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Technical Report
On the
Gathering Lake Lithium Pegmatite Property

Thunder Bay Mining District
Northwestern Ontario, Canada

Prepared for:

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January 20th 2022

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1.0 SUMMARY

The Gathering Lake Lithium Pegmatite Property covers approximately 364 hectars of land in the South Beatty Lake and Gathering Lake Areas. Most of the historic work was accomplished in 1955 to 1957. It is located approximately 173 to 175 km northeast of Thunder Bay and 43km southwest of Geraldton. Geologically, the area is located within the Quetico Subprovince of the Superior Province. The Quetico Subprovince is composed of predominantly metasediments consisting of wacke, iron formation, conglomerate, and siltstone, which deposited between 2.70 and 2.69 Ga. The igneous rocks in the Quetico Subprovince include abundant felsic and intermediate intrusions, metamorphosed rare mafic and felsic extrusive rocks and an uncommon suite of gabbroic and ultramafic rocks. The earlier felsic intrusions occurred 5 to 10 million years after the accumulation of sediments and are interpreted to be I-type intrusions. The later felsic intrusions occurred 20 million years after the sedimentation and are designated as S-type. The pegmatites in the Quetico Subprovince which contain lithium and rare metals (beryllium, tantalum, niobium and tin) are hosted by metasediments and by their parent granite.

The pegmatite dykes, sills and lenses can be subdivided into rare-element pegmatites and granitic pegmatites. The rare-element pegmatites are of economic significance and they contain microcline or perthite, albite, quartz, muscovite and spodumene and minor amounts of beryl, columbite-tantalite and cassiterite. The granitic pegmatites are like the irregular pegmatites described above except that they contain more abundant plagioclase. Some of the pegmatites are parallel to the foliation or bedding of the metasediments, whereas others occur in joints in either the metasediments or granite. Contacts are usually sharp and, except where dykes cut granitic rocks, often found to be marked by a thin border zone of aplite or granitoid composition. A few pegmatites are internally zoned with mica-rich or tourmaline-rich rock along or close to the walls and quartz cores.

In 2008, F. Breaks, O.G.S.'s expert on rare-element pegmatite deposits, in his study of the Georgia Lake rare-element pegmatite field, discovered a new pegmatite group (Gathering Lake Pegmatite Group) that contained beryl-type and albite-type pegmatites. Breaks stated that this new pegmatite group has potential for a rare-element pegmatite deposit. Breaks located, on roads, nine Ta-Nb-Oxide bearing pegmatites within the Gathering Lake pegmatite group. This reports covers the attempt by the author to discover additional dykes using ground geophysics.

2.0 INTRODUCTION

2.1 Purpose of Report

The present report summarizes the ground geophysics performed by the author and contractor on the Gathering lake Lithium Project.

2.2 Sources of Information

This report is based on published assessment reports available from the Ministry of Northern Development, Mines (MNDM) Ontario, and published reports by the Ontario Geological Survey (OGS), the Geological Survey of Canada ("GSC"), various researches, websites, and results of present exploration work. All consulted sources are listed in the References section. The sources of the maps are noted on the figures. The exploration work was carried out under the supervision of the author who worked and supervised on the property in June 2018.

3.0 PROPERTY DESCRIPTION AND LOCATION

The Gathering Lake Lithium Pegmatite Property consists of 1 provincial multi-cell mining claim which covers approximately 365 hectares of land in the South Beatty Lake and Gathering Lake Areas. The center of the claim block is located 173km northeast of Thunder Bay, ON and 43km southwest of Geraldton, ON. Claim data is summarized in the Table 1, while a map showing the claims is presented in Figure 2.

Table 1: Claim Data

Township	Claim ID	Type	Due Date	Status	Work Required	
GATHERING LAKE AREA	570582	Multi-cell Mining Claim	2022-01-23	Active	$7,200	

Figure 1: Property Location Map

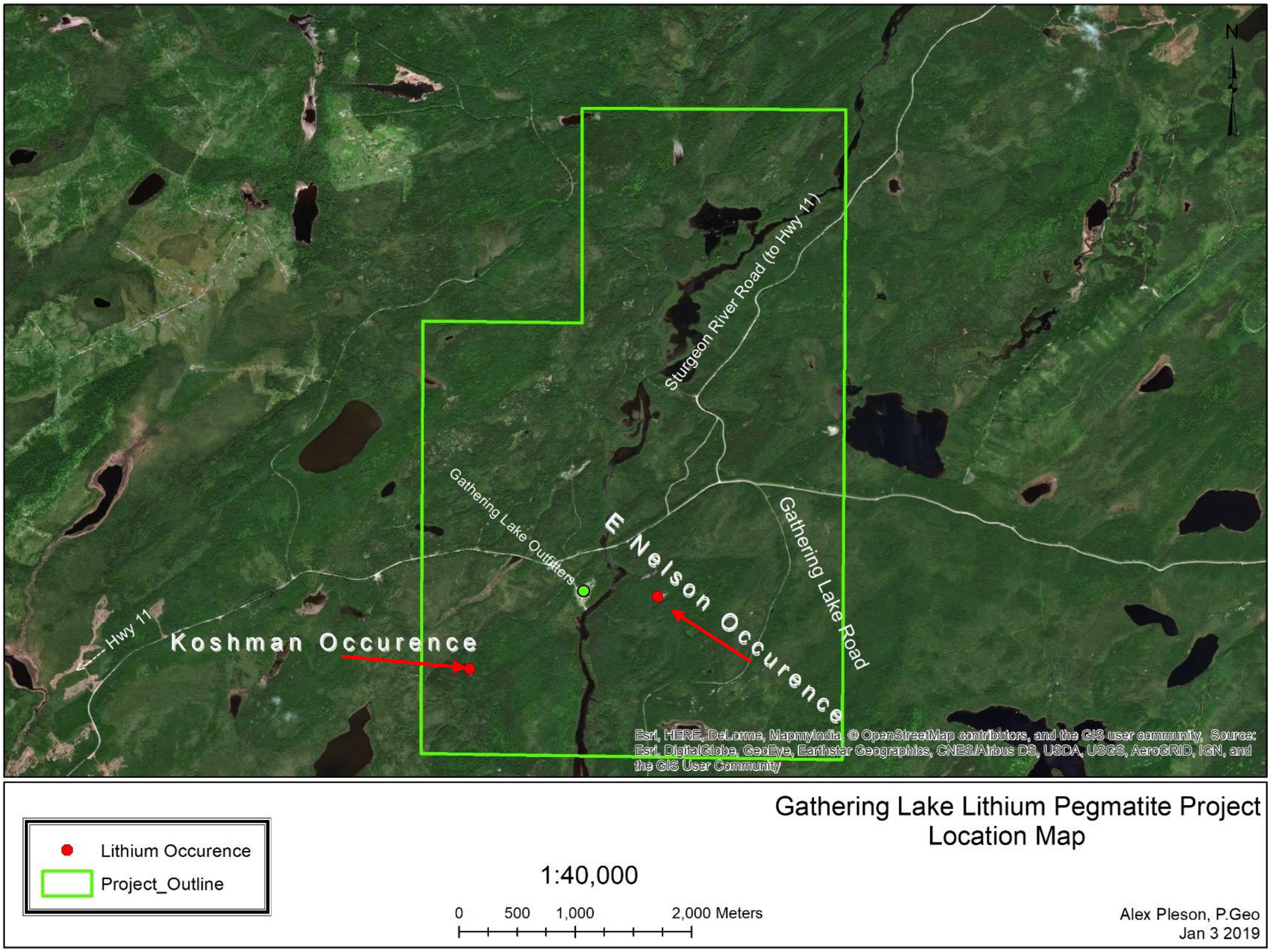
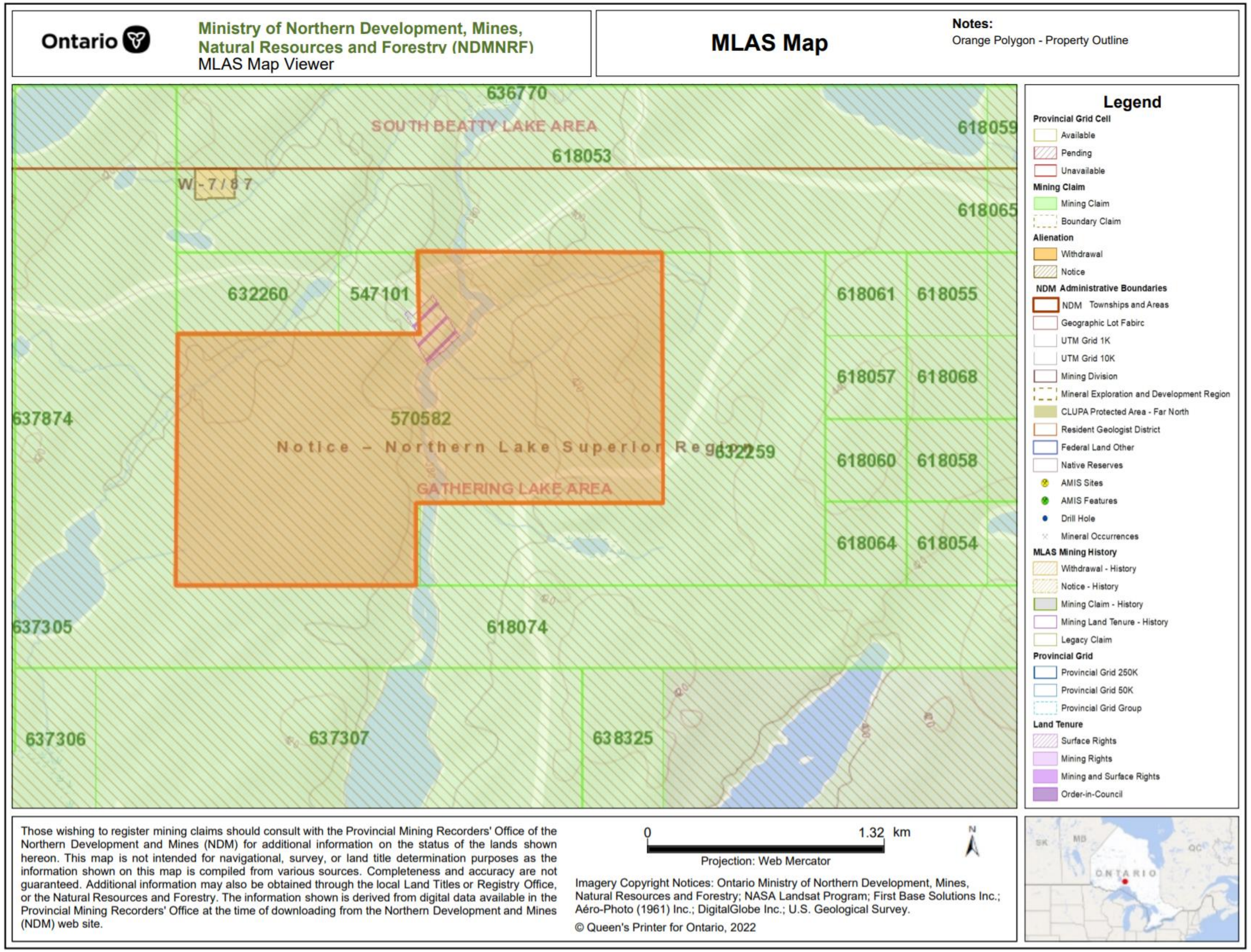


