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XMET INC.

MAG-GPS, RESISTIVITY / INDUCED POLARIZATION IPOWER3D[®] CONFIGURATION & HIGH-RESOLUTION GROUND GRAVITY SURVEYS

BLACKFLAKE PROJECT

PITOPIKO RIVER AND LIMESTONE RAPIDS TOWNSHIPS, HEARST, ONTARIO, CANADA

LOGISTICS AND INTERPRETATION REPORT

14N021 MAY 2014



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ABSTRACT

On behalf of Xmet Inc., a **Resistivity / Induced Polarization** survey using the **IPower3D**[®] configuration as well as a **MAG-GPS** and high resolution ground **Gravity** surveys were conducted on the **Blackflake Property**, located in the Pitopiko River and Limestone Rapids Townships, in Hearst, Ontario.

During the period of March 24 to April 7, 2014, a total of 10 km of Time Domain Resistivity / Induced Polarization survey was completed using the IPower3D[®] configuration. A total of 70 km was surveyed using MAG-GPS and a total of 488 gravity stations were surveyed. Survev specifications, instrumentation controls, data acquisition, processing and interpretation were successfully completed within the Abitibi Geophysics quality system framework.

The objectives of this survey were to identify zones amenable to Graphite and or Kimberlite mineralization on the Blackflake property and to identify potential drill targets.

A follow-up program including prospecting, drilling and a survey extension has been proposed. The recommendations are presented in section 8 of this report.



1. THE MANDATE

Project ID	Blackflake Project (Our reference: 14N021)
GENERAL LOCATION	60 km northwest of Hearst, Ontario
CUSTOMER	Xmet Inc. 120 Adelaide Street West, Suite 2500 Toronto, Ontario M5H 1T1 Telephone: (416) 644-6588
REPRESENTATIVES	Mr. Stephen Stewart President sstewart@xmet.ca
SURVEY TYPES	Time domain Resistivity / Induced Polarization, IPower3D [®] configuration with MAG-GPS and High Resolution Ground Gravity surveys

- GEOPHYSICAL OBJECTIVES
- Identify zones amenable to graphite and or kimberlite mineralization
- Identify targets for further exploration.







2. THE BLACKFLAKE PROJECT

LOCATION	Pitopiko River and Limestone Rapids Townships Hearst, Ontario, Canada Centred on 50.03641°N and 84.26987°W, UTM NAD83, zone 16N: 695 500 mE, 5 546 250 mN NTS sheets: 42K/01
NEAREST SETTLEMENTS	Hearst : 60 km to the southeast. Calstock: 30 km south southeast.
Access	The Blackflake grid was accessed daily from Hearst by taking highway 11 west for ~35 km until route 663. From here, route 663 was taken north, past Calstock, for ~35 km. An access road is then taken west for ~10 km and from there the grid was accessed by ATV.
GEOMORPHOLOGY	Topography on the grid ranged from 120 m to 160 m above sea level. There are a few small streams present, and vegetation consisted of spruce, fir and birch.
CULTURAL FEATURES	Cultural features were limited to the access trail; this is not thought to have any effect on the data.
MINING LAND TENURE	The Blackflake property comprises 16 claims; the claim numbers encompassed by the present survey are illustrated on page 4.
SURVEY GRID	This grid on the Blackflake property consists of 14, 3000 m E/W lines spaced at 150 m and 17, 1350 m N/S lines spaced at 75 m. The base line is located at 0+00 E cutting the grid in half. There are also 2 tie lines located at 15+00W and 14+50E. In total 10 line km were surveyed with IPower3D [®] , 70 line km with MAG-GPS, and 506 gravity stations.
ENVIRONMENTAL HEALTH AND SAFETY	As part of the Abitibi Geophysics EHS program crew members received first aid training and are provided with safety equipment and specialized training for the induced polarization technique. In addition, the crew was provided with a satellite telephone for emergency communication.
COORDINATE SYSTEM	Projection: Universal Transverse Mercator, zone: 16N Datum: NAD83



Figure 2. Index of claims covering the Blackflake Project





3. IPower3D[®] Resistivity / Induced Polarization Survey

Type of survey	Time Domain Resistivi	ity / Induced Polarization
CONFIGURATION	IPower3D [®]	
Personnel	Bruno Tremblay, Éric Vallerand, Darryl Ouellet, Guillaume Nantel, Mario Bertrand, Bruno Tremblay, Carole Picard, tech Pam Coles, G.I.T., Chris Brown, P.Geo.,	Crew Chief & Operator Assistant Assistant Assistant Logistics Production of maps Quality Control, Processing, Interpretation Report Final verification of product conformity
SURVEY COVERAGE	10 km	
DATA ACQUISITION	March 31 to April 5, 2	2014
Transmitter	GDD Instruments Tx Generator: Maximum output: Electrodes: Resolution: Waveform: Pulse Duration: +1	III, s/n 259 Honda 5000 VA 3.6 kW at 10 A at 2,400 V memory-shape alloy rods 1 mA on output current display Bipolar square wave with 50% duty cycle 1 second \leftarrow 1s \rightarrow

Figure 3. Signal transmitted across electrodes C1-C2



□ RECEIVER IRIS Elrec-PRO, (10 input channels), s/n 184 IRIS Switch-PRO 240, s/n 65 Electrodes: Memory-shape alloy rods V_p Primary voltage measurement:

- Input impedance: 100 MΩ
- Resolution: 1 μV
- Typical accuracy: 0.2%

M_A Apparent chargeability measurement:

- Resolution: 0.01 mV/V
- Typical accuracy: **0.4%**
- Semi-log sampling mode, 20 time windows (M₁ to M₂₀).



