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Ontario

Work Assessment Report

on the

JUNIOR LAKE PROPERTY

2014 Fall Diamond Drill Program
(B4-7 Deposit, B4-7 East Area, VW Deposit, VW West Area)

Falcon Lake Area
Thunder Bay North Mines and Minerals Division
Ontario

NTS 52I/08 and 42L/05

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September 9, 2016
Thunder Bay, Ontario

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1 SUMMARY

The Junior Lake property is located approximately 230 kilometres north-northeast of the city of Thunder Bay, Ontario, within the central portion of the Caribou-O'Sullivan Greenstone Belt. The property is host to two NI 43-101 compliant mineral resources – the B4-7 Ni-Cu-Co-PGE deposit and the VW Ni deposit, located 3 kilometres apart. Other occurrences of Ni-Cu-PGE, Cu, Cu-Zn, Cr, Li and Au are known on the property.

This report covers the fall 2014 exploration drilling program conducted on various localities in the central portion of the Junior Lake property; the B4-7 Deposit, the B4-7 East area, the VW Deposit and the VW West area. Drilling followed up on results from both the 2014 3D DCIP + MT and the 2013 Electromagnetic (MaxMin), VLF and Magnetometric ground geophysical surveys, which identified several prospective anomalies through these areas. During September to October 2014, a total of 12 drill holes (0414-493 to 0414-504) for 2,675 metres were drilled, logged and sampled.

Drilling confirmed a down dip extension of the main B4-7 massive sulphide zone 140 metres below the existing B4-7 resource on line 00 as well as intersecting high grade Alpha zone platinum group element (PGE) mineralisation. Drilling on the B4-7 deposit further intersected shallow high-grade nickel and Alpha zone PGEs on the eastern portion of the deposit (line 275E). Exploration drilling southwest of the VW deposit discovered significant near-surface precious metal mineralisation.

Mineralized intersections of fall 2014 drilling include:

- DDH 0414-495: 0.34 metres at 1.21g/t Au, 280g/t Ag, and 2.25% Zn
- DDH 0414-503: 20.15 metres at 0.11% Ni, 0.04% Cu, 0.01% Co, 1.54g/t Pd, 0.64g/t Pt, and 0.01g/t Au
  including 0.55 metres at 3.39g/t Pd and 1.56g/t Pt
  including 0.56 metres at 10.90g/t Pd and 11.50g/t Pt
  including 0.72 metres at 12.85g/t Pd and 2.50g/t Pt
- DDH 0414-503: 3.46 metres at 0.62% Ni, 0.18% Cu, 0.05% Co, 0.29g/t Pd, 0.04g/t Pt, and 0.01g/t Au including 0.60 metres at 1.01% Ni, 0.09% Cu, 0.06% Co, 0.41g/t Pd, 0.03g/t Pt and 0.01g/t Au
- DDH 0414-504: 2.90 metres at 0.78% Ni, 0.14% Cu, 0.04% Co, 1.54g/t Pd, 0.06g/t Pt including 0.58 metres at 1.77% Ni, 0.20% Cu, 0.08% Co, 3.27g/t Pd, 0.09g/t Pt, 0.02g/t Au
- DDH 0414-504: 0.93 metres at 0.68% Ni, 0.22% Cu, 0.04% Co, 2.25g/t Pd, 0.70g/t Pt, 0.02g/t Au
- DDH 0414-504: 0.89 metres at 0.72% Ni, 0.81% Cu, 0.16% Co, 0.34g/t Pd, 0.04g/t Pt, 0.05g/t Au
- DDH 0414-504: 9.32 metres at 0.98% Ni, 0.50% Cu, 0.08% Co, 0.96g/t Pd, 0.24g/t Pt, 0.02g/t Au including 3.29 metres at 0.35% Ni, 1.59% Cu, 0.17% Co, 0.15g/t Pd, 0.13g/t Pt, 0.13g/t Au

The B4-7 combined open pit and underground NI 43-101 compliant resource is 2,695,000 tonnes at 1.24% Nickel equivalent (NiEq) for 33,248 tonnes of contained metal, all within the Indicated category (2013 Mineral Resource estimate and a NI 43-101 Technical Report on the B4-7 Deposit, RPA, Toronto, Canada).

The VW deposit contains 3.73 million tonnes at 0.49% NiEq in the Indicated category at a cut-off grade of 0.25 per cent. nickel (2010 Mineral Resource estimate and a NI 43-101 Technical Report on the VW Deposit, RPA, Toronto, Canada). There is a further 0.72 million tonnes at 0.49% NiEq in the Inferred category at the same cut-off grade giving a combined total of 21,760 tonnes of contained NiEq metal.

Follow-up drilling is required to further establish continuity of the high grade Alpha PGE zone in the B4-7 deposit, and to determine the scope of B4-7 main massive sulphide mineralisation at depth. Geophysics results indicate high potential for further B4-7 massive sulphide and Alpha zone disseminated sulphide mineralisation at depth and along strike to the east and west of the drilled B4-7 deposit, adding considerable tonnage and value to the deposit.

Additionally, further drilling is warranted in the prospective area between the B4-7 and VW deposits as identified by previous ground geophysics.

The fall 2014 exploration drilling program included program preparation (including local grid line cutting), 2,675 metres of NQ size diamond drilling, assaying, and geological analysis of results. The total amount from this exploration program claimed for assessment credit is $568,360.52.

2 INTRODUCTION

This report and accompanying documentation presents the results of the fall 2014 exploration drilling program conducted on Landore Resources Canada Inc.'s Junior Lake property. The Junior Lake property is located approximately 230 kilometres north-northeast of the city of Thunder Bay, Ontario, within the central portion of the Caribou-O'Sullivan Greenstone Belt. It is host to several PGE-Cu-Ni, Cu, Cu-Zn, Li, Au, and Ag occurrences. In the vicinity of the fall 2014 drilling program, the property hosts two NI 43-101 compliant nickel deposits – the B4-7 Ni-Cu-Co-PGE deposit and the VW Ni deposit, located three kilometres apart.

The fall 2014 exploration drilling program was conducted on the B4-7 Deposit, the B4-7 East area, the VW Deposit and the VW West area. During this drilling program a total of 12 drill holes (0414-493 to 0414-504) for 2,675 metres were drilled, logged and sampled. Drilling intersected a down dip extension of the main B4-7 massive sulphide zone 140 metres below the existing B4-7 resource on line 00 as well as intersecting high grade Alpha zone platinum group element (PGE) mineralisation. Drilling on the B4-7 deposit further intersected shallow high-grade nickel and Alpha zone PGEs on the eastern portion of the deposit (line 275E). Exploration drilling southwest of the VW deposit discovered significant near-surface precious metal mineralisation.

Base metals, PGE and gold assaying were undertaken by ALS-Chemex of Vancouver, British Columbia, and Accurassay of Thunder Bay, Ontario.

This report is submitted to the Ontario Ministry of Northern Development and Mines Geoscience Assessment Office to claim assessment credit.

3 PROPERTY DESCRIPTION AND LOCATION

The Junior Lake property is located approximately 230 km north-northeast of Thunder Bay, Ontario, and approximately 75 km east-northeast of the village of Armstrong, Ontario (Figure 2-1). The centre of the property is located at 87º59'4"W longitude and 50º23'9"N latitude; NAD83 UTM coordinates Zone 16, 430,000E and 5,580,000N. The property area is within the NTS 1:50,000 Jackfish Lake and Toronto Lake topographic map sheets NTS 52I/08 and 44L/05, respectively. The Junior Lake property claims and leases are located on the Falcon Lake, Junior Lake, Toronto Lake, Kapikotongwa River, Summit Lake, and Willet Lake claim maps (Thunder Bay Mining Division areas NTS 52I/08NE and SE, 42L/05NW, SE and SW).

LAND TENURE

Landore's Junior Lake property consists of 95 mineral claims (1,145 units) and four mining leases totaling 3,793 hectares (Tables 3-1 and 3-2, Figure 3-2).

Landore held a 100% interest in claims TB1077140 to TB1077142, TB1217179 to TB1217181, and TB1233556 and TB1233557, subject to a 2% net smelter return (NSR) royalty held by Wing Resources Inc. The above claims, except TB1077140, have been taken to lease. The B4-7 Deposit lies on patented claims PA39127, PA39128 and lease CLM460, whereas the VW Deposit lies on lease CLM461.

The exploration work undertaken by Landore prior to 28th August, 2008 was on mining leases in which Landore held a 100% interest: mining claims TB1077142, TB1217179. These claims were taken to lease (CLM 461) on 28th August, 2008.

Figure 3-1: Junior Lake Property Location

Figure 3-2: Junior Lake Property Leases and Claims

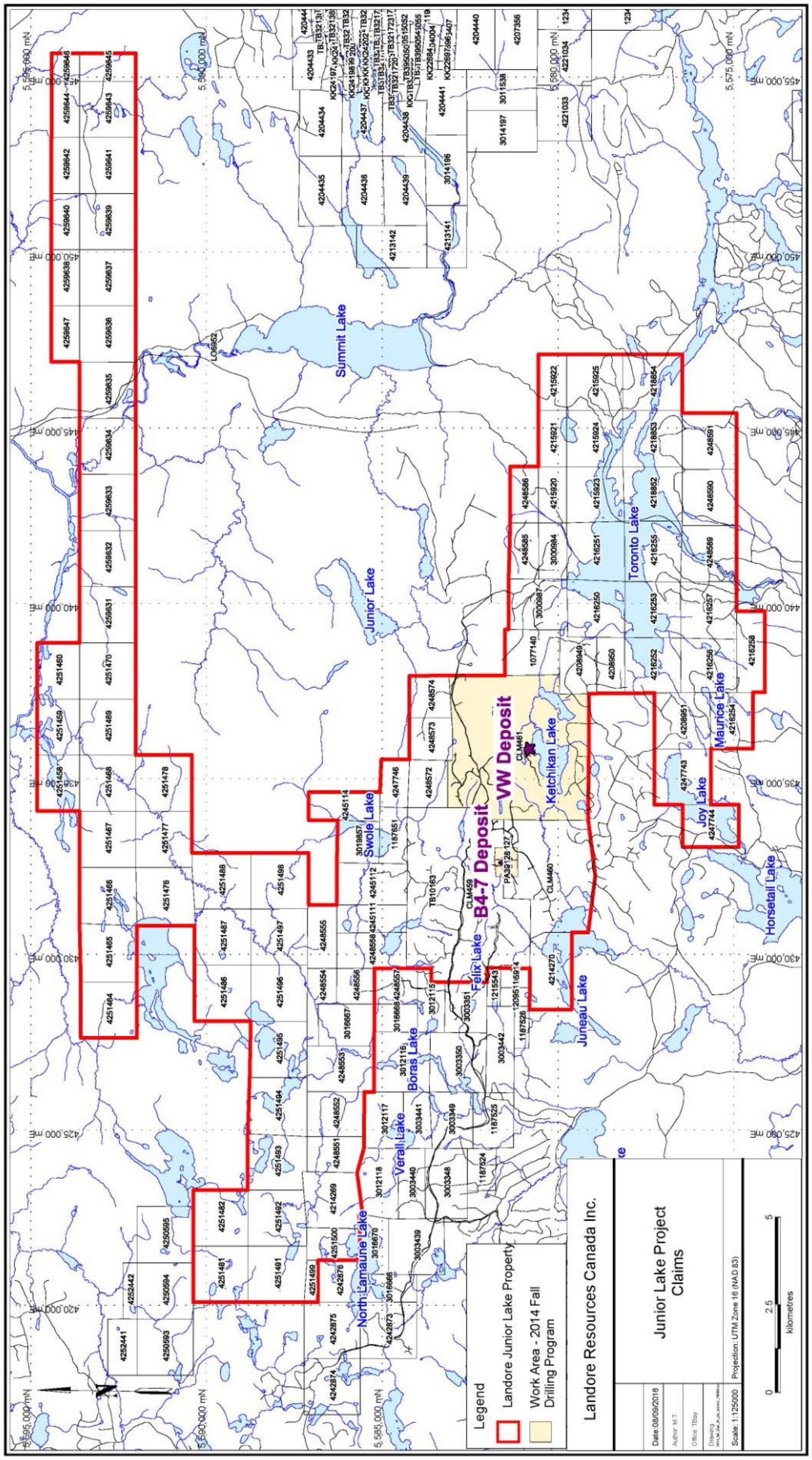


Table 3-1: Landore Mineral Claims (100% Interest)

Claim	Calculated Area (ha)	Units	Area	Claim	Calculated Area (ha)	Units	Area
1077140	201.533	9	Junior Lake	4251487	255.274	16	Junior Lake
1187651	126.417	8	Junior Lake	4251488	262.535	16	Junior Lake
3000984	129.049	8	Toronto Lake	4251491	263.875	16	Falcon Lake
3000987	241.242	14	Toronto Lake	4251492	268.399	16	Falcon Lake
3016667	191.05	12	Falcon Lake	4251493	262.351	16	Falcon Lake
3019857	143.257	9	Junior Lake	4251494	271.898	16	Falcon Lake
4208949	174.532	10	Toronto Lake	4251495	258.004	16	Falcon Lake
4208950	127.147	8	Toronto Lake	4251496	255.947	16	Junior Lake
4208951	252.102	16	Toronto Lake	4251497	257.811	16	Junior Lake
4215920	128.463	8	Toronto Lake	4251458	191.468	12	Kapikotongwa R.
4215921	128.45	8	Toronto Lake	4251459	191.47	12	Kapikotongwa R.
4215922	128.443	8	Willet Lake	4251460	191.467	12	Kapikotongwa R.
4215923	255.325	16	Toronto Lake	4251464	255.269	16	Junior Lake
4215924	255.568	16	Toronto Lake	4251465	262.916	16	Junior Lake
4215925	255.604	16	Willet Lake	4251466	262.901	16	Junior Lake
4216250	262.832	16	Toronto Lake	4251467	255.274	16	Junior Lake
4216251	272.55	16	Toronto Lake	4251468	255.272	16	Junior Lake
4216252	250.524	16	Toronto Lake	4251469	255.284	16	Junior Lake
4216253	252.657	16	Toronto Lake	4251470	255.289	16	Junior Lake
4216254	194.379	12	Toronto Lake	4251476	269.928	16	Junior Lake
4216255	277.21	16	Toronto Lake	4251477	255.274	16	Junior Lake
4216256	244.297	16	Toronto Lake	4251478	255.285	16	Junior Lake
4248585	132.61	8	Junior Lake	4251481	247.747	16	Falcon Lake
4248586	127.648	8	Junior Lake	4251498	262.503	16	Junior Lake
4248589	251.612	16	Toronto Lake	4251499	42.8057	3	Falcon Lake
4248590	246.935	16	Toronto Lake	4251500	135.24	8	Falcon Lake
4248591	246.981	16	Toronto Lake	4259631	255.3975	16	Junior Lake
4216257	259.137	16	Toronto Lake	4259632	255.2966	16	Junior Lake
4216258	184.68	12	Toronto Lake	4259633	255.4375	16	Junior Lake
4218852	257.424	16	Toronto Lake	4259634	254.9172	16	Junior Lake
4218853	269.098	16	Toronto Lake	4259635	255.9382	16	Summit Lake
4218854	269.053	16	Willet Lake	4259636	255.4302	16	Summit Lake
4245111	157.295	10	Junior Lake	4259637	255.5525	16	Summit Lake
4245112	236.437	15	Junior Lake	4259638	127.6468	8	Summit Lake
4245114	164.202	10	Junior Lake	4259639	255.2938	16	Summit Lake
4247743	254.86	16	Toronto Lake	4259640	127.657	8	Summit Lake
4247744	191.684	12	Toronto Lake	4259641	255.3059	16	Summit Lake
4247746	108.215	6	Junior Lake	4259642	127.6557	8	Summit Lake
4248551	160.761	10	Falcon Lake	4259643	255.3034	16	Summit Lake
4248552	206.752	12	Falcon Lake	4259644	127.1522	8	Summit Lake
4248553	248.178	15	Falcon Lake	4259645	127.654	8	Summit Lake
4248554	123.24	8	Junior Lake	4259646	63.16516	4	Summit Lake
4248555	151.516	9	Junior Lake	4259647	127.6481	8	Summit Lake
4248556	107.904	8	Junior Lake	4214269	249.927	16	Falcon Lake
4248558	200.808	10	Junior Lake	4214270	258.207	16	Toronto Lake
4248572	151.125	9	Junior Lake	**95**	**18,495.98**	**1,145**	
4248573	147.231	9	Junior Lake				
4248574	137.856	9	Junior Lake				
4251482	247.088	16	Falcon Lake				
4251486	255.278	16	Junior Lake				

Table 3-2: Landore Leases (100% Interest)

Lease #	Description	G-Number	Anniversary Date	Area (ha)	Annual Rent ($)	Expiry Date	Total Work in Reserve ($)
107421	PA 39127, 39128	4000476	98-Jan-01	52.969	158.91	2019-Jan-01	1,096,271
108257	CLM459[1]	4040218	08-Aug-01	1,460.795	4,382.39	2029-Aug-01	17,284
108258	CLM461[1]	4040217	08-Aug-01	1527.388	4,582.16	2029-Aug-01	2,468,109
108259	CLM460[1]	N/A[2]	08-Aug-01	687.794	2,063.38	2029-Aug-01	0
Totals	**4 Leases**			**3,728.946**	**11,186.84**		**3,581,664**

Notes:

1. Wing Resources holds a 2% NSR on 3 claims within CLM459, 1 claim within 460 and 3 claims within 461.
2. G-number is generated when work reports are filed.

Landore has been granted four mining leases, which include mining and surface rights, over an area encompassing the B4-7 and VW Deposits. The leases cover 23 mineral claims and two patents for a total area of 3,729 ha and have been granted for 21 years renewable for further terms of 21 years (Table 3-2).

Within the mining leases, Landore has the right, subject to provisions of certain Acts and reservations, to:

- sink shafts, excavations, etc., for mining purposes;
- construct dams, reservoirs, railways, etc., as needed; and
- erect buildings, machinery, furnaces, etc., as required, and treat ores.

There are no known environmental liabilities on the property.

4 ACCESSIBILITY

Access to the Junior Lake property from Thunder Bay is via paved provincial highways No. 17 (15 km) and No. 527 to Armstrong, with an overall distance of approximately 235 km. From Armstrong, the Buchanan Forest Products Inc. gravel haulage road (BHR) is taken east to kilometre 105, where a skidder haulage road leads approximately one kilometre to the Landore Junior Lake camp. Skidder and drill roads provide access on the property. The site of the fall 2014 exploration drilling program is located in the central portion of the Junior Lake property, within the prospective area spanning from the B4-7 Nickel-Copper-Cobalt-PGE deposit to the VW Nickel-Copper deposit.

There are no power lines or railway lines on the property; however, the main CNR line is approximately 13 kilometres to the south.

During the summer, most drill sites are accessible by 4-wheel-drive vehicles.

5 HISTORY

Routledge (2010) has summarized the exploration and development history of the Junior Lake property as:

Geological mapping and exploration in the vicinity of the Junior Lake property is recorded as early as 1917. In 1968, Canadian Dyno Mines Limited staked 333 claims in 15 groups to cover conductors picked from an airborne electromagnetic (EM) and magnetic (MAG) survey. Two groups, B3 and B4, included the Junior Lake property. The company merged with Mogul Mines Limited, and the successor, International Mogul Mines Limited, in joint venture with Coldstream Mines Limited, carried out prospecting, mapping, ground MAG and EM surveys, soil sampling, and trenching on the B3 and B4 claim groups. Eight diamond drill holes totaling 674.8 m (2,213.9 ft.) were drilled to test conductors in January 1969, resulting in the discovery of the B4-7 zone. The discovery hole, No. 69-5, intersected 8.26 m (27.1 ft.) of massive pyrrhotite-pyrite-chalcopyrite mineralization grading 0.80% Ni and 0.53% Cu. The B4-7 deposit was delineated by an additional 30 holes (6,850 m, or 22,479 ft.) in 1969. In the same campaign, eight holes for 628.2 m (2,061 ft.) explored other conductors on the property. A detailed MAG and EM survey was also completed over the deposit and petrographic work done on core at that time.

In late 1969, 136.1 kg (300 lbs) of drill core was composited from 71 assay rejects in 11 drill holes, split to 56.7 kg (125 lbs), and submitted to SGS for flotation recovery (metallurgical) testing, which included semi-quantitative spectrographic analysis for 30 elements. A manual tonnage/grade estimate for the B4-7 deposit was carried out, to total 2,282,520 tons (2,070,689 tonnes) averaging 0.87% Ni and 0.59% Cu (Zurowski, 1970). This historical estimate is not NI 43-101 compliant.

Coldstream Mines Limited acquired 100% of the property in 1970 and took two claims to lease in 1976.

In 1983-1986, Québec Cobalt and Exploration Limited staked part of the south portion of the Junior Lake property and carried out mapping, geophysics, and soil and rock sampling. Conwest Exploration Co. Ltd., the successor to Coldstream Mines Limited, optioned the leases covering the B4-7 deposit to Menacorp Limited in 1990, which resampled B4-7 core, and then to Minatco Exploration Ltd. in 1993.

In addition to the B4-7 deposit, exploration in the Junior Lake-Lamaune area prior to Landore work also revealed two low-grade Cu-Ni zones and occurrences of copper, iron, lithium, chrome, asbestos, zinc, and gold-molybdenite. Most of the occurrences are within two kilometres of the VW and B4-7 deposits.

Landore optioned part of the property from North Coldstream Mines Limited in 1998 and additional claims from Brancote Canada in 2000.