Figure 2: Mineral Claim Map



4.0 ACCESS, CLIMATE, PHYSIOGRAPHY, LOCAL RESOURCES, AND INFRASTRUCTURE

4.1 Access

The Gathering Lake project can be accessed by dirt roads off Highway 11 north of the town of Nipigon. The claims are accessed by driving 40 km north of the town of Nipigon on Highway 11, then driving approximately 59 km northeast on the Gorge Creek Road (Camp 75 Rd.) and Camp 51 road. An alternative route can also utilized from Highway 11 between Beardmore and Geraldton, ON along the Camp 51/Sturgeon River Road towards Gathering Lake for approximately 28km southwest.

4.2 Climate

The forest of the Gathering Lake area is mixed growth of spruce, balsam, jack pine, poplar, birch and cedar (Pye, 1965). Vegetation is typical of continental climate a mixture of coniferous (pine and black spruce) and deciduous (primarily birch and minor poplar).

The climate is continental with cold and long winters (from November to late March) and significant snow accumulations. The temperature in the winter months (January and February) can reach -40° C but typically ranges between -10° and -25°C. The Canadian Climate normals for 1971-2000 from Environment Canada (/www.climate.weatheroffice.gc.ca/climate_normals/) for Geraldton (closest weather station to the property) indicate that the daily average temperature ranges from -19ºC in January to 17ºC in July. The highest average accumulation of rain for a month is 112 mm in July. The highest average accumulation of snow for a month is 49 cm in November. The highest average snow depth is 48 cm in February. Drilling can be conducted year-round except for spring thaw in mid-March and April. Geological mapping and outcrop sampling can be conducted May to October when there is no snow on the ground.

4.3 Physiography

Pye (1965) summarized the topography of the Gathering Lake area:

“The area is one of topographic contrasts. The parts of the area in which metasediments are exposed are, for the most part, of low relief. In contrast, the parts underlain by granitic rocks are rugged, with rounded hills rising to about 150 ft. (=45.7 m) above the general level. Most conspicuous, however, are high, imposing vertical or near-vertical cliffs at the boundaries of large exposed sheet-like masses of diabase.”

“Rock exposures in the area are abundant, and between the outcrops there is a thin mantle of glacial deposits. These glacial deposits consist mainly of stratified accumulations of unconsolidated sand and gravel. Some of them represent a ground moraine sorted by the action of glacial meltwaters; others form prominent terraces along the shores of Lake Nipigon and in the valley occupied by Keemle and Wanogu Lakes, and are abandoned beach deposits. Esker ridges also are present but are not high and do not extend for any great distances."

The topography is moderate. The minimum elevation is 250 m and the maximum elevation is 560 m above sea level. Thus, the range is 310 m. The low-lying areas are typically underlain by metasediments and the higher areas are underlain by Nipigon diabase.

4.4 Local Resources and Infrastructure

The towns of Beardmore and Geraldton is the closest community, located approximately 40 km northwest and 44km northeast, respectively, of the project. Beardmore is part of Greenstone, an amalgamated town encompassing Nakina, Geraldton, Longlac, Beardmore, Caramat, Jellicoe, Macdiarmid and Orient Bay. The population of Greenstone is 4,906 people (Statistics Canada, www.statcan.gc.ca) and the population of Beardmore is approximately 150 people (http://www.highway11.ca/ThunderBay/06Beardmore). Beardmore has limited accommodation and restaurants.

The town of Nipigon, located about 50 km to the south of the Property has most of the basic supplies needed for exploration work. Nipigon has grocery stores, a hardware store, restaurants, hotels, a hospital and an OPP station. The population for Nipigon Township is 1,752 people in 2006 (Statistics Canada, www.statcan.gc.ca).

The town of Thunder Bay, located about 130-150 kilometres from the Property, is the largest city in Northwestern Ontario, serving as a regional commercial Centre. The town is a major source of workforce, contracting services, and transportation for the forestry, pulp and paper and mining industry. Thunder Bay is a transportation hub for Canada, as the TransCanada highways 11 and 17 link eastern and western Canada. It is close to the Canada-U.S. border and highway 61 links Thunder Bay with Minnesota, United States. Thunder Bay has an international airport with daily flights to Toronto, Ontario and Winnipeg, Manitoba, and the United States.

