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Technical Report

High-Resolution Heliborne Magnetic and TDEM Survey

*Solitude Lake Project, Savant Lake Area
Patricia Mining Division, Ontario
2021*

*Copper Ridge Exploration Corp.
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Prospectair Geosurveys

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I. INTRODUCTION

Prospectair conducted a heliborne magnetic (MAG) and time-domain electromagnetic (TDEM) survey for the mineral exploration company Copper Ridge Exploration Corp. on its Solitude Lake Property, located in the Savant Lake area, Patricia Mining Division, Province of Ontario (Figure 1). The survey was flown on June 21st, 2021.

Figure 1: General survey location

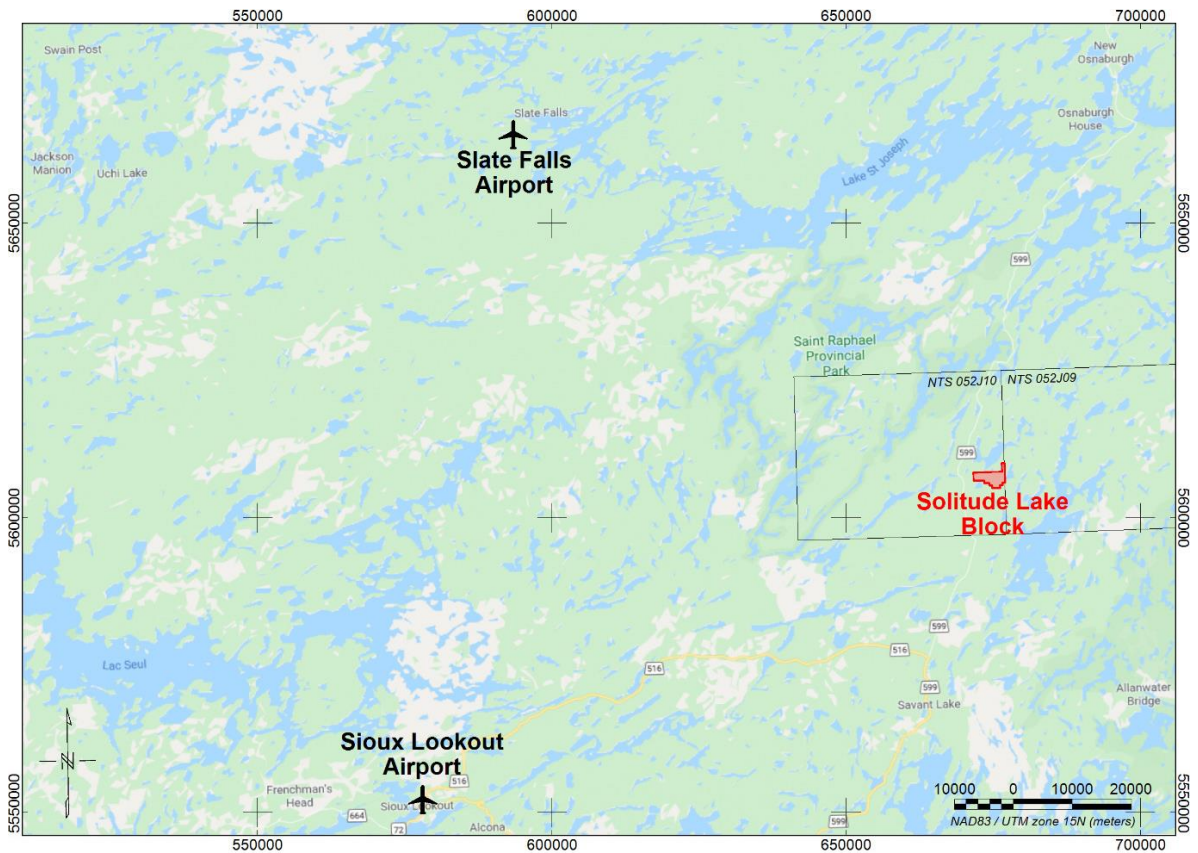


One survey block was flown for a total of 249 l-km (Table 1). A total of 3 production flights were performed using Prospectair's Eurocopter EC120B, registration C-GEDI. The helicopter and survey crew operated out of the Sioux Lookout Airport located about 110 km to the southwest of the block (Figure 2). The block is found approximately 40 km to the north of Savant Lake, the closest village.

Table 1: Survey block particulars

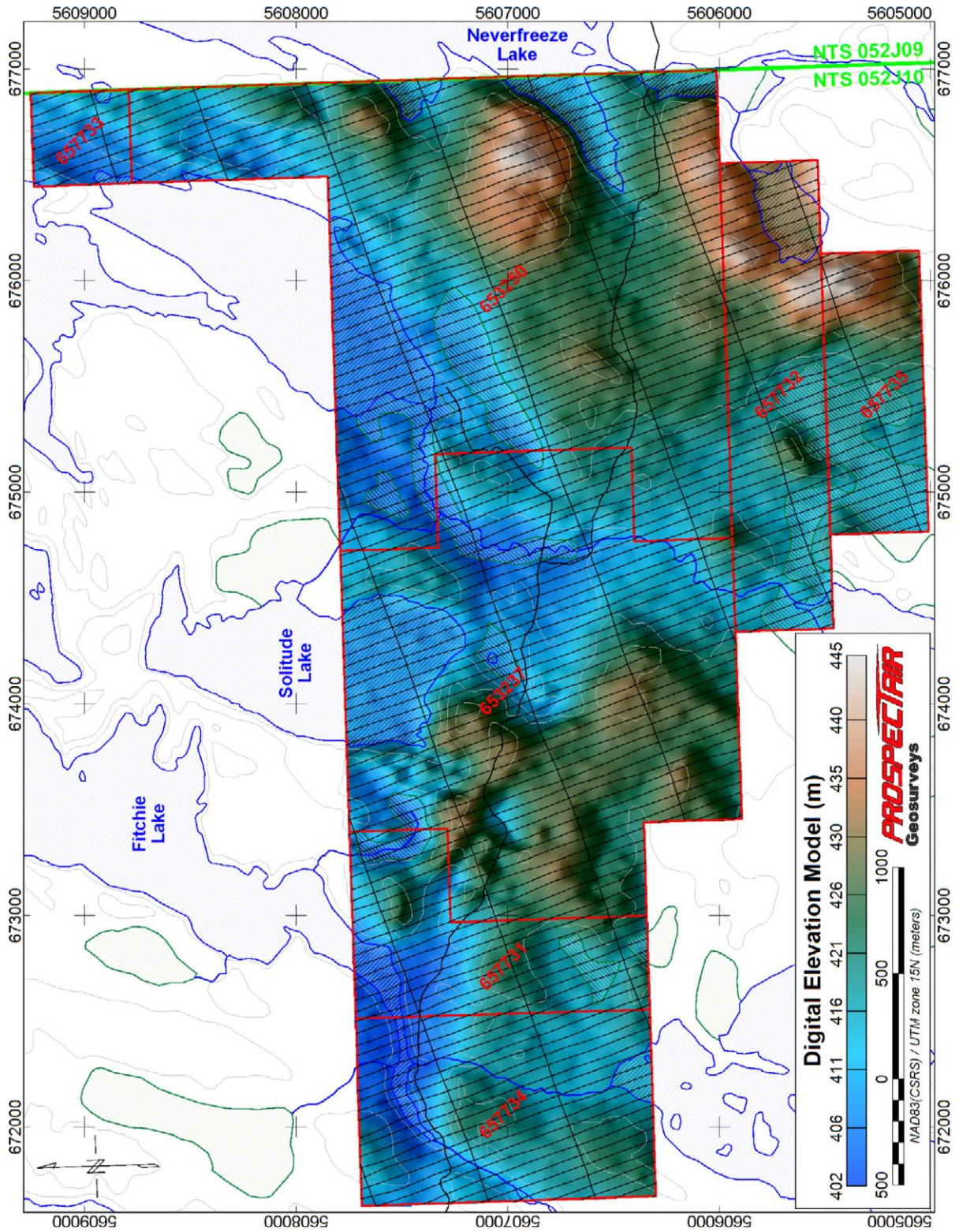
Block	NTS Mapsheets	Line-km flown	Flight numbers	Date Flown
Solitude Lake	052J10	249 l-km	Flt 1 to 3	June 21 st

