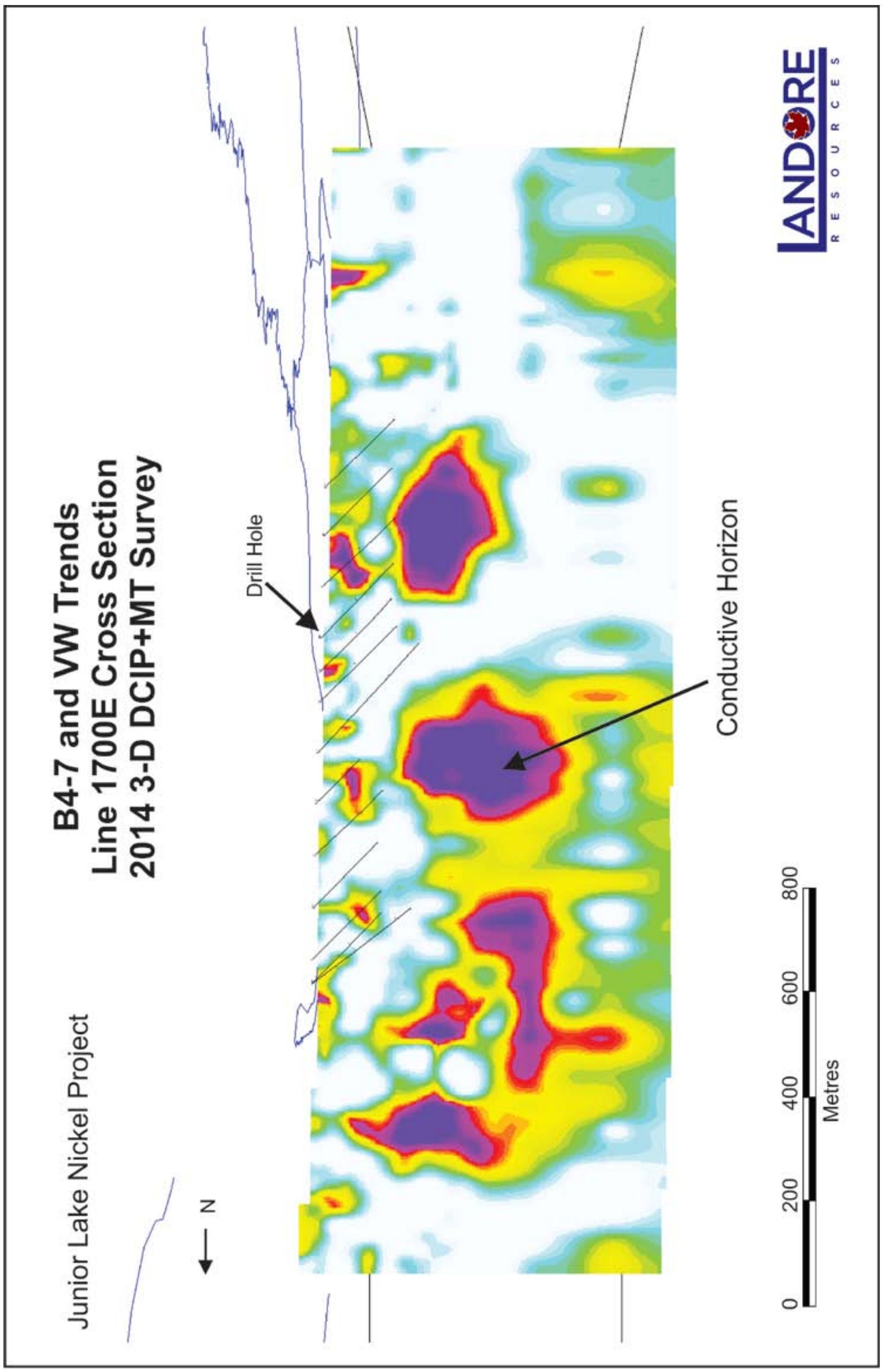


Figure 8-10: E-W Long Section 1 ohm-m DC Resistivity Isosurfaces