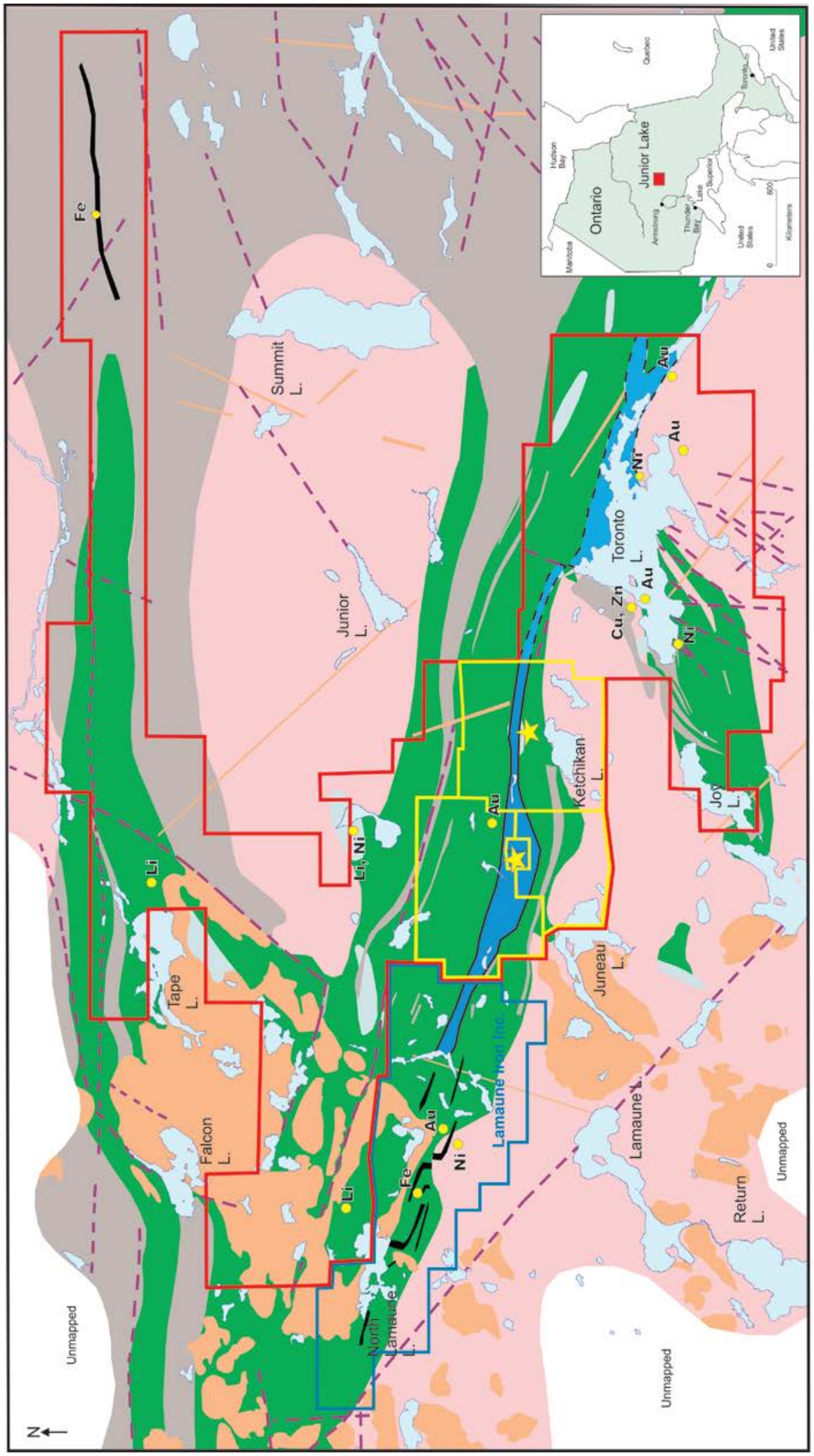
6 GEOLOGICAL SETTING

The regional, local and property geology has been for the most part summarized from Routledge, (2010), Lester (2009b), MacTavish (2004, 2004a), and Routledge (2006). Additional contributions are from various others, including Cooper (2009, 2014), Mungall (2009), and Pressacco (2013).

6.1 Regional Geology

The Junior Lake property is located within the Wabigoon Subprovince of the Superior Province of the Precambrian Shield and within the east-west trending Caribou-O'Sullivan greenstone belt. The belt is flanked to the south by the Robinson Lake Batholith of the Lamaune Batholithic Complex and to the north by a major, east-west-striking shear zone / terrain boundary that marks the southern limit of the English River Subprovince. Northeast of the property the belt is intruded by the elliptical, tonalitic to quartz dioritic Summit Lake Batholith. The western portion of the greenstone belt has been intruded by undulating, flat-lying, NeoProterozoic-age Nipigon diabase sills and localized dykes. These sills are the discontinuous, erosional remnants of laterally extensive sills comprising the Nipigon Plate which is centred on Lake Nipigon, approximately 30 kilometres to the south (MacTavish, 2004, 2004a). The regional geology of the Junior Lake property area is shown in Figure 6-1.

Figure 6-1: Junior Lake Regional Geology



6.2 Local and Property Geology

The supracrustal rocks and associated mafic to ultramafic intrusions of the Caribou-O'Sullivan greenstone belt are subdivided by Berger (1992) into the Archean-age Toronto and Marshall Lake groups. The two lithostratigraphic groups are similar in many respects; however, the Marshall Lake Group (MLG) contains a higher proportion of clastic metasedimentary rocks and apparently lesser amounts of mafic intrusive rocks.

The Toronto Lake Group (TLG) underlies the southern third of the Junior Lake property and consists of a bimodal assemblage of tholeiitic mafic flows and calc-alkaline rhyolitic to dacitic tuff, tuff breccias, and subordinate flows. The assemblage has been intruded by numerous mafic to ultramafic sills, dykes, and small stocks.

Four lithostratigraphic sequences defined within the TLG are as follows:

- The laterally extensive Carrot Top sequence trends west-northwest within the southern portions of the TLG and is comprised of magnetic talc-carbonatechlorite+/-tremolite schists derived from deformed and altered ultramafic rocks and clastic and chemical metasedimentary rocks. This sequence is 300 to more than 600 metres thick and hosts the D-Z iron occurrence, and several Ni-PGE (including Carrot Top and Zap Zone), Cu, Zn-Cu and Ag occurrences. Strong centimetre to metre scale folding is evident in the iron formation, and as such likely exists on a larger scale, possibly causing thickening and thinning along the main trends.
- The west-northwest trending Grassy Pond Sill intrudes the top of the TLG at its contact with the Marshall Lake Group (MLG) through the centre of the Junior Lake property. The Grassy Pond sill is a thick (100m to 500 metre wide), deformed, laterally continuous, gabbroic to locally anorthositic intrusive. The sill's most identifying characteristic is the presence of large (up to 10 cm in diameter) subhedral to euhedral plagioclase phenocrysts that often collect to form leucogabbro and anorthositic intervals of highly variable thicknesses. The Grassy Pond Sill hosts PGE, Cu and Ni occurrences, and is interpreted as being on the same geophysical structure as the B4-7 zone to the east.
- The B4-7 Sequence is a composite sequence, 1.9 kilometres long and up to 400 metres thick, of primarily mafic metavolcanic flows, intrusives and clastic and chemical metasediments that host the B4-7 Ni-Cu-Co-PGE deposit including the B4-7, Alpha and Beta Zones. The B4-7 sequence lies between the Carrot Top Sequence and the Grassy Pond Sill.
- The BAM Sequence is a composite sequence composed of mafic metavolcanic flows, mafic dykes and sills, and intermediate dykes. The BAM sequence is estimated to be 1.65 kilometres long and up to 160 m thick, possibly associated with an oblique structure. Archean Lamprophyre Dykes cut the TLG rocks.

In the north portions of the Junior Lake property, the Marshall Lake Group (MLG) includes tholeiitic, amphibolite mafic flows and calc-alkalic dacitic tuff, minor tuff breccias, and intercalated greywacke, chert and sulphide iron formation. Thin, discontinuous intermediate to felsic metavolcanic rock units also occur in the MLG. A higher portion of metasedimentary rocks and fewer mafic intrusives occur in the MLG compared to the TLG. Most of the rocks observed on the property are finely amphibolites, pillowed, mafic metavolcanic flows with well-defined pillow selvedge and a greater occurrence of plagioclase phenocrysts than observed within mafic flows south of the Grassy Pond Sill. Some outcrops exhibit an irregular, pervasive alteration, characterized by large, acicular actinolite porphyroblasts contained within a fine-grained matrix of chlorite, sericite, actinolite/tremolite, and epidote. This alteration is very similar to localized alteration observed within the Toronto Lake Group.

Pye (1968) interprets the presence of a large-scale fold on the western portion of the Junior Lake property southeast of Lamaune Lake and east-northeast-trending syncline in the vicinity of Toronto Lake to the east. The east-southeast trending, north-dipping North Lamaune Lake anticline is interpreted from magnetometer surveys tracing Iron Formation.

Structural Geology

Regional deformation rotated the supracrustal packages into near vertical orientation and developed a large west-northwest trending deformation zone (local portion referred to as the Junior Lake Shear Zone) north and west of Toronto Lake. This zone is the most prominent structural feature in the area and is characterized by narrow discrete zones of intensely sheared rock displaying dextral rotation separated by relative undeformed rock packages (Larouche, 1999). The deformation zone is evident as an aeromagnetic lineament which extends east and west of the Junior Lake property and appears to join the regional 450 km long Sydney Lake-Lake St. Joseph (SL-LSJ) Fault zone to the north, which also coincides with the boundary of the English River (ERT) and East Wabigoon subprovinces (EWT). The brittle-ductile fault zone of the SL-LSJ is steeply dipping, one to four kilometres wide, and is estimated to have accommodated about 30 km of right-lateral transcurrent displacement and 2.5 km of north vergent thrust movement (Percival, 2007).

A second, more local deformation in the east part of the property is confined to the supracrustal rocks around the periphery of the Robinson Lake Batholith, with deformation expressed as crenulation cleavage, northeast trending faults, and lineations which clearly post-date the regional deformation (Larouche, 1999).

Junior Lake Shear Zone and Associated Geology

Narrow, discrete zones of intense shearing (Junior Lake Shear Zone) form a corridor up to 800 m wide along the contact between the TLG and MLG. This shearing roughly follows the north contact of the Grassy Pond Sill. The evidence for the shear zone at Junior Lake is based on known geology and textures in drill holes and from limited exposures with deformation textures found from the micro to the macro level encompassing mylonites, cataclasites, sharp thin failure planes, and pressure-solution features such as stylolites. The widespread occurrence of pseudotachylite veinlets and infill demonstrates localized melting on failure planes.

Within the shear zone, the TLG is dominated by a large gabbro intrusive centred in the Grassy Pond Sill to VW area. It is a long linear intrusive and possibly split into several individual units. It is intruded into a mafic volcanic pile consisting of submarine pillow lavas and volcaniclastics. Cooper (2009) speculates that the gabbro has been the feeder for the volcanism and has then intruded its own lava pile.

Although the shear zone is slightly sinuous through Junior Lake, three of the mineral occurrences, Carrot Top, B4-7 and VW, fall on a straight line and Grassy Pond is only slightly to the north of this line. The length of the shear zone is uncertain, however, a length of at least 10 km has been defined. Along this length, there are variations in intensity with local domains of low deformation surrounded by high deformation zones as a result of competency contrast, general heterogeneity through the zone and lithology types. The rock succession in Junior Lake was deformed within a mobile greenstone belt and all geology became subvertical and with continued deformation within a deep ductile-regime, shear zones developed. During and post to shearing, gabbroic intrusive episodes occurred with a final pulse of very extensive vertical gabbro dikes. Major hydrothermal mineralizing events post-dated the gabbro dike swarm possibly as the result of heat from the post-tectonic sanukitoid style granites, such as high-Mg granitoid found in convergent margin settings (Cooper, 2009).

Less obvious at surface but no less voluminous are ultramafic lithologies such as peridotite, dunite, serpentinite, and their derivatives as talc dominated schistose metamorphic rocks. The ultramafic lava and/or intrusive suite was probably coeval with the basic suite but has suffered much more degradation of original texture and mineralogy within the mobile belt and shear environment. Variably textured granite and quartz diorite to tonalite gneiss and migmatite mapped along the south property boundary are part of the Robinson Lake Batholith.

Metamorphism

Metamorphism on the property is characterized by staurolite-cordierite-garnet, and rare sillimanite, in clastic metasediments; garnet-aluminosilicates-amphibole and rarely staurolite in the felsic and intermediate metavolcanic rocks; and garnet and amphibole in mafic meta-volcanic rocks. Most of the supracrustal rocks attained lower amphibolite grade metamorphic conditions, and greenschist grade metamorphism is only locally present (Larouche, 1999).

B4-7 DEPOSIT

The B4-7 Deposit is located in the south central area of the Junior Lake property. The B4-7 Deposit consists of polymetallic Ni-Cu-Co+PGE+Au mineralization hosted in massive sulphide (vein) and disseminated sulphides in a gabbro-basic volcanic setting coinciding with the Junior Lake shear zone. Strike length attains at least 600 m. Widths are up to approximately 18 m but are usually less than five metres. The B4-7 massive sulphide vein system appears to be a fairly simple dilational structure with marked pinch and swell in the vertical plane with an apparent plunge to the west. The mineralization was possibly introduced rapidly along pre-disposed failure planes under conditions of shearing. B4-7 consists of a continuous tabular body of semi-massive pyrrhotite-rich sulphides hosted in an assemblage of mafic volcanics and mafic intrusive. The contacts of the massive sulphide are typically very sharp and linear with minor wall rock contamination. Host rocks include leucogabbro, melanogabbro, gabbro, as well as mafic metavolcanics at the east end and may play an important role in hosting mineralization. Proximal to the sulphide mineralized zone are mafic schists, shear zones, metasedimentary rocks, locally iron formation, amphibolite, as well as pyroxenite, particularly at the east end.

The Alpha Zone occurs parallel to and approximately 50 m up dip of the B4-7 hanging wall in the gabbro and amphibolitized mafic volcanics. The zone, as described by MacTavish (2004a), is a broad envelope of narrow, moderate to high grade erratic, sulphide-rich carbonate and quartz carbonate veinlets and mineralized shears that generally trend oblique to stratigraphy. This hanging wall mineralization is localized in small fractures, foliations, and gashes and was probably introduced over a much longer time frame than the massive sulphides as the deformation progressed from foliation, through shear failure and into cataclasis. Disseminated sulphide mineralization consists of pyrrhotite, pyrite, and chalcopyrite, occurring as blebs and clots at the margin of gabbro intrusives against the host basic volcanics. Such mineralization is widespread but the better and more contiguous contact style mineralization occurs on the major gabbro contact. Thickness varies from <5 m to >15 m (Pressacco, 2013).

MacTavish (2004a) reported the ancillary Beta Zone, identified during 2001 drilling, as a net textured extension of the massive B4-7 zone and concluded that it was the probable magmatic source for the massive sulphides injected/emplaced into the fault/breccia zone.

B4-7 EAST AREA – Grassy Pond Sill

The B4-7 East Area is located immediately east of the B4-7 deposit and extends for at least 1.5 kilometres roughly along Landore's local grid baseline 00. This vicinity is dominated by gabbroic rocks ranging from anorthosite to gabbro cut by melagabbro dykes, interpreted to be the southern margin of the Grassy Pond sill. The geophysical anomaly (as seen in airborne magnetic, ground DCIP+MT, maxmin and vlf data) which delineates the B4-7 deposit extends eastwards through this area defining the boundary of the Grassy Pond sill with the mafic volcanic (typically mafic flows/pillow units) and metasedimentary rocks (commonly metapelites, with some cherty iron formation) located immediately to the south.

As interpreted by C. Cooper, the Grassy Pond sill is the largest of a cluster of gabbro sills in the centre of the Junior Lake greenstone belt. These sills are interpreted as palaeo-magma chambers which originally fed sub-aerial and submarine volcanoes with tholeiitic lava.

As reported by Cooper (2014):

The Grassy Pond is a sill or lopolith of basic to ultrabasic composition intruded into a basaltic lava and meta-sediment package that was possibly still presenting a high thermal gradient. The sill was most likely to have been horizontal or sub-horizontal at time of emplacement as it presents conformable contacts with the host rocks and is not particularly chilled near contacts. The sill is a composite intrusion consisting of several differing compositions but all in the proximity to the gabbro field. The feeder for the sill was a dyke or series of dykes (that may have been eroded and in fact a good part of the sill has also been eroded but we do not know how much). Composite magma chambers are sills kept molten by repeated magma supply by dykes. This is particularly true of magma chambers at divergent spreading centres such as mid-ocean ridges. It is possible that the sill could be the result of a magma chamber intruding its own lava sequence.

The feeder dykes are interpreted as the melanorite dykes seen on the long drill fence 1700E. They should be found south of the sill contact if they are true feeder dykes.

Considering all the known evidence so far it is likely that there was a primary genetic relationship between the basaltic lava piles and the Grassy Pond Sill and the latter is a fossil magma chamber within a volcanic pile at a spreading centre.

In the area of B4-7 the Grassy Pond sill is at its thickest and also most differentiated compositionally with a range from anorthosite to gabbro. Several of the smaller gabbro bodies may in fact be later dykes and the largest gabbro, the Grassy Pond Sill is more likely to be a complex nest of individual sills or an interfingered sill/host succession.

It can readily be appreciated that all significant nickel mineralisation is sited around the southern margin of the largest sill but other occurrences adjacent to some of the smaller sills cannot be ruled out.

The Grassy Pond sill, compared to other sills observed on the property, is the most differentiated in terms of composition. But the high degree of differentiation seems to be limited to the central, thicker part of the sill. This could be a critical point for the deposition of PGEs, and thus the development of Alpha PGE-enriched zones (delineated in the adjacent B4-7 deposit).

VW DEPOSIT

The VW Deposit lies 50 m to 150 m north of Ketchikan Lake near the southeastern end of the Junior Lake claim group on claim TB1077142. The deposit consists of a series of five mineralized subzones hosted by deformed assemblage of sheared mafic and ultramafic metavolcanics, gabbros, and chemical and clastic metasedimentary rocks. Host rocks are mafic volcanics and, to a lesser extent, mafic intrusives and metasedimentary rocks. The subzones are contained in a 125 m to 200 m wide shear (Junior Lake Shear) that dips steeply north. The VW deposit itself has been drilled over a strike length of 620 m, and dips subvertically or steeply to the north, with the deepest intersection in the most southerly subzone at 320 m (22 m elevation).

Stratigraphy of the VW deposit consists of a mixed sequence of mafic volcanics (2A, 2A F2), ultramafic volcanics (1A, 1C), mineralized volcanics (MZ 2A and 2A MZ), and mineralized gabbro (9C MZ) and gabbro (9C). This sequence youngs from south to north, with rock units described by McKay (2006) as follows:

- **Mafic-pillowed volcanics (2A F2)** are fine- to coarse-grained, grey, greenish grey, with locally well-developed pillow selvages up to three centimetres wide. Elongated, pale green to beige porphyroblasts occur locally. Chlorite alteration is moderate. Biotite alteration occurs locally. Quartz veins up to 50 cm wide are scattered throughout. The unit is locally sheared. Red-brown garnets occur but are rare. Sulphides consist of minor disseminated pyrrhotite.

- **Ultramafic volcanic (1A)** is fine- to medium-grained, steel grey, soapy and massive to well foliated. Serpentinization is pervasive. Chlorite, fuchsite and epidote alterations are moderate. White, locally beige, quartz veins up to 10 cm across are scattered throughout the unit.

- **Mafic volcanics (2A)** are fine- to coarse-grained, grey, greenish grey, foliated to massive. Chlorite and biotite alterations are moderate. Silicification is weak and patchy. Quartz veins up to seven centimetres across are scattered throughout. Sulphides, up to 1% pyrite and 0.5% chalcopyrite, occur predominantly in shear zones (?) or possibly poorly defined pillow selvages.

- **Mineralized mafic volcanics (2A MZ and MZ 2A)** are fine- to coarse-grained, grey, greenish grey, locally foliated to massive. Chlorite and biotite alterations are moderate to strong and epidote alteration is locally weak. Quartz veins are scattered throughout. The unit is sheared and fractured throughout. A 30 m intersection in hole 0405-36 contains poorly defined, very fine grained, cherty magnetite iron formation. The magnetite banding is up to 30 cm thick and rarely has sharp contacts with the hosting volcanics. The sulphide content of this magnetite iron formation is up to 15% pyrrhotite, 3% chalcopyrite, and 1% pyrite. The overall sulphide content of the unit is up to 5% pyrrhotite, 4% pyrite, and 2% chalcopyrite. In addition to the banding in the iron formation, the sulphides also occur as fracture filling, scattered blebs associated with quartz veins, and as disseminations.

- **Mineralized gabbro (9C MZ)** at the collar of hole 0405-35 is fine- to coarse-grained, grey, greenish grey and well foliated to massive. Biotite alteration and silicification are moderate to strong. Amphibole patches occur throughout. Moderate chlorite alteration occurs locally. Creamy white and beige quartz veining occurs throughout the unit. The unit is locally sheared, fractured and faulted, with minor gouge. Disseminated and blebby sulphides, up to 3% pyrite and 2% pyrrhotite, locally enhance the foliation.

VW WEST AREA

The VW West area is located immediately adjacent to the VW nickel deposit, and extends for approximately 1.5 kilometres towards the B4-7 Ni-Cu-Co-PGE deposit. Drilling activities in this area during 2010 and 2011 focused on a prominent east-west trending geophysical anomaly and several coincident EM conductors. Five fence lines were drilled at 200 metre intervals across a 1.5 km portion of the VW West area to test the potential for a westward strike extension of the VW deposit.

The area lithology resembles that of the VW deposit itself, dominated by sequences of ultramafic rocks with ultramafic-mafic flows. Additional units include metasediments and iron formation.

Lithology:

Cross section interpretations for VW West indicate a similar lithological and laterally correlative succession as the VW nickel deposit; with a series of upturned (near vertical) interbedded and fault bounded sequences of north facing ultramafic (peridotite) (as slices and larger lithon-like masses), juxtaposed(?) or as a stratigraphic succession against a suite of thinly interbedded ultramafic-mafic flows (locally biotitic or amphibolized) hosting pelitic and metavolcanic horizons and local sulphidic iron formation chemical sediments (chert). The stratigraphy is cut by gabbro and leucogabbro intrusions (locally porphyritic to anorthositic), and the latest gabbro intruded as very fine grained (mafic) dikes often emplaced as swarms. The stratigraphic sequence appears to repeat itself to the south of the magnetic high; but limited drilling tested this area. In the entire lithologic suite of VW West one particularly important unit of the VW deposit the 2Af1 "grey rock' (Cooper, 2011) was not noted in any appreciable abundance in the logging of drill core at VW West.

Structure and Alteration:

Numerous shear zones transect the drilled area sub parallel to the strike of the stratigraphy and in essence the entire width of much of the stratigraphy is highly strained to ductile deformed slices of varying but laterally correlative lithologic units. The well-bedded and hydrous nature of both the metavolcanic/pelite/BIF and olivine-rich peridotite act as slip planes in the stratigraphic stack and the deformation appears to be focused into these units deflected by the competency contrast of ridged lithons of either massive peridotite or massive gabbroic intrusions or dike swarms. Most strained zones are frequently accompanied by quartz-carbonate veining and alteration. The peridotite when thin or sandwiched between mafic and or pelitic or BIF is often intensely altered to carb-rich talc-schist.

Ophicarbonate is common in the upper part of the stratigraphic sequences (roughly coinciding with the northern limit of the magnetic high) and is hosted in highly strained talc-schist unit. The ophicarbonate horizon is of ultramafic aspect and is the result of carbonatisation of ultrabasic rocks and is directly associated with listvenites, [listwaenites]. These rocks are frequently associated with terrain boundaries and are a common feature in shear zones along terrain boundaries (Cooper, 2009). Pseudotachylite was observed in the upper portion of drill hole 0410-285 indicating fault movements involving mylonitization and/or partial melting. An approximate 30m true thickness of the structural zone hosting the ophicarbonates is defined in an east-west linear trend across the drill area.

Overall the stratigraphy is fairly linear along strike however a discontinuity is apparent in the magnetic data at line 1700E on the local grid. This discontinuity is apparent in the drilling as a pinching out of the massive peridotite and greater abundance of mafic flow along this section. Perhaps there is a NE-SW regional fault crossing the area here as well.

7 MINERALIZATION

7.1 B4-7 Deposit – Nickel, Copper, Cobalt, PGE, Gold

B4-7 deposit mineralization has been summarized from Pressacco (2013), with contributing information from MacTavish (2004).

Mineralization in the B4-7 zone is composed of semi-massive to massive sulphides primarily composed of pyrrhotite (25%-75%), with minor pentlandite and chalcopyrite (1%-5%). Other sulphides include pyrite, violarite, sphalerite, and covellite. Magnetite concentrations are minor (1%-5%) to moderate (5%-25%) and generally occur where oxide iron formation is proximal to the zone walls.