Figure 2: Survey location and base of operation



The Solitude Lake block was flown with traverse lines at 50 m spacing and control lines spaced every 500 m. The survey lines were oriented N160. The control lines were oriented perpendicular to traverse lines. The average height above ground of the helicopter was 83 m, with the mag sensor and receiver coil at 58 m, and the transmitter loop at 33 m above the ground. The average survey flying speed (calculated equivalent ground speed) was 32.5 m/s. The survey area is covered by forest, lakes and a few wetlands, and the topography is generally gentle, with only a few low-level hills. The elevation is ranging from 402 to 445 m above mean sea level (MSL). The survey block is approximately located to the south of the Fitchie and Solitude Lakes, and to the west of the Neverfreeze Lake. Road 599, which links Savant Lake to the northern community of Pickle Lake, passes less than 3 km to the west of the block. Coordinates outlining the survey block are given in Appendix A, with respect to NAD-83 datum, UTM projection zone 15N. The location of the Solitude Lake Property claims (in red) and of the survey lines is shown on Figure 3. The Property claims numbers covered by the survey are listed in Appendix B.

Figure 3: Survey lines and Solitude Lake Property claims



II. HISTORY

The property is currently in the early stages of exploration and has been historically sparse and has not been the focus of widespread work. The earliest regional-scale mapping efforts were undertaken by the Ontario Department of Mines in 1927 (Moore, 1927). Prospecting in the Savant Lake area revealed several gold occurrences which were subsequently explored from the 1960s onwards.

A rash of activity began in the area north and west of Savant Lake in the 1960s. The iron ore prospect at North Kashaweogama Lake (about 15 km SW of the Property) was discovered in the early 1960s; the strike extension of this iron formation (to within 3 km of the Solitude Lake Property boundary) was drilled by Hanna Mining in 1967. Hanna Mining held claims that came within 600 m of the present Property boundary, to the west, and completed a detailed geologic mapping program on this part of their property in the same year (Hogg 1967). The Hanna mapping program revealed an east-west- striking complex of mafic volcanics interspersed with sulphide and oxide iron formations, “basic tuffs” (likely tectonized/schisted mafic volcanics), greywackes, and chert- and graphite-bearing quartzites. A grab assay of 0.05% Cu is noted on the Hanna maps about 1,500 m west of the Solitude Lake Property boundary.

The first well-recorded exploration within Solitude Lake Property limits dates to 1968 when Canadian Nickel Co (Canico) completed ten drillholes. These were presumably aimed at the east-west conductive and magnetic targets, also visible in later surveys. Many assessment files refer to airborne surveys completed by Canico and/or Inco in “the late 1960s” but the Authors were unable to locate any actual survey documents.

The most promising of the Canico drillholes was #37663 which encountered several sulphidic and graphitic schists containing up to 5% cpy over a core length of 9.6 m (Thomson 1968). Most drillholes encountered zones of semi massive sulphides hosted by iron formations, graphitic shales and volcanics, many of which contained at least trace chalcopyrite, and occasional sphalerite, according to the logs. No assay information is reported in any of the drill logs, nor are any plans or drill sections presented in any of the available assessment files. A later prospectus by Redstone Resources refers to drillhole 37663 (as part of a discussion of properties adjacent to theirs, which covered the western extreme of the current Solitude Lake Property) and states that an assay interval of 0.7% Cu over 70 ft (= 21.3 m) was returned (Paterson 1976). The Authors have been unable to locate the assay information upon which this figure is based.

In the years following the Canico work the central part of the current Solitude Lake Property fell into the ownership of Durnin and Read, local prospectors, while the western, unexplored portion was held by the aforementioned Redstone Resources.

Both of these claim blocks were acquired by Surveymin Ltd in 1976 who then completed ground resistivity and magnetic surveys along 400 ft-separated gridlines covering an area of about 250 Ha, corresponding to the area around the best Canico drillhole and the westward strike extensions. This work outlined a series of broadly east-west conductors, notably the closely spaced “A” and “B” conductors which can be traced over about 2.5 km. The conductors were interpreted to be offset by a late, cross-cutting fault which lines up with local topographic lineaments (Kuryliw 1976). Based on the Canico drilling, the “A” conductor represents a 1-2 m thick massive pyrrhotite horizon within andesites, while the “B” conductor is caused by a pyrite- pyrrhotite-chalcopyrite-graphite schist unit. The more southerly “E” conductor has the strongest coincident magnetic response and may

represent an oxide iron formation. Kuryliw (1976) mentions that outcrops are common in the western portion of the Property, but this had never been comprehensively mapped until the recent program.

In 1978, Surveymin completed two drillholes testing the “B” conductor. Both confirmed the presence of the py-po-cpy-bearing graphite schist; the brecciated texture of the horizon was also noted (Surveymin 1978). The interval in hole S-79-1 contained “>1% chalco” over 40 ft (= 12.2 m) and “minor - >1% chalco” over 97 ft (29.6 m) in hole S-79- 2. The logs indicate that samples were taken but assays are not reported. Lineations were high against the core (80-87°) indicating that these intervals, while not quantitative, are close to true width.

Also in that year, local prospector George Armstrong completed a winter drill program on Neverfreeze Lake. This program was beyond the Solitude Lake Property boundary but the southernmost hole, #6, was within 100 m of the boundary. Sulphidic graphite schists were seen in all six drillholes of which most, including #6, carried “some chalcopyrite”.

Maps of the drillhole locations and historic claim fabrics are difficult to reconstruct. No collars were identified in the field during the 2021 program or the Minroc site visit. The Authors caution that the drillhole locations given in this Report come with potential errors in the order of 50 mm.