Figure 4. Semi-log windows (1 s pulse)

- All gates are normalized with respect to a standard decay curve for field quality assurance.
- Final chargeability values were normalized to the 2 second pulse Newmont standard.
- APPARENT
 RESISTIVITY
 CALCULATION

$$\rho_{a} = 2\pi \cdot \frac{V_{p}}{I} \cdot \frac{1}{\left(\frac{1}{C_{1}P_{1}} - \frac{1}{C_{2}P_{1}}\right) - \left(\frac{1}{C_{1}P_{2}} - \frac{1}{C_{2}P_{2}}\right)} \quad (\Omega \cdot m)$$

Cumulative error: **5% max**, mainly due to chaining accuracy.



QUALITY CONTROLS (RECORDS AVAILABLE UPON REQUEST)

Before the survey:

- ✓ Transmitters & motor generators were checked for maximum output using calibrated loads.
- ✓ Receivers were checked using the Abitibi Geophysics SIMP™ certified and calibrated V_P & M_a signal simulator.

During data acquisition:

- ✓ Rx & Tx cable insulation was verified every morning.
- ✓ Proprietary Software ProsysControl[™] allowed a daily thorough monitoring of data quality and survey efficiency.
- ✓ Sufficient pulses were stacked: 8 pulses for every reading.

At the Base of Operations:

- ✓ Field quality assurance inspected & validated.
- ✓ Each IP decay curve was analyzed with *ProsysControl*[™]. The few gates that were rejected were not included in the calculation of the plotted M_a.

QUALITY STATISTICS

Table 1. Quality statistics – IPower3D[®]

Blackflake Project IPower3D [®]	
Average contact resistance across R_X dipole (P_1 - P_2)	1.4 kΩ
Average current applied to T_X dipole (C ₁ -C ₂)	4042 mA
Average V_p measured across R_x dipole (P ₁ -P ₂)	134.6 mV
Observed windows found to fit a pure electrode polarization relaxation curve	89.5%
Average deviation of the validated, normalized windows with respect to the mean chargeabilities.	0.24 mV/V



4. GPS INTEGRATED GROUND MAGNETIC SURVEY

Type of survey	Measurement of the Total Magnetic Field (TMF) with GPS readings recorded every second. The plotted values were corrected for diurnal variations using readings from a synchronized MAG base station.	
Personnel	Sylvain Brousseau, Bruno Tremblay, Carole Picard, tech Pam Coles, G.I.T., Chris Brown, P.Geo.,	Geophysical operator Logistics Production of maps Quality Control, Processing, Interpretation Report Final verification of product conformity
DATA ACQUISITION	March 24 to 31, 2014	
Survey coverage	70 km	
FIELD MAGNETOMETER	GEM Systems GSM-19 Proton precession mag built-in GPS. Resolution: Absolute accuracy: Gradient tolerance: TMI sensor: Sensor:	W , s/n 71191 gnetometers with Overhauser effect and 0.01 nT / 1 m 0.2 nT / 2-5 m >10 000 nT/m at a height of 1.8 m above ground 14140
BASE STATION	GEM Systems GSM-19 Proton precession mage Resolution: Absolute accuracy: Cycle time: Sensor s/n: Location (UTM NAD83) Reference field:	9, s/n 61519 netometer with Overhauser effect 0.01 nT 0.2 nT 10 seconds 42231 : 16U, 5 547 334 mN, 696 794 mE 57 100 nT



QUALITY CONTROLS (RECORDS AVAILABLE UPON REQUEST)

Before the survey:

✓ All magnetometers were successfully field-tested on Abitibi Geophysics' private control line.

Every day during data acquisition:

- ✓ Every morning, the operator had to successfully test for any magnetic contamination.
- ✓ In the evening, the geophysical operator reviewed the base station and the mobile unit recordings using our proprietary MAGneto[®] processing and QC software.
- ✓ The geophysical operator ensures no active geomagnetic activity would be encountered during the survey by visiting the Space Weather Canada website (www.spaceweather.gc.ca).

At the Base of Operations:

- ✓ Field QCs were inspected & validated.
- \checkmark All profiles were inspected and only a few noisy readings were removed from the database.



5. GRAVITY DATA ACQUISITION

Type of survey	Semi-detailed gravity survey with expected Bouguer anomaly accuracy better than 0.05 mGal (before terrain corrections).	
Personnel	Dennis Palos, Daniel Duquette, Sandro Paquette, Luc Houle, Martin Dubois, Carole Picard, Tech., Madjid Chemam, P.Geo., Chris Brown, P. Geo.,	Crew chief, gravimeter operator Gravimeter operator GPS operator GPS operator Logistics Plotting QC, processing and interpretation Final validation of product conformity
SURVEY COVERAGE	506 stations, including one visited stations.	e base station but excluding all the re-
DATA ACQUISITION	March 28 to April 6, 2014	L
INSTRUMENTATION	Type: Reading resolution: Stand. Field Repeatability: Operating Range: Residual Long Term drift: Operating Temperature: Serial numbers: Calibration constant: Seismic filter: Continuous tilt correction: Auto rejection: Tide correction: Read time:	Scintrex CG-5u Autograv Fused Quartz 0.001 mGal (digital recording) < 0.005 mGal 8000 mGal without resetting < 0.02 mGal/day (static) -40°C to +45°C 080740418 (418) 980890430 (430) K ₄₁₈ = 0.999642 mGal / unit K ₄₃₀ = 0.999645 mGal / unit Enabled Enabled Enabled (threshold is 6 x standard deviation with seismic filter on) Enabled 2 to 4 cycles of 30 seconds
Software	SCTutil (Scintrex) for data Xcelleration (Oasis Mor remaining gravity processi	a transfer to a PC. ntaj module from Geosoft) for all ng.
BASE STATION	The survey is tied to Cana (CGSN) gravity base 922 office of Hearst. Two seco 8888 were established re Hotel and on the survey description of the gravity b	adian Gravity Standardization Network 29-1976 located at the Canada Post ondary control bases HJH-9999 & BF- espectively at the Howard Johnson / grid, and used daily (refer to the ase BF-8888 at page 13).