The pyrrhotite is fine grained and carries pentlandite and violarite exsolved as very fine grains (20 microns to 40 microns). Pentlandite occurs as relatively coarse grains as well as fine pentlandite flames. Nickel contents of up to 3% occur in the B4-7 lenses, whereas in the disseminated sulphides the nickel content is usually less than 1%.

The disseminated style of blebby and network style mineralization is, however, generally much richer in PGEs, with up to 10 g/t Pd and up to 1 g/t Pt. Both Cu and Co occur as sulphides in both styles of mineralization. PGE mineral species are represented by kotulskite and borovskite, both palladium tellurides.

MacTavish (2004) described the B4-7 zone mineralization as composed of massive, fine-grained pyrrhotite, highly variable disseminated, stringer, locally semi-massive, fine-grained chalcopyrite, and localized coarse-grained secondary pyrite. Most of the chalcopyrite is concentrated as fine to medium grains near contacts, within fragment-rich portions of the zone, or as veins and stringers within the adjacent hanging wall and footwall rocks. Pyrite occurs as coarse cubes, up to three centimetres in diameter, that usually concentrate near the upper and lower contacts of the massive sulphide zone. Cobalt content correlates directly with the amount of coarse-grained pyrite. The B4-7 zone is locally rich in disseminated magnetite, which, in most cases is correlatable to the presence of oxide facies iron formation in the adjacent hanging wall or footwall rocks.

7.2 B4-7 East Area – Nickel, Copper, Cobalt, PGE, Gold

Sulphide mineralization consisting of pyrite, pyrrhotite and chalcopyrite was observed in the gabbro units, with minor sulphides sometimes present in the metasedimentary rocks and to a lesser extent the mafic volcanic flows (though these were typically devoid of sulphides). These sulphides are disseminated through the host rock in relatively low abundances; trace to 2%, but in various intervals reached as high as around 20% as blebs, laminae, and veinlet infill semi massive and network with po the dominant sulphide typically. There were few instances of centimeter-scale high abundances (>50%) of pyrite and pyrrhotite (massive sulphide) but the dominant style of mineralization is as blebs and laminae.

7.3 VW Deposit – Nickel, Copper

VW deposit mineralization has been summarized by Routledge, R.E. (2010b):

There are three styles of sulphide mineralization in the VW deposit. The most important is thin lamina and veinlets following the foliation of the volcaniclastic rocks. The hydrothermal fluids appear to have been constrained by the impermeable gabbro dikes in that the highest concentrations of sulphides and grades are found immediately adjacent to dike contacts with a gradual diminution away from them.

The pyrrhotite, pyrite and chalcopyrite mineralized volcaniclastic host unit itself is usually thick, up to 25 m, but of lower overall grade than the mineralization ponded and channelled along the dikes. Mineralization in the centre of the Katrina zone may be of magmatic origin.

The third style is low grade mineralization up to 0.4% Ni occurring as fine blebs of pyrrhotite and pentlandite within ultramafic schist, peridotite and serpentinite.

The subzones within the VW deposit are composed generally of 1% to 5% sulphides consisting of pyrrhotite-pyrite-pentlandite-chalcopyrite-magnetite ± sphalerite. Locally sulphides can reach 40-50%, however, no massive mineralization has been noted. In addition to blebs/clots, lamina on foliation planes and veinlets, sulphides also occur to a lesser extent as breccia matrix, as replacement style net texture and as fracture filling. Weak to moderate sulphide mineralization at less than 2% sulphides and stronger mineralization containing more than 2% sulphides was distinguished in logging the mineralized zones.

Pyrrhotite is fine grained and carries minor pentlandite exsolved as very fine flames as well as occluded pentlandite as discrete fine grains. Free pentlandite appears to be rare. Chalcopyrite and pyrite occur as fine to medium grains.

7.4 VW West – Nickel, Copper, PGE, Chromium

Sulphide mineralization consisting of pyrite, pyrrhotite and minor chalcopyrite was observed in the gabbro, mafic volcanic and ultramafic units. These sulphides are disseminated through the host rock in relatively low abundances; trace to 2%, but in various intervals reached as high as 25% as stringers, bands, and veinlet infill semi massive to massive and network with po the dominant sulphide typically. There were few instances of centimeter-scale high abundances (>50%) of pyrite and pyrrhotite (massive sulphide) but the dominant style of mineralization is as stringers and wisps. The highest potential for Ni- bearing sulphide mineralization is that hosted in the mafic volcanic and peridotite units.

Up to 25% magnetite is disseminated through the mafic volcanic and ultramafic sequences, a contributing factor to the magnitude of the east-west trending geophysical anomaly through the area. Locally the magnetite has been shown to replace the chromium spinel and is indicative of progressive metamorphic intensity (Barnes, 1998). The degree of replacement of chromite by magnetite can be used as indicator of metamorphic grade. Locally the greenschist facies appears conducive to low grade Ni-Cr mineralization hosted in massive peridotite as reported from 0411-285, 0411-289, 0411-294, 0411-297 indicating continuity along strike.

Petrography of the mineralized peridotite confirms the presence of antigorite and chrysotile (fibrous asbestos) as up to 30% of the sample composition in the mineralized zones. Important associated sulphides include millerite, chromite, and magnetite after chromite (Kjarsgaard, 2011). A few rare species of supergene Ni-sulphides were noted and serve as an indication of late Ni enrichment. Results from drilling indicate Ni-Cr mineralization in the lower and least altered portions of the peridotite unit with up to 0.499% Cr over 1m (0411-289); and 0.18% Cr over 31.5m reporting 147.50m in 0411-297.

7.5 Mineralization Elsewhere on the Property

Prior to Landore ownership, exploration in the Junior Lake–Lamaune Lake area that located the B4-7 deposit in 1969 also revealed two low-grade Cu-Ni zones and occurrences of copper, iron, lithium, chromite, asbestos, zinc, and gold-molybdenite. Most of these are within two kilometres of the VW Zone.

From 1990 to 2003, Landore found nine PGE-Cu-Ni occurrences, one Cu-Pd zone, one gold zone, and Zn-Au-Ag and Zn-Co occurrences in old trenches and boulders bearing base and precious metal or arsenic mineralization. The VW deposit was discovered in 2005.

Four lithostratigraphic sequences favourable for nickel mineralization on the Junior Lake property have been identified by MacTavish (2004b) as follows:

- VW Sequence: a 1.9 km long, up to 400 m thick package of mafic metavolcanic flows, mafic intrusive dikes and sills, and clastic and chemical metasedimentary rocks that host the VW Zone.

- B4-7 Sequence: 1.9 km long and up to 400 m thick, is composed of primarily mafic metavolcanic flows (2AF1), gabbroic intrusive (9A,B,C), and clastic and chemical metasediments (6P) that lies between the Carrot Top Sequence and the Grassy Pond Sill. This sequence hosts the B4-7 Ni-Cu-Co-PGE deposit including the B4-7 massive sulphide zone and the Alpha and Beta zones.

- Grassy Pond Sill, a laterally extensive 100 m to >500 m thick gabbroic sill that hosts Cu-Ni-PGE mineralization near its base.

- Carrot Top Sequence: a complex laterally extensive 300 m to >600 m thick sequence of mafic metavolcanic flows, ultramafic schists, and clastic and chemical metasedimentary rocks that host several Ni-Cu-PGE occurrences. This sequence is located in the west portion of the Junior Lake property.

- BAM Sequence: a 1.65 km long, up to 165 m wide assemblage composed of mafic metavolcanic flows, mafic dikes and sills, and intermediate dikes that host the BAM gold occurrence. The BAM sequence is located northwest of the VW deposit in the north central portion of the Junior Lake property.

8 EXPLORATION

Cheatle (2010a) outlined the exploration history of the Junior Lake property:

Landore optioned part of the property from North Coldstream Mines Limited in 1998 and additional claims from Brancote Canada in 2000. Since then, Landore exploration has found nine PGE-Cu-Ni occurrences, one Cu-Pd zone, one gold zone, and Zn-Au-Ag and Zn-Co occurrences in old trenches and boulders bearing base and precious metals or arsenic mineralization. Landore has successfully delineated several deposits and other potential areas of significant mineralization throughout the Junior Lake property including two Ni+PGE deposits (B4-7 and VW).

Landore initial work in 2000 involved data compilation, Landsat image interpretation, prospecting, mapping, and resampling of the 1969 core, and followed up an Ontario Geological Survey (OGS) airborne EM and MAG survey flown over the area.

Ground magnetometer MaxMin II EM surveys, in addition to drilling, were completed in 2001. In 2003, Landore conducted drilling, stripping, trenching and channel sampling. All drilling data were digitized and reinterpreted, 856 core samples were assayed to fill in unsampled runs in the B4-7 deposit, in its hanging wall mineralization known as the Alpha Zone as well as in mineralization in the east extension of the B4-7 zone known as the Beta Zone.

A low level helicopter AeroTEM time-domain electromagnetic and magnetometer survey was flown in 2004. Principal geophysical sensors utilized in this survey included AeroQuest's AeroTEM© time domain helicopter electromagnetic system and a high sensitivity cesium vapour magnetometer. Bedrock EM anomalies were interpreted and graded according to the conductance.

The VW deposit was discovered in 2005 by follow-up prospecting of an AeroTEM conductor where 0.45% Ni was returned in a surface grab sample. Landore subsequently drilled the new VW deposit, as well as the Whale, NO and BAM zones, and other areas on the Junior Lake and Lamaune projects.

In 2006, Landore drilled the VW deposit, B4-7 zone, and other exploration targets including the Junior Lake, Pichette, and Lamaune claims. The 2006 campaign at the VW deposit included two surface trenches which were excavated and channel sampled. Metallurgical work included preliminary flotation and work indexes were carried out at Lakefield in September–October. Scott Wilson RPA also prepared a technical report (NI 43-101) on the B4-7 zone in 2006.

During 2007, diamond drilling of the VW and B4-7 deposits was the main focus of exploration activity. The following work was completed on the Landore property:

- Relogging of pre-2007 VW deposit drill core was initiated.
- Drill collars of the VW and B4-7 deposits and topographic control areas of the Junior Lake property were surveyed by an Ontario Land Surveyor.

- Minor line cutting was completed near Ketchikan Lake and the B4-7 deposit area to support the drilling operations.

- Baseline environmental studies were initiated and conducted by or under the guidance of Golder Associates Ltd. (Golder), of Sudbury, Ontario:
    - These studies were started in March 2007 and include quarterly sampling and analysis of lake and stream waters
    - Lake and stream sediment sampling was completed during the summer.
    - A benthic study, bathymetric study, and a fisheries study of Ketchikan Lake were completed.

- A weather station was installed at the Landore Junior Lake camp to record wind speed and direction, temperatures and three seasons of precipitation data.

- Sampling of the VW deposit drill core (quarter-cut core) was completed for metallurgical purposes.

- Claim lines were rehabilitated and the claim boundary surrounding an area to be leased was cut and surveyed in advance of filing the application to the Mining Recorder to lease the claims. Four leases have subsequently been granted.

- The land package was expanded to the southeast by staking an additional 24 claims totalling 5,056 ha.

- Aerial photography (stereo) was completed over the lease area by KBM Forestry Consulting in late 2007 to produce an air photo mosaic for exploration and infrastructure planning. The photographic data were processed to establish a detailed digital terrain topographic model (DTM).

- Golder commenced baseline aquatic studies in February 2007 on lakes and drainage tributaries in the vicinity of Junior Lake. These studies, repeated three monthly, are proceeding well and will continue through to economic studies. In addition, Golder completed a "Fish community and Fish habitat" survey of Ketchikan Lake, immediately south of the VW deposit, in addition to a bedrock resistivity survey on the northern side of the lake to determine depth of silt and evaluate bedrock competence.

- The camp was expanded and core storage was improved to hold the Junior Lake drill core on site.

- Core from previous Landore drilling in the VW deposit was relogged with a view to better understanding the controls on mineralization and identifying the disposition of mafic intrusives (dikes and sills) in the zone. In addition, further petrographic investigation was carried out on the VW deposit (Mungall, 2007). The drill hole collars were resurveyed to the Ontario base.

- In early 2007, a resource estimate was carried out by Scott Wilson RPA on the VW deposit.

In May 2008, Scott Wilson RPA prepared an updated resource estimate and NI 43-101 compliant technical report for the VW deposit. Scott Wilson RPA updated the VW deposit estimated resources to reflect 2008 to 2009 drilling and prepared a separate NI 43-101 compliant technical report (Routledge and Scott, 2009).

A non-NI 43-101 compliant mineral resource estimate to JORC standards was carried out by the Snowden Group (Snowden) on the B4-7 deposit in 2008. Scott Wilson RPA prepared resource estimates for the B4-7 deposit in 2006 and 2009.

Exploration efforts in 2009 included drilling, mapping and prospecting throughout the contiguous claims covering approximately 10 km^2, with work concentrated in the Lamaune Iron, BAM and VW areas. Additional exploration completed included prospecting and mapping at Swole Lake and Toronto Lake as well as east and west of the VW deposit.

To 2009, the VW deposit has been delineated and tested by 141 drill holes with 2,766 analyzed intervals over 2,838.36 m completed in the deposit subzones. Scott Wilson RPA has updated the VW deposit estimated resources to reflect 2008 to 2009 drilling and has prepared a separate NI 43-101 compliant technical report (Routledge and Scott, 2009).

Other exploration efforts in 2009 included mapping and prospecting throughout the contiguous claims covering approximately 10 km^2, with work concentrated in the Lamaune Iron, BAM and VW areas. Additional exploration completed included prospecting and mapping at Swole Lake and Toronto Lake as well as east and west of the VW deposit.

Overview of Recent Exploration

Recent exploration activity at Junior Lake from 2006 to 2014 has seen drilling focused on several areas including additional resource drilling at VW and B4-7 deposits, Lamaune area exploration drilling, the Whale Zone, Felix Lake, Swole Lake and B4-7 West and East exploration drilling.

Other recent work, in 2007-2014, included detailed geologic mapping (B4-7, VW, BAM, Lamaune), 55 trenches over approximately 13km (Lamaune Iron, Grassy Pond, Felix Lake, Juno Lake, BAM Zone, Toronto Lake), additional geophysical work (impulse EM survey, ground magnetic, and reinterpretation and integration with historic magnetic data), as well as approximately 70 km of line cutting. Regional scale prospecting, regional reconnaissance and geologic mapping, including an airborne geophysical coverage (AeroTEM electromagnetic and magnetic) of the Toronto Lake area (various Ni, Au, PGE potential), and Swole Lake (pegmatite lithium) prospecting were also undertaken. Numerous consultant reviews and studies have been completed, including detailed Scanning Electron Microscope (SEM) and petrography studies of the VW and B4-7 deposits; relogging, resampling and reinterpretation of geology for the VW, B47, and BAM sites; as well as reviewing of regional exploration potential. Surveying of drill collars, claim lines, additional claim staking, initiation of environmental baseline study, aerial photography, and metallurgical testing were also undertaken.

In June 2011, the Lamaune block, comprised of 23 claims, for 4,096 hectares, containing the Lamaune Iron deposit as well as the Lamaune Gold prospect, was transferred into a separate private company ('Lamaune Iron Inc.').

In October 2012, a deep penetrating ORION 3D 'Direct Current Induced Polarization' (DCIP) and Magnetotellurics (MT) survey was performed over the Scorpion zone of the Junior Lake property by Quantec Geoscience Ltd. This survey encompassed the western portion of the Scorpion zone, from line 1400W eastwards to line 400E in the B4-7 deposit.

Tuomi (2013) describes the 2012 DCIP+MT survey:

This survey acquired three sets of data in multi-directions; DC (direct current), IP (induced polarization) and MT (magnetotellurics), and is a true three dimensional survey. Sophisticated digital signal processing was utilized to obtain high resolution imaging at depths up to 1000+ metres below surface. This survey utilized DC resistivity to identify prospective nickel mineralization, and used IP chargeability to investigate potential copper and PGE targets.

The survey identified three areas of interest, located in the central, eastern and northern parts of the survey area, which appear to be interconnected and geologically controlled by fault lines. A portion of the eastern survey area is drill tested and hosts the B4-7 deposit.

The DCIP + MT survey results indicate that the conductive horizon which harbours the B4-7 massive sulphide mineralization extends to the west through the Exploration Target, an area identified west along strike and down dip from the B4-7 resource containing a potential 1.5 Mt to 2.0 Mt of sulphide mineralization of similar grade range to that which has been outlined to-date (Pressacco, 2013).

Additionally, DC resistivity results have identified potential sulphide mineralization along a 1.53 km wide corridor encompassing the western Scorpion zone. IP chargeability and MT results further support the potential of the Scorpion zone area for massive and disseminated sulphide mineralization.

Subsequent drilling in winter 2013 has tested the DC resistivity and IP chargeability results at various localities along the western portion of the Scorpion zone. Drilling in the Exploration Target area between lines 175W and 300W successfully intersected B4-7 massive sulphide mineralization as well as Alpha zone disseminated sulphide mineralization.

In December 2013, an Electromagnetic (MaxMin), VLF and Magnetometric ground geophysics program was completed over the VW deposit and VW West areas, from line 900E to line 4000E and covering 35.7 line kilometres. The survey was conducted by Geosig Inc., Québec, for Landore Resources.

Results from these surveys have been highly encouraging, identifying multiple near-surface conductor anomalies along the VW Nickel deposit trend with similar signatures to the VW deposit conductive anomaly itself.

From January 17 to February 14 2014, a 3-Dimensional (3D) Direct Current Induced Polarization and Magnetotellurics (DCIP + MT) ground geophysics program was completed over the VW deposit, VW West, and B4-7 East areas, from line 300E to line 3700E and from 700N to 1500S. The survey, covering 739.02 hectares, was conducted by Quantec Geoscience Ltd., Toronto, for Landore Resources. This survey is located directly adjacent to the east of Landore Resources' 2012 Orion 3D DCIP + MT survey block covering the B4-7 West zone.

The DC resistivity displays a good correlation with the B4-7 deposit while the IP chargeability model is best correlated with the VW deposit. Generally, across the survey grid the areas of the highest chargeability values offset the locations of the highest conductivities. This observation is consistent throughout the entire IP chargeability model when compared with the DC resistivity model. Considering the fact that the B4-7 mineralized zone is characterized as conductive target and the VW mineralized zone is characterized as chargeable target, it manifests that these two deposits are of different sources and depositional processes and should be interpreted accordingly. Therefore, for follow-up exploration program it is important to incorporate geological conceptual model for a given deposit type and use the corresponding DC or IP models as guideline for drill targeting (Gharibi/McGill, 2014).

Results from the 2014 survey have been highly encouraging, delineating nine significant new zones ranging from approximately 400m to 1,200m in length of potential nickel sulphide mineralization along strike and adjacent to the existing B4-7 Nickel-Copper-Cobalt-PGEs resource and the VW Nickel resource. Numerous targets are at depth and below existing Landore exploration drilling.

In the survey report, Gharibi/McGill (2014) notes:

The resistivity varies sharply over short lateral distance, indicating a complex subsurface dominated by faults and dike-like structures (Figure 8-1). The conceptual fault system drawn using the resistivity map at 50 m elevation indicates that these deformations are likely the result of a number of tectonic episodes, dominated by an E-W stress field system.

A number of high conductivity areas are resolved across the survey grid; anomalies DC1 through DC9 (Figure 8-2). The conductive anomalies display a general E-W elongated trend. Main conductivity feature, which is observed across the survey grid, is a large conductive body comprising anomalies DC1, DC2, DC3, and DC4 (Figure 8-2). The large conductive body appears to construct a large E-W elongated conductive structure and shows a number of lateral displacements in N-S direction. Displacements are evident along grid line L1300E between DC1 and DC2 and grid line L2000E between DC2 and DC3 and grid line L2800E between DC3 and DC4.

Additionally, resistivity cross-sections indicate that the B4-7 mineralized zone partially correlates with the anomaly DC1 to the south of the baseline (L0N). Existing drill holes in this area do not probe the possible deep extension of B4-7 to the east, south of the baseline L0N. The anomaly DC1 shows a drop of more than 100 m in depth in this area. This behavior is similar to a more than 100 m drop of the conductive anomaly associated with the B4-7 in the west, as observed in the 2012 Orion 3D survey over the same deposit. This suggests that the B4-7 deposit may be an uplifted conductive feature along a very long conductor with an E-W orientation, which is observed across the previous and present Orion survey areas. Therefore, the deep extension of anomaly DC1 south of gridline L0N resolved in the present survey is a potential target of interest to extend the existing B4-7 deposit eastward.

A number of large chargeability anomalies are resolved in the west and northwest of the survey grid (Figure 8-3, anomalies IP4, IP5, and IP6). At greater depth these three anomalies merge together and construct a large chargeable body with chargeability as high as 85 milliradians. The source of this large chargeable body is unknown but the scale and geometry of the anomaly suggest that it could be structural and formational IP responses.

Anomaly IP2 appears to correspond with the B4-7 deposit (Figure 8-3). This anomalous zone displays connectivity with large chargeability anomalies in the east.

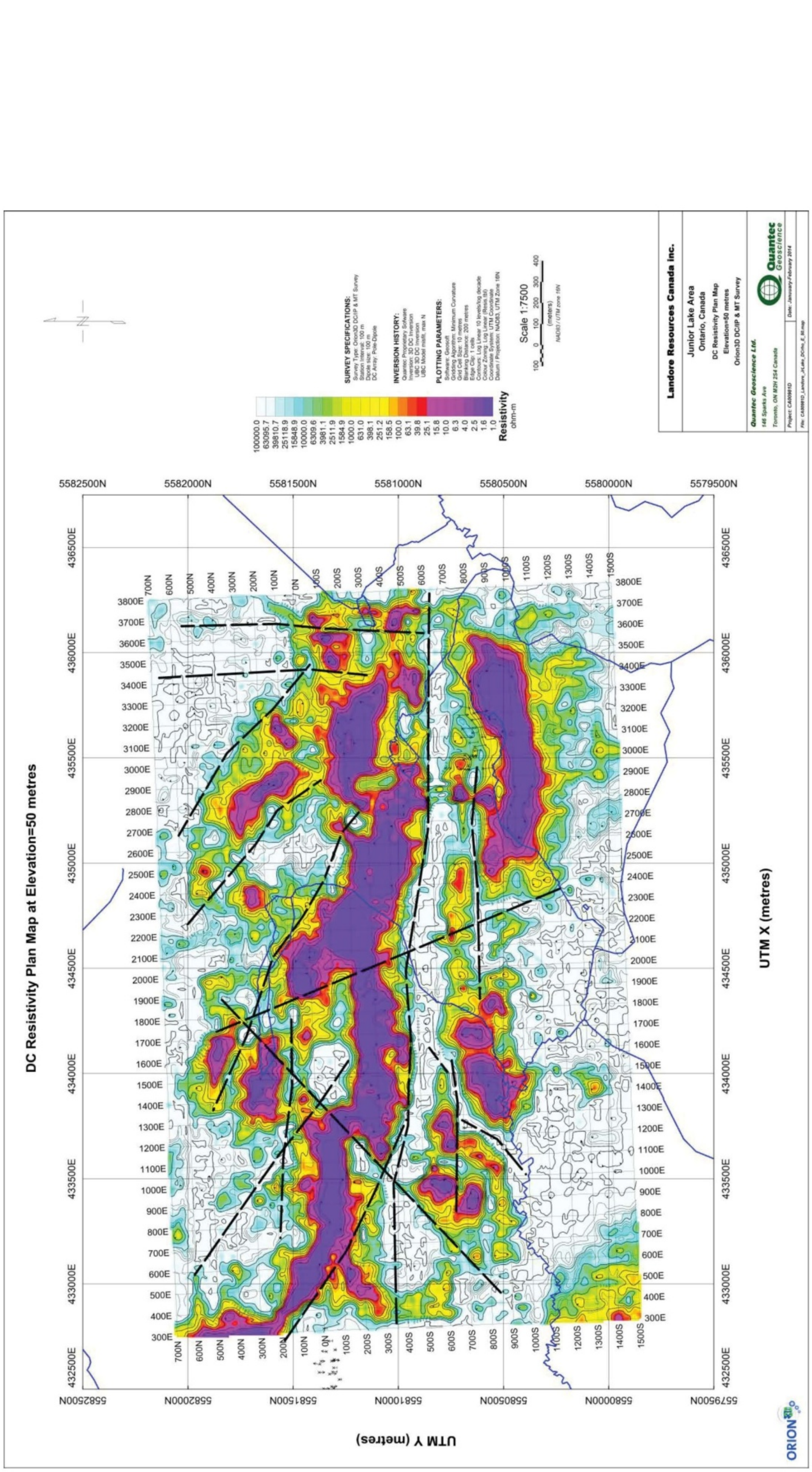
During July to August 2014, a total of 16 drill holes (0414-477 to 0414-492) for 4,201 metres were drilled on the B4-7 East area, a prospective zone extending eastwards 1.5 kilometres from the B4-7 Nickel-Copper-Cobalt-PGEs deposit. Drilling followed up on results from the 2014 3D DCIP + MT ground geophysical survey, which identified several prospective anomalies through this area.