Table 2 Table of Drillholes on the Solitude Lake Property

DDH	Company	Year	Claim	UTME	UTM N	Dip	Azimuth	Length (m)	Reference
34500	Canadian Nickel	1968	653237	673944	5607002	-50	360	53.4	Thomson 1968
37630	Canadian Nickel	1968	657732	675242	5605809	-50	360	38.1	Scheibler 1968
37631	Canadian Nickel	1968	657732	675626	5605703	-50	360	43.6	Scheibler 1968
37632	Canadian Nickel	1968	657735	675606	5605426	-50	360	52.7	Scheibler 1968
37661	Canadian Nickel	1968	653237	674226	5607010	-55	360	53.7	Thomson 1968
37663	Canadian Nickel	1968	653237	674287	5606811	-55	360	57	Thomson 1968
37664	Canadian Nickel	1968	653237	673868	5606425	-50	360	56.4	Sheibler & Fimmerman 1968
37665	Canadian Nickel	1968	653237	673891	5606251	-50	360	53.7	Sheibler & Fimmerman 1968
37666	Canadian Nickel	1968	653237	673850	5606114	-50	360	53.3	Sheibler & Fimmerman 1968
37667	Canadian Nickel	1968	653237	673979	5606805	-50	360	61.6	Thomson 1968
S-79-1	Surveymin	1978	653237	674335	5606884	-50	360	71.6	Surveymin 1978
S-79-2	Surveymin	1978	653237	674213	5606863	-50	360	80.8	Surveymin 1978

III. GEOLOGIC SETTING

Regional Geology

The Solitude Lake Property lies within the Wabigoon Subprovince, an Archean terrane consisting of greenschist-grade volcanic-sedimentary sequences (“greenstone belts”) and voluminous granitoid batholiths. In age the Wabigoon units range from 2.77 to 2.72 Ga (Sanborn-Barrie & Skulski 2006). To its south, the Wabigoon is in contact with the Quetico Subprovince, which consists of slightly younger (-2.69 Ga; Davis et al 1990) clastic sediments derived in part from Wabigoon units.

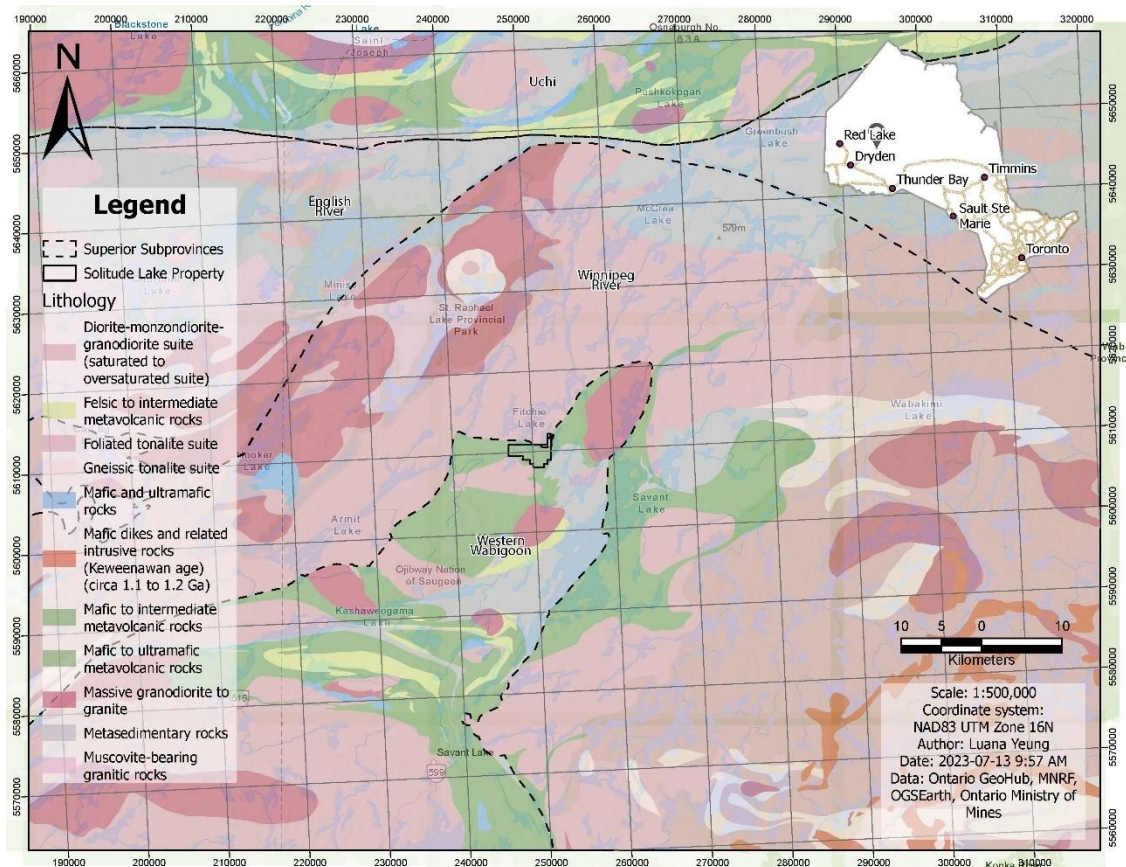
To the north of the Wabigoon Subprovince lie the Winnipeg River and English River Subprovinces. Respectively these consist of tonalitic gneisses of 3.2-2.8 Ga, and clastic

sediments of 2.72 Ga derived from the Uchi volcanic sequences further north (Corfu et al 1995). All of the above are components of the Superior Province and were accreted in a series of late Archean orogens. To the east, the Archean terranes are covered by Phanerozoic sequences of the Hudson Bay / James Bay platform. Around Lake Nipigon, the Wabigoon Subprovince is intruded by Proterozoic diabase sills of substantial size and affiliated with the Midcontinent Rift system.

The Solitude Lake Property is situated in the northernmost portion of the Sturgeon Lake Greenstone Belt (SLGB). The SLGB consists of mid- to late-Archean volcanic sequences and minor sedimentary sequences, originating in oceanic terranes and continental margins and is shaped like an elongated “S” surrounded on all sides by late Archean granitoids, the largest of which is the Lewis Lake Batholith which fills the southern lobe of the “S”. The bulk of the SLGB, around Sturgeon Lake itself, consists of the South Sturgeon, Handy Lake, Fourbay Lake, Quest Lake and Central Sturgeon mafic to felsic volcanic assemblages of oceanic island arc origin.

The northern portion of the SLGB, around Savant Lake, consists of a basin roughly bounded by the Kashaweogama and Savant Lake Faults. Jutten Group tholeiitic mafic- intermediate volcanics comprise the majority of the northern portion of the Belt. The core of the basin is filled with coarse clastic and chemical sediments of the West Shore Group.