with Drilling

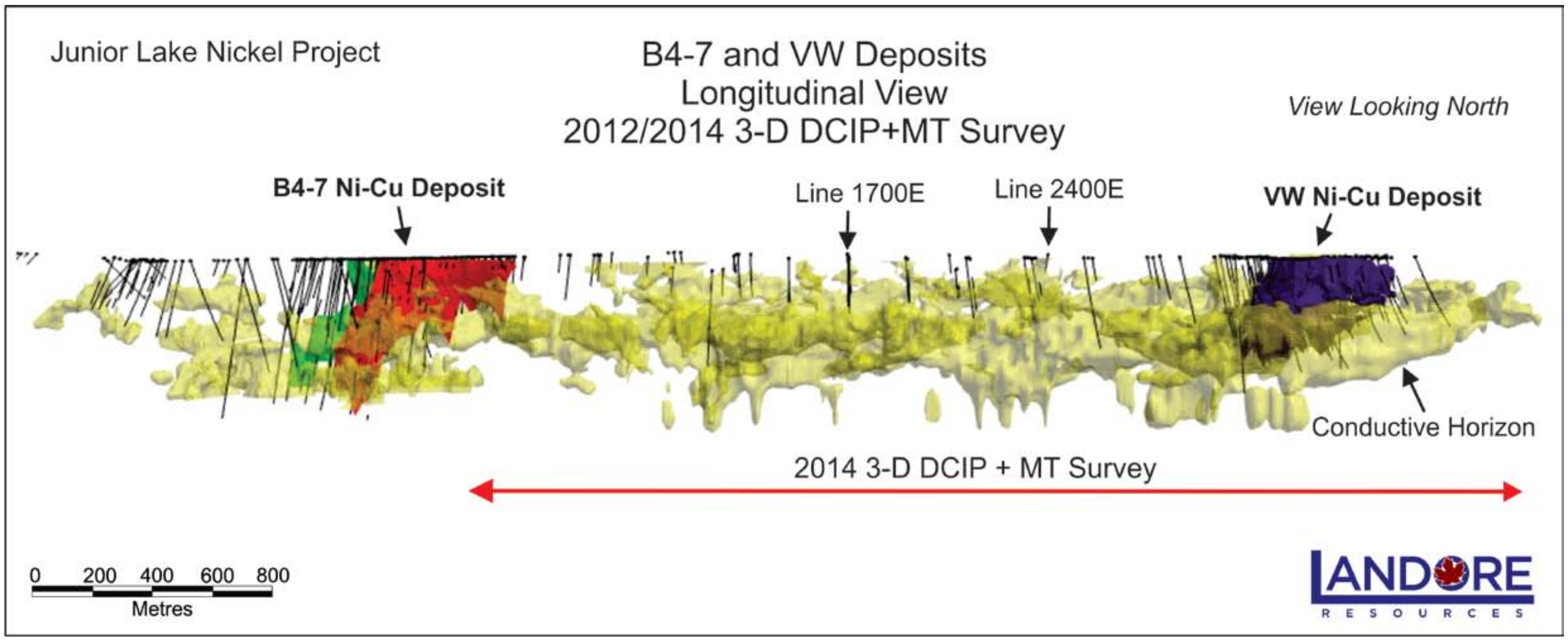
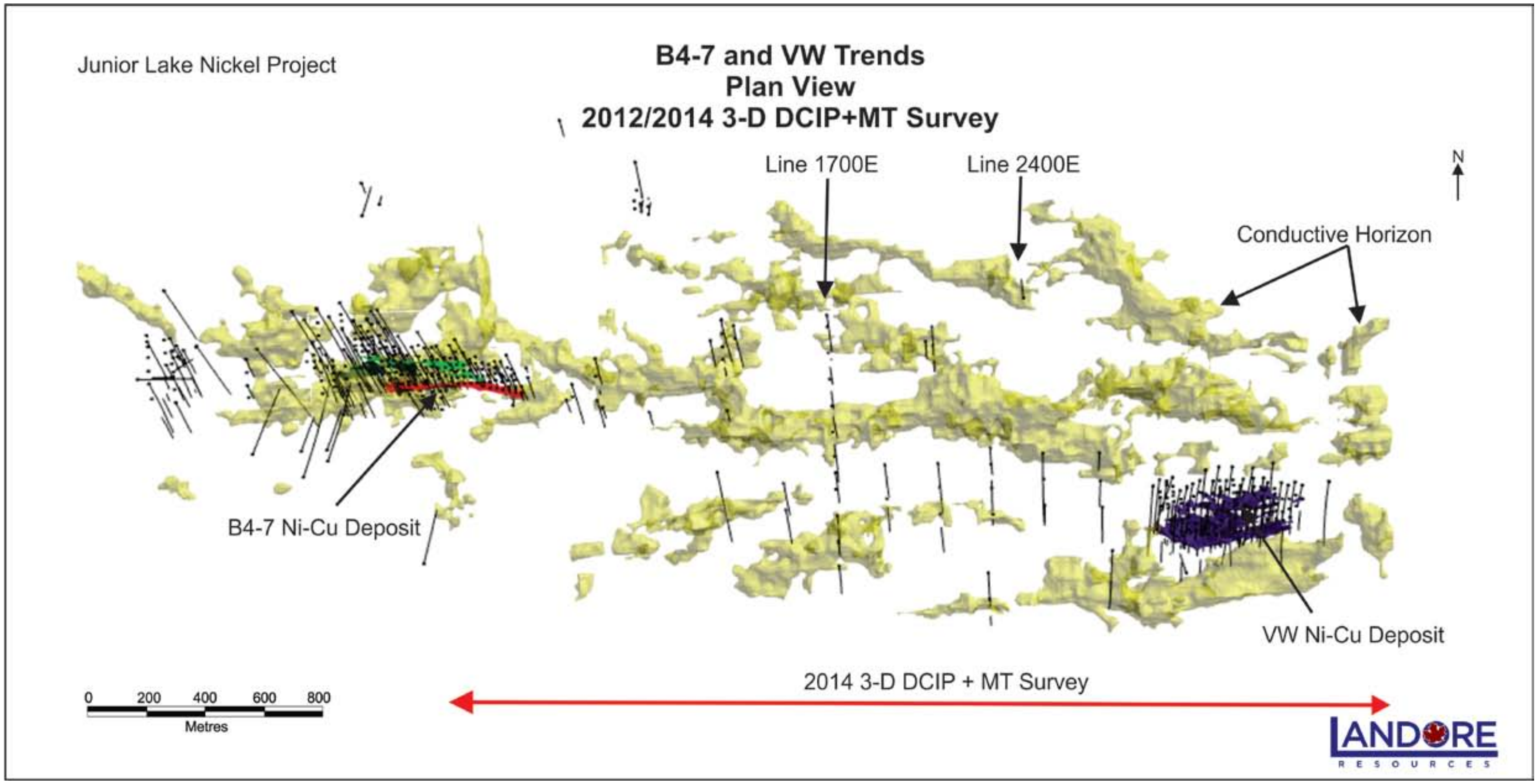