Drilling confirmed the extension of a significant copper/gold mineralized structure previously delineated from line 900W in the B4-7 West zone eastwards adjacent to the B4-7. Values as high as 5.49% copper over 0.77 metres in drill-hole 0412-368 and 26.1g/t gold over 0.75 metres in drill-hole 0406-252 had been intersected by previous drilling. The drilling in July-August 2014 extended this copper/gold trend through to line 1600E, with a further 500 metres potential strike length indicated by historical drill hole S-5 on line 2100E with 0.21 metres at 8.97% Cu, giving an overall potential strike extension of 3.0 kilometres.

The drilling also intersected elevated polymetallic mineralization on line 1200E with drill-hole 0414-485 returning 5 metres at 0.25% Ni, 0.33% Cu, 0.01% Co, 497ppb Pd, 100ppb Pt, and 48ppb Au from 61 metres down-hole which together with previous highly encouraging trench results on line 1350E holds potential for near-surface, economic polymetallic mineralization which would provide added value to the B4-7 deposit.

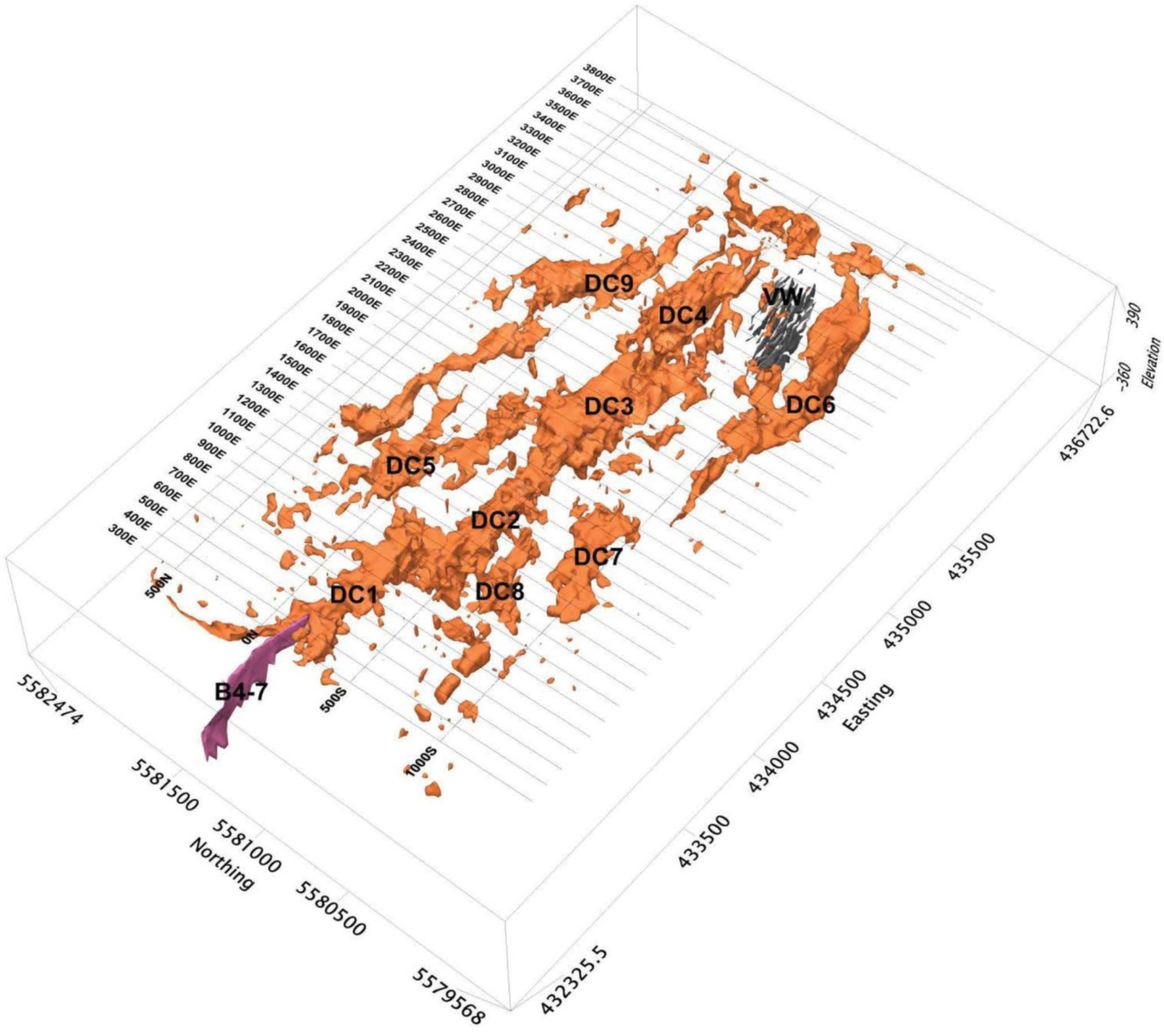
Further drilling is necessary to ascertain the scope of the near-surface polymetallic mineralization, as well as to further define the promising copper/gold mineralized structure.

Figure 8-1: DC resistivity plan at 50m elevation



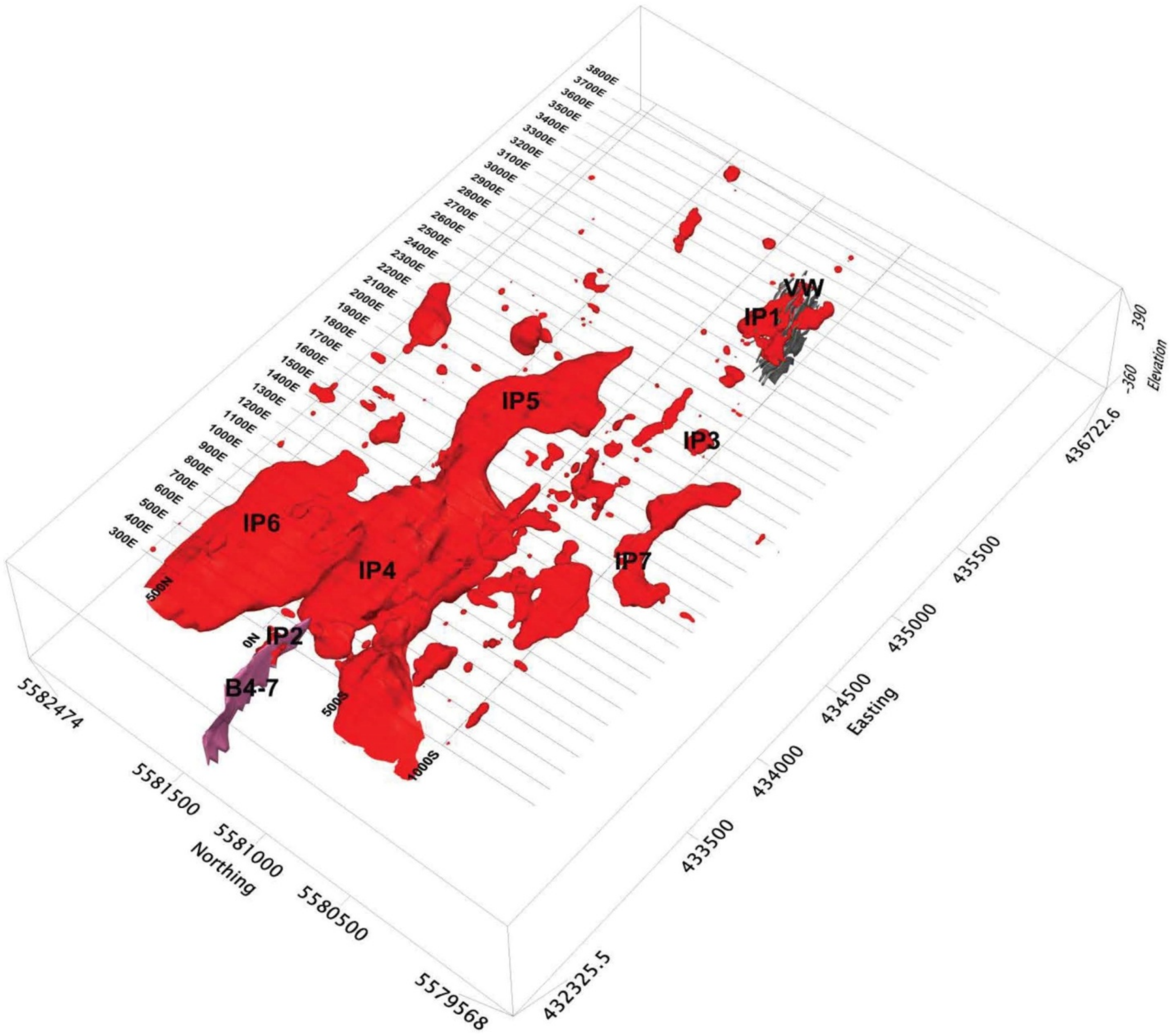
DC resistivity plan map at 50 m elevation with conseptual faults and structures.

Figure 8-2: 2 ohm-m DC Resistivity Isosurface



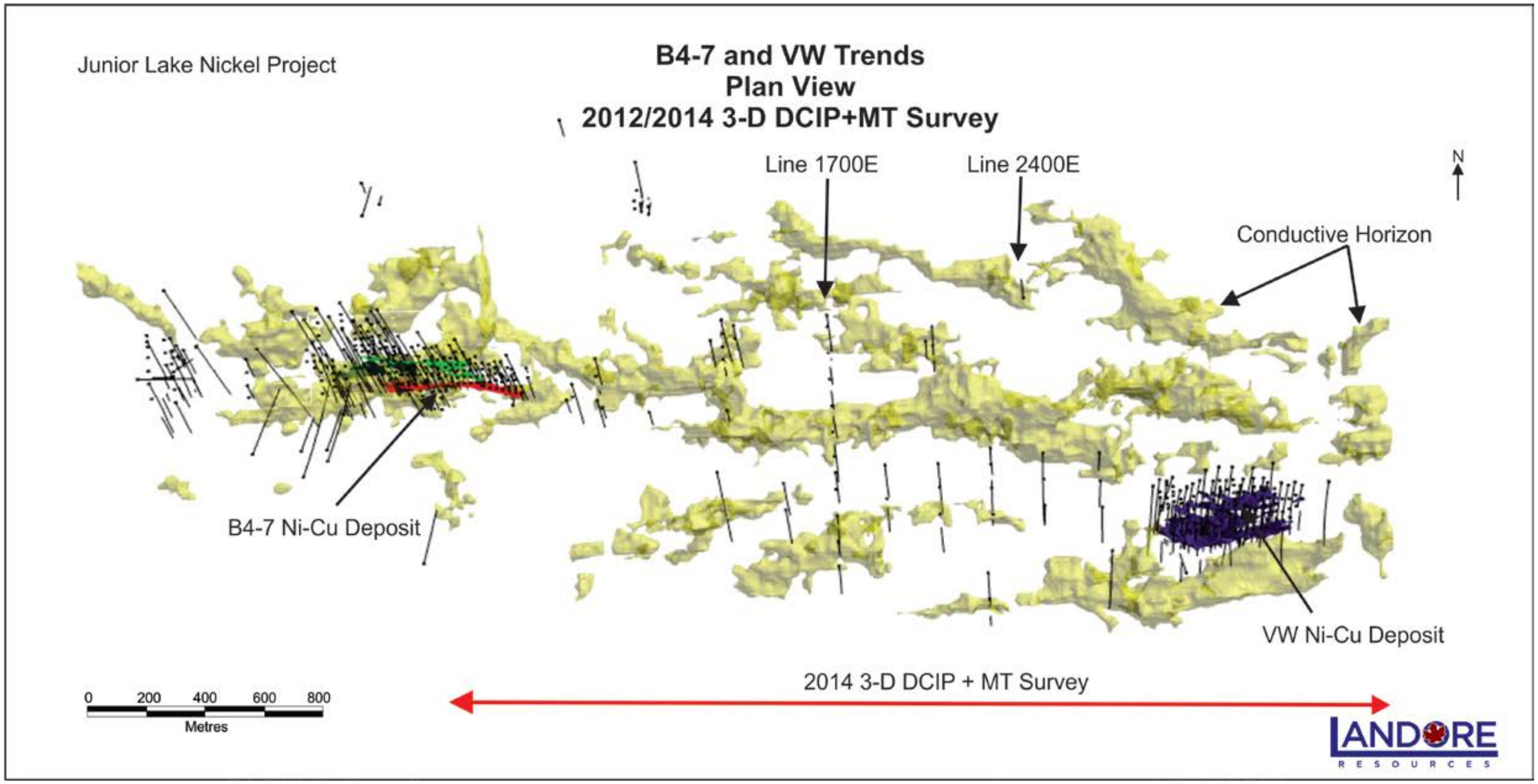
2 Ω·m DC Resistivity isosurface.

Figure 8-3: 70 milliradian Chargeability Isosurface



70 milliradian chargeability isosurface.

Figure 8-4: Plan View 1 ohm-m DC Resistivity Isosurfaces with Drilling



9 SURVEY DESIGN AND PROCEDURES

The fall 2014 exploration drilling program was conducted from September to October 2014. A total of 12 NQ size drill holes, composed of drill-holes holes 0414-493 to 0414-504 inclusive, for 2,675 metres, were drilled, logged and sampled. Drilling is summarized in Table 9-1.

9.1 2014 Fall Exploration Drilling Program (B4-7 Deposit, B4-7 East Area, VW Deposit, VW West Area)

Exploration drilling tested several geophysical anomalies identified by the 2014 DCIP+MT survey and the 2013 Electromagnetic (MaxMin), VLF and Magnetometric ground geophysical survey between the B4-7 deposit and VW Deposit and explored the base and precious metal mineral potential of various localities along this trend (Table 9-2, Appendix B and C).

Mineralized intersections of fall 2014 drilling include:

- DDH 0414-495: 0.34 metres at 1.21g/t Au, 280g/t Ag, and 2.25% Zn
- DDH 0414-503: 20.15 metres at 0.11% Ni, 0.04% Cu, 0.01% Co, 1.54g/t Pd, 0.64g/t Pt, and 0.01g/t Au
  including 0.55 metres at 3.39g/t Pd and 1.56g/t Pt
  including 0.56 metres at 10.90g/t Pd and 11.50g/t Pt
  including 0.72 metres at 12.85g/t Pd and 2.50g/t Pt
- DDH 0414-503: 3.46 metres at 0.62% Ni, 0.18% Cu, 0.05% Co, 0.29g/t Pd, 0.04g/t Pt, and 0.01g/t Au including 0.60 metres at 1.01% Ni, 0.09% Cu, 0.06% Co, 0.41g/t Pd, 0.03g/t Pt and 0.01g/t Au
- DDH 0414-504: 2.90 metres at 0.78% Ni, 0.14% Cu, 0.04% Co, 1.54g/t Pd, 0.06g/t Pt including 0.58 metres at 1.77% Ni, 0.20% Cu, 0.08% Co, 3.27g/t Pd, 0.09g/t Pt, 0.02g/t Au
- DDH 0414-504: 0.93 metres at 0.68% Ni, 0.22% Cu, 0.04% Co, 2.25g/t Pd, 0.70g/t Pt, 0.02g/t Au
- DDH 0414-504: 0.89 metres at 0.72% Ni, 0.81% Cu, 0.16% Co, 0.34g/t Pd, 0.04g/t Pt, 0.05g/t Au
- DDH 0414-504: 9.32 metres at 0.98% Ni, 0.50% Cu, 0.08% Co, 0.96g/t Pd, 0.24g/t Pt, 0.02g/t Au including 3.29 metres at 0.35% Ni, 1.59% Cu, 0.17% Co, 0.15g/t Pd, 0.13g/t Pt, 0.13g/t Au

B4-7 Nickel-Copper-Cobalt-PGE Deposit:
The drilling on line 00 intersected the down dip extension of the main B4-7 massive sulphide zone with drill-hole 0414-503 returning 0.60 metres at 1.01% Ni. The zone was intersected approximately 140 metres below the last drilled B4-7 main zone intercept and into one of the five geophysical resistivity anomalies identified in the 2012 Quantec survey.

Due to the pinch and swell nature of the B4-7 main vein, there is significant potential for a blowout of nickel mineralisation in the intervening 140 metres. Future planned drilling will target this area for infill and extension with the aim of expanding the existing 'Exploration Target' located west along strike and down dip from the B4-7 resource containing a potential 1.5 Mt to 2.0 Mt of sulphide mineralisation of similar grade range to that which has been outlined to-date.

Alpha Zone:
The Alpha zone, identified by drilling in 2001, consists of stringers, net-textured and disseminated sulphides containing significant palladium with elevated nickel, copper, cobalt, platinum and gold mineralisation and is found in both the open pit and the underground portions of the B4-7 deposit.

The Alpha zone mineralisation is interpreted to be a zone or zones located along the southern margin of the Grassy Pond Sill, a succession of inter-fingered sills composed of rocks ranging from coarse anorthosite to fine grained gabbro with a strike length of over 7 kilometres occurring through the B4-7 deposit to the VW deposit.

The Alpha zone mineralisation has not been included in the B4-7 National Instrument 43-101 (NI 43-101) compliant resource due to its complexity along strike.

Recent drilling and a review of existing data have established that the Alpha zone is far more significant than previously determined. Drilling on line 00 intersected significant Alpha zone mineralisation with drill-hole 0414-503 reporting 20.15 metres at 0.11% Ni, 1.54g/t Pd, and 0.64g/t Pt including 0.56 metres at 10.90g/t Pd and 11.50g/t Pt, and also including 0.72 metres at 12.85g/t Pd and 2.50g/t Pt. These platinum results are the highest reported on the Junior Lake property to-date, and the palladium results are among the highest, all of which are located within the B4-7 deposit.

Infill drilling on line 275E further established the continuity of near-surface Alpha zone polymetallic mineralisation with drill-hole 0414-504 returning 2.9 metres at 0.78% Ni, 0.14% Cu, 0.04% Co, 1.54g/t Pd, and 0.06g/t Pt including 0.58 metres at 1.77% Ni, 0.20% Cu, 0.08% Co, 3.27g/t Pd, and 0.09g/t Pt.

An in-house review of existing drilling on the B4-7 deposit is being conducted to identify potential Alpha zone mineralisation previously not sampled due to low sulphide content. PGE mineralisation is not necessarily tethered to sulphide content, as is the case with the recent high-grade PGEs intersected in drill-hole 0414-503.

Development of the Alpha zone can supply significant credits for the B4-7 deposit and form an integral part of the B4-7 Resource.

The B4-7 combined open pit and underground NI 43-101 compliant resource is 2,695,000 tonnes at 1.24% Nickel equivalent (NiEq) for 33,248 tonnes of contained metal, all within the Indicated category (2013 Mineral Resource estimate and a NI 43-101 Technical Report on the B4-7 Deposit, RPA, Toronto, Canada).

VW Deposit-Southwest:
The fall drilling campaign intersected highly encouraging near-surface precious metal mineralisation with drill-hole 0414-495 reporting 0.34 metres at 1.21g/t gold, 280g/t silver and 2.25% Zinc. This mineralisation, intersected in a sulphide-dominated vein located against a mafic dyke cutting through carbonate facies iron formation, represents a very similar mineralisation setting as the Lamaune Iron property located approximately 10 kilometres to the west-northwest. The Lamaune Iron property, owned by Lamaune Iron Inc., contains an initial 'exploration target' of between forty to fifty thousand ounces of gold at a cut-off grade of 0.3g/t.

Further exploration work is required to ascertain scope and continuity of this mineralisation.

Table 9-1: Summary of Drilling Campaigns at the Junior Lake Property

Landore Resources Canada Inc. - Junior Lake Project, Ontario				
Year	**Sector**	**# Drill Holes**	**No. Metres[1]**	**Drilled Holes**
1969	Exploration	8	720	S1 to S8[2]
1969	B4-7	31	6,941	69-5, 69-9 to 383[3]
1969	Exploration	7	583	69-1, 69-4, 69-6 to 8[3]
2001[4]	B4-7	21	5405	0401-07 to 24; 0401-01 to 03
2001	Exploration	3	600	0401-04 to 06
2003	B4-7	4	480	0403-07 to 10
2003	BAM	6	438	0403-01 to 06
2005	VW	15	4,730	0405-29 to 30; 0405-35 to 47
2005	Exploration	12	1,959	0405-25 to 34; 44, 45
2005[5]	Lamaune	17	2,599	1105-01 to 17
2006	VW	38	8,288	0406-48 to 64; 0406-71 to 88; 0406-97 to 98: 52A
2006	B4-7	7	1,562	0406-89 to 95
2006	Exploration	12.3	2,398	0406-61 to 70; 0406-96, 1506-01(part), well
2006	Lamaune	3.7	499	1106-18 to 20, 1506-01 (part)
2007	B4-7	16	3,580	0407-162 to 0407-177
2007[5]	VW	68	16,843	0407-99 to 161, 113A, 117A, 124A, 151A, 151B, 178
2008	VW	19	4,823	0408-179 to 195; 0407-114RE, 0407-136RE
2008	Exploration	4	795	0408-196 to 0408-199
2008	Lamaune	20	1,034	1108-21 to 40 Carrot Top/Zap Grassy Pond
2008	Lamaune	14	2,040	1108-41 to 54 Lamaune Iron
2009[6]	B4-7	44	9,286	0409-200 to 28; 0409-232 to 243; 0409-255 to 257
2009	VW	3	1,350	0409-229 to 231
2009	Exploration	12	2,277	0409-244 to 254, 258 (Whale Zone and B4-8 Zone)
2009	Lamaune	30	7,133	1109-55 to 83 (incl 59A), extension of 0408-41 to 1108-43 and 1108-53 Lamaune Gold, Iron
2010	Lamaune	69	10,605	1105-05ext, 1110-84 to 151 Lamaune Gold/Carrot Top Zone
2010	Exploration	27	4,422	0410-259 to 285 (Felix, West Ladle, VW West)
2011	Exploration	10	1,441	0411-304 to 313 (Swole Lake)
2011	Exploration	63	13,907	0410-285, 0411-286 to 0411-297 (VW West); 0406-69, 0411-298 to 0411-303, 0411-314 to 0411-357 (B4-8)
2011	B4-7	10	4,911	0411-358(A and B) to 0411-366
2012	B4-7	86	15,783	0412-367 to 0412-448, 0409-202RE, 0409-237RE, 0409-238RE, 0411-359RE
2012	Scorpion/B4-7	15	7,138	0412-449 to 0412-462, 0411-317RE
2013	Scorpion	14	5,778	0413-463 to 0413-476
2014	B4-7 East	16	4,201	0414-477 to 0414-492
2014	**Exploration, B4-7, VW**	**12**	**2,675**	**0414-493 to 0414-504**
	Total	**737**	**157,224**	

Notes:
1) Rounded to nearest metre.
2) AX core, 30.2 mm diameter.
3) BQ? core, 36.5 mm diameter.