Figure 4: Regional Geology of Solitude Lake



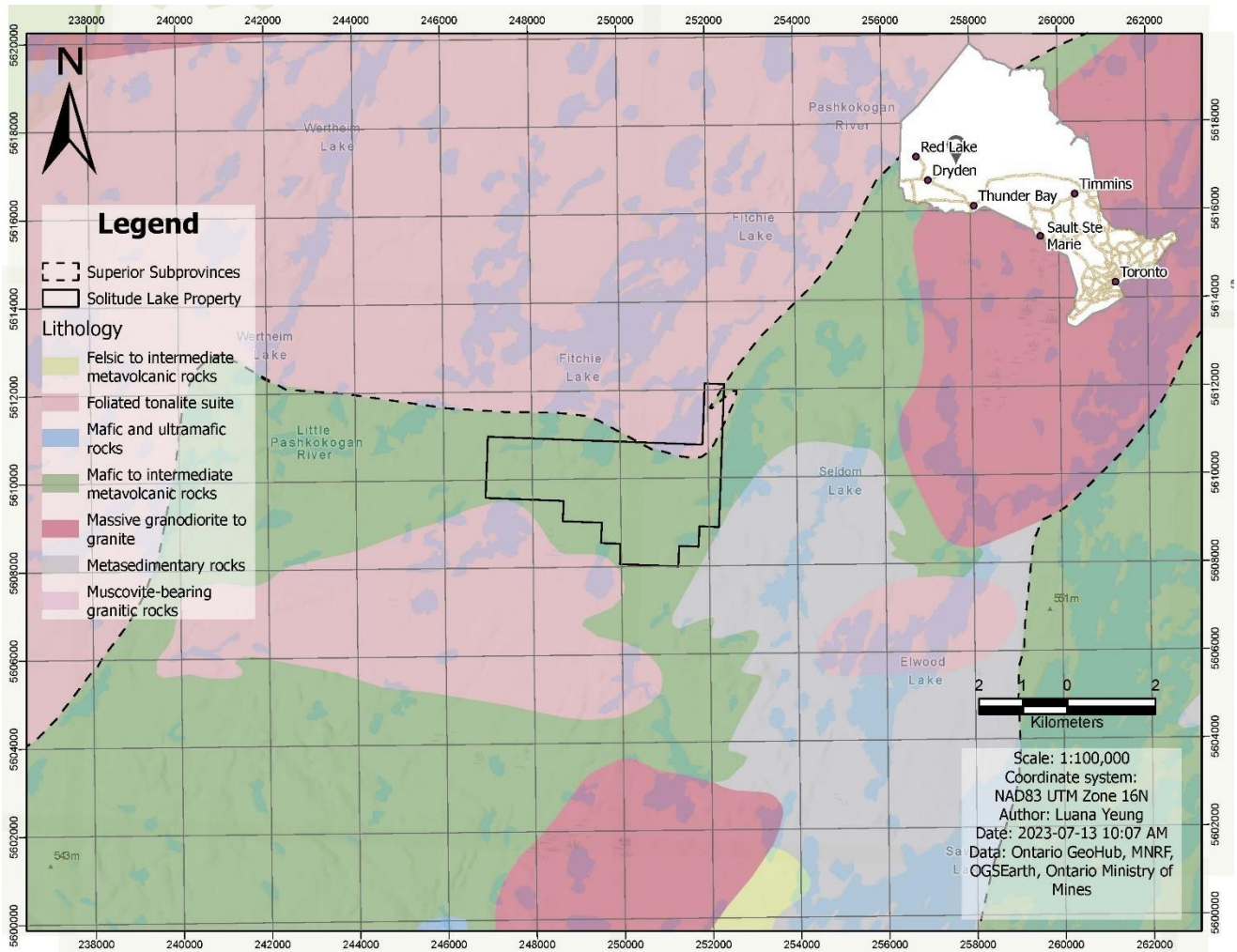
Local and Property Geology

The Solitude Lake Property lies at the northern extreme of the SLGB, on a tongue of volcanic and sedimentary units stretching east-west (Breaks 1976). This tongue is bounded to the north by a substantial gneissose granodiorite mass (the Lewis Creek Batholith), while the southern edge abuts the granodiorite Tersha Lake Stock. Both bodies of granodiorite are known from drilling on the Property. Based on outcrops a short distance west of the Property, the east-west sequence consists mostly of schistose basalts, with “tuffs” (possibly more strongly deformed mafics), iron formations and gabbros in the centre of the belt (Hogg 1967). Dips are subvertical. These units are attributed to the Jutten Group (Sanborn-Barrie 2000). The strike of this volcanicsedimentary sequence abruptly turns north-northeast close to the eastern edge of the Property. The area around Neverfreeze Lake is underlain by a synform striking north-northeast, with the Jutten Group volcanics on the outer limb, surrounding a band of conglomerates of the Narrows Formation, with West Shore Formation greywackes and iron formations in the core of the synform (Sanborn-Barrie 2000). These units likely form limited subcrop in the eastern extreme of the Property.

Drillhole logs on the Property frequently mention “strongly banded” appearances in the mafic-intermediate “tuffs”. Graphite schists, quartz-carbonate vein systems and breccia zones are mentioned frequently. Dykes and lenses of quartz feldspar porphyry are frequently seen; one example appears to fill the gap between the “A” and “B” conductors on the east side of the late cross-cutting fault. Drillholes alternatively mention “silicified zones” and horizons of quartzite; given the brevity of some of the drill log information it is not clear if these refer to the same or different lithologies.

Quaternary deposits of glaciofluvial sand and boulder clay cover much of the Property. Overburden in drillholes can reach depths of 50 feet (about 10 vertical metres).

Figure 5: Property Geology of Solitude Lake



Mineralization

Mineralization, as noted in historic drillholes, consists of disseminated, stringer and semimassive pyrite, pyrrhotite and chalcopyrite, with minor sphalerite, set within graphitic schists, silicified and/or chloritic mafic volcanics, and “quartzites”. The best mineralized of these zones are described as being brecciated and welded with quartz- carbonate veining (Surveymin 1978).

Descriptions in the Canadian Nickel drill logs are sparse. The two Surveymin DDH are better logged and contain information such as dip angles (generally 70-80° to core) which is absent in the Canadian Nickel logs. The log for DDH S-79-1 mentions 40 ft of “>1% chalco” alongside other sulphide mineralization; the true width of this interval would therefore be in the region of 11 m. Paterson (1976) mentions that drillhole 37663 was assayed, returning an interval of “0.7% copper over a core length of 70 feet” (21 m). The Authors have not been able to corroborate this figure in any other historic reports. In the Ontario Mineral Deposit Inventory this copper occurrence is referred to as the “Armstrong Option (1978)”.

“Trace pyrite” is mentioned, but not sampled, within the feldspar porphyry in S-79-2 (Surveymin 1978). This, and any other similar units which may be present on the Property, should be considered a potential target for gold mineralization.

The Property is at too early a stage of exploration for any meaningful comments to be made regarding the grade, tenor, strike length or vertical extent of the mineralization. The Authors also caution that the width figures described above are based on a very limited amount of drilling.