The city of Thunder Bay has most of the required supplies for exploration work including drilling and geophysical survey companies, grocery stores, hardware stores, exploration equipment supply stores, restaurants, hotels, and a hospital. The population of the city of Thunder Bay was 109,140 people in 2006 (Statistics Canada, www.statcan.gc.ca). Many junior exploration and mining companies are based in Thunder Bay, and thus the city is a source of skilled mining labour.

The Gathering Lake Outfitters lodge is in the center of the property and will provide room and board at a reasonable price during fishing and hunting seasons (see Figure 1).

There are several lakes, rivers and creeks in and around the Property area which can be a source of water. Power lines are also within a 30-kilometer range.

(Source: http://www.thunderbaydirect.info/about_thunder_bay

http://www.thunderbay.ca/Doing_Business/About_Thunder_Bay.htm)

5.0 HISTORY

The discovery of spodumene in the Georgia Lake area was summarized by Pye (1965):

"One of the topics featured on the program of the annual convention of the Prospectors and Developers Association in spring 1955 was the lithium deposits of the Preissac-Lacorne area in Quebec (Latulippe and Ingham 1955). Samples of the lithium-bearing mineral spodumene were on display. Many years ago, Eric W. Hadley of Auden had discovered a body of pegmatite forming a reef in Georgia Lake (now known as Island Deposit). He noted that the pegmatite contained a prismatic mineral, which he could not identify and which he considered then to be of no value. At the convention, however, he observed that the spodumene on display was very like the mineral in the pegmatite at Georgia Lake. He immediately contacted Gordon Miller of Conwest Exploration Company Limited. An examination was made at once, and impressed with the occurrence, Mr. Miller submitted samples to E.G. Pye for positive identification. Pye, in turn, presented the samples to Dr. H. Quackenbush, a Fort William dentist and amateur mineralogist, who as part of his hobby, had built a spectroscope. With this spectroscope, Dr. Quackenbush confirmed that the mineral was spodumene, and immediately Mr. Miller proceeded to stake a large group of claims for his company."

"As news of Hadley's discovery was publicized, prospectors entered the area. About 3,200 claims were staked and within a short time numerous additional lithium deposits were located. Many of these deposits were tested by diamond drilling in 1955 and 1956. Due to lack of adequate markets, however, none of these have been developed. Except for some limited diamond drilling by the Ontario Lithium Company Limited to test the original discovery in July 1957, the area has remained inactive since 1956" (as of Pye's 1965 report).

Detailed prospecting and diamond drilling to the west of the project was completed by Rock Tech Lithium Inc. (Rock Tech), Infinite Lithium Corporation and Ultra Lithium Inc. on several of their properties in the Georgia Lake area has lead to the discovery of undocumented lithium-bearing pegmatite dikes.

Historic exploration was carried out by E. MacVeigh, E. Nelson, and Standard Lithium Corp. from 1955-1957 which included prospecting, mapping, trenching, and drilling.

Modern exploration work was completed by John Scott (2012-2013) on behalf of Ken Fenwick. The results are listed in the table below.

Table 2: Historic Exploration Summary

Historic Work (after Fenwick, 2017)	
Period	Description of Work
1955-1957	Two lithium occurences were located and explored. Referred to as the Koshman and Nelson occurences
Sep-55	Diamond drilling on Nelson occruence (42E06NW0002) Sixteen samples were split from the core, totalling more than 100 feet of pegmatite and aplite. Because of the obvious barren nature of the bulk of the rock, only 3 samples, considered representative of all the pegmatite and aplite cored, were selected for assay. These ran: Hole No. / Footage 0.06% Li over 15.0 feet — 4 — 25.0-40.0 0.18 " " 5.0 " — 6 — 40.0-45.0 0.36% " " 5.8 " — 7 — 64.0-69.8 In the district, 1.0% Li is considered marginal ore. The remaining 13 samples will be retained at the Blind River office.
1957	Geological map of Koshman was made describing the local geology and prior trenching work completed. This indicated up to 15% spodumene within a pegmatite dyke and a larger sill or dyke with widespread disseminated spodumene
2012	John Scott found spodumene bearing boulders on property considered to be associated to the Koshman mineralization, where they procuded assays up to 9250 ppm Li.
2013	John Scott located a large outcrop, white pegmatite (non-spodumene bearing) intrusion in proximity to the Nelson occurrence, portions of the large pegmatite samples up to 203 ppm Li. The 1955 drilling of this occurrence indicates much higher lithium grades, which intersected the spodumene zone of the intrusion.

Rock Tech has been active in the Georgia Lake area since 2010 and has completed over 12,100 m of diamond drilling. This work has lead to the discovery of a NI 43-101 resource consisting of 1.89 Mt grading 1.04% Li2O (measured), 4.68 Mt grading 1.00% Li2O (Indicated) and an Inferred resource of 6.72 Mt grading 1.16% Li2O on the Nama Creek Zone (See Rock Tech's news release dated August 2, 2018). This resource is located 7 km northwest of Bold's Jean claim group.

Two diamond drill holes completed by Rock Tech in 2011 intersected the No.4 Dike on the eastern side of the Parole Lake patented claims. Hole PL-11-01 and PL-11-02 were located approximately 250 and 300 m respectively from the boundary with Bold's newly acquired claims (See figure 3 in the Maps and Charts section). Hole PL-11-01 returned 7.29 m @ 1.76% Li2O (including 5.15 m of 2.29% Li2O) and Hole PL-11-02 returned 5.41 m @ 1.25% Li2O (including 3.0 m @ 1.77% Li2O). Reference: Caracle Creek International Consulting Inc., Author Adrian Peshkepia, M.Sc., P. Geo., Drill Report For 2010-2011 Winter Drilling Program, June 14, 2011, prepared for Rock Tech Lithium Inc.

6.0 GEOLOGICAL SETTING AND MINERALIZATION

6.1 Regional Geology

The Georgia Lake area is located within the Quetico Subprovince of the Superior Province. The Quetico Subprovince is bounded by the granite-greenstone Wabigoon Subprovince to the north and Wawa Subprovince to the south (Williams, 1991). The Quetico Subprovince is composed of predominantly metasediments consisting of wacke, iron formation, conglomerate, ultramafic wacke and siltstone, which deposited between 2.70 and 2.69 Ga. The igneous rocks in the Quetico Subprovince include abundant felsic and intermediate intrusions, metamorphosed rare mafic and felsic extrusive rocks and an uncommon suite of gabbroic and ultramafic rocks. The earlier felsic intrusions occurred 5 to 10 million years after the accumulation of sediments and are interpreted to be I-type intrusions. The later felsic intrusions occurred 20 million years after the sedimentation and are designated as S-type (White and Chapell, 1983).

The Quetico Subprovince was subjected to four deformational events between approximately 2700 and 2660 million years (Williams, 1991). The predominant stratigraphic-facing direction is north. Regional schistosity is variably developed and oriented and is interpreted to be the result of regional shortening and dextral shearing.