4) Landore drilling 2001-2014 is all NQ core, 47.6 mm diameter.
5) Two holes deepened in 2008 campaign. Excludes 2008 abandoned holes.
6) Includes three metallurgical test sample holes not included in resource estimate.
7) The highlighted 2014 B4-7 East drilling forms the basis for this assessment report.

Table 9-2: Summary of 2014 Fall Exploration Drilling (B4-7 Deposit, B4-7 East Area, VW Deposit, VW West Area)

DDH	Start Date	Completion Date	Lease No	Final Depth (m)
0414-493	Sept 13, 2014	Sept 14, 2014	CLM461	101
0414-494	Sept 15, 2014	Sept 16, 2014	CLM461	100
0414-495	Sept 16, 2014	Sept 17, 2014	CLM461	115.06
0414-496	Sept 18, 2014	Sept 21, 2014	CLM461	280.50
0414-497	Sept 21, 2014	Sept 24, 2014	CLM461	331.61
0414-498	Sept 24, 2014	Sept 25, 2014	CLM461	100.53
0414-499	Sept 26, 2014	Sept 29, 2014	CLM461	346.5
0414-500	Sept 29, 2014	Oct 3, 2014	CLM461	400
0414-501	Oct 3, 2014	Oct 5, 2014	CLM461	100.91
0414-502	Oct 5, 2014	Oct 6, 2014	CLM461	129.05
0414-503	Oct 7, 2014	Oct 12, 2014	PA39128	537
0414-504	Oct 12, 2014	Oct 13, 2014	PA39127	123.09

9.2 Diamond Drilling Operations

Landore's Junior Lake camp, located at kilometre 105 on the East Road / Jackfish Road from Armstrong, was used as a base of operations. During dry seasons when access trail conditions permit, drill sites can be accessed by 4-wheel drive truck and all-terrain vehicle (ATV).

Drill holes were positioned and oriented by chaining from previous casings along cut lines of the established grid or by GPS and compass where there was no grid. Upon completion of each hole, the casing location was recorded using a Geneq Inc. SkyBlue II handheld Trimble GPS in UTM projection NAD 83 for Zone 16. All casings were left in the holes and capped. The water sources for this drilling were drill casing 0409-207, the north and east ends of Ketchikan Lake, and the west end of Ladle Flats.

Drilling was conducted by Chibougamau Diamond Drilling, of Chibougamau, QC. Drill core from this program is stored on covered core racks at Landore's Junior Lake camp.

Landore's core is stored at Junior Lake and is available for review.

9.3 Down-hole Surveys and Deviation

Down-hole deviation was minimized by the use of NQ size drill rods, hexagonal core barrel and long (18”) reaming shell.

Inclination deviation was monitored as the holes progressed using a Reflex Instruments EZ-Shot down-hole survey instrument and upon completion of each hole a Reflex Instruments Maxibor II instrument (optical method) was used to survey the hole to obtain reliable information on both inclination and azimuth deviation. Both instruments digitally record the down-hole survey data. Survey data is presented on the header page of each drill log in Appendix C.

9.4 Drill Core Logging Procedures

Drill core was aligned, measured and logged for geology. Logging records major and minor rock units (grain sizes, texture structural information: core angles of geological contacts, foliation and bedding, fractures, faults, veins, joints etc.), alteration and sulphide species, content and mode of occurrence. Logging and sampling information was recorded by hand on paper and/or in Microsoft Word and Excel software, then edited as required. Access and MapInfo GIS databases are maintained for drilling information. A copy of Landore’s geological legend is presented in Appendix F.

All drill core is digitally photographed and photos maintained on file in Landore’s Thunder Bay office.

Specific gravity (SG), RQD and magnetic susceptibility (MS) measurements of the mineralized zones and surrounding host rocks in the core were also recorded (Appendix E). The methodology for testing SG and magnetic susceptibility is summarized below:

9.4.1 Specific Gravity (SG) Methodology

- SG measurements were taken where there was visible mineralization, and at 3 metres intervals in select holes for background measurements.
- SG was measured utilizing a Denver Instrument Model PI-2002 scale, accurate to 0.01 gram. The scale was securely setup on a sturdy table, and levelled. A plastic weighting basket was suspended beneath the scale so that it is completely submerged in a pail of water (at room temperature) and then the scale is calibrated to read zero.
- The dry sample is weighted on the scale and the dry weight (DW) recorded. The sample is then placed in the basket, completely submerged in the water and the wet weight (WW) is recorded.
- All dry and wet weights are entered into an Excel spreadsheet and the specific gravity is calculated using the following formula:

  SG=DW/DW-WW.

9.4.2 Magnetic Susceptibility (MS) Methodology

MS measurements were taken where there was visible mineralization, and at 3 metres intervals in select holes for background measurements. MS was measured utilizing a Kappameter, model KP-6 magnetic susceptibility meter. The measurements were entered into an Excel spreadsheet either directly or after they had been recorded by hand on paper.

9.4.3 Rock-Quality Designation and Core Recovery

Rock-Quality Designation (RQD) and core recovery was determined over 3 metre intervals. RQD is calculated using the following formula:

RQD= (Sum of all pieces over 0.1m/ Metres recovered) *100

Core recovery is typically +80% except in rare cases over narrow intervals of highly sheared, foliated intervals. As such it is considered that samples accurately reflect drilled widths sampled. Core recovery is calculated using the following formula:

Core recovery= (Metres recovered/metres drilled)*100

Longest and smallest piece of drill core in the 3 metre interval was measured and recorded, as well as the fracture density. The fracture density is the visual inspection of the intensity of natural fractures in a given 3 metres, and is a numerical value on a scale of 0 to 9 (0 being no fractures, 9 being very intensely fractured). RQD data is available at Landore's Thunder Bay Office.

10 SAMPLING METHOD AND APPROACH

Sampling for the 2014 fall exploration drilling program has consisted entirely of drill core sampling. Cheatle (2010b) outlines the sampling methodology utilized by Landore:

> Core is logged and sampled in the Landore field camp on site, with occasional additional logging and sampling done on mineralized core in the Landore warehouse in Thunder Bay. Logging records major and minor rock units (grain sizes, texture structural information: core angles of geological contacts, foliation and bedding, fractures, faults, veins, joints, etc.), alteration and sulphide species, content and mode of occurrence. Geotechnical measurements including core recovery, rock quality designator (RQD) and fracture density have been taken. Specific gravity tests were carried out on. Sampling was conducted in the visibly mineralized zones and continuous sampling was undertaken throughout the hole length when the target was PGE's.
>
> Industry standard core sampling procedures were employed:
> - all drill core is aligned and measured prior to sampling
> - samples for assay are selected and marked for sampling on the basis of sulphide geology/mineralogy and rock units
> - sample intervals avoid crossing geological contacts
> - samples are sawn in half with a diamond saw blade
> - one half of the sample is placed in a standard, numbered transparent plastic bag with an identifying sample tag and
> - the remaining half returned to the core box with a corresponding tag placed at the beginning of the sample interval
> - the halved drill core is retained in core racks on site.
>
> All core sample bags are sealed with plastic sequentially numbered security tags and eight to ten of these sample bags are placed in larger rice bags also sealed with a numbered security tag. All security tag numbers are recorded prior to shipping and checked upon delivery at the lab.
>
> Sample intervals are typically 1.0m to 1.5m in length.

Only the gold, platinum and palladium are analyzed by fire assay (with AA finish). Nickel, copper, cobalt and silver are digested by aqua regia, then analyzed by AAS.

11 SAMPLE PREPARATION, ANALYSES AND SECURITY

Taken from Cheatle (2010b):

> Core samples are secured in the logging/sampling building at site. The samples are then transported directly from the site to the Accurassay or ALS Chemex lab in Thunder Bay by Landore or Chibougamau Diamond Drilling personnel. There have been no samples lost and no indications of sample tampering.
>
> Prior to 2007, Landore's Lamaune core was stacked outdoors on site with some mineralized intersections stored in a secure warehouse at Landore's office in Thunder Bay. New core racks were constructed on site during 2007 and stacked core was placed on the racks to improve its longevity, storage and accessibility.

11.1 ALS Chemex Laboratories Analytical Procedures

ALS Chemex is an independent, commercial mineral laboratory accredited by the Standards Council of Canada (SCC) under ISO 17025 guidelines. Each ALS lab has a Quality Management System (QMS) to ensure the production of consistently reliable data, and ensures that standard operating procedures are in place, and are being followed. The QMS is monitored by global and regional Quality Control teams. ALS participates in a number of proficiency tests, such as those managed by Geostats and CANMET.

The rock samples are first entered into ALS Chemex Laboratories Local Information System (LIMS), then bar-coded and weighed. The samples are dried, riffled split, then pulverized to better than 70% -2mm. Silica sand is used to clean out the pulverizing dishes between each sample to prevent cross contamination. The homogeneous sample then receives final preparation and analyzed as per the required methods. Assay results are checked by the lab manager before the hard copy is sent in the mail, and/or emailed to the client.

Analysis descriptions below are verbatim from ALS Chemex website: www.alsglobal.com

> *Aqua regia digestion:*
> *The standard aqua regia digestion consists of treating a geological sample with a 3:1 mixture of hydrochloric and nitric acids. Nitric acid destroys organic matter and oxidizes sulphide material. It reacts with concentrated hydrochloric acid to generate aqua regia:*
>
> $3\ HCl + HNO_3 = 2\ H_2O + NOCl + Cl_2$
>
> *Aqua regia is an effective solvent for most base metal sulphates, sulphides, oxides and carbonates.*

Atomic Absorption Finish:

In atomic absorption spectroscopy, an element in its atomic form is introduced into a light beam of appropriate wavelength causing the atom to absorb light (atomic absorption) and enter an excited state. At the same time there is a reduction in the intensity of the light beam which can be measured and directly correlated with the concentration of the elemental atomic species. This is carried out by comparing the light absorbance of the unknown sample with the light absorbance of known calibration standards.

A typical atomic absorption spectrometer consists of an appropriate light source (usually a hollow cathode lamp containing the element to be measured), an absorption path (usually a flame but occasionally an absorption cell), a monochromator (to isolate the light of appropriate wavelength) and a detector.

The most common form of atomic absorption spectroscopy is called flame atomic absorption. In this technique, a solution of the element of interest is drawn through a flame in order to generate the element in its atomic form. At the same time, light from a hollow cathode lamp is passed through the flame and atomic absorption occurs. The flame temperature can be varied by using different fuel and oxidant combinations; for example, a hotter flame is required for those elements which resist atomization by tending to form refractory oxides.

Lithium Borate fusion:

At ALS Chemex, lithium metaborate fusions are carried out in an automated fashion using a Claisse-type fluxer. The fusion melts can be poured into disks in preparation for X-ray fluorescence (XRF) analysis or they can be dissolved in acid for subsequent ICPMS analysis.

XRF:

In X-ray fluorescence spectroscopy, a beam of electrons strikes a target (such as Mo or Au) causing the target to release a primary source of X-rays. These primary X-rays are then used to irradiate a secondary target (the sample), causing the sample to produce fluorescent (secondary) X-rays. These fluorescent X-rays are emitted with characteristic energies that can be used to identify the nucleus (i.e. element) from which they arise. The number of X-rays measured at each characteristic energy can therefore in principle be used to measure the concentration of the element from which it arises.

The fluorescent X-rays are then dispersed and sorted by wavelength using a selection of different diffraction crystals, hence the term wavelength-dispersive X-ray fluorescence. The dispersed X-rays are then detected with a thallium-doped sodium iodide detector or a flow proportional counter. Each X-ray striking the detector causes a small electrical impulse which *can be amplified and measured using a computer-controlled multichannel analyzer. Samples of unknown concentration are compared with well-known international standard reference materials in order to define precise concentration levels of the unknown sample.*

Detection limits for the principal metals are:

Metal	Detection limit
Pd	10 ppb
Pt	15 ppb
Au	5 ppb
Ag	1 ppm
Cu	1 ppm
Ni	1 ppm
Co	1 ppm
Pb	1 ppm
Zn	1 ppm

11.2 Accurassay Laboratories Analytical Procedures

Accurassay is an independent, commercial mineral laboratory accredited by the Standards Council of Canada (SCC) under ISO/IEC 17025 guidelines for PGM, Cu, Ni, and Co analysis by atomic absorption spectroscopy (AA). The laboratory undergoes proficiency testing PTP-MAL through the SCC and participates in Round Robin testing through the Society of Mineral Analysts (SMA).

Accurassay Laboratories analytical procedures are as follows (Moore, J., 2008):

The rock samples are first entered into Accurassay Laboratories Local Information System (LIMS). The samples are dried, if necessary and then jaw crushed to -8mexh, riffle split, a 250 to 400 gram cut is taken and pulverized to 90%-150mesh, and then matted to ensure homogeneity. Silica sand is used to clean out the pulverizing dishes between each sample to prevent cross contamination. The homogeneous sample then receives final preparation and analyzed as per the analysis required require.

<u>Precious Metal Fire Assay:</u>

The sample is mixed with a lead based flux and fused for an appropriate length of time. The fusing process results in a lead button, which is then placed in a cupelling furnace where all of the lead is absorbed by the cupel and a silver bead, which contains any gold, platinum and palladium, is left in the cupel. The cupel is removed from the furnace and allowed to cool. Once the cupel has cooled sufficiently, the silver bead is placed in an appropriately labeled small test tube and digested using a 1:3 ratio of nitric acid to hydrochloric acid. The *samples are bulked up with 1.0mls of distilled deionized water and 1.0mls of 1% digested lanthanum solution. The total volume is 3.0mls. The samples cool and are vortexed. The contents are allowed to settle. Once the samples have settled they are analyzed for gold, platinum and palladium using atomic absorption spectroscopy. The atomic absorption spectroscopy unit is calibrated for each element using the appropriate ISO 9002 certified standards in an air-acetylene flame. The results for the atomic absorption are checked by the technician and then forwarded to data entry by means of electronic transfer and a certificate is produced. The Laboratory Manager checks the data and validates it if it is error free. The results are then forwarded to the client by fax, email, floppy or zip disk, or by hardcopy in the mail. NOTE: This method may be altered according the client's demands. All changes in the method will be discussed with the client and approved by the laboratory manager.*

Base Metals-Geochemical:
Base metal samples are prepped in the same was as precious metals but are digested using a multi acid digest (HNO_3, HF, HCl). The samples are bulked up with 2.0mls of hydrochloric acid and brought to a final volume of 12.0mls with distilled deionized water. The samples are vortexed and allowed to settle. Once the samples have settled they are analyzed for copper, nickel and cobalt using atomic absorption spectroscopy.

Base Metals-Full Assay:
Full assay samples are prepped the same way as geochemical base metals. They are weighed at 2.5g instead of 0.25g and digested using a combination of acids (nitric, hydrochloric and/or hydrofluoric). The samples are bulked up with 30mls of hydrochloric acid and brought to a final volume of 250mls with distilled deionized water using a 250ml volumetric flask. The samples are capped and inverted several times in the volumetric flask until the contents are homogeneous. A portion of the solution is transferred to a labelled test tube and then analyzed for the required elements using absorption spectroscopy.

In Landore's opinion, the sampling, assaying and security protocols, procedures and standards in place for the exploration drilling are industry standard and adequate for mineral resource and mineral reserve estimation.

12 DATA VERIFICATION

Drill hole and assay data entered or imported into Landore's Microsoft Access database is checked by the software and Senior Geologist for data entry errors.

To validate the drill hole database is checked for potential problems such as:

1) Intervals exceeding the hole length (from-to problem).
2) Negative length intervals (from-to problem).
3) Zero length intervals (from-to problem).
4) Inconsistent downhole survey records.
5) Out of sequence and overlapping intervals (from-to problem; additional sampling/QAQC/check sampling included in table).
6) No interval defined within analyzed sequences (not sampled or missing samples/results).

12.1 Quality Control and Quality Assurance

Upon receiving assay results, Landore checks that all standards and blanks are within +/- 3 standard deviations from their certified mean. Landore has in place and follows a standard procedure to ensure that failed assay batches are re-run.

Certified standards used include various standards from Geostats Party Ltd, Australia. Also, certified standards from CDN Resource Laboratories Ltd. were used.

The silica sand blank was obtained from ALS Chemex laboratory in Thunder Bay, Ontario.

The base metal standards are inserted every 20th submitted sample. A precious metal standard is inserted in every sample batch.

The silica sand blank was inserted every 20th submitted sample. Landore ensured that at least 2 standards and 1 blank were placed in every batch.

As part of the QAQC regimen, rejects and split pulps for 5% of the samples (selection at geologist's discretion) are submitted to Accurassay (with one portion of the split pulps going to ALS) for confirmation. Original assay results are reported unless the check assay results question the original assays. In addition to this, other results that may be questionable (i.e. low value amongst high values) are check assayed.

12.1.1 ALS Chemex Quality Control

ALS employs an internal quality control system that tracks certified reference materials and in-house quality assurance standards. ALS uses a combination of reference materials, including primary, certified reference, or in-house reference materials. Should any of the standards not fall within an acceptable range, re-assays will be performed with a new certified reference material. The number of re-assays depends on how far the certified reference material falls outside its acceptable range. Additionally, ALS verifies the accuracy of any measuring or dispensing device (i.e. scales, dispensers, pipettes, etc.) on a daily basis and is corrected as required.

12.1.2 Accurassay Quality Control

Accurassay Laboratories employs an internal quality control system that tracks certified reference materials and in-house quality assurance standards. Accurassay uses a combination of reference materials, including reference materials purchased from CANMET, standards created in-house and tested in round robin analyses with laboratories across Canada, and ISO certified calibration standards purchased from suppliers. Should any of the standards fall outside the warning limits (mean $\pm 2\sigma$), re-analysis is performed on 10% of the samples analyzed in the same batch and the new values are compared with the original values. If the values from the re-analysis match original assays, the data is certified. If they do not match, the entire batch is re-analyzed. Should any of the analyses for standards fall outside the control limit (mean $\pm 3\sigma$), all analyses in that batch are rejected and all of the batch samples are re-analyzed prior to returning results to Landore.

Accurassay also re-assays every 10^{th} sample as a duplicate and inserts a blank control sample in the batch as part the internal laboratory QA/QC process.

13 INTERPRETATION AND CONCLUSIONS

The 2014 fall exploration drilling program was completed over the B4-7 Deposit, the B4-7 East area, the VW Deposit and the VW West area. Drilling followed up on promising geophysics anomalies from the 2014 DCIP + MT survey together with the 2013 Electromagnetic (MaxMin), VLF and Magnetometric ground geophysical survey. Drilling investigated the base and precious mineral potential at various localities eastwards between the B4-7 deposit and VW deposit, a trend which is interpreted to be highly prospective for further economic mineralisation.

The drilling program was successful in confirming a down dip extension of the main B4-7 massive sulphide zone 140 metres below the existing B4-7 resource on line 00 as well as intersecting high grade Alpha zone platinum group element (PGE) mineralisation. Drilling on the B4-7 deposit further intersected shallow high-grade nickel and Alpha zone PGEs on the eastern portion of the deposit (line 275E). Exploration drilling southwest of the VW deposit discovered significant near-surface precious metal mineralisation.

More drilling is required to delineate additional mineralization and further test the high potential of this area as indicated by geophysical anomalies from recent surveys.

14 RECOMMENDATIONS

The 2014 fall exploration drilling program was successfully completed on the area between the B4-7 Ni-Cu-Co-PGE deposit and the VW Ni deposit, a distance of three kilometres. Drilling tested several geophysical anomalies identified by the 2014 DCIP+MT survey and the 2013 Electromagnetic (MaxMin), VLF and Magnetometric ground geophysical survey and explored the base and precious metal mineral potential of various localities along this trend.

B4-7 deposit drilling intersected the down dip extension of the main massive sulphide zone approximately 140 metres below the last drilled B4-7 main zone intercept and into one of the five geophysical resistivity anomalies identified by the 2012 Quantec survey. Drilling also intersected significant Alpha zone mineralisation with the highest reported platinum results on the Junior Lake property to-date, and palladium results among the highest. Additional drilling is required to expand the B4-7 resource, and to further establish continuity of the Alpha zone to bring in to the formal resource.

The fall drilling campaign also intersected highly encouraging near-surface precious metal mineralisation southeast of the VW deposit in a similar setting as other known gold occurrences in the region, such as the Lamaune Iron property located approximately 10 kilometres to the west-northwest. Further exploration work is required to ascertain scope and continuity of this mineralisation.

More work is warranted to further test the base and precious metal potential between the B4-7 and VW deposits which harbours promising geophysics anomalies untested to-date by drilling.

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16 SIGNATURE PAGE

This report titled "Work Assessment Report on the Junior Lake Property – 2014 Fall Diamond Drill Program (B4-7 Deposit, B4-7 East Area, VW Deposit, VW West Area) – September 9, 2016" was prepared by M. Tuomi and signed by the following Author:

Michele Tuomi, P.Geo.
Landore Resources Canada Inc.

Thunder Bay, Ontario
September 9, 2016

17 CERTIFICATE OF QUALIFIED PERSON

Michele Tuomi, P.Geo.
Landore Resources Canada Inc.
555 Central Avenue, Suite 1
Thunder Bay, ON
P7B 5R5

Tel: +1 807 623 3770

I, Michele Tuomi, am a Professional Geoscientist, employed as a VP Exploration of Landore Resources Canada Inc.

This certificate applies to the geological report titled "Work Assessment Report on the Junior Lake Property – 2014 Fall Diamond Drill Program (B4-7 Deposit, B4-7 East Area, VW Deposit, VW West Area)" dated September 9, 2016.

I am a member of the Association of Professional Geoscientists of Ontario. I graduated with a BSc. degree in Geology from Lakehead University in 1999.