IV. DEPOSIT TYPES

The Solitude Lake Property is considered fertile for Volcanogenic Massive Sulphide (VMS) and orogenic gold, or greenstone-hosted gold deposits.

Volcanogenic Massive Sulphide (VMS)

VMS deposits typically consist of semi massive to massive lenses of sulphide, constrained by stratigraphy and spatially associated with vein stockworks and distinctive alteration patterns, including zones of carbonate, silica, chlorite, sericite and potassic alteration. VMS deposits are widely understood to be formed by hydrothermal activity in marine environments with extensional tectonic settings and are frequently found in Archean “greenstone” terranes hosted within mafic-felsic volcanic and/or volcanic-sedimentary cycles. Major sulphides present include pyrite, pyrrhotite, sphalerite and galena in the lenses, and chalcopyrite is typically present within the stockwork “pipe” or “feeder zone” in the footwall. The hangingwall is often accompanied by horizons of chert or iron formation. These types of deposits are significant economic sources of zinc, lead, silver and copper. Following the original deposition, later tectonic and hydrothermal activity may remobilize part of the mineralization.

Known zinc and copper occurrences, consisting of “VMS style” mineralization are well known and well studied in the southern portion of the Sturgeon Lake Greenstone Belt (the Mattabi, Lyon Lake, F Group and Sturgeon Lake mines, all about 85 km south of the Property), as well as the Sabin deposit, 35 km south of the Property.

The Sturgeon Lake VMS deposits, combined from 1970 to 1991, produced a total of 18,663,000 tonnes with average grades of 8.06% Zn, 1.09% Cu, 119.6 g/t Ag and 0.5 g/t Au (Franklin 1995).

Significant examples of VMS deposits elsewhere include the Kidd Creek deposit near Timmins, Ontario, as well as the Noranda and Matagami camps in Quebec. “Gold-rich VMS” deposits form a distinct subclass, an example being LaRonde in Cadillac, Quebec.

Orogenic Gold

Orogenic gold deposits are common in Archean greenstone terranes of the Canadian Shield. These deposits generally consist of a system of auriferous quartz-carbonate veins, which

have a strong spatial association with crustal-scale, compressional or transpressional shear zones with mixed brittle-ductile expression (or second or third order deformation zones). Further, there is commonly an association with particular lithologies, which are theorized to create favourable rheological or chemical environments for vein emplacement and/or gold precipitation. In many camps there is an affinity with porphyritic intermediate-felsic intrusives, iron formations and “Timiskaming-type” conglomerates; in the Sturgeon Lake area a common association is with porphyry sills and horizons of iron formation.

The shear zone is generally theorized to act as a pathway for hydrothermal fluids. These fluids are then emplaced as veins in dilated portions of ductile-deformed units, in brecciated portions of more brittle units, or in pore spaces of more porous units. Gold, which is often in solution with sulphur or arsenic in these fluids, will then be precipitated wherever the sulphur or arsenic can react with minerals in the country rock. Orogenic gold deposits can have highly complex geometries due to the intricate interplay of faults and favourable host units, continued tectonic activity on the shear zone after the emplacement of the mineralized veins, and disruption by later tectonic events.

The SLGB is home to one past producing gold mine, the St Anthony mine on Sturgeon Lake’s North Arm. St Anthony produced 63,310 oz Au in several periods from 1919 to 1942 (Evans 2009). Here, and at the King Bay gold occurrences in central Sturgeon Lake, gold is found in coarse, native form in dense stockworks of blue quartz associated with porphyry sills within chlorite and sericite altered mafic volcanics.

Major examples of iron formation associated orogenic gold deposits from elsewhere in northwest Ontario include the Musselwhite mine north of Pickle Lake, and the historic Central Patricia and Dona Lake mines in Pickle Lake.

Magmatic Massive Sulphide (MMS)

While less likely, the potential for the Solitude Lake property to host magmatic massive sulphide deposits should not be discounted, particularly given the relative lack of intensive exploration in the property area. MMS deposits are primary sulphide deposits which form pseudo stratigraphic horizons of massive, net-textured, stringer and/or disseminated sulphide within voluminous layered intrusive bodies, deposited as a result of their fractional crystallization. These deposits are major sources of copper, nickel, titanium, vanadium, chromium and platinum-group elements.

Major examples of MMS deposits from Archean terranes in Ontario include the Texmont, Langmuir, Alexo and Montcalm deposits, all in the Timmins-Cochrane area.

The metamorphic reworking of Archean MMS deposits can result in the remobilization of part of the mineralization, and re-emplacment of copper sulphides in veins and shears. This is theorized to have occurred at the Thierry deposit, north of the Property near Pickle Lake (Patterson & Watkinson 1984). Given the paucity of detailed geologic information, it is possible that the chalcopyrite mineralization at Solitude Lake has been remobilized from a magmatic source.

Due to the lack of data and poor quality of historic geophysical work on the property, a modern airborne survey would give clarity to geophysical anomalies which would subsequently allow for high-confidence target generation.

V. SURVEY EQUIPMENT

Prospectair provided the following instrumentation for this survey.

Airborne Magnetometers

Geometrics G-822A

Both the ground and heliborne systems used a non-oriented (strap-down) optically-pumped Cesium split-beam sensor. These magnetometers have a sensitivity of 0.005 nT and a range of 15,000 to 100,000 nT with a sensor noise of less than 0.02 nT. The heliborne sensor was mounted in a bird made of non-magnetic material located 25 m below the helicopter when flying. Total magnetic field measurements were recorded at 10 Hz in the aircraft. The ground system was recording magnetic data at 1 sample every second.

Time-Domain Electromagnetic Transmitter and Receiver

ProspecTEM

Prospectair Geosurveys significantly modified and improved the *Emosquito II* that was built by THEM Geophysics of Gatineau (Québec) to develop ProspecTEM. It is a powerful light-weight system adapted for small size helicopters and easy manoeuvrability enabling the system to be flown as close to the ground as safely possible and ensuring maximum data resolution. Advanced signal processing technique and a full processing package was developed in house to optimize the ProspecTEM data. The technical specifications are listed below in Table 2.

ProspecTEM system employs a transient or time-domain electromagnetic transmitter that drives an alternating current through an insulated electrical coil system. The towing bridle is constructed from a Kevlar rope and multi-paired shielded cables. It is attached to the helicopter by a weak link assembly. An onboard harness with outboard connectors mounted on a plate allows for quick disconnection or connection of the exterior elements. The system uses a 4 KW generator and a large condenser to transmit alternating 2.75-ms half sine pulses with intervening off-times of 13.916 ms electric pulse, 60 pulses per second.