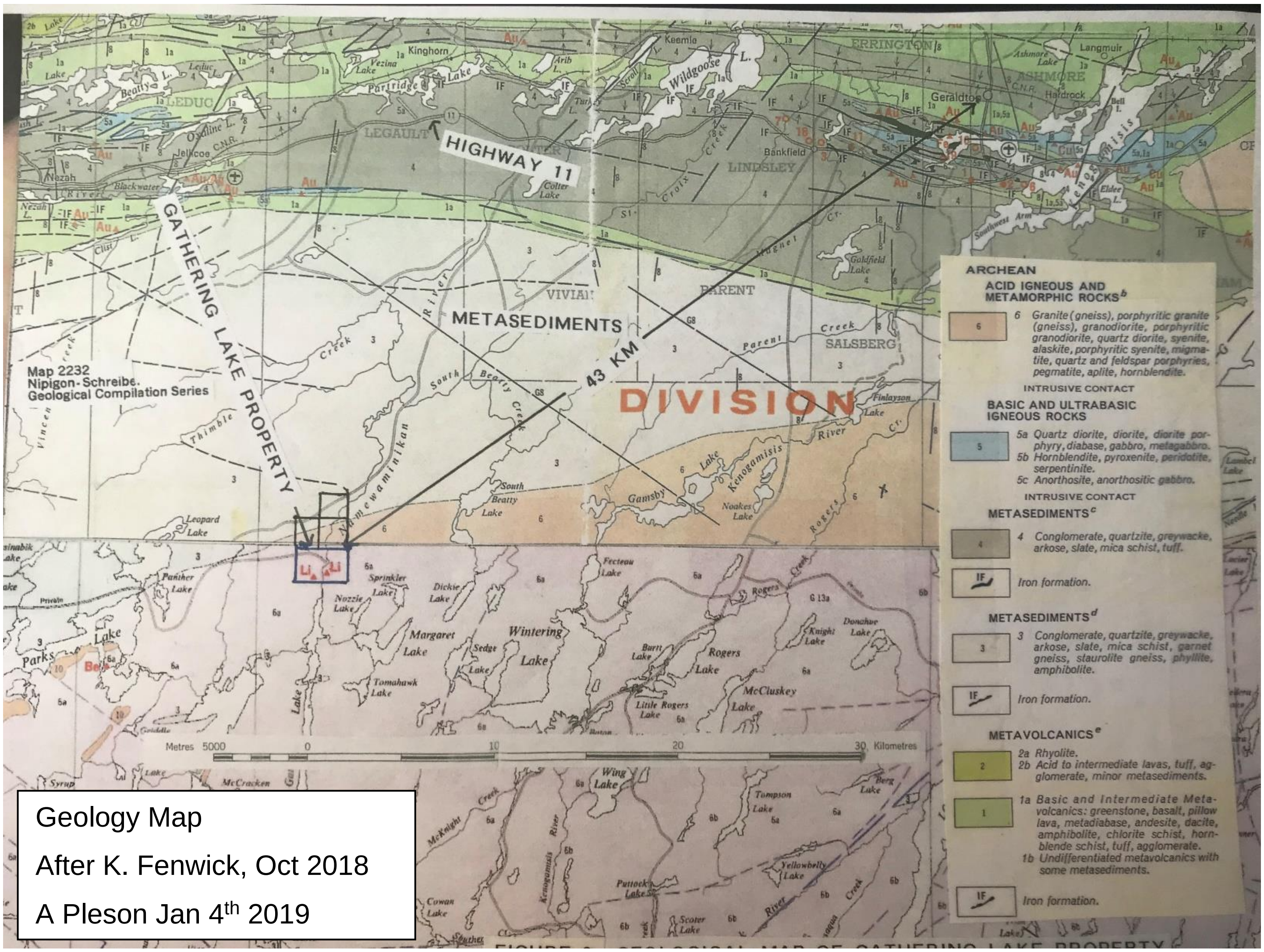
Four major faults cut through the Quetico Subrpovince: the easterly trending Quetico fault the Rainy Lake-Seine River fault, the northeasterly trending Gravel River fault (Williams, 1989) and the Kapuskasing Structural Zone (Selway 2011).

Metamorphism, migmatite formation and granite intrusion occurred between 2.67 and 2.65 Ga (Williams, 1991). The grade of metamorphism ranges from lower greenschist to amphibolite facies and tends to be lower in the marginal rocks of the subprovince and higher in the core regions.

Widespread economic mineralization within the Quetico Subprovince is generally lower than in the adjacent greenstone dominated terranes (Williams, 1991). Minor gold mineralization is associated with veining along the Quetico Fault (Poulsen, 1983). Molybdenite occurs in biotite leucogranites in the Dickinson Lake area. The only potentially important ore deposit type consists of the late-stage pegmatites that contain the rare elements lithium, beryllium, tantalum, niobium and tin (Williams, 1991). The rare-element pegmatites have widespread distribution in the Quetico Subprovince covering at least a 540-km strike length from west to east and a large percentage of pegmatites occur in the centre of the subprovince (Breaks, Selway and Tindle, 2006).

The pegmatites in the Quetico Subprovince are hosted by metasediments and by their parent granite (Pye, 1965; Breaks, Selway and Tindle, 2003a, 2003b).

Figure 3a: Regional geological map



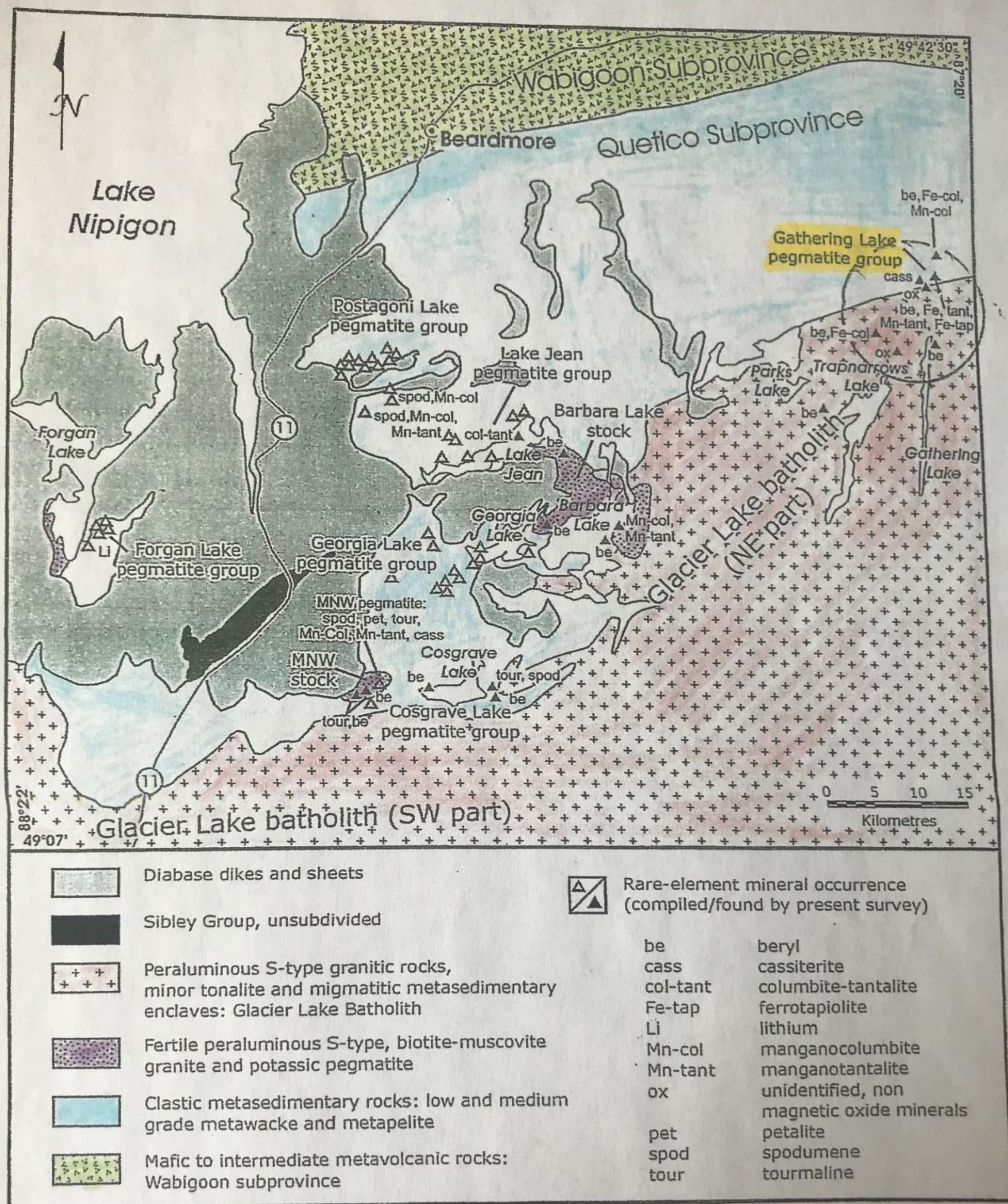


FIGURE 3: General geological setting of rare-element pegmatites of the Georgia Lake field and adjacent, S-type, peraluminous granitic rocks of the Glacier Lake batholith. (BREAKS, 2008)

Twelve new Ta-Nb oxide mineral occurrences were found in Break's survey area (see Figure 3 for the size of the survey area). The Gathering Lake pegmatite group contains nine (9) of these Ta-Nb-oxide-bearing pegmatites.