□ *FIELD PROCEDURES* Each loop began and ended with a reading at the **HJH-9999** and **BF-8888** base stations. Only gravity readings gathered at the local base **BF-8888** were used during the processing to correct for the residual instrument drift. The closure errors (drift) correction was linearly applied to the data included within the loop.

At each station, the following parameters were recorded in the field notebooks:

- Station identification;
- Local time in UTC;
- Standard deviation of each gravity reading;
- Instrument readings;
- Instrument base height above ground;
- Slopes to the centre of each of the four Hammer zone sectors:B (2-12 m) and C (12-50 m).

The following data were recorded in the instrument's memory:

- Station identification;
- Instrument readings;
- The standard deviation of each reading;
- Tilt X axis in arc-seconds, at the end of the reading;
- Tilt Y axis in arc-seconds, at the end of the reading;
- The gravity sensor temperature compensation factor;
- The tide correction for each reading;
- The number of samples that were read;
- The number of rejected samples;
- Local time in UTC.

The first step in data acquisition is to locate the most appropriate site for both GPS and gravity readings. The gravity meter is then placed on the tripod; the meter is levelled with the tripod levelling screws. Before the reading is to be initiated the operator stands still to avoid inducing ground motion that could affect the gravity meter. Once the above requirements have been met, the reading is taken.

The gravity meter was put in its padded and insulated carrying bag immediately after each reading.

At the end of each day, the data were downloaded from the CG-5u Autograv gravimeters through an USB cable to a computer. The Scintrex *SCTutil* program was used to facilitate the data transfer.



	Xmet Inc.	Hearst 9229-1976 CGSN Gravity Base	
Station number:	9229-1976	9229-1976 base	
Name:	Hearst		
Area:	Town of Hearst, Ontario	Front St. Concessor	
NTS sheet:	42G/12		
Geodetic system:	NAD 83, UTM Zone 16N		
Longitude:	83°39' 54.41" W		
Latitude:	49° 41' 16.00" N		
Easting:	740 523.694 ± 0.010 mE		
Northing:	5 509 257.800 ± 0.010 mN		
Elevation (HTv2):	242.994 ± 0.016 m		
g _{ABS} (fixed):	980 953.970 mGal		
Primary station:	IGSN71	Image agrandie	
Positioning:	Leïca Viva DGPS	99 Aue 31 35	
The CGSN gravity base station is located at the Canada Post Office of Hearst at the northern entrance on 9 th street, 76 cm from the NW corner of the building. The station is marked by an aluminum disc.		Bureau de Poste Post Office	

Figure 5. Hearst 9229-1976 CGSN gravity base description



Xmet Inc.		BF-8888 Secondary Control Base
Station number:	BF-8888	BF-88888 base
Name:	Black Flake	
Area:	MNR District, Hearst, Ontario	
NTS sheet:	42K/01	
Geodetic system:	NAD 83, UTM Zone 16N	
Longitude:	83°15' 45.09" W	
Latitude:	50° 02' 28.43" N	
Easting:	696 005.933 ± 0.010 mE	L450 N
Northing:	5 546 805.129 ± 0.010 mN	L150 N L150 N
Elevation (HTv2):	137.045 ± 0.016 m	
g _{ABS} (calculated):	980 990.379 mGal	11° 18° 218° 218° 218° 218° 218° 218° 21
Primary station:	Tied to 9229-1976 CGSN	
Positioning:	Leïca Viva DGPS	
The gravity base station is located in the survey grid near the line 1500N, 17.5 M from the station 2161. The station is not monumented.		N° 21° 21° 21° 21° 21° 21° 21° 21° 21° 21

Figure 6. BF-8888 gravity base description



GRAVITY QC'S

Before the survey:

- The gravity meter had been heated and stabilized for more than a week.
- ✓ All the internal constant values had been checked.
- ✓ Temperature compensation had been fine-tuned to \pm 0.1 mGal/mK.
- The X and Y tilt sensitivities had been checked and did not require any adjustments.

During the survey:

✓ The drift was controlled after each loop. The absolute average closure error was found to be:

Drift	Contract	Absol	ute drift
(mGal)	Requirements	Meter 418	Meter 430
Per loop	0.100	0.024	0.023
Per hour	0.010	0.004	0.004

Table 2. Absolute average closure error

Chart of the daily closure errors (drift) of the gravimeters 418 and 430 are presented on page 15.

- ✓ While readings were recorded, the operator observed the tiltmeter to ensure that the instrument stayed levelled well within \pm 20 arc sec range.
- ✓ At least two readings were taken at every station. A third and even a fourth one was recorded if the first two were more than 0.005 mGal apart.

At Operation Headquarters:

- Dump files (YYMMDD_#Grav.txt) were inspected to detect any spurious readings, for instance:
 - reading time: < 30 seconds
 - standard deviation: > 0.1 mGal
 - rejected samples: > 5

506 unique stations were surveyed, **20** readings (4%) were randomly repeated for quality control. The average repeat error of the absolute gravity g_A was **0.006** mGal and the standard deviation was **0.008** mGal.