The current in the coil produces an electromagnetic field. Termination of the current flow is not instantaneous but occurs over a very brief period of time (a few microseconds) known as the ramp time, during which the magnetic field is time-variant. The time-variant nature of the primary electromagnetic field, which propagates downward and outward into the subsurface, induces eddy currents which characteristics are governed by rocks conductivity distribution. These eddy currents generate a secondary electromagnetic field, in accordance with Faraday's Law. This secondary field immediately begins to decay in the process. Measurements of the secondary field are made only during the time-off period by a vertical component receiver located almost halfway between the helicopter and the

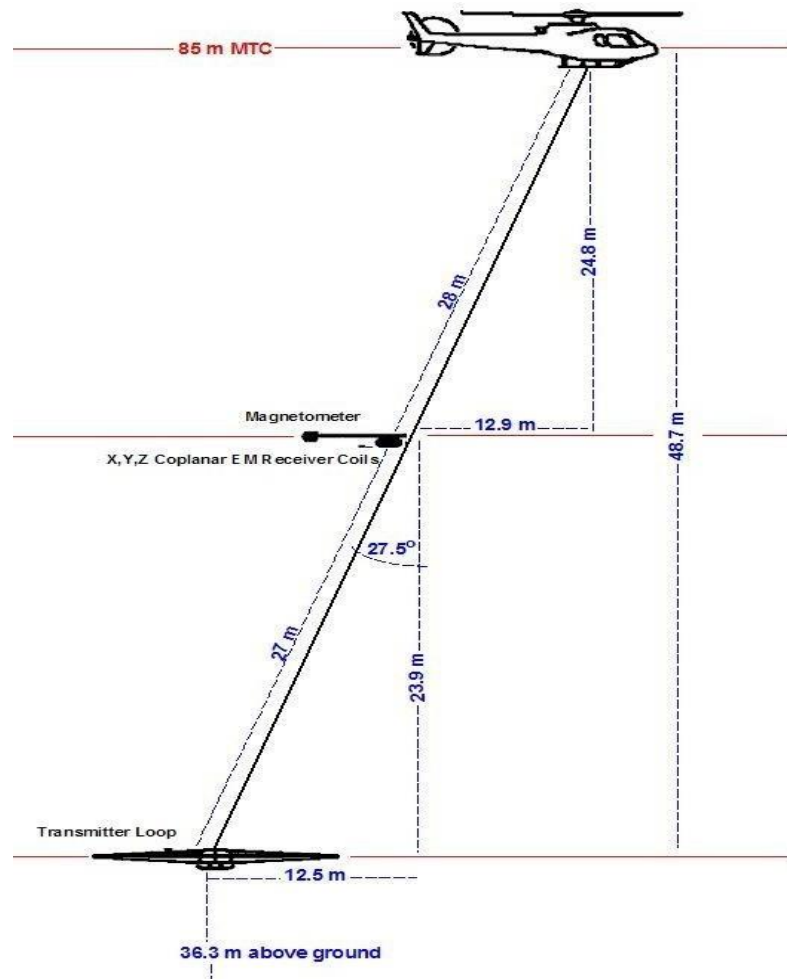
transmitter loop. It is placed with the magnetometer taped to a horizontal boom which supports the receiving coils tear-drop shape vessel at its end. The boom has an elastic suspension. A proprietary suspension system protects the orthogonal coils assembly and limits the total field excursions. The tear-drop vessel acts as a vane and maintains the mast in the line of flight.

Depth of investigation depends on the time interval after shutoff of the current, since at later times the receiver is sensing eddy currents at progressively greater depths. The intensity of the eddy currents at specific times and depths is determined by the bulk conductivity of subsurface rock units and their contained fluids.

Table 3: Technical specifications of the ProspecTEM Time-Domain system

Item	Specification
Transmitter:	
Loop Diameter:	5.6 meters
Current Waveform:	Half-Sine
Turns:	2
Pulse Length	2.75 ms
Frequency	30 Hz
Loop Area	25 m ²
Peak Current	3000A
Tow Cable Length	65 meters
Self-Powered	13HP Honda coupled with 28 Volts Alternator
Receiver:	
Coils axis	Z
Configuration	Coaxial (Z)
Two channels	Current and Z
Max Sampling rate	1000 points per half cycle at 90 Hz
Survey sampling rate	1000 per half cycle at 30Hz
Sampling	Full waveform
Gates	Programmable
On time signal	Recorded
Mechanical:	
Maximum survey speed:	120 km per hour
Transmitter height	30 meters AGL
Receiver height	55 meters
Weight (Total)	200 kg

Figure 6: ProspecTEM system configuration



Real-Time Differential GPS

Omnistar DGPS

Prospectair uses an OmniStar differential GPS navigation system to provide real-time guidance for the pilot and to position data to an absolute accuracy of better than 5 m. The *Omnistar* receiver provides real-time differential GPS for the Agis on-board navigation system. The differential correction data set was relayed to the helicopter via the appropriate Omnistar network satellite for the survey location. The receiver optimizes the corrections for the current location.

Airborne Navigation and Data Acquisition System

Pico-Envirotec AGIS-XP system

The Airborne Geophysical Information System (AGIS-XP) is advanced, software driven instrument specifically designed for mobile aerial or ground geophysical survey work. The AGIS instrumentation package includes a GPS based navigation system, real-time flight path information that is displayed over a map image of the area, and reliable data acquisition software. Thanks to simple interfacing, the radar and barometric altimeters, the TDEM system and the Geometrics magnetometer are easily integrated into the system and digitally recorded. Automatic synchronization to the GPS position and time provides very close correlation between data and geographical position. The AGIS is equipped with a software suite allowing easy maintenance, upgrades, data QC, and project and survey area layout planning.

Magnetic Base Station

GEM GSM-19

A GEM GSM-19 Overhauser magnetometer, a computer workstation and a complement of spare parts and test equipment serve as the base station. Prospectair establish the base station in a secure location with low magnetic noise. The GSM-19 magnetometer has resolution of 0.01 nT, and 0.2 nT accuracy over its operating range of 20,000- to 100,000 nT. The ground system was recording magnetic data at 1 Hz.

Altimeters

Free Flight Radar Altimeter

The Free Flight radar altimeter measures height above ground to a resolution of 0.5 m and an accuracy of 5% over a range up to 2,500 ft. The radar altimeter data is recorded and sampled at 10 Hz.

Prospectair Digital Barometric Pressure Sensor

The barometric pressure sensor measures static pressure to an accuracy of ± 4 m and resolution of 2 m over a range up to 30,000 ft above sea level. The barometric altimeter data are sampled at 10 Hz.

Survey helicopter

Eurocopter EC120B (registration C-GEDI)

The survey was flown using Prospectair's EC120B helicopter that handles efficiently the equipment load and the required survey range. Table 3 presents the EC120B technical specifications and capacity, and the aircraft is shown in Figure 5.

Table 4: Technical specifications of the EC120B Eurocopter helicopter

Item	Specification
Powerplant	One 376kW (504hp) Turbomeca Arrius 2F
Rate of climb	1,150 ft/min
Cruise speed	223 km/h – 120 kts
Service ceiling	17,000 ft
Range with no reserve	710 km
Empty weight	991 kg
Maximum takeoff weight	1,715 kg

Figure 7: C-GEDI Eurocopter EC120B



VI. SURVEY SPECIFICATIONS

Data Recording

The following parameters were recorded during the course of the survey:

In the helicopter:

- GPS positional data: time, latitude, longitude, altitude, heading and accuracy (PDOP) recorded at intervals of 0.1 s.
- Total magnetic field: recorded at intervals of 0.1 s.
- Terrain clearance as measured by the radar altimeter at intervals of 0.1 s.
- Z and Current TDEM channels at 90000Hz.