Figure 4b: Regional geological map after *Breaks 2008*

6.2 Local Geology

The geology is of Precambrian age and is discussed by Pye (1965).

Metasediments

The oldest rocks are the Archean metasediments. The metasediments strike east-northeast and dip steeply, in general, to the north. The dominant metasedimentary rock is biotite-quartz-feldspar schist or gneiss. It is a grey, rather dark colored rock, having a distinct banded appearance due to compositional variations reflecting an original sedimentary stratification, with individual layers less than an inch to several feet thick. There is a distinct foliation due to parallel alignment of biotite crystals. Microscopic examination of the biotite-quartz-feldspar schist shows that it is made up of: 15-40 vol.% biotite, 20-35 vol.% quartz, 25-45 vol.% plagioclase, 1-3 vol.% magnetite, trace amounts of zircon and rare hornblende. Secondary minerals include chlorite, sericite and epidote. The plagioclase shows myrmekite texture. The most abundant texture in the biotite-quartz-feldspar schist or gneiss is granoblastic, but porphyroblastic rocks are also present with porphyroblasts of garnet, staurolite and cordierite.

Metagabbro

The metagabbro has intrusive relationships and have been metamorphosed and intruded by granitic rocks. East of Cosgrave Lake and south of Barbara Lake, the metasediments were intruded by metagabbro. The metagabbro bodies range in size from a few hundred feet across to 9,500 feet (=2.9 km) across. The metagabbro is dark-colored (mesocratic), medium- to coarse-grained with a brownish weathered surface. For the most part, it is massive, but it is gneissic near its contacts with metasediments. The major minerals are: green hornblende and plagioclase (sodic andesine). The minor minerals include: microcline and biotite and trace amounts of magnetite and apatite. The alteration minerals are chlorite, epidote and sericite.

The porphyritic metagabbro differs from the metagabbro only in the presence of feldspar phenocrysts (usually microcline). The feldspar phenocrysts are pale-pink to red, stubby, rectangular, subhedral to euhedral and range in size from ¼ by 1/8 inch (=0.6 by 0.3 cm) to 2 by 1 inches (5 by 2.5 cm). The porphyritic metagabbro is best developed near the margins of the metagabbro bodies close to the granites.

Metagabbro dykes and sills cross cut the metasediments near Dump and Pawky lakes and near Blay, Georgia and Conner lakes. All the dykes and sills are small with thicknesses of 3 feet or less (=0.9 m). They are thought to be genetically related to the metagabbro, as they are similar in appearance and composition. They are cross cut by pegmatite and feldspar porphyry dykes.

Granite

The metasediments were also intruded by large masses of granitic rocks and by numerous sills and dykes of genetically-related porphyry, pegmatite and aplite. The granitic rocks are pale-grey or pale-pink in colour and their essential components are: 45-65 vol.% feldspar (microcline and plagioclase), 40 vol.% quartz, and one or both of muscovite and biotite and rarely little hornblende. The plagioclase has a composition of albite. Minor components of the granites include magnetite, zircon, and garnet, and secondary minerals: chlorite, sericite and epidote. For the most part the granites are equigranular, but porphyritic phases with microcline phenocrysts also occur. The contacts between the equigranular granitic rocks and the metasediments are generally abrupt.

Pegmatite

There is an abundance of pegmatites close to and within the large masses of granitic rocks. A regional zoning is apparent and a genetic association of pegmatites and granite is indicated. The pegmatites occur in two geometries: as irregular-shaped bodies and as thin dykes, sills and attenuated lenses. The irregular bodies of pegmatite are intimately associated with the granite bodies often within a few hundred feet of the contact zone. They typically are medium- to coarse-grained, up to very coarse-grained and are made up of quartz, microcline, perthite and little muscovite. These would be classified as potassic pegmatites. Accessory minerals include biotite, tourmaline and garnet.

The pegmatite dykes, sills and lenses can be subdivided into rare-element pegmatites and granitic pegmatites. The rare-element pegmatites are of economic significance and they contain microcline or perthite, albite, quartz, muscovite and spodumene and minor amounts of beryl, columbite-tantalite and cassiterite. The granitic pegmatites are like the irregular pegmatites described above except that they contain more abundant plagioclase. Some of the pegmatites are parallel to the foliation or bedding of the metasediments, whereas others occur in joints in either the metasediments or granite. Contacts are usually sharp and, except where dykes cut granitic rocks, often found to be marked by a thin border zone of aplite or granitoid composition. A few pegmatites are internally zoned with mica-rich or tourmaline-rich rock along or close to the walls and quartz cores.