The maximum error value was found to be **0.020** mGal at station **1022.**





Figure 7. Daily gravity loop closure errors of the gravimeter 418



Figure 8. Daily gravity loop closure errors of the gravimeter 430



6. GPS DATA ACQUISITION

□ *TYPE OF SURVEY* Real Time Kinematic (RTK) GPS surveying from an unknown geodetic point with an expected accuracy better than 5 cm in elevation and horizontal positioning.

INSTRUMENTATION	Base station:	Leica Viva GS15:	s/n: 1508148
	Mobiles:	Leica Viva GS15: Leica Viva GS15:	s/n: 1501746 s/n: 1506143

- GNSS PERFORMANCE
- Advanced measurement engine.
 - Jamming resistant measurements.
- Excellent low elevation tracking.
- Very low Noise GNSS carrier phase measurements with < 0.5 mm precision.
- 120 channels.

.

- Up to 60 Satellites simultaneously on two frequencies.
- Tracking Satellite signals:
 - GPS: L1, L2, L2C, L5
 - GLONASS: L1, L2
 - Galileo (Test): Glove-A, Glove-B
 - Galileo: E1, E5a, E5b, Alt-BOC
 - GNSS measurements:

- Fully independent code and phase measurements of all frequencies.
 - GPS: Carrier phase full wave length.
 - GLONASS: Carrier phase full wave length, Code (C/A, P narrow Code).
- Reacquisition time: < 1 sec.

Accuracy (rms) with Real-Time (RTK):

- Rapid static (phase): Horizontal: 5 mm + 0.5 ppm
- Static mode after initialization: Vertical: 10 mm + 0.5 ppm
- Kinematic (phase): Horizontal: 10 mm + 0.5 ppm
- Moving mode after initialization: Vertical: 20 mm + 0.5 ppm

Accuracy (rms) with Post-Processing:

- Static (phase) with log observations:
 - Horizontal: 3 mm + 0.1 ppm
 - Vertical: 3.5 mm+0.4 ppm
- Static and rapid static (phase):
 - Horizontal: 5 mm + 0.5 ppm
 - Vertical: 10 mm + 0.5 ppm
- Kinematic (phase):
 - Horizontal: 10 mm + 1 ppm
 - Vertical: 20 mm + 1 ppm

□ SOFTWARE

LEICA Geo-Office 8.2



COORDINATE SYSTEM	Projection:	UTM Zone 16N
	Type:	Transverse Mercator
	Datum:	NAD 83
	Central meridian:	87°00' W
	Central scale factor:	0.9996
	False easting:	500 000 m
	False northing:	0 m
	Ellipsoid:	GRS 1980
	Geoïd model:	HTv2.0

□ *GPS BASE STATIONS* The DGPS survey is tied to the geodetic point **HRST RACS GLCORS**, Two non-monumented control bases (Check-10 & Check-11) were established on the grid and used for the survey duration. Positions of these two GPS base stations were tied to the geodetic point **HRST RACS-GLCORS** located at Hearst.

Base	Easting (mE)	Northing (mN)	Elevation
HRST RACS GLCORS	751739.16	5507501.89	267.362
Check-10	695995.03	5546789.19	137.365
Check-11	696001.11	5546804.90	136.836

□ *GPS PROCESSING* The coordinates were processed and recorded in real-time (in the field) with the radio information transmitted by the base station. Data were downloaded from the GPS SD cards to a computer and checked after each survey day. Secondary processing was performed at head office in Val-d'Or, QC.

□ GPS QC's

Before the survey:

- The GPS units were verified against three provincial calibration benchmarks in Val-d'Or, QC.
- Level and graduated staff accuracies were checked.

During the survey:

- The distance between the antenna and the ground was measured to a precision of 1 millimetre and was then entered on the keypad.
- The gravity station number was manually incremented on the keypad-display unit and then checked against the number recorded in the gravimeter.
- The base station unit provided real-time corrections of the satellites time-tagged information to the mobile unit every second.
- ✓ The base station unit also took a recording automatically every 5 seconds for static post processing.



At Operation Headquarters:

- ✓ The average precision of the elevation values as determined by *LEICA Geo-Office* 8.2 is **1.6 cm** with a standard deviation of **0.7 cm**. The maximum value was found to be **4.8 cm** at station **2008**.
- The average precision on the horizontal positions (X & Y) is
 1.0 cm with a standard deviation of **0.4 cm**.
- Re-visited stations were processed daily to identify any possible weakness in the survey method. 22 stations out of 506 were re-visited (4.3%). The average absolute error of the elevation was calculated to be 0.4 cm with a root mean square (RMS) error of 0.5 cm. The maximum error values were found to be 1.1 cm at station 2081.



7. DATA PROCESSING AND DELIVERABLES

- □ TOTAL MAGNETIC FIELD CONTOURS The total magnetic field was gridded using a bi-directional gridding method (BIGRID GX) with grid cell size of 25 m. This method is ideal for line-oriented data as it inherently tends to strengthen trends perpendicular to the survey line direction. One pass of a 3 x 3 Hanning filter was then applied to improve the overall appearance of the final *Total Field Contours* map (1.2). For the Blackflake grid, the Geosoft colour table was used with linear intervals of 5 nT, from 57 050 to 57 370 nT.
- MAGNETIC GRADIENT (FIRST VERTICAL DERIVATIVE)
 The first vertical derivative was calculated from the total magnetic field grid. The Geosoft colour table was used with linear intervals of 2 nTm, from -4 to 4 nT/m.
- □ *IPOWER3D[®] QUALITY CONTROL* The first step in processing IPower3D[®] data is quality control. The IPower3D[®] configuration takes a large number of readings using different electrode orientations to thoroughly investigate the subsurface in 3D. Because of the varying geometry used there are a small number of readings that are not at favourable dipole orientations. IPower3D[®] incorporates a high degree of redundancy, so a moderate percentage of readings can be rejected, without compromising survey coverage.