At the base and remote magnetic ground stations:

- Total magnetic field: recorded at intervals of 1 s.
- GPS time recorded every 1 s to synchronize with airborne data.

Technical Specifications

The data quality control was performed on a daily basis. The following technical specifications were adhered to:

- *Height* – 85m target terrain clearance for the MAG-TDEM survey except in areas where Transport Canada regulations prevent flying at this height, or as deemed necessary by the pilot to ensure safety. Traverse lines and control lines must be flown at the same altitude at points of intersection; the altitude tolerances are limited to no more than 30 m difference between traverse lines and control lines.
- *Airborne Magnetometer Data* - The noise envelope not to be exceeded 0.5 nT more than 500 m line-length without a reflight.
- *Diurnal Specifications* – A maximum tolerance of 5.0 nT (peak to peak) deviation from a long chord of one minute at the base station.
- *EM data* – No spikes on Z channel and constant current confirmed.
- *Flying Speed* – The average ground speed for the survey aircraft shall be 120 kph. The acceptable high limit is 160 kph over flat topography.
- *Radar Altimeter* – minimal accuracy of 5%, minimum range of 0-2500 m.
- *Barometer* – Absolute air pressure to 0.1 kPa.
- *Flight Path Following* – Maximum deviation of 30% of line spacing allowed over a maximum line distance of 300 m.

VII. SYSTEM TESTS

Magnetometer System Calibration

The survey configuration using a bird towed 25 m below any magnetic piece of the helicopter allows the simplification of the magnetic calibration requirement. Consequently, heading error and aircraft movement noise was considered negligible and no correction was applied to the data.

Instrumentation Lag

The data lag is a combination of two factors: 1) the time difference between when a reading is sensed, and when that value is recorded by the acquisition system, and 2) the time taken for the sensor to arrive at the location of the GPS antenna. The second factor is defined by the physical distance between the GPS antenna and any given sensor, and the speed of the aircraft. The total magnetic lag value for the AGIS acquisition system has been calculated to 0.19 s for this survey. The TDEM lag has been calculated to 1.11 s.

VIII. FIELD OPERATIONS

The survey operations were conducted out of the Sioux Lookout Airport on June 21st, 2021. The MAG-TDEM data acquisition required 3 flights. At the end of the production day, the data were sent to Dynamic Discovery Geoscience's office via internet. The data were then checked for Quality Control to ensure they fulfilled contractual specifications. The full dataset was inspected prior to provide authorization for the field crew to demobilize. The GEM-19 magnetic base station was set up in a magnetically quiet area close to the block, at latitude 50.5898783°N, longitude 90.6075191°W. The survey pilot was Guy Labelle and the survey system technician was Pascal St-Denis Mercier.

Figure 8: Example of a magnetic base station setup



IX. DIGITAL DATA COMPILATION

Data compilation including editing and filtering, quality control, and final data processing was performed by Joël Dubé, P.Eng. Processing was performed on high performance desktop computers optimized for quick daily QC and processing tasks. Geosoft software Oasis Montaj version 9.10 and Matlab R2018a were used.

Magnetometer Data

The airborne magnetometer data, recorded at 10 Hz, were plotted and checked for spikes and noise on a flight basis. A 0.19 second lag correction was applied to all data to correct for the time delay between detection and recording of the airborne data.

Ground magnetometer data were recorded at 1 sample per second and interpolated by a spline function to 10 Hz to match airborne data. Data were inspected for cultural interference and edited where necessary. Some low-pass filtering was deemed necessary on the ground station magnetometer data to remove minor high frequency noise. The diurnal variations were removed by subtracting the ground magnetometer data to the airborne data and by adding back the average of the ground magnetometer value.

The levelling corrections were applied in several steps. First of all, a correction for altitude was applied by multiplying the First Vertical Derivative of the pre-levelled data by the difference between the actual survey altitude and the average survey altitude. Standard levelling corrections were then performed using intersection statistics from traverse and tie lines. After statistical levelling was considered satisfactory, decorrugation was applied on the data to remove any remaining subtle non-geological features oriented in the direction of the traverse lines.

Once the Total Magnetic Intensity (TMI) was gridded, its First Vertical Derivative (FVD) and Second Vertical Derivative (SVD) were calculated to enhance narrower geological features. Finally, the component of the normal Earth's magnetic field, described by the International Geomagnetic Reference Field (IGRF), has been removed from the TMI to yield the residual TMI. This ensures that the very long wavelength signal within the block is indeed originating from the local geology and not from the Earth's expected regional gradient.

In order to enhance the subtle magnetic features some more, the Tilt Angle Derivative (TILT) was also computed for this project.

It has been shown that it is possible to use the Tilt Angle Derivative to estimate both the location and depth of magnetic sources (Salem et al., 2007).

When two body of different magnetic susceptibility are in contact, the vertical and horizontal gradients along a horizontal line perpendicular to the vertical contact are governed by the following equations:

$$\delta M/\delta h = 2KFc(z_c/(h^2+z_c^2))$$

$$\delta M/\delta z = 2KFc(h/(h^2+z_c^2))$$

where

K = susceptibility contrast

F = magnetic field's strength

c = $1 - \cos^2(\text{field Inclination})\sin^2(\text{field Declination})$

h = location along an horizontal axis perpendicular to the contact

z_c = contact depth

$$\delta M/\delta h = \text{sqrt}((\delta M/\delta x)^2 + (\delta M/\delta y)^2)$$

The Tilt Angle (θ) is defined as

$$\theta = \tan^{-1}[(\delta M/\delta z)/(\delta M/\delta h)]$$

By substitution of the gradients we get

$$\theta = \tan^{-1} [h/z_c]$$

This has two main implications for any given anomaly:

- 1- The 0° angle line is located directly above the contact between a magnetic source and the surrounding rock. This allow for accurate estimation of source location.
- 2- The distance between the 0° and the $+45^\circ$ lines as well as the distance between the -45° and the 0° lines are equal to the depth of the source at the contact. This allow for a direct estimation of the depth of the source of the anomaly. The depth estimated with this method is actually the distance between the magnetic sensor and the top of the source. Knowing that the sensor was 58 m above the ground in average enables direct depth estimates.

In practice, the signal originating from multiple sources at different depth within a same area will cause juxtaposition of the Tilt Angle values, and complicate location and depth estimation. Nevertheless, the method remains an excellent tool for rapid assessment of sources characteristics, without the need for complex assumptions to be made or heavy computer requirements, as is the case with 3D Euler deconvolution or 3D data inversions.

Radar Altimeter Data

The terrain clearance measured by the radar altimeter in metres was recorded at 10 Hz. The data were filtered to remove high frequency noise using a 1 sec low pass filter. The final data were plotted and inspected for quality.