Diabase

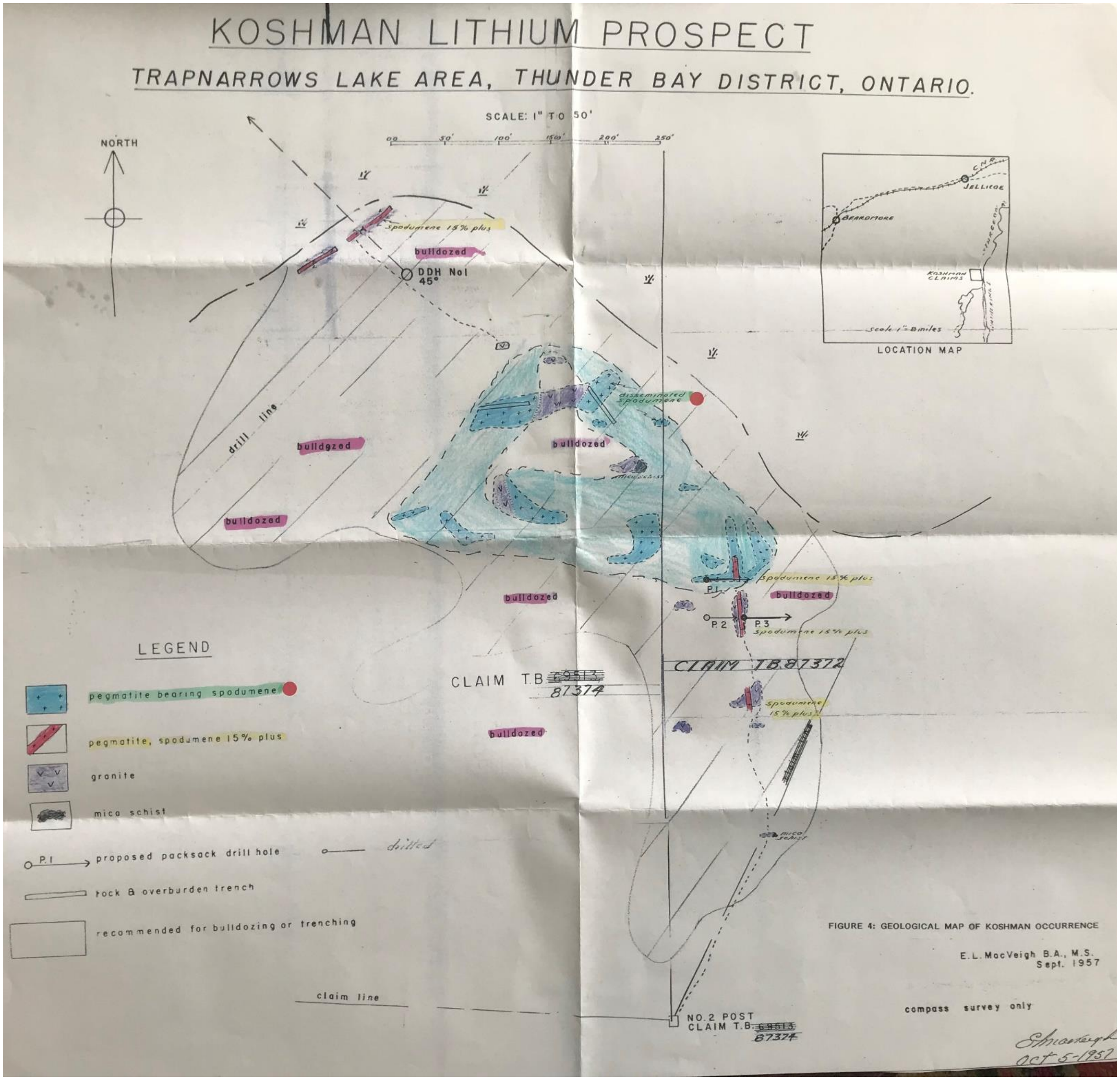
Intrusive into the Proterozoic sedimentary rocks and the older formations are bodies of diabase. The largest occur as flat sheets (Logan sills), up to about 650 ft. (=198.1 m) in thickness, and as dykes of vertical or near-vertical attitude. Most of the dykes are related closely to the sheets and are Keweenawan age. The gently dipping diabase sheets are dark colored and massive. The diabase sheets are well-jointed and most of the joints are vertical or steeply dipping. In outcrop, the diabase shows poorly-formed columnar structure.

There are two types of diabase dykes: one is equigranular and the other is porphyritic. The equigranular dykes are more abundant. Some of the dykes along or close to the contact zone of the large granite mass strike easterly; most dykes in other localities strike north or within 20º of north. With few exceptions, the dykes are vertical or dip steeply. The porphyritic diabase dykes are massive medium-grained, dark-colored rock characterized by many pale-greenish yellow phenocrysts of highly altered plagioclase. Porphyritic diabase dykes are found near the Jackpot deposit.

6.3 Property Geology

The property is dominated by granite related to the Glacier Lake Batholith. Based on the limited outcrop close to the historic workings, the general property geology is best described by the historic 1955 drilling and a small amount of observations by Pye 1956, Breaks 2008, and from the authors time on the project. Overall, the granite is mainly coarse-grained consisting of quartz, feldspar, and muscovite. The metasediments observed are typically meta-sandstone described as a muscovite schist with observed bedding and various stages of metamorphism are present, including migmatization imparting a gneissic texture. The pegmatite dykes observed at the Koshman occurrence are simple, non-zoned, albite-type spodumene pegmatites with trace amounts of oxides and apatite. The pegmatites at the Nelson occurrence display partial zoning, with a spodumene zone observed on the western portion of the large pegmatite outcrop/hill. The geology of the Koshman occurrence was mapping in detail by E. MacVeigh in 1957 and is presented in figure 4.

Figure 4: 1957 Mapping of Koshman Lithium Prospect (MacVeigh, 1957)



7.0 EXPLORATION WORK

The exploration tasks were to identify areas of where a prospective grid for VLF could be conducted using a Geonic's model EM16 VLF handheld survey unit and based GPS surveying for station accuracy to ~3-meters. The work consisted of surveying a planned grid using the instrument and GPS stations (No lines were cut on the property as a permit has not been applied for). A total of 11 lines (1500 meters each) were surveyed duration the 6-day duration of the project between January 3rd to January 8th 2022. Stations were selected to be 50m apart as well as each line was spaced 50m in order to cover a large area around the Koshman occurrence but still be able to identify any major changes or intrusions/conductors in the vicinity.