To ensure consistent and efficient quality control Abitibi Geophysics has developed $ProSysControl^{TM}$ This application analyses the normalized decay curve for each reading within the data set. Only readings that successfully pass quality control will be used to calculate the final chargeability. This software also allows the user to view each decay curve for additional manual quality control.

Figure 9 is a screen grab from $ProSysControl^{TM}$ showing an alarm for high contact resistance (red box) that has been accepted (green bar) and plot, showing the decay curve (red) and normalized decay curve (blue) for the selected reading, highlighted in blue.

□ IPower3D[®] INVERSION Apparent resistivity and chargeability values were inverted using RES3DINV x 64 version 3.04.98 from GEOTOMO (www.geoelectrical.com). This software calculates three dimensional patterns of resistivity and chargeability of the subsurface that best explain the values recorded at surface. The software generates a model consisting of rectangular prisms and applies a nonlinear algorithm to minimise the difference between the calculated model and field measurements.



 $\square ABSOLUTE GRAVITY (g_A) The observed gravity values were first imported into Geosoft Xcelleration along with the elevation data. In view of the high accuracy of the gravity readings, the absolute RMS error of the observed gravity repeat values was$ **0.006**mGal.

Field observations corrected for earth tides, long term instrument drift and corrections made for instrument height and residual instrument drift. These values were estimated as following:

- ✓ Tidal correction is re-calculated in Xcelleration using the station location and measuring time. The accuracy of this correction is better than 0.001 mGal.
- ✓ New drift correction factors of 0.296 and 0.143 were calculated for the gravimeters 418 and 430, respectively to automatically correct the long term instrument drift during the survey.
- ✓ In presence of the snow, the value of the instrument height is typically accurate within \pm 3.0 cm which is equivalent to an error of \pm 0.009 mGal.
- ✓ The residual drift of the gravimeter is assumed linear between two readings at the base station. The closure error was linearly distributed on the data in the corresponding loop, taking into account the instrument height and the tidal correction. Non-linear drifting could introduce an error of up to 0.005 mGal in the middle of a loop.

The maximum error on g_{A} is therefore estimated to be $\pm\,0.021\,\,\text{mGal.}$

No re-visited station has exceeded that value.

Density " ρ ": a mean crustal rock density of **2.75** g/cm³ was taken for the Bouguer correction.

Latitude Correction: both the rotation of the Earth (generating centripetal acceleration) and its slight equatorial bulge produce an increase of gravity with latitude. This effect is corrected when computing the theoretical value of gravity at the surface of the reference ellipsoid using the Geodetic Reference System GRS67:

 $g_{L}(\phi) = 978\ 031.846\ \{1.0+0.005278895\ \sin^{2}(\phi) + 0.000023462\ \sin^{4}(\phi)\}\ mGal$

For local surveys, only the gradient due to latitude is important:

 $g_{L} = 0.0000812132 \sin(2\phi) Y_i$ mGal/m

With dual frequency DGPS, the horizontal (latitude) position is accurate to better than ± 5.0 cm, so the error is less than **0.001** mGal.



 $\square BOUGUER ANOMALY (g_B) (CONT'D) Free distant$

(*Free Air Anomaly*: since gravity varies with the inverse of distance squared, it is necessary to correct for changes in elevation between stations so that all field readings are reduced to a datum surface. The Free Air anomaly at elevation "h" is:

 $g_F = g_A - g_L + 0.308596 * h$

Simple Bouguer Anomaly: it accounts for attraction of bedrock of density " ρ " between the station and the datum plane, which was ignored in the Free Air calculation. The Bouguer anomaly is:

 $g_B = g_F - 0.0419088 * \rho * h$

Using a density of 2.67 g/cm³ and a typical accuracy of \pm **1.6 cm** in elevation, the error on the combined Free Air and Bouguer correction is \pm **0.003** mGal.

\Box ACCURACY OF g_{B}

Table 4. Accuracy of g_B

Ston	Source / Evaluation	Error (mGal)				
Step	Method	Maximal	Typical			
g _A	Re-occupied stations	0.021	0.009			
g∟	Latitude correction	0.001	0.000			
g в	Free Air and Bouguer	0.009	0.003			
Estimated err	or \mathcal{E}_{g_B} on g_B^*	0.023	0.009			
Contract requi	rement *	0.050	0.050			

excluding terrain correction,

$$\varepsilon_{_{g_B}}=\pm\sqrt{\varepsilon_{_{g_A}}^2+\varepsilon_{_{g_L}}^2+\varepsilon_{_{g_B}}^2}$$
 , mGal

□ *TERRAIN CORRECTION* Terrain correction allows for surface irregularities in the vicinity of the station, not terrain accounted for in the Simple Bouguer anomaly calculation. An excess mass in the form of a positive topographic feature adjacent to a gravity station exerts an upward pull on the gravity meter, thus lowering the observed reading. A mass deficiency in the form of a valley adjacent to a gravity station causes the gravity field to be lower than it would be where the terrain is flat. Terrain corrections compensate for these effects and are therefore always positive.