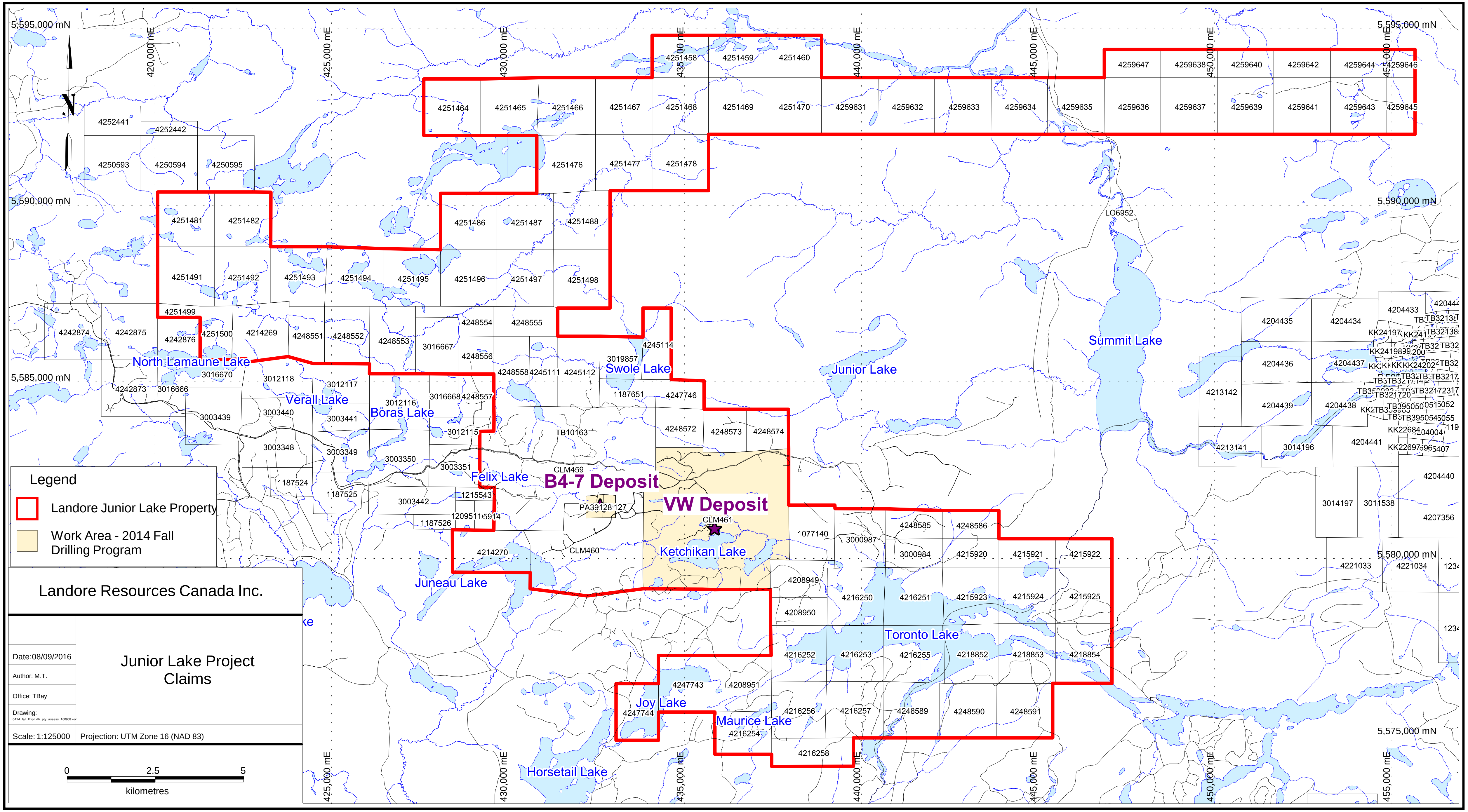
I have practiced my profession for 16 years. I have been directly involved in mineral exploration and mineral project assessment, as well as mineral resource estimations.

I have visited the Junior Lake property in northern Ontario, Canada on numerous occasions, the most recent being August 15, 2016.

I am responsible for all items of the assessment report "Work Assessment Report on the Junior Lake Property – 2014 Fall Diamond Drill Program (B4-7 Deposit, B4-7 East Area, VW Deposit, VW West Area) – September 9, 2016".

As of the date of this certificate, to the best of my knowledge, information and belief, the technical report contains all scientific and technical information that is required to be disclosed to make the assessment report not misleading.

Michele Tuomi, P.Geo.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Landore Resources Canada Inc.
Junior Lake Project Claims
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Author: M.T.
Office: TBay
Drawing:
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Projection: UTM Zone 16 (NAD 83)
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2.5
5
kilometres
N
B4-7 Deposit
VW Deposit
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Verall Lake
Boras Lake
Swole Lake
Felix Lake
Ketchikan Lake
Juneau Lake
Junior Lake
Summit Lake
Toronto Lake
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Maurice Lake
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Legend
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Diamond Drill Hole
MaxMin Conductor
VLF Anomaly
Cut grid line
Landore Junior Lake Property
Lease/Claim
Road/Trail
Lake
Stream
Landore Resources Canada Inc.
Map 1
Junior Lake Property
B4-7 Deposit to VW Deposit
2014 Fall Drilling Program
Date: 09/09/2016
Author: GIS
Office: TBay
Scale: 1:5000
Projection: UTM Zone 16 (NAD 83)
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metres
Magnetic Declination 2013
5°33' West
Landore Exploration Camp
CLM459
CLM460
CLM461
PA39127
PA39128
Ladle Flats
Ketchikan Lake
TL 600N
TL 300N
TL 300S
TL 600S
0414-493
0414-494
0414-495
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0414-502
0414-503
0414-504

Legend
2014 Diamond Drill Hole submitted for assessment credit
Diamond Drill Hole
MaxMin Conductor
VLF Anomaly
Cut grid line
Landore Junior Lake Property
Lease/Claim
Road/Trail
Lake
Stream
Landore Resources Canada Inc.
Date: 09/09/2016
Author: GIS
Office: TBay
Map 2
Junior Lake Property
B4-7 Deposit to VW Deposit
2014 Fall Drilling Program
Mag 1vd Base
Scale: 1:5000
Projection: UTM Zone 16 (NAD 83)
0 75 150 300
metres
Magnetic Declination 2013
5°33' West
Landore Exploration Camp
TL 600N
TL 300N
TL 300S
TL 600S
CLM459
CLM460
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PA39127
PA39128
Ladle Flats
Ketchikan Lake
0414-493
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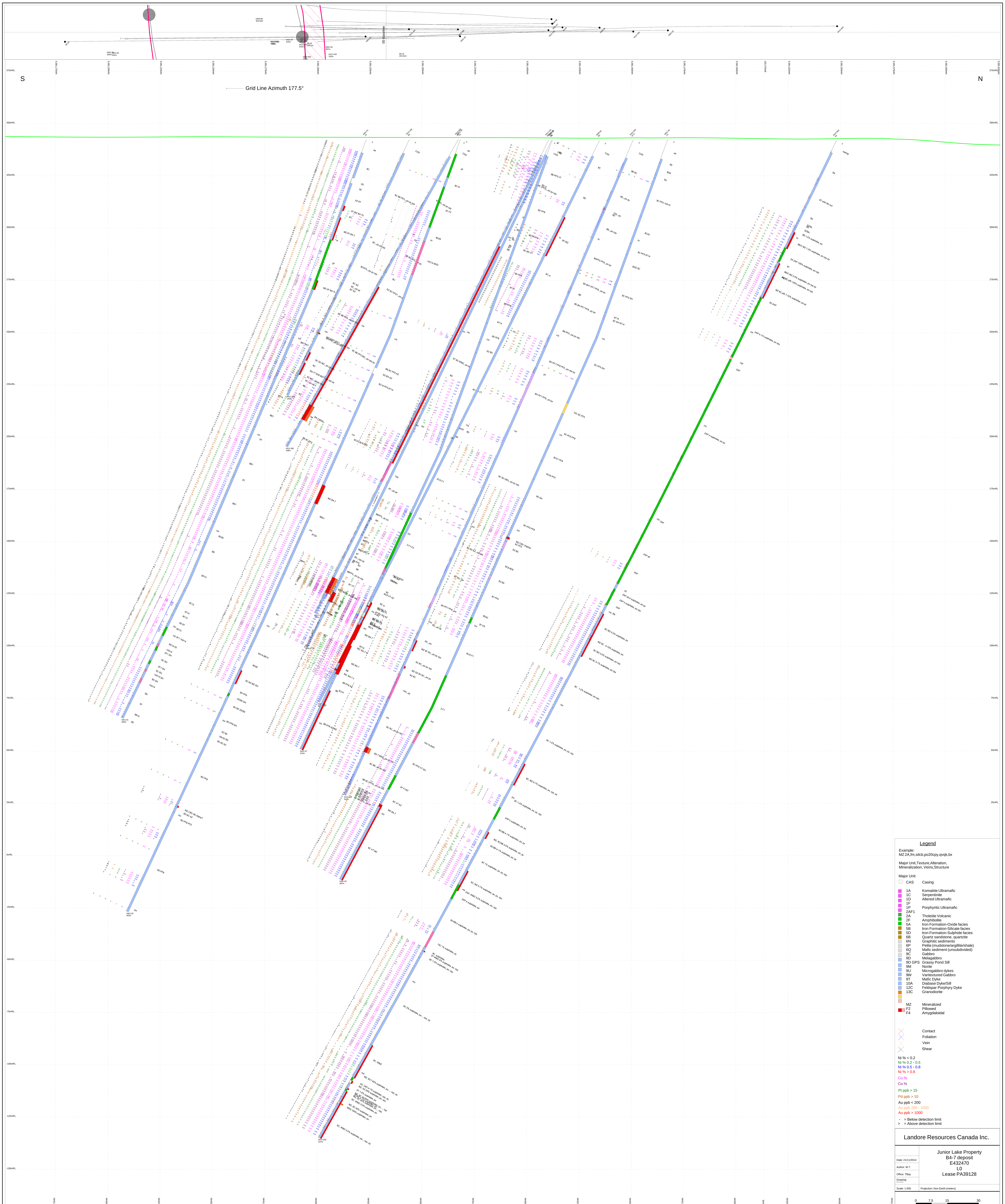
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Major Unit,Texture,Alteration,
Mineralization, Veins,Structure
Major Unit
CAS Casing
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1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
> = Above detection limit
Landore Resources Canada Inc.
Junior Lake Property
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Author: M.T.
Office: TBay
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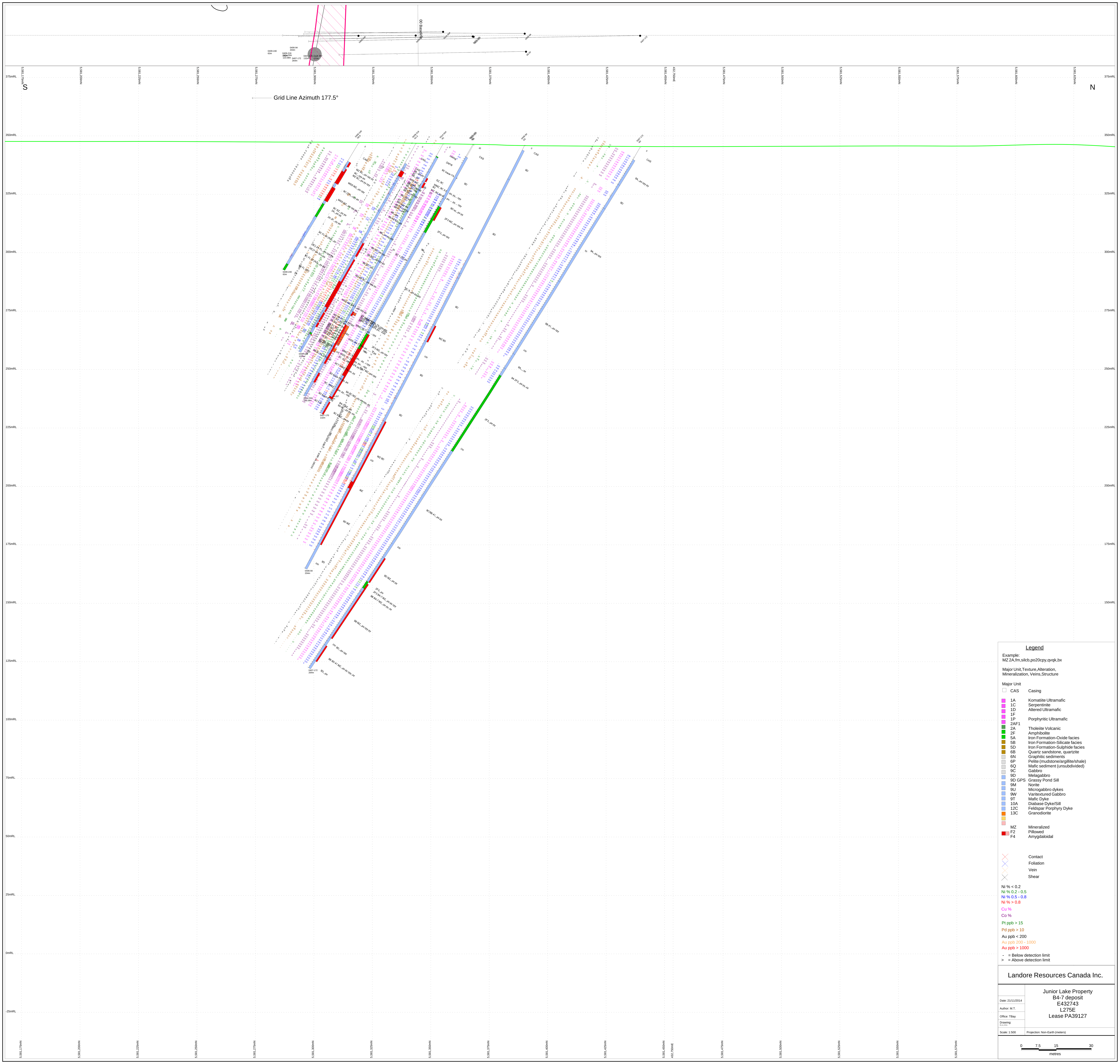
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Grid Line Azimuth 177.5°
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,
Mineralization, Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
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6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
> = Above detection limit
Landore Resources Canada Inc.
Junior Lake Property
B4-7 deposit
E432743
L275E
Lease PA39127
Date: 21/11/2014
Author: M.T.
Office: TBay
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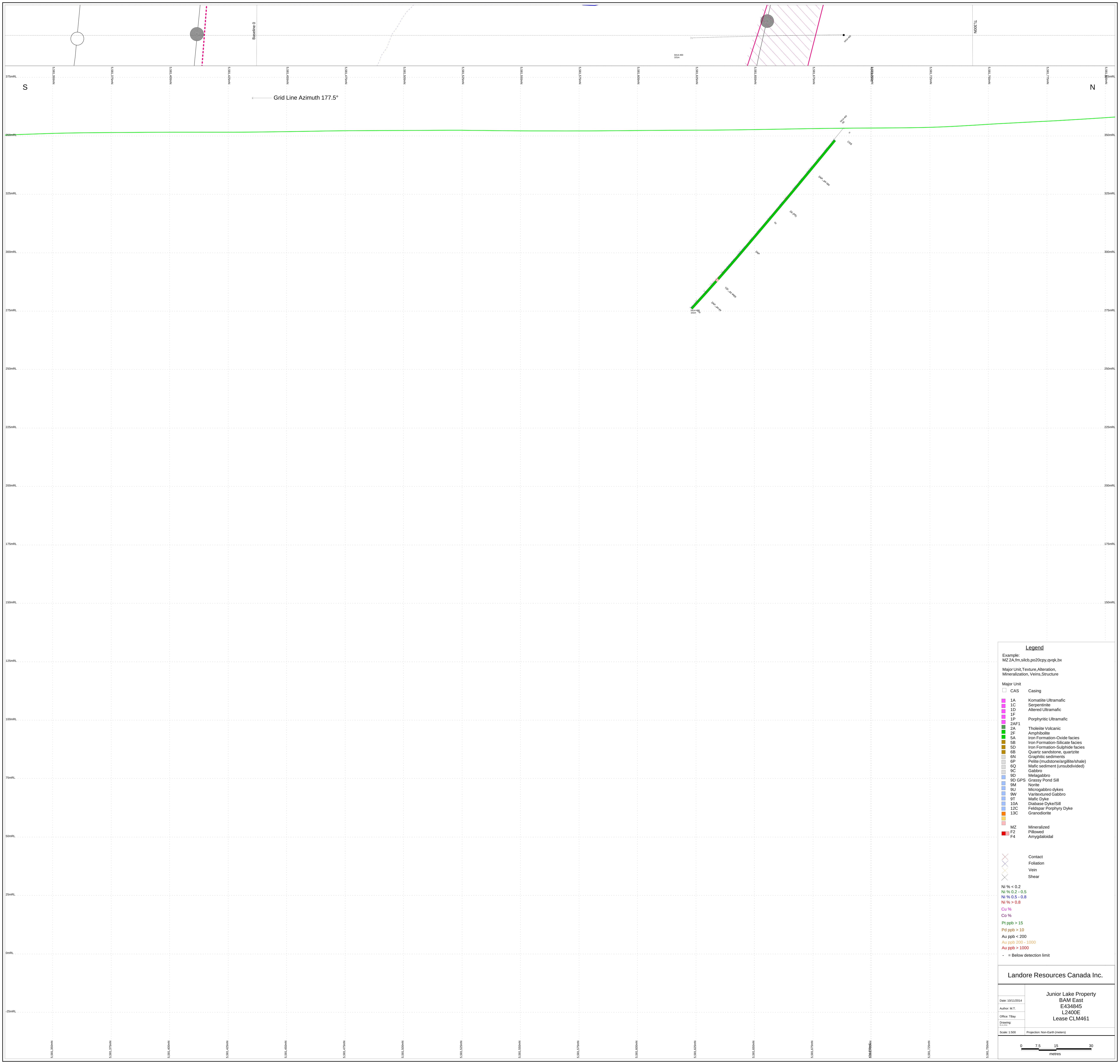
metres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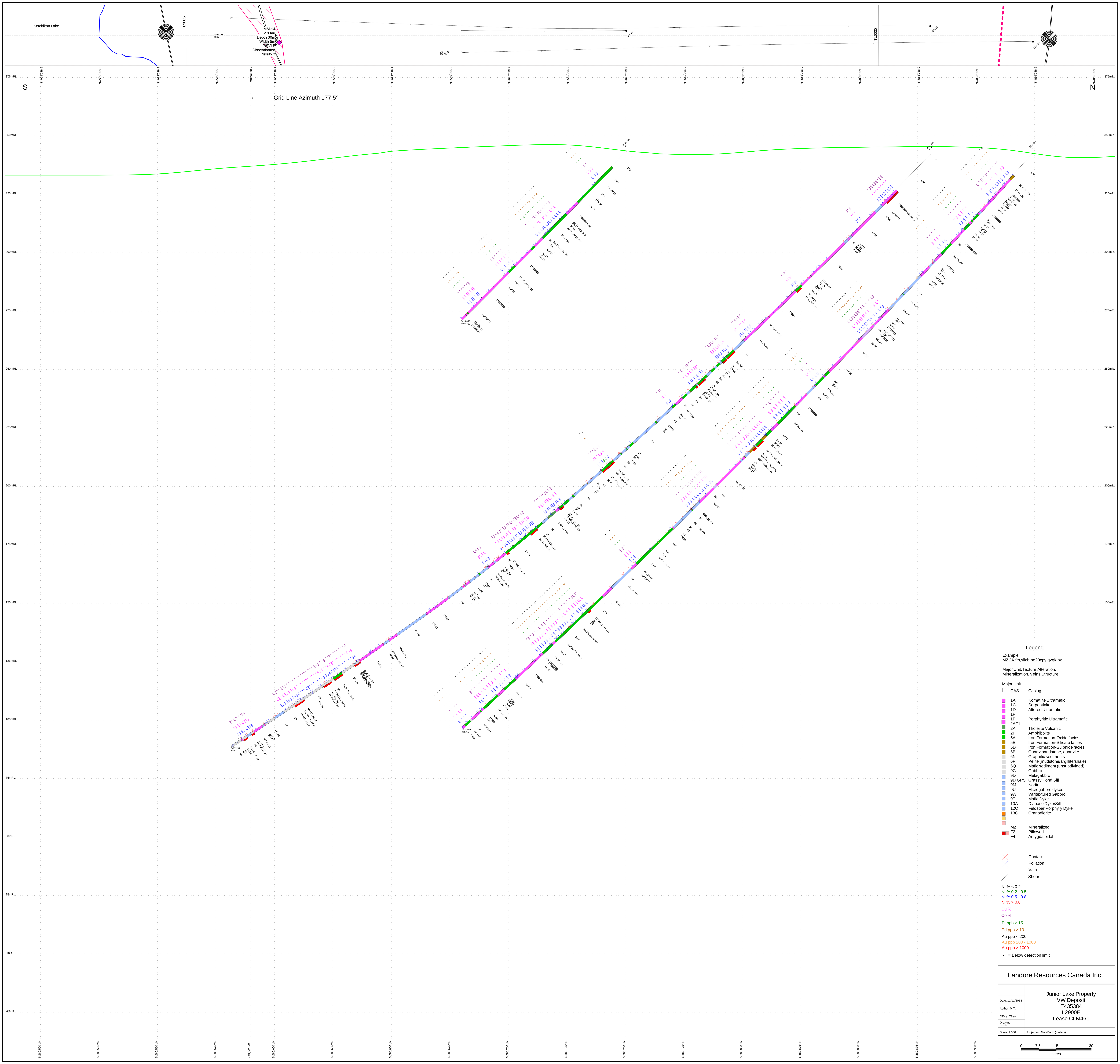
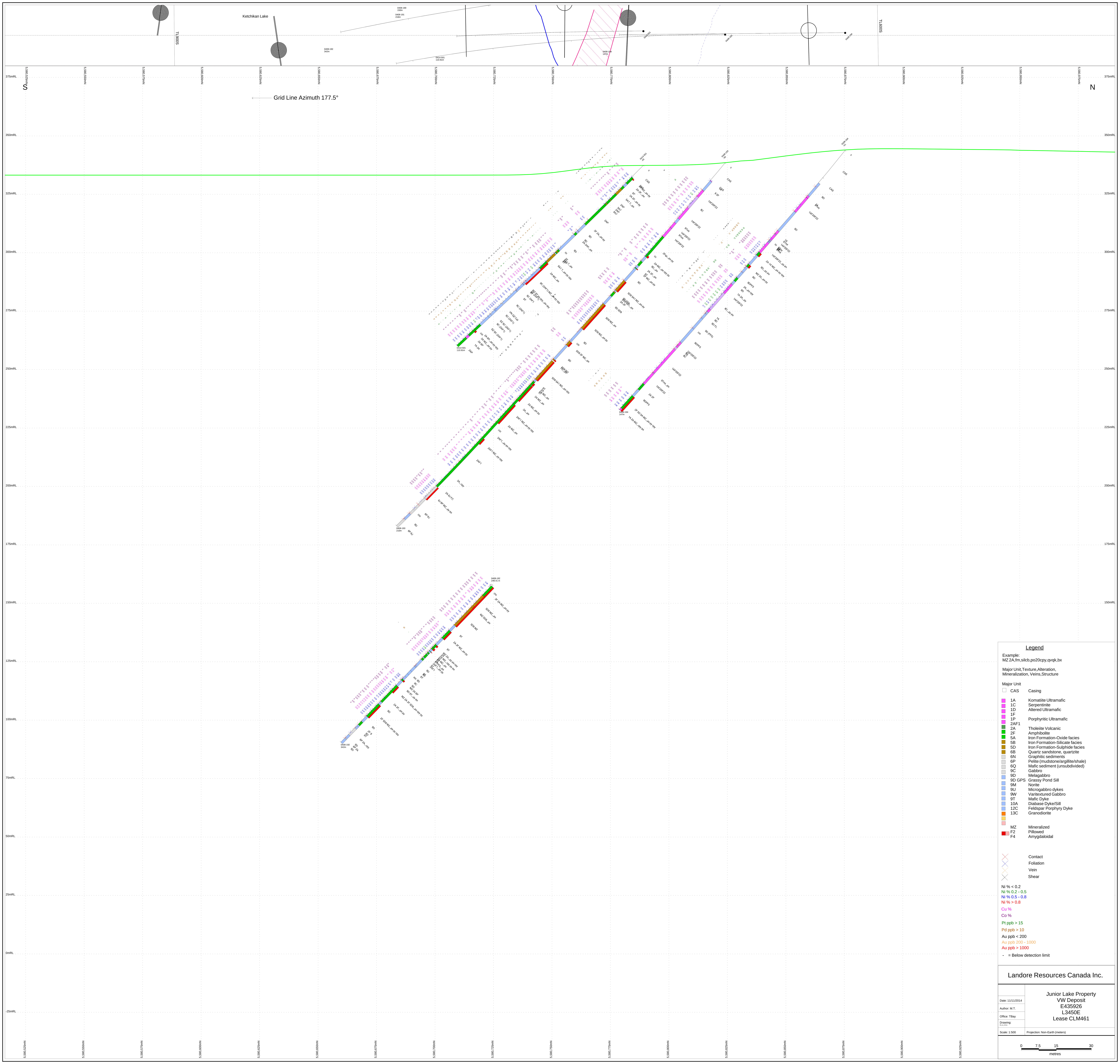
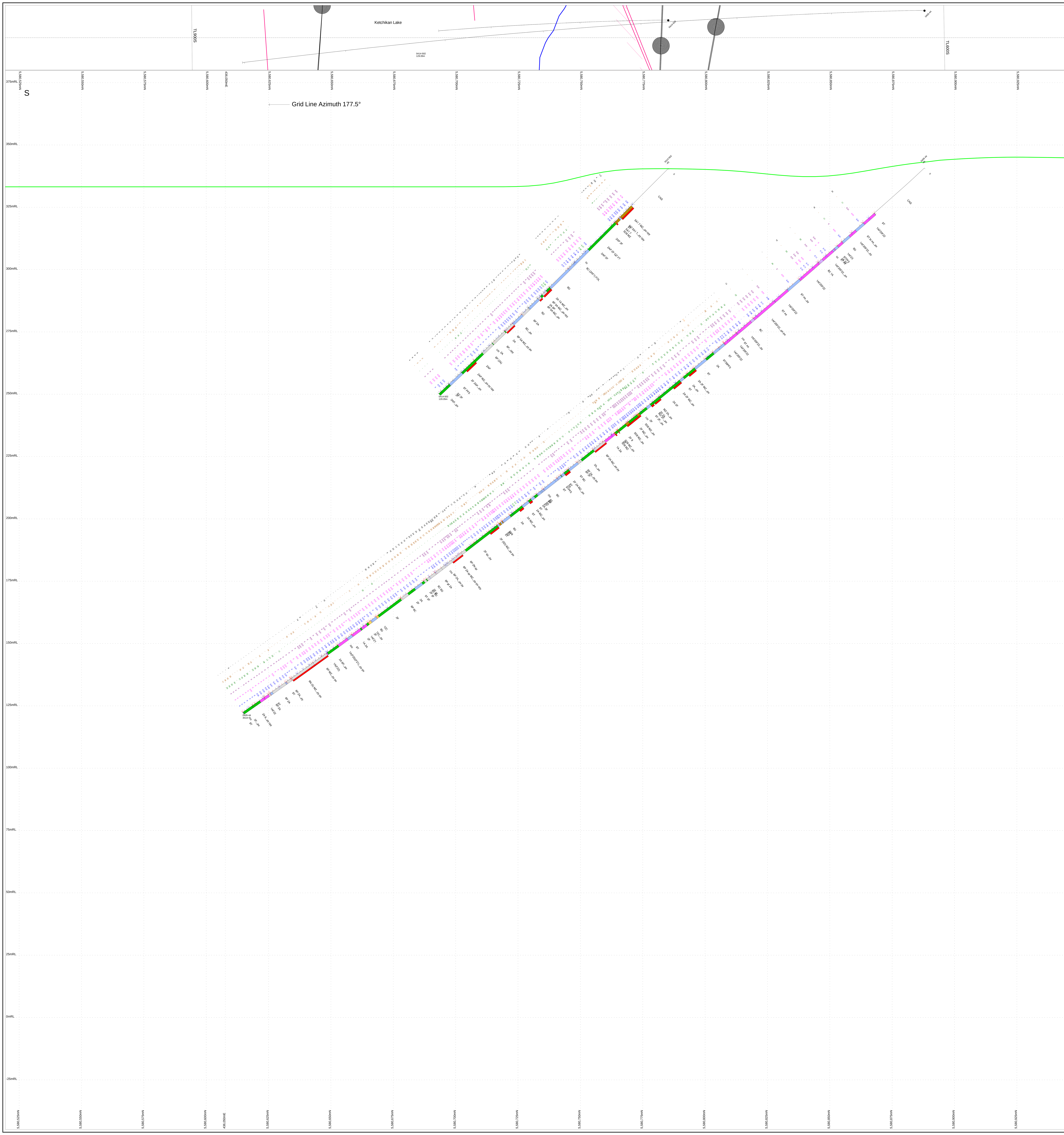
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,
Mineralization, Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
Landore Resources Canada Inc.
Junior Lake Property
BAM East
E434845
L2400E
Lease CLM461
Date: 10/11/2014
Author: M.T.
Office: TBay
Drawing:
Scale: 1:500
Projection: Non-Earth (meters)
0 7.5 15 30
metres