Figure 5: Koshman VLF Grid Location

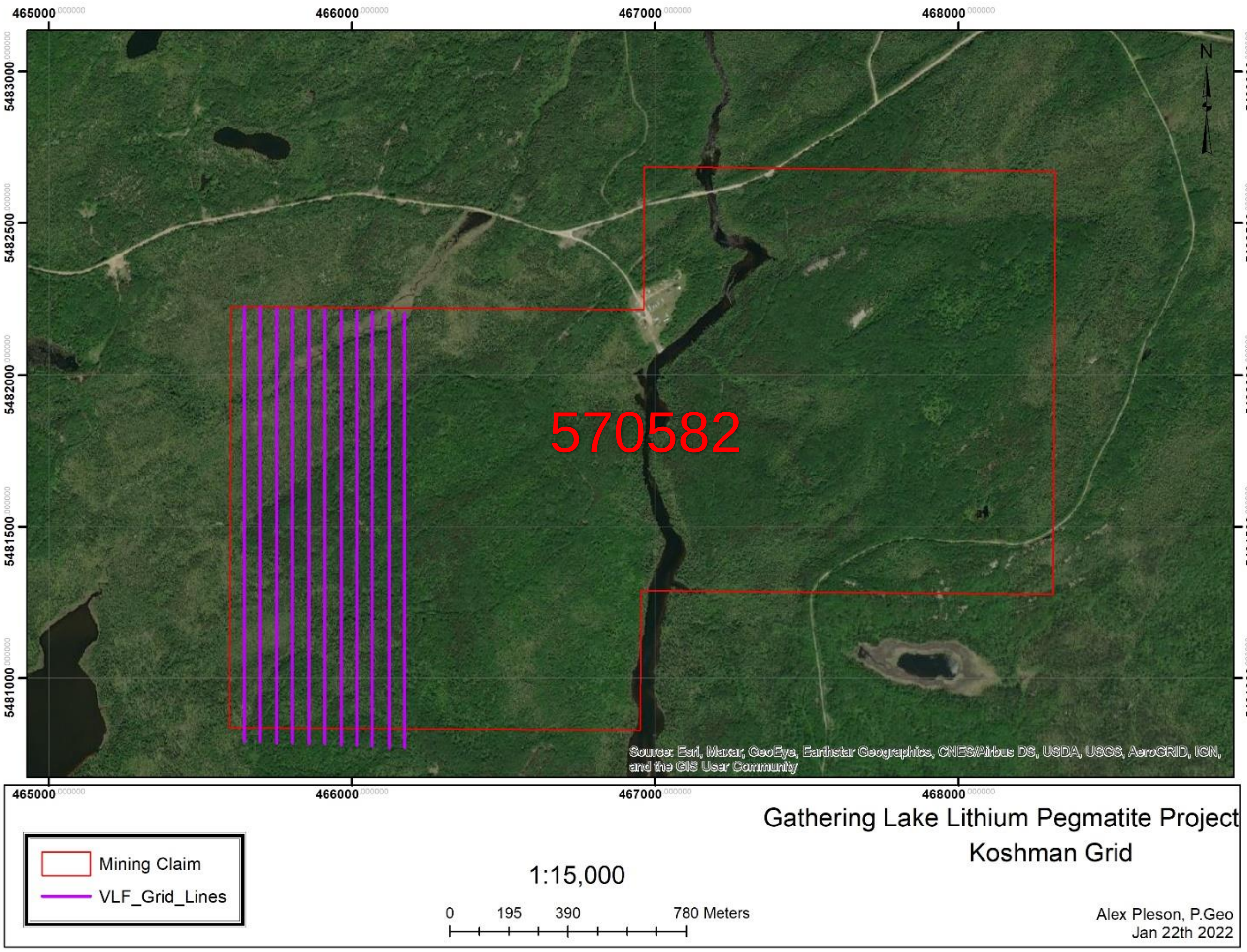
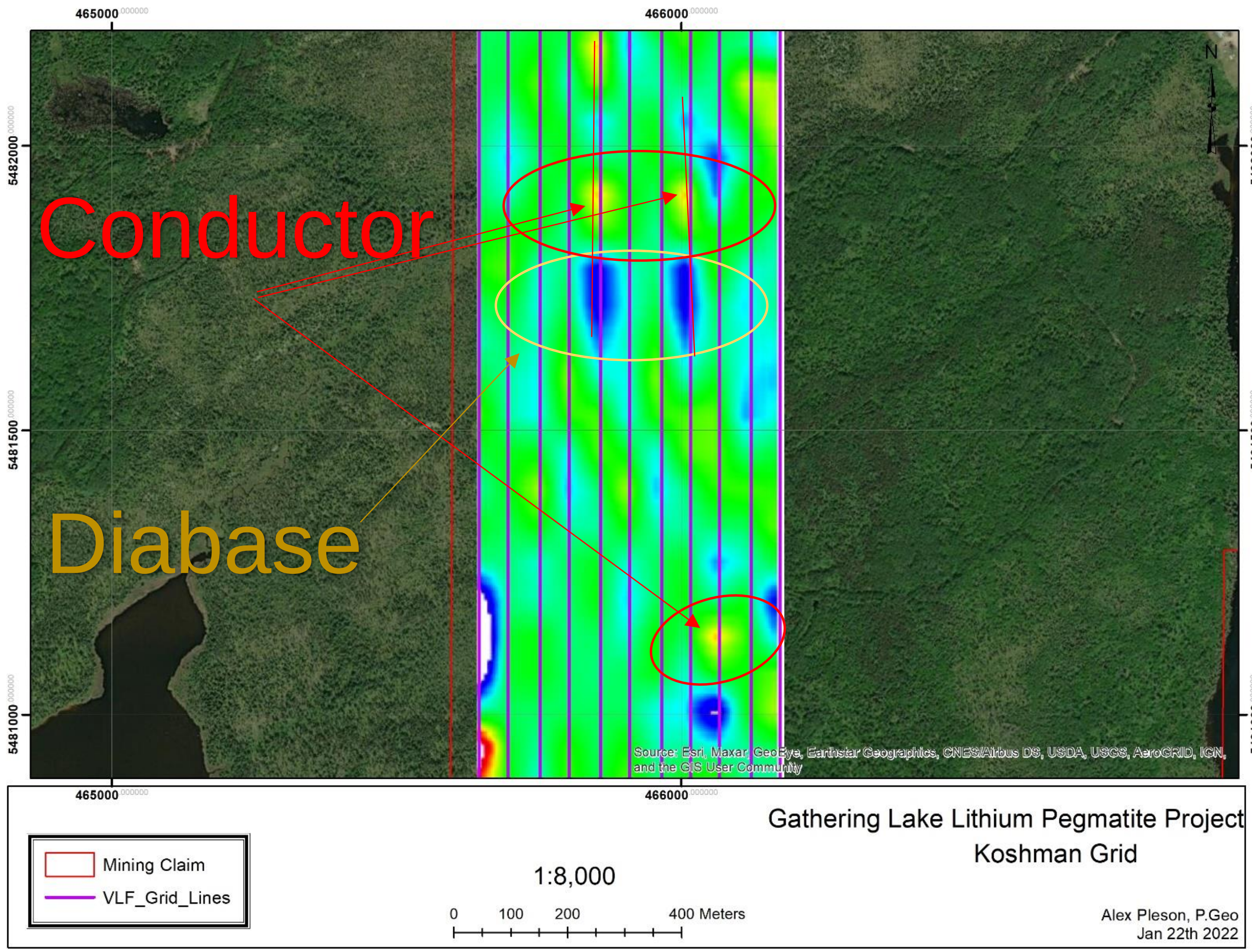


Figure 6 : VLF Conductors



8.0 EXPLORATION RESULTS

The initial geophysical investigation produced encouraging results. The author was able to confirm a slight conductor near historic lithium mineralization which occurs in a spodumene bearing pegmatite dyke and has identified similar conductors in the survey boundary which bear similar structures and size to that of the known Koshman Occurrence. The larger conductors in the northern portion of the survey are most likely large interference due to diabase dykes which were noted in the 2018 prospecting campaign but the smaller "red" conductors are the most interested as they do correspond to the Koshman Occurrence. Due to the low-quality processing of the data, it would also be recommended to have the data reprocessed.

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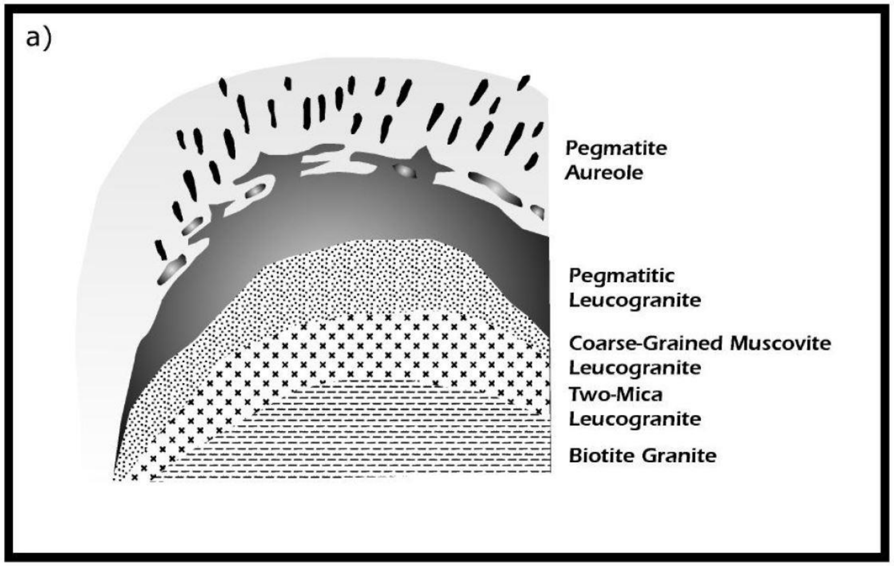
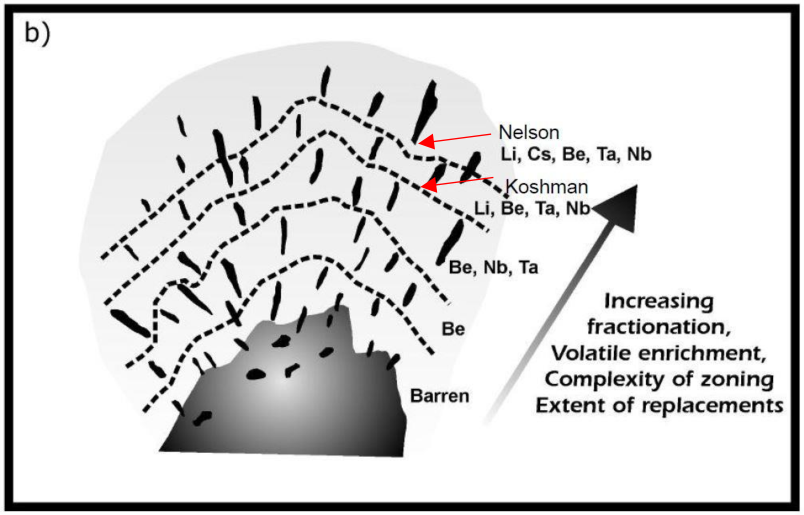
Figure 7: Pegmatite Diagram with proposed locations of occurrences on Gathering Lake Property (after *Cerny 1991*)





9.0 CONCLUSIONS

The Gathering Lake Lithium Property represents an area with high economic potential for lithium mineralization. The Koshman occurrence showcases potential for spodumene bearing dykes. Based on the current program and information from the historic reports, the Koshman area may host large dykes or sills that are lithium bearing to the north of the known Koshman occurrence. A prospecting campaign based on the highlighted conductors in the geophysics map will be completed in the summer of 2022. Follow up soil sampling along the grid and line cutting the grid would also be recommended.