TERRAIN CORRECTION
 (CONT'D)

Inner Terrain Correction (ITC):

The ITC, which corrects for local topographic features (terrain effect up to 50 m from the station), is performed using the Hammer method (Geophysics, 1939). This method involves creating a set of concentric rings, or zones as they are referred to here, each of which is divided into a specified number of equal-sized segments. The average elevation for each of these segments is entered relative to the elevation specified for a given station. In this way, the topography immediately surrounding each station can be defined and the inner terrain correction calculated:

$$g_{T} = \gamma * 0 * \theta * \left(r_{2} - r_{1} + \sqrt{r_{1}^{2} + z^{2}} - \sqrt{r_{2}^{2} + z^{2}} \right)$$

Where $\gamma =$

 $\gamma = 6.67 \times 10^{-8} \text{ dyne cm}^2/\text{g}^2$

- θ = sector angle (in radians)
- $\rho = density (in g/cm^3)$
- $r_1, r_2 =$ inner and outer sector radius (in cm)
- z = average height difference between the station and the sector (in cm)

	Zone B	Zone C
r ₁ :	2 m	12 m
r ₂ :	12 m	50 m
Number of sectors:	4 (θ = 1.57)	4 (θ = 1.57)

The slopes of hammer Zones B and C were collected at each gravity station during the fieldwork operations and the inner terrain correction was calculated using our Excel code and added in the final processing.

Outer Terrain Correction (OTC):

The OTC, which corrects for more distant topographic features, was not performed in this project because the DEM grid of 50 m resolution (cell size) provided by the client is not wide enough to perform this correction properly.

In this project only near (inner) terrain correction was used to calculate the Complete Bouguer anomaly. The obtained near terrain effect values range from 0 to 0.257 mGal for an average of 0.004 mGal on the Black Flake property.



 COMPLETE BOUGUER ANOMALY
 The Bouguer anomaly was calculated employing standard reductions using a mean crystal rock density value of 2.67 g/cm³.

The reduced data are interpolated to a square grid of 35 m cell size using a minimum curvature method (RANGRID) and followed by a second gridding (regrid) process with a cell size of 5 m to improve the overall appearance of the final data (map 2.2). The Oasis Montaj color table Clra64.tbl was used with linear intervals of 0.05 mGal from -55.25 to -56.95 mGal.

 RESIDUAL ANOMALY CONTOURS
 The residual anomaly was generated by subtracting the regional field from the Bouguer gravity anomaly (map 2.5).

The Oasis Montaj color table (Clra64.tbl) was used with linear intervals of 0.2 mGal from -0.3 to 0.5 mGal.

□ VERTICAL GRADIENT CONTOURS Using a convolution filter method, the first vertical derivative (vertical gradient) of the Bouguer anomaly was calculated (map 2.7). Two passes of a Hanning 3 x 3 filter were applied to the grid to improve the overall appearance of the final contours map.

The Oasis Montaj color table (Clra64.tbl) was used with linear interval of 0.0001 mGal/m from -0.003 to 0.003 mGal/m.

MAPS PRODUCED
 The following colour maps (next page) are bound or inserted in pouches at the end of this report.

Our Quality System requires every final map to be inspected by at least two qualified persons before being approved and included within a final report.



Table 5. Maps produced

Map Number	Description	Scale
L6+00N to L12+00N (5 plates)	IPower3D [®] Induced Polarization Survey – Vertical sections	1:10 000
1.1	GPS-positioned Magnetic Field Survey – Total Magnetic Intensity Anomaly Profiles (nT)	1:5 000
1.2	GPS-positioned Magnetic Field Survey – Total Magnetic Intensity Anomaly Contours (nT)	1:5 000
1.4	GPS-positioned Magnetic Field Survey – Calculated Vertical Gradient Contours (nT/m)	1:5 000
2.2	High-Resolution Ground Gravity Survey – Bouguer Anomaly Contours (mGal) (Slab density: 2.67 g/cm ³)	1:5000
2.3	High-Resolution Ground Gravity Survey – Elevation Contours (m)	1:5000
2.5	High-Resolution Ground Gravity Survey – Residual Anomaly Contours (mGal) (1 st order polynomial trend)	1:5000
2.7	High-Resolution Ground Gravity Survey – First Vertical Derivative Contours (mGal/m)	1:5000
8.2_150	IPower3D [®] IP Survey – Inverted Resistivity at a depth of 150 m (Ohm-m)	1:5 000
8.2_250	IPower3D [®] IP Survey – Inverted Resistivity at a depth of 250 m (Ohm-m)	1:5 000
8.3_150	IPower3D [®] IP Survey – Inverted Chargeability at a depth of 150 m (mV/V)	1:5 000
8.3_250	IPower3D [®] IP Survey – Inverted Chargeability at a depth of 250 m (mV/V)	1:5 000
8.6_150	IPower3D [®] IP Survey – Calculated Gold Index at a depth of 150 m	1:5 000
8.6_250	IPower3D [®] IP Survey – Calculated Gold Index at a depth of 250 m	1:5 000
10.0	Geophysical Interpretation	1:5 000

DIGITAL DATA

The above-described maps are delivered in the Oasis Montaj map file format on DVD-Rom.

A copy of all survey acquisition data (ASCII text format) and processed data (Geosoft Montaj databases) are also delivered on DVD-Rom.



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Figure 9. Screen grab from ProSysControl, Abitibi Geophysics' proprietary QC software



8. RESULTS AND RECOMMENDATIONS

□ MAGNETIC

The analysis of the *Total Magnetic Field profiles* (1.1), *contours* (1.2) and the *Calculated Vertical Gradient Contours* (1.4) from the Blackflake survey has shown several magnetic trends throughout the grid. There are **4** magnetic highs, labelled **MH-01** to **MH-04** as well as **3** magnetic low bodies, coloured in green on the *Geophysical Interpretation* map (10.0). The magnetic high trends correlate in strike and location with known magnetie rich diabase dikes that cut through the grid at ~130°. There are several circular magnetic low regions throughout the grid as well. Two of these, located near the center of the grid, between lines 6+00N and 12+50N also correlate with anomalous gravity lows.