Figure 8-11: Plan View 1 ohm-m DC Resistivity Isosurfaces with Drilling



9 SURVEY DESIGN AND PROCEDURES

The design and procedures for the Orion 3-Dimensional (3D) Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) survey are summarized by Gharibi/McGill (2014):

9.1 Survey Design

The DCIP survey over the Junior Lake Project - Phase II was completed using two Orion 3D spreads. Two spreads include 84 multi-channel Data Acquisition Units (DAU's) with a total of 504 receiver dipoles of 100 m length, oriented in an orthogonal grid. The two southernmost receiver rows in N-S direction were 150 m in length. The same DAU's were used simultaneously for the magnetotelluric survey, using the same dipole array and 2 set of magnetic induction coil sensors.

A total of 897 current injection locations were used, resulting in more than 240,000 DC and IP data samples over the survey area. Additionally, 168 MT stations (2 per DAU) with a nominal spacing of 100 m and 300 m were surveyed using the same receiver dipoles.

9.2 Survey Procedure

After mobilization to the survey site, the initial setup of each spread was completed in 2 days. Data acquisition for both spreads took a total of 17 days, averaging 52 current injections per day. MT data was acquired at night throughout the survey period.

Initial data QC and processing took place at the field processing office, and processed results were transmitted to Quantec's Toronto office daily.

9.3 DCIP + MT Survey

The Orion 3D system acquires three types of geophysical data – magnetotelluric (MT), direct current resistivity (DC), and induced polarization (IP).

The MT and DC methods are used to resolve the resistivity distribution of the subsurface by measuring the electric potential (DC) and the variation of natural source electric and magnetic fields (MT). Resistivity can be an indicator of metallic mineralization, but is more often than not controlled by rock porosity and is therefore an indirect indicator of alteration and mineral grain fabric.