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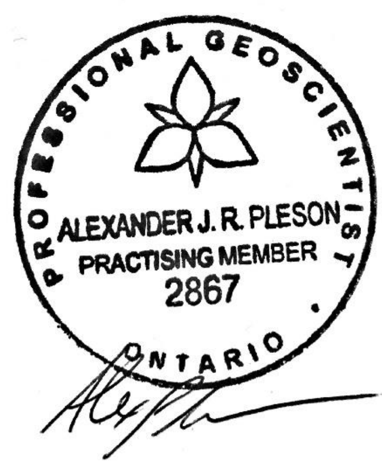
12.0 CERTIFICATE OF AUTHOR

I, Alexander Pleson, P.Geo., as an author of this report regarding the exploration project in the Thunder Bay Mining District, Northwestern Ontario, Canada; do hereby certify that:

1. I am a consulting geologist at Pleson Geoscience of Nipigon, ON, CA P0T 2J0
2. I have B.Sc. degree in Geology from Lakehead University.
3. I am registered as a Professional Geologist in Ontario (License #: 2867).
4. I have been practicing as a professional since 2017, and have 10 years of experience in mineral exploration.
5. The exploration work was carried out under my supervision and I was on site through the duration of the project.
6. The tasks outlined in this report were completed, in part, to gain a portion of ownership in the property

Dated: January 22nd 2022

Signed and Sealed:



APPENDIX A

LIST OF PERSONNEL WORKED ON EXPLORATION WORK

List of Personnel / Contractors Involved on the Project

1. **Alexander Pleson, P.Geo., - Geologist of Nipigon, ON (Pleson Geoscience)**
2. **Ted Cox – Geophysicist of Beardmore, Ontario (Pleson Geoscience)**

APPENDIX B

STATEMENT OF EXPENDITURES

Item	Cost
VLF Survey and Access	$10,959
Data Processing and Report	$850
TOTAL	$11,809

APPENDIX D

ACTIVITY LOG

	Task	Location	Cost
January 3rd 2022	Break Trail, Flag first line	Gorge Ck Rd.	$1459 (2 day wages + travel + Skidoos)
January 4th 2022	Survey L0, L50	Koshman Occurrence	$2000
January 5th 2022	Survey L100, part L150	Koshman Occurrence	$1500
January 6th 2022	Survey L200, 250, finish L150	Koshman Occurrence	$2000
January 7th 2022	Survey L300, 350 and portion of L400	Koshman Occurrence	$2000
January 8th 2022	Finish L450, L500 and walk back to finish L400	Koshman Occurrence	$2000
January 22nd Report and Map	Report and Data	Nipigon	$850

APPENDIX E

VLF DATA

Quadrature											
	Line ID										
	0	50E	100E	150E	200E	250E	300E	350E	400E	450E	500E
Station ID											
0	0	0	0	0	0	-4	2	-4	6	0	0
50	2	4	2	-6	0	-8	6	-6	8	0	-6
100	2	8	4	-8	6	-6	6	0	0	0	0
150	8	0	0	0	8	-8	0	0	6	-8	4
200	6	4	6	-8	0	0	0	0	2	-2	4
250	0	10	0	0	5	0	4	0	0	0	0
300	2	8	4	-8	6	-6	6	0	0	0	0
350	8	0	0	0	8	-8	0	0	6	-8	4
400	6	4	6	-8	0	0	0	0	2	-2	4
450	2	8	4	-8	6	-6	6	0	0	0	0
500	8	0	0	0	8	-8	0	0	6	-8	4
550	6	4	6	-8	0	0	0	0	2	-2	4
600	2	8	4	-8	6	-6	6	0	0	0	0
650	8	0	0	0	8	-8	0	0	6	-8	4
700	6	8	4	-8	6	-4	-2	-8	-4	4	4
750	2	20	4	-4	4	-8	6	-6	6	0	2
800	-8	6	-6	6	0	0	0	-8	0	0	4
850	0	8	-8	0	0	6	-8	8	-8	0	4
900	-8	0	0	0	0	8	-8	0	0	6	-4
950	-8	6	-6	6	-8	6	-4	-2	-8	-4	6
1000	0	8	-8	0	-4	4	-8	6	-6	6	0
1050	-8	6	-4	-2	6	0	0	0	-8	0	0
1100	0	8	-8	0	0	0	6	-8	8	-8	4
1150	-8	0	0	0	0	0	2	-2	0	0	4
1200	0	5	0	4	6	0	0	0	6	-6	4
1250	-8	6	-6	6	0	0	6	-8	8	-8	2
1300	0	8	-8	0	0	6	0	0	6	-8	8
1350	-8	0	0	0	0	2	0	0	8	-8	0
1400	-8	6	-6	6	0	0	6	-8	6	-4	-2
1450	0	8	-8	0	0	6	0	-4	4	-8	6
1500	-8	0	0	0	0	2	-2	6	0	0	0

In-Phase											
	Line ID										
	0	50E	100E	150E	200E	250E	300E	350E	400E	450E	500E
Station ID											
0	0	-5	2	0	0	-2	0	2	-5	0	2
50	5	-5	4	-8	0	0	0	5	-10	-8	8
100	10	0	5	-10	10	0	-5	10	-14	-10	16
150	20	-10	-5	0	0	2	0	8	5	0	2
200	20	-5	5	-2	0	0	2	6	5	-2	0
250	20	-10	-5	0	0	2	0	8	5	0	2
300	20	-5	5	-2	0	0	2	0	5	-2	0
350	5	-5	4	-8	0	0	0	5	-10	4	8
400	10	-2	0	0	2	6	0	8	5	0	2
450	50	0	0	2	0	8	2	6	5	-2	0
500	25	-2	0	0	2	6	-8	2	-20	2	12
550	30	0	0	2	0	8	0	0	-8	0	10
600	0	2	6	0	0	6	0	0	2	6	16
650	-5	0	0	0	2	0	0	2	0	8	-8
700	-8	0	0	0	5	0	0	5	-10	4	8
750	0	0	-2	0	2	-2	0	2	-5	2	2
800	-8	0	10	-5	0	10	-5	0	-5	10	12
850	-5	5	8	-10	5	8	-10	5	0	8	8
900	0	0	5	-5	10	5	-5	10	5	2	8
950	0	0	2	0	8	2	0	8	5	0	2
1000	-2	0	0	2	0	0	2	0	5	-2	0
1050	-5	0	0	2	0	0	2	0	2	-5	0
1100	2	6	-8	2	-20	-8	2	-20	8	-10	-2
1150	0	8	0	0	-8	0	0	-8	10	-5	-5
1200	2	6	0	0	2	0	0	2	6	-2	0
1250	0	0	0	2	0	0	2	0	-6	-4	2
1300	0	0	0	5	-10	0	5	-10	0	-8	6
1350	2	0	8	0	-2	0	2	-5	-5	10	18
1400	0	2	6	0	10	-5	0	-5	-5	0	10
1450	2	0	8	5	8	-10	5	0	0	0	4
1500	0	0	6	0	5	-5	10	5	-5	10	5