Ketchikan Lake
TL900S
MM-14
2.8 fair
Depth 30m
Width 5m
VLF
Disseminated
Priority 3
TL600S
S
N
Grid Line Azimuth 177.5°
375mRL
350mRL
325mRL
300mRL
275mRL
250mRL
225mRL
200mRL
175mRL
150mRL
125mRL
100mRL
75mRL
50mRL
25mRL
0mRL
-25mRL
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,
Mineralization, Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
Landore Resources Canada Inc.
Date: 11/11/2014
Author: M.T.
Office: TBay
Drawing: 04_cs_1411a
Scale: 1:500
Junior Lake Property
VW Deposit
E435384
L2900E
Lease CLM461
Projection: Non-Earth (meters)
0 7.5 15 30
metres
5,580,900mN


Ketchikan Lake
TL900S
TL600S
S
N
Grid Line Azimuth 177.5°
375mRL
350mRL
325mRL
300mRL
275mRL
250mRL
225mRL
200mRL
175mRL
150mRL
125mRL
100mRL
75mRL
50mRL
25mRL
0mRL
-25mRL
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,
Mineralization, Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
Landore Resources Canada Inc.
Junior Lake Property
VW Deposit
E435926
L3450E
Lease CLM461
Date: 11/11/2014
Author: M.T.
Office: TBay
Drawing:
04_cs_1411a
Scale: 1:500
Projection: Non-Earth (meters)
0
7.5
15
30
metres



Legend

Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx

Major Unit,Texture,Alteration,
Mineralization, Veins,Structure

Major Unit

☐	CAS	Casing
■	1A	Komatiite Ultramafic
■	1C	Serpentinite
■	1D	Altered Ultramafic
■	1F	
■	1P	Porphyritic Ultramafic
■	2AF1	
■	2A	Tholeiite Volcanic
■	2F	Amphibolite
■	5A	Iron Formation-Oxide facies
■	5B	Iron Formation-Silicate facies
■	5D	Iron Formation-Sulphide facies
■	6B	Quartz sandstone, quartzite
■	6N	Graphitic sediments
■	6P	Pelite (mudstone/argillite/shale)
■	6Q	Mafic sediment (unsubdivided)
■	9C	Gabbro
■	9D	Melagabbro
	9D GPS	Grassy Pond Sill
■	9M	Norite
■	9U	Microgabbro dykes
■	9W	Varitextured Gabbro
■	9T	Mafic Dyke
■	10A	Diabase Dyke/Sill
■	12C	Feldspar Porphyry Dyke
■	13C	Granodiorite
■		
■		
	MZ	Mineralized
■■	F2	Pillowed
	F4	Amygdaloidal

╳	Contact
╳	Foliation
╳	Vein
╳	Shear

Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8

Cu %
Co %

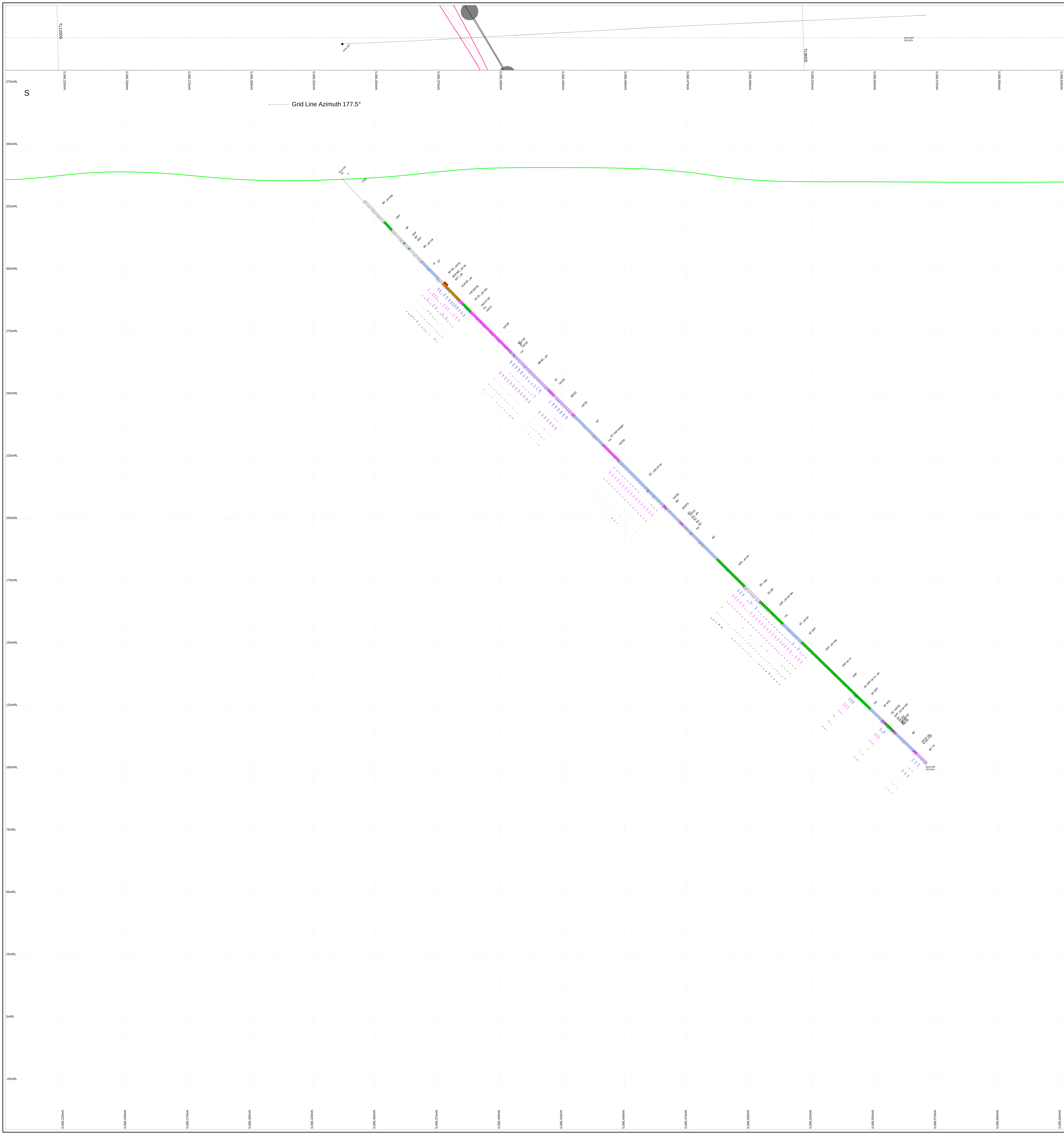
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000

- = Below detection limit
> = Above detection limit

Landore Resources Canada Inc.

Date: 21/11/2014	Junior Lake Property VW East E436033 L3550E Lease CLM461
Author: M.T.	
Office: TBay	
Drawing: 04_cs_1411a	
Scale: 1:500	Projection: Non-Earth (meters)

0 7.5 15 30
metres



Legend

Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx

Major Unit,Texture,Alteration,
Mineralization, Veins,Structure

Major Unit

CAS	Casing
1A	Komatiite Ultramafic
1C	Serpentinite
1D	Altered Ultramafic
1F	
1P	Porphyritic Ultramafic
2AF1	
2A	Tholeiite Volcanic
2F	Amphibolite
5A	Iron Formation-Oxide facies
5B	Iron Formation-Silicate facies
5D	Iron Formation-Sulphide facies
6B	Quartz sandstone, quartzite
6N	Graphitic sediments
6P	Pelite (mudstone/argillite/shale)
6Q	Mafic sediment (unsubdivided)
9C	Gabbro
9D	Melagabbro
9D GPS	Grassy Pond Sill
9M	Norite
9U	Microgabbro dykes
9W	Varitextured Gabbro
9T	Mafic Dyke
10A	Diabase Dyke/Sill
12C	Feldspar Porphyry Dyke
13C	Granodiorite
MZ	Mineralized
F2	Pillowed
F4	Amygdaloidal

Contact
Foliation
Vein
Shear

Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8

Cu %

Co %

Pt ppb > 15

Pd ppb > 10

Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000

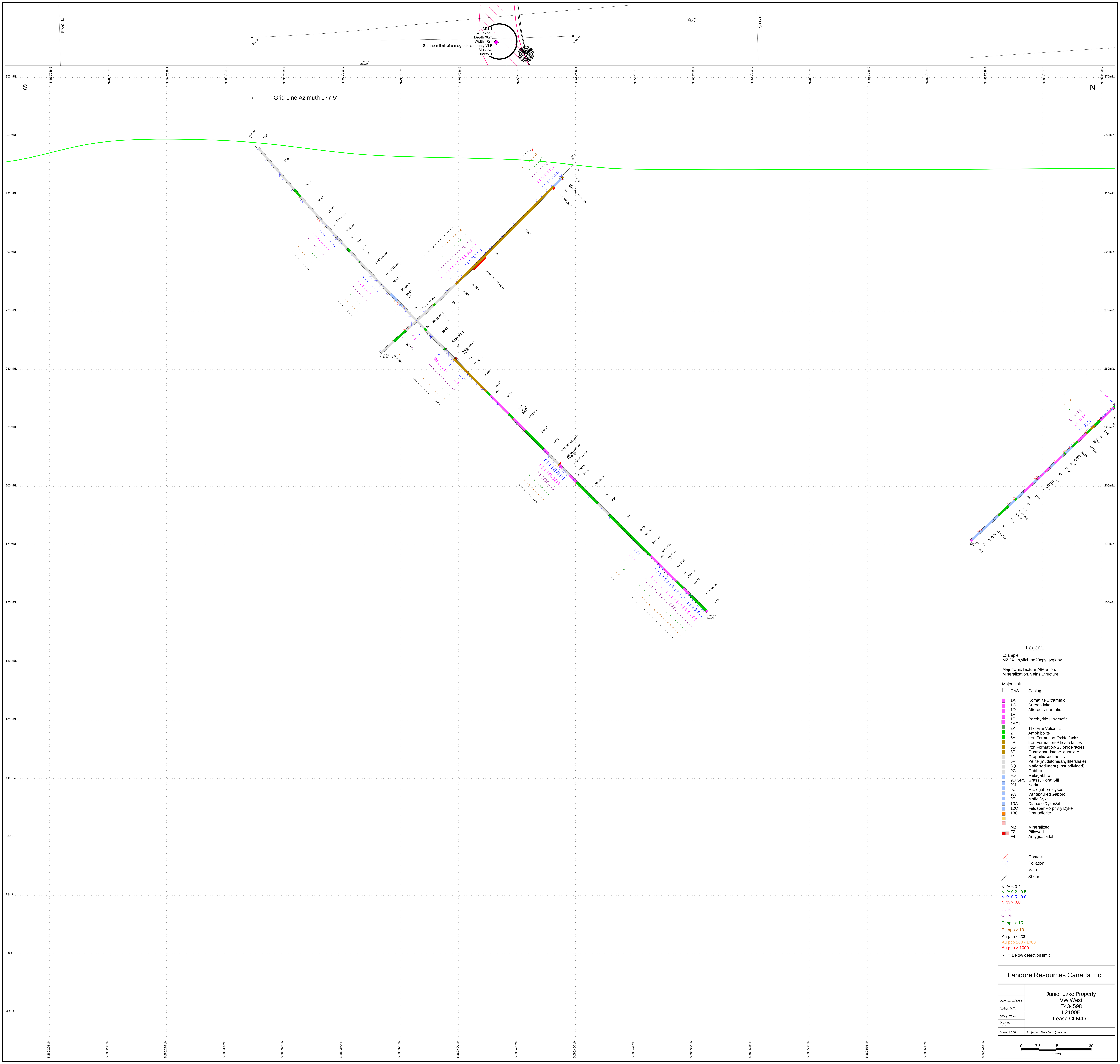
- = Below detection limit

Landore Resources Canada Inc.

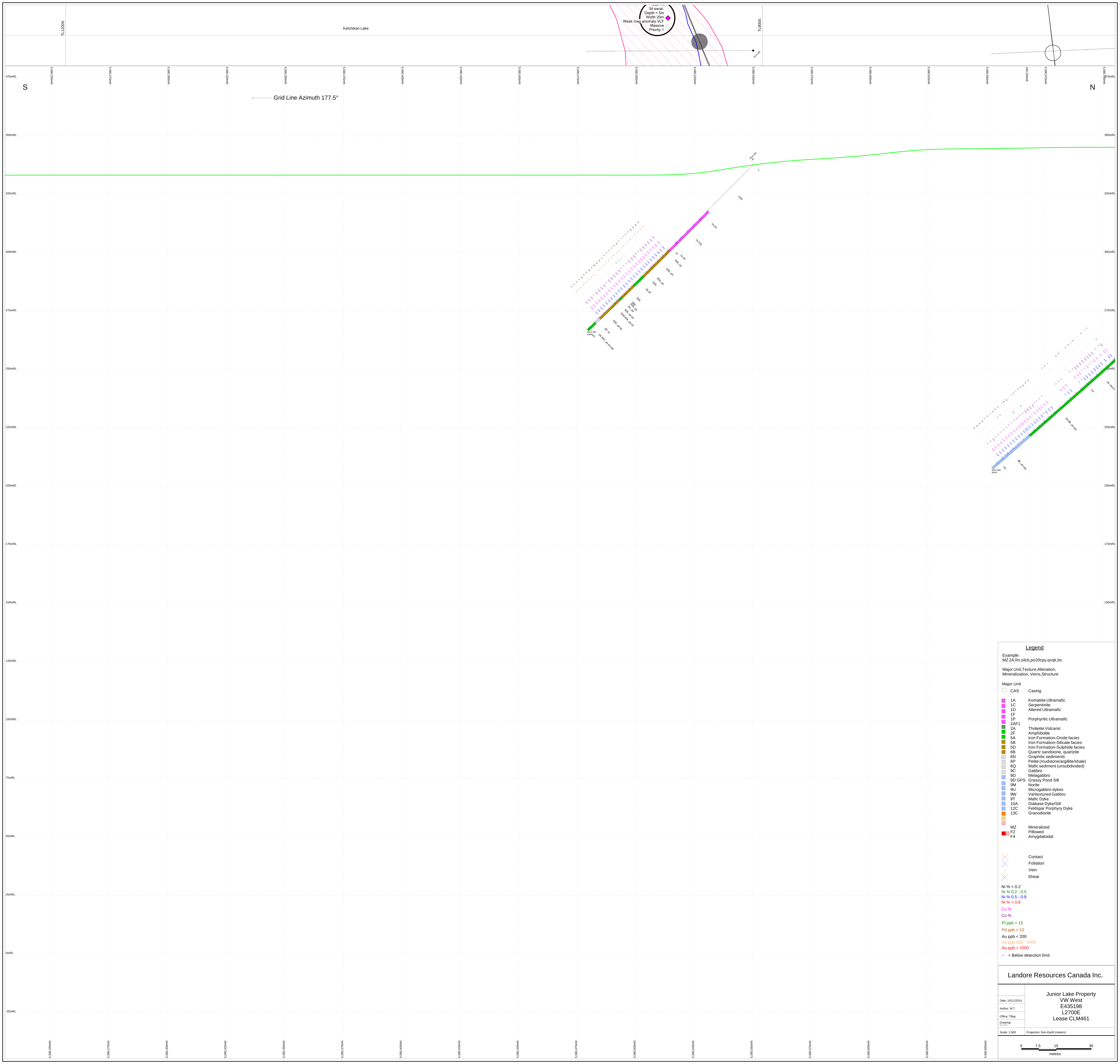
Junior Lake Property
VW West
E434439
L1950E South
Lease CLM461

Date: 11/11/2014
Author: M.T.
Office: TBay
Drawing: 04_cs_1411a
Scale: 1:500
Projection: Non-Earth (meters)

0 7.5 15 30
metres



S
N
Grid Line Azimuth 177.5°
MM-1
40 excel.
Depth 30m
Width 10m
Southern limit of a magnetic anomaly VLF
Massive
Priority 1
TL1200S
TL900S
375mRL
350mRL
325mRL
300mRL
275mRL
250mRL
225mRL
200mRL
175mRL
150mRL
125mRL
100mRL
75mRL
50mRL
25mRL
0mRL
-25mRL
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,
Mineralization,Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
Landore Resources Canada Inc.
Junior Lake Property
VW West
E434598
L2100E
Lease CLM461
Date: 11/11/2014
Author: M.T.
Office: TBay
Drawing:
Scale: 1:500
Projection: Non-Earth (meters)
0
7.5
15
30
metres



S
N
Grid Line Azimuth 177.5°
Ketchikan Lake
Weak mag anomaly VLF
Depth < 5m
Width 20m
Massive
Priority 1
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,Mineralization,Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
Landore Resources Canada Inc.
Junior Lake Property
VW West
E435198
L2700E
Lease CLM461
Date: 10/11/2014
Author: M.T.
Office: TBay
Drawing: 04_cs_1411a
Scale: 1:500
Projection: Non-Earth (meters)
0 7.5 15 30
metres