In the induced polarization method, the electrical capacitance or chargeability of the subsurface is measured. Chargeability is a near-direct indicator of the presence of sulphide mineralization, in both massive and disseminated forms. Chargeable mineralization is most commonly various sulphides and graphite, but also includes clay-type minerals potentially making it a useful tool for base-metals exploration.

The DC-IP survey utilized a pole-dipole configuration with 100m receiver dipoles (except for the two southernmost receiver rows) and current injection points designed to provide complete sampling of the survey area.

The MT survey utilized the same DC-IP dipoles, plus magnetic coils. A remote reference site with the same magnetic sensor configuration was used to improve processing of the MT data.

9.4 Survey Methodology, Logistics

Detailed information on the 2014 DCIP + MT survey methodology, logistics and acquisition parameters are provided in the Quantec 2014 Orion 3-D DCIP + MT survey report (Gharibi/McGill, 2014).

10 DATA VERIFICATION

Drill hole and assay data entered or imported into Landore's Microsoft Access database is checked by the software and Senior Geologist for data entry errors.

To validate the drill hole database is checked for potential problems such as:

1) Intervals exceeding the hole length (from-to problem).
2) Negative length intervals (from-to problem).
3) Zero length intervals (from-to problem).
4) Inconsistent downhole survey records.
5) Out of sequence and overlapping intervals (from-to problem; additional sampling/QAQC/check sampling included in table).
6) No interval defined within analyzed sequences (not sampled or missing samples/results).

11 INTERPRETATION AND CONCLUSIONS

The Orion 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) ground geophysics program was completed over the VW deposit, VW West, and B4-7 East areas, from line 300E to line 3700E and from 700N to 1500S. The survey, covering 739.02 hectares, was conducted by Quantec Geoscience Ltd., Toronto, for Landore Resources. This survey is located directly adjacent to the east of Landore Resources' 2012 Orion 3D DCIP + MT survey block covering the B4-7 West zone.

The survey was used to identify prospective strike extension to the east of the B4-7 Deposit and west of the VW NI 43-101 compliant resource as well as to depth of both deposits, and to investigate the base and precious mineral potential in the intervening 3 kilometre distance between the two deposits.

The DCIP + MT geophysics survey on the VW Deposit, VW West and B4-7 East areas was successful in identifying potential massive sulphide mineralization as well as disseminated sulphide mineralization containing economic grades of nickel, copper, cobalt, PGEs and gold.

12 RECOMMENDATIONS

The 2014 Orion 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) ground geophysics program was successfully completed on the VW Deposit, VW West, and B4-7 East areas, located in the central portion of the Junior Lake property. Survey results from the 2014 survey have delineated nine significant new zones ranging from approximately 400m to 1,200m in length of potential nickel sulphide mineralization along strike and adjacent to the existing B4-7 Nickel-Copper-Cobalt-PGEs resource and the VW Nickel resource. Numerous targets are at depth and below existing Landore exploration drilling.

Follow-up work is recommended to drill prospective geophysical targets along the VW West and B4-7 East areas utilizing the results of this survey together with Landore's geophysical information from earlier surveys. Deposit drilling is recommended to delineate strike extension of the B4-7 and VW deposits to bring further mineralized zones into the NI 43-101 compliant resource.

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14 SIGNATURE PAGE

This report titled "Work Assessment Report on the Junior Lake Property – 2014 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) Ground Geophysics Program (VW Deposit, VW West, B4-7 East areas) – January 5, 2016" was prepared by M. Tuomi and signed by the following Author:

Michele Tuomi, P.Geo.
Landore Resources Canada Inc.

Thunder Bay, Ontario
January 5, 2016

15 CERTIFICATE OF QUALIFIED PERSON

Michele Tuomi, P.Geo.
Landore Resources Canada Inc.
555 Central Avenue, Suite 1
Thunder Bay, ON
P7B 5R5

Tel: +1 807 623 3770

I, Michele Tuomi, am a Professional Geoscientist, employed as a VP Exploration of Landore Resources Canada Inc.

This certificate applies to the geological report titled "2014 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) Ground Geophysics Program (VW Deposit, VW West, B4-7 East areas) – January 5, 2016" dated January 5, 2016.

I am a member of the Association of Professional Geoscientists of Ontario. I graduated with a BSc. degree in Geology from Lakehead University in 1999.