Figure 10 below shows the 3D inversion results of the magnetic data. Negative susceptibility values are represented by the blue isosurfaces and range from -0.003 to -0.005 and positive susceptibility values are represented in hot colors and range from 0.002 to 0.007. The two positive features striking NW/SE correlate to the diabase dikes in the area. Negative values represent values below the average magnetic susceptibility.



Figure 10. Magnetic 3D model with susceptibility isosurfaces at 1100 m depth, cool colors are negative susceptibility and hot colors are positive susceptibility (looking northeast)



□ THE IPOWER3D[®] SYSTEM

The IPower3D[®] configuration has been designed to maximize the sensitivity of the induced polarization survey and is especially effective under conditions of high conductance overburden (conductance is the product of conductivity and thickness). The values of apparent chargeability measured in the field by the IPower3D[®] configuration can be several times greater than the values measured by a conventional electrode configuration, however the background response is not amplified as it occupies the entire range of sensitivity (both positive and negative). The advantage of the IPower3D[®] configuration is that it is able to detect anomalies that would otherwise be within the noise envelope of conventional IP arrays.

The IP response of this grid is very low; with apparent chargeability anomalies having small amplitudes up to 5.68 mV/V. The inversion results of the IPower3D[®] data show that the intrinsic chargeabilities are within the range that would be expected with the inversion of data collected using a conventional electrode configuration such as dipole-dipole. The disadvantage of conventional configurations is that sensitivity is often not sufficient to allow the inversion to resolve the source of a polarizable anomaly beneath thick conductive overburden.

Resistivity and chargeability anomalies have been interpreted by studying the 3D inversion models, the true-depth sections and the inverted resistivity and chargeability maps. A total of **3 anomalous trends** have been interpreted, the inferred surface projection of the resistivity / chargeability sources are shown along the survey lines on the *Geophysical Interpretation* map (10.0) and on the true-depth section plates. The anomalies have been correlated from line-to-line and are fully described in appendix B, found at the end of this report.

□ RESISTIVITY

The resistivity inversion indicates a thick layer of conductive overburden. This can be seen clearly on the vertical sections included at the end of the report. It's variable thickness varies, ranging from ~50 m in the west and generally thickening to the east, up to ~200 m in some areas. Underneath this thick layer the ground is more resistive, with values increasing to the east.

A resistive zone has been interpreted from the data and is defined by the 1000 Ω -m contour and can be seen on the *Geophysical Interpretation* map (10.0) in blue. The chargeable anomaly **BF-01** is associated, at least partially with this wide zone of high resistivity.

A resistive low zone has been interpreted as well. This zone has a resistivity of up to 200 Ω •m and is very small. This zone is associated slightly with chargeable anomaly **BF-03**.

CHARGEABILITY

The IPower3D[®] survey has revealed 3 regions of anomalous chargeability. The chargeable anomaly **BF-01** is nearer to surface (~30 m below surface) in the south at L6+00N and starts to dip towards the north (~300 m below surface). Chargeable anomaly **BF-02** has the same signature as **BF-01**; however because of its close proximity to **BF-01** it is only distinguishable as a separate anomaly from L9+00N to L12+50N. The anomaly **BF-03** has the strongest chargeable response of the three (3) anomalies present on the Blackflake grid, however caution should be exercised in the interpretation of anomalies occurring at the ends of lines as the inversion is not as well constrained at the edge of the grid and spatial relationships may not be accurate.





Figure 11. IPower3D[®] model with chargeability isosurfaces sitting on a resistivity surface at 490 m depth, cool colors are resistive and hot colors are conductive (looking northeast)

GOLD INDEX

In addition to resistivity and chargeability, the sections also display the calculated Gold Index, this value is the product of the squared chargeability multiplied by the resistivity ($M^2 * R$). This highlights regions that are both resistive and chargeable, helping to localize the areas with a high potential for hosting gold mineralization. The Gold Index can be a useful tool in exploration for disseminated mineralization associated resistive material such as quartz veins, or silicified zones. In the case of the Blackflake property, most of the survey area is associated with high resistivity, once past the conductive overburden. The Gold Index highlights anomaly **BF-01** nicely throughout its entire trend.

DRILLING RECOMMENDATIONS

IPower3D[®] Drilling Recommendations

Diamond drilling is recommended in order to test the **BF-01** anomaly. Figure 11 below illustrates the drilling target; this information is shown in table 6 and in appendix A.

	DDH	Target (<mark>not collar, cen</mark>	ter of target)	
Anomaly	Easting	Northing	Estimated vertical depth below surface (m)	Priority
BF-01	694 850	5 546 150	250	1

Table 6. IPower3D[®] Diamond drilling targets



PRIORITY 1 TARGET



Figure 12. Proposed DDH testing anomalies **BF-01** (looking north)

GRAVITY

The Bouguer gravity anomaly is predominantly negative, ranging from -55.25 to -56.95 mGal, (average -56.10 mGal). The most dominant feature on the *Bouguer* map (2.2) is the gravity high trending NW/SE on the western side of the grid. This feature correlates well with the diabase dikes in this area and also with the magnetic high trends. Note that there is no correlation between the pattern of the Bouguer anomaly and its associated topography as seen on the *Elevation Contours* map (2.3).

Upward residual gravity anomaly with the calculation of the first vertical derivative is performed in order to extract the local gravity anomalies from the regional field.