MM-14
MM-3
Ketchikan Lake
TL600S
S
N
Grid Line Azimuth 177.5°
375mRL
350mRL
325mRL
300mRL
275mRL
250mRL
225mRL
200mRL
175mRL
150mRL
125mRL
100mRL
75mRL
50mRL
25mRL
0mRL
-25mRL
Legend
Example:
MZ 2A,fm,silcb,po20cpy,qvqk,bx
Major Unit,Texture,Alteration,
Mineralization, Veins,Structure
Major Unit
CAS Casing
1A Komatiite Ultramafic
1C Serpentinite
1D Altered Ultramafic
1F
1P Porphyritic Ultramafic
2AF1
2A Tholeiite Volcanic
2F Amphibolite
5A Iron Formation-Oxide facies
5B Iron Formation-Silicate facies
5D Iron Formation-Sulphide facies
6B Quartz sandstone, quartzite
6N Graphitic sediments
6P Pelite (mudstone/argillite/shale)
6Q Mafic sediment (unsubdivided)
9C Gabbro
9D Melagabbro
9D GPS Grassy Pond Sill
9M Norite
9U Microgabbro dykes
9W Varitextured Gabbro
9T Mafic Dyke
10A Diabase Dyke/Sill
12C Feldspar Porphyry Dyke
13C Granodiorite
MZ Mineralized
F2 Pillowed
F4 Amygdaloidal
Contact
Foliation
Vein
Shear
Ni % < 0.2
Ni % 0.2 - 0.5
Ni % 0.5 - 0.8
Ni % > 0.8
Cu %
Co %
Pt ppb > 15
Pd ppb > 10
Au ppb < 200
Au ppb 200 - 1000
Au ppb > 1000
- = Below detection limit
Landore Resources Canada Inc.
Junior Lake Property
VW West
E435333
L2850E
Lease CLM461
Date: 11/11/2014
Author: M.T.
Office: TBay
Drawing:
04_cs_1411a
Scale: 1:500
Projection: Non-Earth (meters)
0
7.5
15
30
metres

LANDORE RESOURCES INC. GEOLOGICAL LEGEND (ROCK CODES)

16 Code	TECTONITES
A	Protomylonite
B	Mylonite
C	Ultramylonite
D	Blastomylonite
E	Phyllonite
F	Cataclasite
G	Tectonic breccia

15 Code	MIGMATITES/GNEISSES/GRANULITES
A	Unsubdivided Gneiss
B	Orthogneiss (Igneous rock-derived)
C	Granite gneiss
D	Granodiorite gneiss
E	Tonalite gneiss
F	Paragneiss (sediment-derived)
G	Ortho-/Paragneiss Composite
H	Mafic Gneiss (unsubdivided)
J	Intermediate gneiss (unsubdivided)
K	Quartzo-feldspathic gneiss
L	Straight (layered) gneiss
M	Injection gneiss
N	Augen gneiss
P	Anorthosite
Q	Amphibolite
R	Unsubdivided Migmatite
S	Schist
T	Unsubdivided Granulite
U	Granulite
V	Granulite gneiss
W	Lit-par-lit Migmatite

14 Code	ALKALINE HYPABYSSAL ROCKS
A	Lamprophyre (undifferentiated)
B	Felsic Lamprophyre
C	Intermediate lamprophyre
D	Mafic lamprophyre
E	Ultramafic lamprophyre
F	Lamproite (olivine-rich, phlogopite bearing)

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

13 Code	FELSIC-INTERMEDIATE PLUTONIC ROCKS
A	Alkali-feldspar granite
B	Granite
C	Granodiorite
D	Monzonite
E	Quartz Monzonite
F	Tonalite
G	Gneiss
H	Diorite
J	Quartz diorite
K	Granophyre
S	Schist

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

12 Code	FELSIC-INTERMEDIATE HYPABYSSAL ROCKS
A	Quartz-feldspar porphyry dykes
B	Quartz porphyry dykes
C	Feldspar porphyry dykes
D	Aplite/Felsite dykes
E	Pegmatite dykes
F	Felsic dykes (undifferentiated)
H	Intermediate dykes (undifferentiated)

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

11 Code	ALKALINE PLUTONIC ROCKS
A	Quartz syenite (nordmarkite, sil oversat)
B	Syenite (Pulaskite, silica saturated)
C	Nepheline syenite (foyaite, sil undersat)
D	Alkalic Gabbro (Essexite)
E	Carbonatite (undifferentiated)
F	Phoscorite (magnetite-apatite rock)
G	Fenite (contact alkali metasomatism)
H	Kimberlite

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

10 Code	DIABASE DYKES & SILLS
A	Diabase (unsubdivided)
B	Quartz Diabase
C	Olivine Diabase

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

9 Code	MAFIC PLUTONIC ROCKS
A	Anorthosite(90-100% plag)
B	Leucogabbro (10-35% cpx)
C	Gabbro (35-65% cpx)
D	Melagabbro (65-90% cpx)
E	Troctolite (ol-plag cumulate)
F	Hornblende gabbro (35-65% hbl)
G	Gneiss
H	Leucogabbronorite (10-35% cpx+opx)
J	Gabbronorite(35-65% cpx+opx)
K	Melagabbronorite (65-90% cpx+opx)
L	Leuconorite (10-35% opx)
M	Norite (35-65% opx)
N	Melanorite (65-90% opx)
S	Schist
T	Mafic dykes (undifferentiated)
U	Microgabbro dykes
V	Gabbro dyke

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

8 Code	ULTRAMAFIC PLUTONIC ROCKS
A	Dunite
B	Serpentinite
C	Peridotite (unsubdivided)
D	Wehrlite
E	Harzburgite
F	Lherzolite
G	Gneiss
H	Pyroxenite (unsubdivided)
J	Clinopyroxenite
K	Orthopyroxenite
L	Websterite
M	Hornblendite
S	Schist
T	Ultramafic dykes (undifferentiated)

Code	Intrusive Modifiers	Code		Code		Code	
1	Layered (unsubdivided)	14	Biotite	27	Ocellae	40	Taxitic
2	Modal/density grading	15	Phlogopite	28	Rheomorphic Dyke	41	Varitextured
3	Grain-size grading	16	Muscovite	29	Chilled Margins	42	Dyke/Sill
4	Igneous lamination	17	Quartz	30	Intrusion Breccia	43	Mottled
5	Flow differentiated	18	Apatite	31	Breccia dyke		
6	Pegmatitic	19	Feldspathic (<10% Plag)	32	Magma mixing		
7	Glomerophyric	20	Leucocratic (10-35% mafics)	33	Xenoliths [xeno lithology]		
8	Poikilitic	21	Melanocratic(65-90% mafics)	34	Autoliths [auto lithology]		
9	Oikocrystic	22	Granophyric	35	Amoeboid Inclusions		
10	Olivine	23	Graphic	36	Diffuse margins		
11	Pyroxene	24	Ophitic	37	Scalloped margins		
12	Hornblende	25	Subophitic	38	Rheomorphic textures		
13	Magnetite	26	Aphyric	39	Xenocrysts		

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

7 Code	ALKALINE METAVOLCANIC ROCKS
A	Trachyte
B	Phonolite
C	Nephelinite
D	Leucitite (pyroxene+leucite)
E	Trachyandesite (latite)
F	Trachybasalt
G	Gneiss
S	Schist

Code	Flows (F)	Code	
F1	Massive	F13	Spinifex
F2	Pillowed	F14	Cumulate
F3	Flow banding	F15	Perlitic
F4	Amygdaloidal	F16	Ignimbrite
F5	Variolitic	F17	Debris Flow/Lahar (Mudflow)
F6	Spherulitic	F18	Autoclastic breccia
F7	Vesicular	F19	Flow lobe toe(s)
F8	Hyaloclastite		
F9	Flow top breccia		
F10	Pillow breccia		
F11	Polygonal Jointing		
F12	Bladed		

Code	Pyroclastics (P)
P1	Pyroclastic Breccia (>64mm)
P2	Tuff breccia (>64mm)
P3	Lapilli tuff (2-64mm)
P4	Lapillistone (2-64mm)
P5	Tuff (<2mm)
P6	Vitric tuff
P7	Crystal tuff
P8	Lithic tuff

Code	Pyroclastics Modifiers
A	Felsic-in-mafic breccia
B	Mafic-in-felsic breccia

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

6 Code	CLASTIC METASEDIMENTARY ROCKS
A	Conglomerate
A1	Orthoconglomerate
A2	Paraconglomerate (tillites, matrix-rich)
B	Quartz sandstone, quartzite (quartz arenite)
C	Sandstone
D	Feldspathic sandstone
E	Lithic sandstone
F	Arkose
G	Gneiss
H	Wacke (greywacke)
J	Siltstone
K	Argillite
L	Claystone
M	Shale
N	Graphitic sediments
P	Pelite (mudstone/argillite/shale)
Q	Mafic sediment (unsubdivided)
R	Ultramafic sediment (unsubdivided)
S	Schist

Code	Sedimentary Modifiers	Code	
1	Interbedded	18	Clast supported
2	Intercalated	19	Matrix supported
3	Interlaminated	20	Polymictic
4	Imbricated	21	Monomictic
5	Interflow	22	Graphitic
6	Crossbedded	23	Argillaceous
7	Ripple cross-lamination	24	Tuffaceous
8	Ripple marks		
9	Graded bedding		
10	Sole marks (unsubdivided)		
11	Groove marks		
12	Flute casts		
13	Load casts		
14	Flame structures		
15	Slump folds		
16	Laminated		
17	Banded		

5 Code	CHEMICAL METASEDIMENTARY ROCKS
A	Iron Formation-Oxide facies
B	Iron Formation-Silicate facies
C	Iron Formation-Carbonate facies
D	Iron Formation-Sulphide facies

Code	Iron Formation Assemblages	Code	
1	Banded chert-magnetite	9	Banded chert-sulphide
2	Banded chert	10	Banded oxide-wacke
3	Banded siderite-ankerite-chert	11	Banded oxide-siltstone
4	Banded grunerite-hornblende	12	Banded oxide-pelite
5	Amphibole-garnet-biotite	13	Sulphidic pelite
6	Pyritic graphitic pelite	14	Banded chert-wacke
7	Amphibolitized	15	Banded magnetite
8	Banded chert-carbonate		

4 Code	FELSIC METAVOLCANIC ROCKS
A1	Tholeiitic rhyolite
A2	High-silica rhyolite
A3	Calc-alkalic rhyolite
A4	Alkaline rhyolite
B1	Tholeiitic rhyodacite
B2	Calc-alkalic rhyodacite
C1	Tholeiitic dacite
C2	Calc-alkalic dacite
G	Gneiss
S	Schist

Code	Flows (F)	Code	
F1	Massive	F13	Spinifex
F2	Pillowed	F14	Cumulate
F3	Flow banding	F15	Perlitic
F4	Amygdaloidal	F16	Ignimbrite
F5	Variolitic	F17	Debris Flow/Lahar (Mudflow)
F6	Spherulitic	F18	Autoclastic breccia
F7	Vesicular	F19	Flow lobe toe(s)
F8	Hyaloclastite		
F9	Flow top breccia		
F10	Pillow breccia		
F11	Polygonal Jointing		
F12	Bladed		

Code	Pyroclastics (P)
P1	Pyroclastic Breccia (>64mm)
P2	Tuff breccia (>64mm)
P3	Lapilli tuff (2-64mm)
P4	Lapillistone (2-64mm)
P5	Tuff (<2mm)
P6	Vitric tuff
P7	Crystal tuff
P8	Lithic tuff

Code	Pyroclastics Modifiers
A	Felsic-in-mafic breccia
B	Mafic-in-felsic breccia

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

3 Code	INTERMEDIATE METAVOLCANIC ROCKS
A	Andesite - Tholeiitic (<16% Al2O3)
B	Andesite - Calc-alkaline
C	Icelandite (>16% Al2O3,>0.35% P2O5)
D	Shoshonite
E	Amphibolite (amphibole-plagioclase)
G	Gneiss
S	Schist

Code	Flows (F)	Code	
F1	Massive	F13	Spinifex
F2	Pillowed	F14	Cumulate
F3	Flow banding	F15	Perlitic
F4	Amygdaloidal	F16	Ignimbrite
F5	Variolitic	F17	Debris Flow/Lahar (Mudflow)
F6	Spherulitic	F18	Autoclastic breccia
F7	Vesicular	F19	Flow lobe toe(s)
F8	Hyaloclastite		
F9	Flow top breccia		
F10	Pillow breccia		
F11	Polygonal Jointing		
F12	Bladed		

Code	Pyroclastics (P)
P1	Pyroclastic Breccia (>64mm)
P2	Tuff breccia (>64mm)
P3	Lapilli tuff (2-64mm)
P4	Lapillistone (2-64mm)
P5	Tuff (<2mm)
P6	Vitric tuff
P7	Crystal tuff
P8	Lithic tuff

Code	Pyroclastics Modifiers
A	Felsic-in-mafic breccia
B	Mafic-in-felsic breccia

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

2 Code	MAFIC METAVOLCANIC ROCKS
A	Tholeiite
B	Fe Tholeiite
C	Mg Tholeiite
D	Calc-alkalic basalt
E	Picrite
F	Amphibolite (amphibole-plagioclase)
H	Mafic Hornfels
G	Gneiss
S	Schist

Code	Flows (F)	Code	
F1	Massive	F13	Spinifex
F2	Pillowed	F14	Cumulate
F3	Flow banding	F15	Perlitic
F4	Amygdaloidal	F16	Ignimbrite
F5	Variolitic	F17	Debris Flow/Lahar (Mudflow)
F6	Spherulitic	F18	Autoclastic breccia
F7	Vesicular	F19	Flow lobe toe(s)
F8	Hyaloclastite	F20	Talc schist
F9	Flow top breccia	F21	Talc-chlorite schist
F10	Pillow breccia	F22	Talc -carbonate schist
F11	Polygonal Jointing	F23	Tremolite schist
F12	Bladed		

Code	Pyroclastics (P)
P1	Pyroclastic Breccia (>64mm)
P2	Tuff breccia (>64mm)
P3	Lapilli tuff (2-64mm)
P4	Lapillistone (2-64mm)
P5	Tuff (<2mm)
P6	Vitric tuff
P7	Crystal tuff
P8	Lithic tuff

Code	Pyroclastics Modifiers
A	Felsic-in-mafic breccia
B	Mafic-in-felsic breccia

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

1 Code	ULTRAMAFIC METAVOLCANIC ROCKS
A	Komatiite
B	Basaltic Komatiite
C	Serpentinite
G	Gneiss
S	Schist

Code	Flows (F)	Code	
F1	Massive	F13	Spinifex
F2	Pillowed	F14	Cumulate
F3	Flow banding	F15	Perlitic
F4	Amygdaloidal	F16	Ignimbrite
F5	Variolitic	F17	Debris Flow/Lahar (Mudflow)
F6	Spherulitic	F18	Autoclastic breccia
F7	Vesicular	F19	Flow lobe toe(s)
F8	Hyaloclastite	F20	Talc schist
F9	Flow top breccia	F21	Talc-chlorite schist
F10	Pillow breccia	F22	Talc -carbonate schist
F11	Polygonal Jointing	F23	Tremolite schist
F12	Bladed		

Code	Pyroclastics Modifiers
P1	Pyroclastic Breccia (>64mm)
P2	Tuff breccia (>64mm)
P3	Lapilli tuff (2-64mm)
P4	Lapillistone (2-64mm)
P5	Tuff (<2mm)
P6	Vitric tuff
P7	Crystal tuff
P8	Lithic tuff

Code	Pyroclastics Modifiers
A	Felsic-in-mafic breccia
B	Mafic-in-felsic breccia

Code	Porphyritic (PF)
PF1	Feldspar>quartz
PF2	Quartz>feldspar
PF3	Feldspar
PF4	Quartz
PF5	Plagioclase
PF6	K-feldspar
PF7	Amphibole
PF8	Pyroxene
PF9	Olivine
PF10	Biotite
PF11	Muscovite
PF12	Crowded

LANDORE RESOURCES INC. LEGEND MODIFIERS

Code	TEXTURAL & COMPOSITIONAL MODIFIERS
aph	Aphanitic
aut	Autoliths
bnd	Banded
cg	Coarse-grained
cs	Clast supported
cum	Cumulate
fg	Fine-grained
frg	Fragments
fsp	Feldspatic
glm	Glomerophyric
hl	Homeolithic
ht	Heterolithic
ib	Interbeded/ interlaminated
lam	Laminated
lrd	Layered
m	massive
mg	Medium-grained
mm	Monomictic
mxs	Matrix supported
oik	Oikocrystic
oph	Ophitic
pf	Porphyritic
pg	Pegmatitic
pm	Polymictic
tax	Taxitic
vcg	Very coarse-grained
vfg	Very fine-grained
vt	Varitextured
xen	Xenoliths

Code	ALTERATION MODIFIERS
abz	Albitized
ac	Actinolitized
am	Amphibolitized
ank	Ankeritization
az	Alteration zone
bi	Biotization
bl	Bleached
cam	Calcium metasomatism
cb	Carbonatization
cc	Calcitic
ch	Chloritization
ep	Epidotization
fec	Fe-carbonate
fen	Fenitization
fp	Feldspathized
gcb	Green carbonate
gm	Green mica
gf	Graphitic
hem	Hematization
lim	Limonitized
lx	Leucoxene
k	Potassic alteration/metasomatism
rst	Rusty
ser	Sericitized
sil	Silicification
splz	Spilitization
srp	Serpentinization
sz	Saussuritization
tc	Talcose
tcl	Talc-chlorite
tour	Tourmalinization
ur	Uralitization

Code	MINERALIZATION
asp	Arsenopyrite
bn	Bornite
cp	Chalcopyrite
cr	Chromite
gf	Graphite
gn	Galena
hem	Hematite
ilm	Ilmenite
mo	Molybdenite
mt	Magnetite
pn	Pentlandite
po	Pyrrhotite
py	Pyrite
s	Sulphides (uncharacterized)
sc	Scheelite
stb	Stibnite
sp	Sphalerite
td	Tellurides
vg	visible gold
tm	Titanomagnetite

Code	Mineralization Characterization
t	Trace
b	Bleb/Blebby
bd	Bedded
bnd	Banded
cf	Clasts/fragments
ds	Disseminated
f	Fracture-controlled/fill
MZ	Mineralized zone
nt	Net-textured
pc	Patchy/Clustered
st	Stringered
m	Massive
sm	Semi-massive
sw	Stockwork
v	Vein
wp	Wispy

Code	Intensity Codes
w	Weak(ly)
m	Moderate(ly)
s	Strong(ly)
I	Intense(ly)

Code	VEINING
am	Amphibole
ank	Ankerite
c	Calcite
cb	Carbonate
chl	Chlorite
ep	Epidote
hem	Hematite
mt	Magnetite
q	Quartz
qank	Quartz-ankerite
qc	Quartz-calcite
qcb	Quartz-carbonate
qcbch	Quartz-carbonate-chlorite
qcbep	Quartz-carbonate-epidote
qchl	Quartz-chlorite
qchhm	Quartz-chlorite-hematite
qfl	quartz-fluorite
qfsp	quartz-feldspar (quartzo-feldspathic)
qt	Quartz-tourmaline
srp	Serpentine
tour	Tourmaline

Code	Vein Morphology
V1	Non-mineralized
V2	Non-mineralized with min'd haloes
V3	Mineralized with non-min'd host
V4	Mineralized with mineralized host
V5	Non-mineralized with min'd host
cs	Crack-seal
f	Fracture fill
fld	Flooded/flooding
isw	Incipient stockwork
m	Massive
qsw	Quartz stockwork
rxl	Recrystallized
sht	Sheeted
st	Stringer (<1cm)
sw	Stockwork
v	Vein (>10cm)
vbx	Vein breccia
vt	Veinlet (1-10cm)
xc	Crosscutting foliation/bedding

Code	STRUCTURAL MODIFIERS
bg	Boudinaged
bx	Breccia/brecciated
bxd	Breccia dyke
agn	Augen
fld	Fold/folded
fr	Fracture/fractured
frz	Fracture zone
fz	Fault zone
gg	Gouge
jt	Jointed
my	Mylonite
sh	Shear/sheared
sz	Shear zone
fl	Foliated
ln	Lineated
sl	Slickenside
ps	Pseudotachylite

Code	METAMORPHIC MODIFIERS
ss	Schist/schistose
gn	Gneissic
ph	Phyllite
hf	Hornfels
pb	Porphyroblastic
gc	Granoblastic
am	Amphibolite
sg	Sugary-textured

Code	OTHER ABBREVIATIONS
BC	Broken core
BIF	Banded Iron Formation
BVD	beaver dam
C	Compilation data
D	Diamond drill data
IF	Iron Formation
BIF	Banded Iron Formation
LC	Lost core
MA	Magnetic attraction
o/c	outcrop
ob	overburden
T	Temiskaming

Code	ASSEMBLAGE MINERALS (1)
ab	Albite
ac	Actinolite
ad	Andalusite
ah	Anhydrite
ahb	Alkali-hornblende
ak	Ankerite
al	Almandine
als	Alumino-silicates
am	Amphibole
anl	Analcite
ant	Anthophyllite
ap	Apatite
asb	Asbestos
aug	Augite
bar	Barite
bi	Biotite
cb	Carbonate(s)
chd	Chloritoid
chl	Chlorite
cm	Cummingtonite
cpx	Clinopyroxene
crdi	Chrome Diopside
ct	Cordierite
dol	Dolomite
dp	Diopside
en	Enstatite
ep	Epidote
fc	Fuchsite (Cr-muscovite)
flr	Fluorite
fsp	Feldspar
gr	Grossularite
gt	Garnet
gyp	Gypsum
hb	Hornblende
hd	Hedenbergite
hy	Hypersthene
ksp	K-feldspar
ky	Kyanite
le	Leucite
lm	Limonite
lp	Lepidolite (Li-muscovite)
lx	Leucoxene
mag	Magnetite
mal	Malachite
mc	Marcasite
mi	Mica
mk	Microcline
mu	Muscovite
ne	Nepheline
ol	Olivine
opx	Orthopyroxene
or	Othoclase
orp	Orpiment
phl	Phlogopite
pl	Plagioclase

Code	ASSEMBLAGE MINERALS (2)
prv	Perovskite
px	Pyroxene
pyr	Pyrope
qtz	Quartz
rc	Roscoelite (V-muscovite)
rea	Realgar
rut	Rutile
sd	Siderite
ser	Sericite
slm	Silliminite
spd	Spodumene
spl	Spinel
sps	Spessartite
srp	Serpentine
st	Staurolite
tan	Tantalite
ta-cl	Tantalite-Columbite
tlc	Talc
tnt	Titanite
tre	Tremolite
wol	Wollastinite
zrn	Zircon

Code	TREE TYPES
At	Trembling aspen
Bs	Silver Birch
Bw	White bitch
By	Yellow birch
Ce	Cedar
Fb	Balsam Fir
H	Hemlock
L	Larch/Tamarack
Mh	Hard maple
Mm	Mountain/Moose/Striped Maple
Mr	Red Maple
Ms	Silver Maple
Ob	Burr Oak
Or	Red Oak
Ow	White Oak
Pj	Jack Pine
Pr	Red Pine
Pw	White Pine
Sb	Black Spruce
Sw	White Spruce
Wt	Tag alder