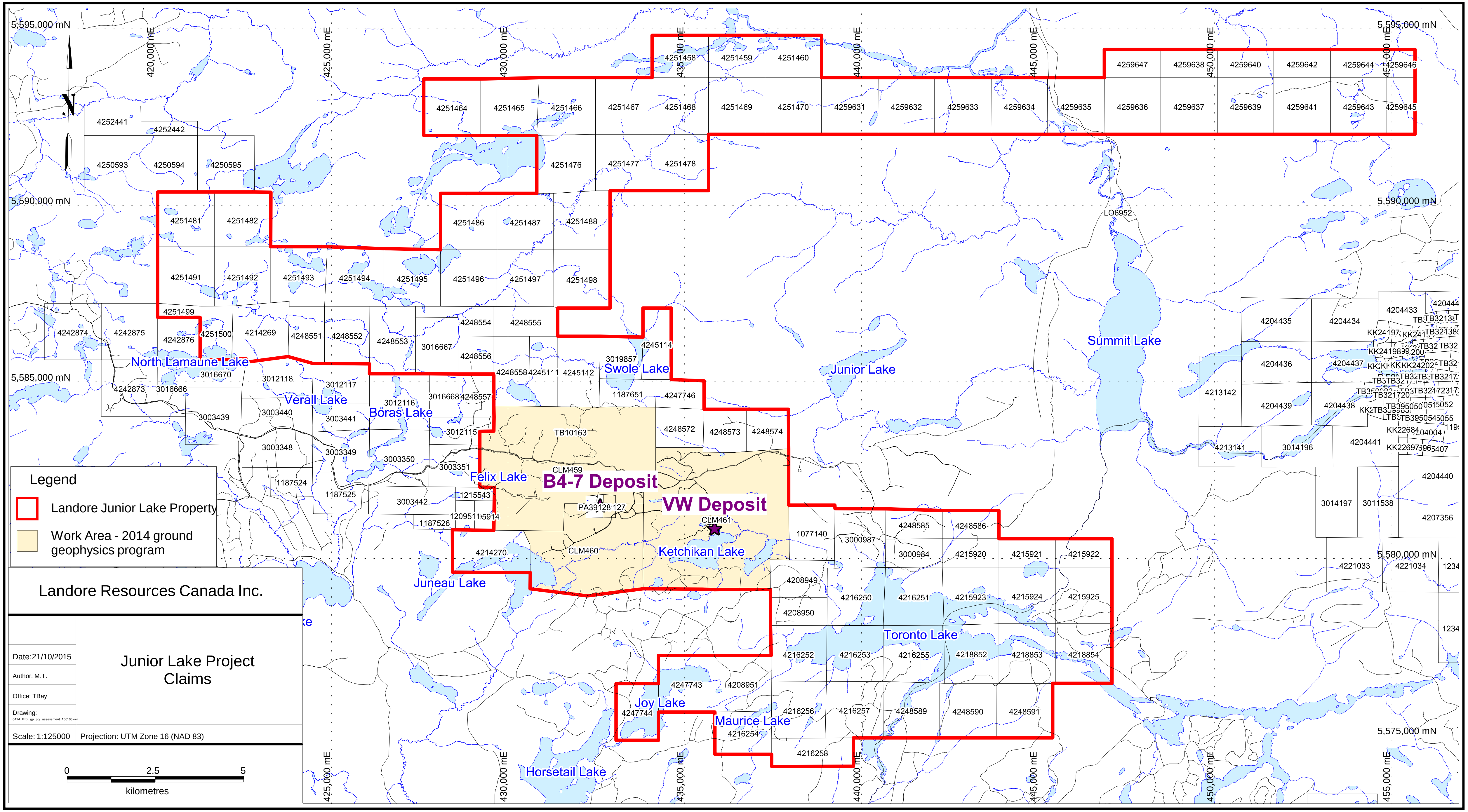
I have practiced my profession for 16 years. I have been directly involved in mineral exploration and mineral project assessment, as well as mineral resource estimations.

I have visited the Junior Lake property in northern Ontario, Canada on numerous occasions, the most recent being December 6, 2015.

I am responsible for all items of the assessment report "2014 3-Dimensional Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) Ground Geophysics Program (VW Deposit, VW West, B4-7 East areas) – January 5, 2016".

As of the date of this certificate, to the best of my knowledge, information and belief, the technical report contains all scientific and technical information that is required to be disclosed to make the assessment report not misleading.