Feature A is a gravity high trend of 0.3 mGal in amplitude striking NW-SE, for the length of the grid and up to 440 m width. Referring to the geological information of the Blackflake property, anomaly **A** corresponds to known diabase dikes at this location. The *magnetic* map of the property (1.2) shows that the feature **A** is also magnetic and the chargeable anomaly **BF-01** is also associated with this anomaly.

Feature B is gravity high of 0.26 mGal in amplitude, open ended to the northeast, and ~300 m wide.

Feature C is a weak, circular gravity low of -0.1 mGal in amplitude, open ended to the west, and ~200 m wide.

Feature D is a weak gravity low of -0.08 to -0.12 mGal in amplitude, open ended to the west and south.

Feature E is a group of small, circular and weak gravity anomalies varying from -0.08 to 0.12 mGal in amplitude. There is a circular magnetic low association with the southern circular gravity low in this group and can be seen on the *Geophysical Interpretation* map (10.0).

Feature F is a small, circular and weak gravity anomaly of 0.12 mGal in amplitude, and ~250 m wide. There is a circular magnetic low associated with this gravity low and can be seen on the *Geophysical Interpretation* map (10.0).



Discussion

The objectives of this survey were to identify zones amenable to graphite and/or kimberlite mineralization. The type of graphite mineralization being explored for is hydrothermal vein style located within breccia pipes. A circular gravity low paired with a circular magnetic anomaly would be the response expected. Based on the response seen from gravity features **E** and **F** an IPower3D[®] survey was conducted over this region, L6+00N to L12+50N. This survey would yield a strongly conductive and chargeable feature in this same locale if the mineralization was graphite. However the chargeability response over this region was very low, indicating a lack of graphite mineralization.

A similar sort of response would be expected from to a kimberlite pipe, a circular gravity low anomaly paired with a circular magnetic anomaly. Typically the magnetic response would be a high due to the presence of magnetite; however, some kimberlites do not contain magnetite. Again, we would expect a circular conductive feature due to the clay depression in the upper section of the pipe. This grid does display a thick layer of conductive overburden that is potentially masking this response. Gravity anomalies **E** and **F** may be attributed to an overburden trough over a kimberlite pipe and are recommended for follow up. However it should be noted that this gravity low region may be due to sediment accumulation.

The chargeable anomaly **BF-01** is located within a resistive region. This chargeable and resistive feature may indicate disseminated sulphide mineralization within a quartz vein or silicified region. This is a geological environment that may be amenable to gold mineralization.

□ SURVEY EXTENSIONS

It is recommended to extend the IPower3D[®] survey to the north, south and east to further delineate the chargeable anomalies discussed above.



The interpretation of the geophysical data embodied in this report is essentially a geophysical appraisal of the Blackflack Project. As such, it incorporates only as much geoscientific information as the author had on hand at the time. Geologists thoroughly familiar with the area may be in a better position to evaluate the geological significance of the various geophysical signatures. Moreover, as time passes and data provided by follow-up programs are compiled, the priority and significance of exploration targets reported in this study may be downgraded or upgraded.

Respectfully submitted, Abitibi Geophysics Inc.



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PC/mw

Cim

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DESCRIPTION OF THE IP / RESISTIVITY ANOMALIES INTERPRETED ON THE BLACKFLAKE PROJECT

		Location		Cont	rast	Strike		Rec					
Anomaly	Line	From	То	Charg.	Res.	Length & Orientation	Comments	P: Prospect DDH: Drillin X: Survey e	Priority				
	6+00N	5+25W	6+00W	1	(R)			X:	survey extension to the north and				
	7+50N	5+25W	6+00W	1	(R)	600 m N-S (Open to the north and south)	(R)600 m N-S(R)(Open to the north and south)	600 m N-S (Open to the north and south)	600 m N-S (Open to the north and south)	Weakly chargeable anomaly		south	
BF-01	9+00N	5+50W	6+00W	1	(R)					associated partially with the resistive zone in the south. This source has	DDH: Northing:	Proposed DDH 5 546 150	1
	10+50N	5+75W	6+25W	1	-					south)	- south)	mineralization.	Easting: Azimuth:
	12+00N	5+75W	6+50W	1 -					Dip: Depth:	-45° 675 m			
	9+00N	3+50W	4+00W	1	-	300 m							
BF-02	10+50N	3+50W	4+00W	1	-	N-S (Open to the	N-S (Open to the	N-S (Open to the	Weakly chargeable anomaly.	X:	survey extension to the north	2	
	12+00N	3+50W	4+25W	1	-	north)							
	6+00N	9+25E	East end	1	-		Moderately to weakly chargeable						
	7+50N	9+50E	East end	1	-	600 m	association. This source is located at	y .	survey extension				
BF-03	9+00N	9+25E	East end	1	(C)	N-S (Open to the	the eastern extent of the survey area.	^ .	to the north, south	3			
	10+50N	9+25E	East end	2	-	north, south and east)	edge of the survey are not reliable		and east				
	12+00N	9+25E	East end	2	-		delineate this target.						

*pending favourable prospecting results

Legend									
Chargeability	Increase:	? = Marginal 1 = Weak 2 = Moderate 3 = High 4 = Very High							
Resistivity	Increase:	 ↑ = Resistive ↑↑ = Very Resistive (R) = Wide Resistive Zone 							
roolouvity	Decrease:	$\begin{array}{l} \downarrow = \text{Conductive} \\ \downarrow \downarrow = \text{Very Conductive} \\ (C) = \text{Wide Conductive zone} \end{array}$							



APPENDIX B

IPOWER3D[®] INDUCED POLARIZATION SURVEY - VERTICAL SECTION