Michele Tuomi, P.Geo.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Office: TBay
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Landore Junior Lake Property
Work Area - 2014 ground geophysics program
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VW Deposit
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Verall Lake
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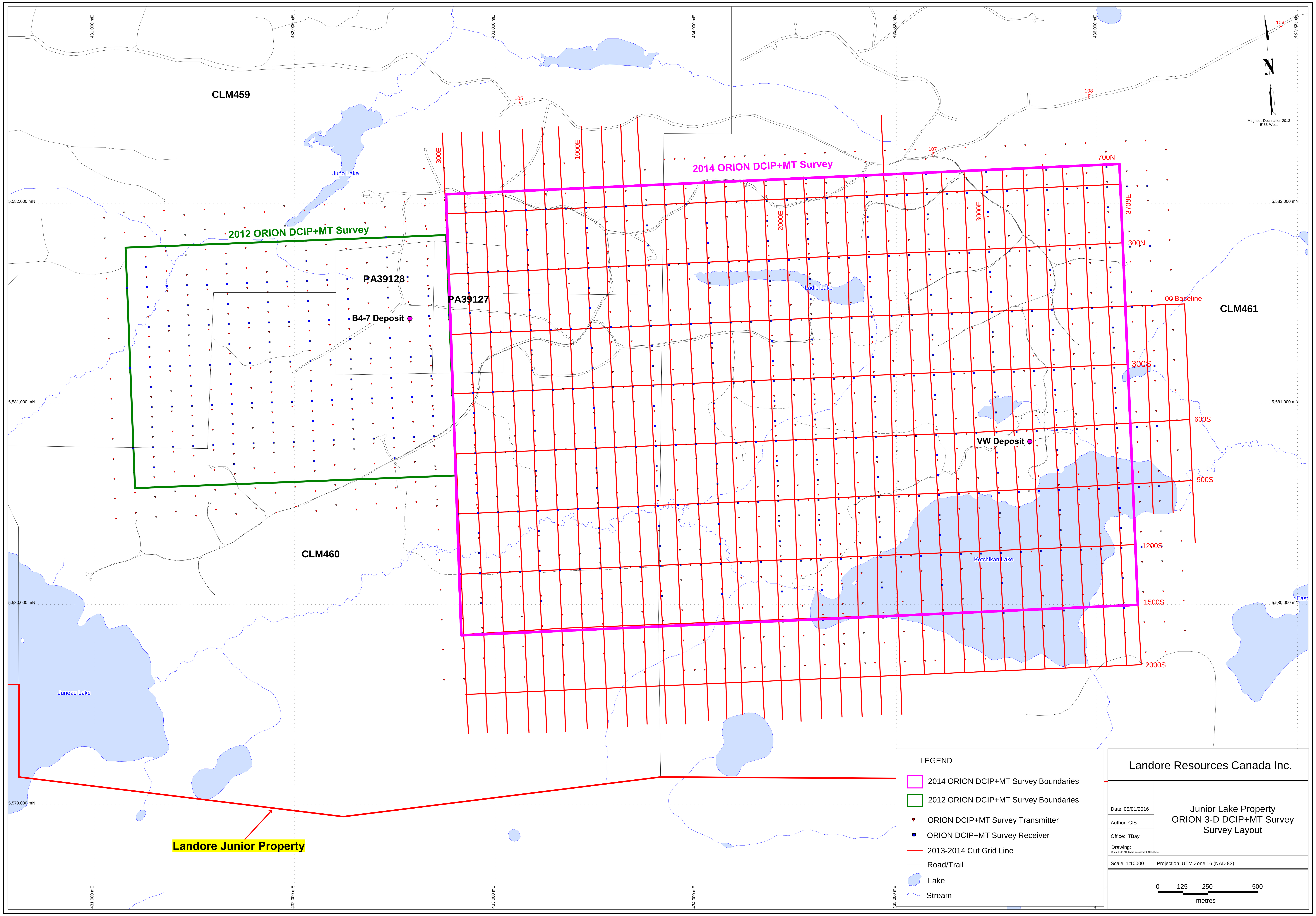

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CLM460
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PA39128
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B4-7 Deposit
VW Deposit
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Ladle Lake
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2014 ORION DCIP+MT Survey
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300N
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300S
600S
900S
1200S
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2000S
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Lake
Stream
Landore Resources Canada Inc.
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Author: GIS
Office: TBay
Drawing:
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