Positional Data

Real time DGPS correction provided by Omnistar was applied to the recorded GPS positional data.

Positional data (Lat, long, UTM X, UTM Y, geoid height) were recorded at 10 Hz sampling rate and all data processing was performed in the WGS-84 datum. The delivered data are provided in X, Y locations in UTM projection zone 15 North, with respect to the NAD-83 (CSRS) datum. Altitude data were initially recorded relative to the GRS-80 ellipsoid, but are delivered as orthometric heights (MSL elevation).

Terrain Data

Terrain elevation data (also referred to as digital elevation model, or DEM) are computed from the altitude of the helicopter, given by DGPS recordings, and the radar altimeter data.

TDEM Data

The PicoEnvirotec EM Digital Acquisition System records the vertical component (Z) of the receiver coils at a sampling rate of 90000Hz. There are 30 full cycles (60 half cycles) of the full waveform (Tx ON and OFF time) every second.

The first data manipulation involves a stacking procedure where each half cycle is weighted with respect to the previous cycle ($\pm\frac{1}{4}$), the next cycle ($\pm\frac{1}{4}$) and its own value ($\pm\frac{1}{2}$). The positive and negative signs of the respective multiplication coefficients are used to make positive all negative half cycles. The next step is the half cycle averaging corresponding to the desired sampling rate. In the present case, from the 60 stacked positive half cycles per second, 6 consecutive half cycles are averaged to produce one sample every 0.1 sec.

The windowing settings for the 40 different channels are presented in Table 4. Channels 1 to 11 correspond to the ON-time measurements and channels 12 to 40 correspond to the OFF-time. Channel 12 isn't used for interpretation and mapping as some 'ramp-off' effects remain that alters the data quality. Each window is filtered with a median filter removing spikes and with a finite impulse response (FIR) selective filter of the 251th order improving the signal to noise ratio. An average lag correction of 1.11 sec was applied to the data after being empirically determined by flying a sharp anomaly in two opposite direction.

Table 5: Setting used in the windowing of the full waveform

Channel #	Starting time (msec)	Width (msec)	Pulse	Channel #	Starting time (msec)	Width (msec)	Pulse
1	0.16667	0.01667	ON	21	3.15000	0.53333	OFF
2	0.25000	0.01667	ON	22	3.26667	0.53333	OFF
3	0.33333	0.01667	ON	23	3.40000	0.53333	OFF
4	1.30000	0.01667	ON	24	3.40000	1.10000	OFF
5	1.31667	0.01667	ON	25	3.45000	1.10000	OFF
6	1.33333	0.01667	ON	26	3.65000	1.10000	OFF
7	2.58333	0.01667	ON	27	3.88333	1.10000	OFF
8	2.66667	0.01667	ON	28	4.13333	1.10000	OFF
9	2.80000	0.08333	ON	29	4.43333	1.10000	OFF
10	2.81667	0.08333	ON	30	4.76667	1.10000	OFF
11	2.83333	0.08333	ON	31	5.16667	1.10000	OFF
12	2.85000	0.16667	RAMP	32	5.20000	2.20000	OFF
13	2.86667	0.18333	OFF	33	5.55000	2.20000	OFF
14	2.86667	0.25000	OFF	34	6.13333	2.20000	OFF
15	2.86667	0.36667	OFF	35	6.78333	2.20000	OFF
16	2.91667	0.36667	OFF	36	7.51667	2.20000	OFF
17	2.91667	0.53333	OFF	37	8.36667	2.20000	OFF
18	2.95000	0.53333	OFF	38	9.33333	2.20000	OFF
19	3.00000	0.53333	OFF	39	10.4500	2.20000	OFF
20	3.03333	0.53333	OFF	40	11.7000	2.20000	OFF

As for the magnetic data, levelling corrections were applied to the TDEM data using intersection statistics from traverse and tie lines, as well as light decorrugation based on gridded information, in order to remove base line offsets. The levelled TDEM data are delivered in the database.

Gridding

The magnetic, early off-time TDEM (channel 13), mid off-time TDEM (channel 20), and late off-time TDEM (channel 27) data were interpolated onto a regular grid using a bi-directional gridding algorithm to create a two-dimensional grid equally incremented in x and y directions.

The final grids were created with 10 m grid cell size, appropriate for the survey lines spaced at 50 m. Traverse lines were used in the gridding process.

X. RESULTS AND DISCUSSION

Magnetic data

The Residual Total Magnetic Intensity (TMI) of the Solitude Lake block, presented in Figure 7 together with TDEM anomalies, is relatively active and varies over a range of 2,270 nT, with an average of -129 nT and a standard deviation of 196 nT.

The survey block can be subdivided in two distinct domains based on their magnetic responses. The first domain, which dominates most of the surveyed area, occurs as a wide central band depicting a “flattened U” or “banana” shape, with one end pointing to the northwest tip of the block and the other end pointing towards the northeast. The magnetic signal seen in this domain is very dynamic and organized as series of deformed linear magnetic features characteristic of alternating sequences of mafic volcanics with sedimentary or intermediate to felsic volcanic rocks, with possibly some small size intrusive stocks locally. The second domain is seen at the southwest and southeast tips of the block, as well as along its northern edge, sandwiching the first domain. It is characterized by mostly homogeneous magnetic textures and decreased signal variability, typical of large size felsic to intermediate intrusions or of areas dominated by sedimentary rocks. The strongest anomalies of the survey are found at the south end of the block and possibly relate to iron formations or ultramafic rocks. Stronger anomalies are best seen on Figure 8 which shows the residual TMI data with a linear color distribution. Other weaker anomalies that are still relatively strong are all located within the central magnetic domain and are likely associated to mafic volcanic or intrusive rocks.

Most magnetic lineaments are found within the central magnetic domain. They are generally trending WNW-ESE in the western part of the block, and are gradually changing to a NE-SW orientation towards the east, or even to a NNE-SSW strike at the northeast corner of the block. Several lineaments are locally curved, and even possibly heavily folded locally, mostly in the central part of the surveyed area, attesting that the area underwent strong deformation events in the past. In general terms, magnetic lineaments are related to rock formations that are enriched in magnetic minerals (magnetite and/or pyrrhotite).

Throughout the block, it is possible to detect structural features offsetting observed magnetic lineaments and causing abrupt interruption or changes of the magnetic response. These features are typically caused by faults, fractures and shear zones. If they are thought to be favorable structures in the exploration context of the Solitude Lake project, they should be paid particular attention and should be the object of a comprehensive structural interpretation, which is beyond the scope of this report.

Shorter wavelength anomalies are greatly enhanced on the FVD (Figure 9) and on the TILT (Figure 10) products. Since the FVD attenuates longer wavelength anomalies, and the TILT enhances very weak amplitude anomalies, they are the preferred products for structural interpretation. As well, a joint analysis of these results with the topography data (Figure 11) can help in the interpretation process of geological structures.

Figure 9: Total magnetic intensity with equal area color distribution and TDEM anomalies

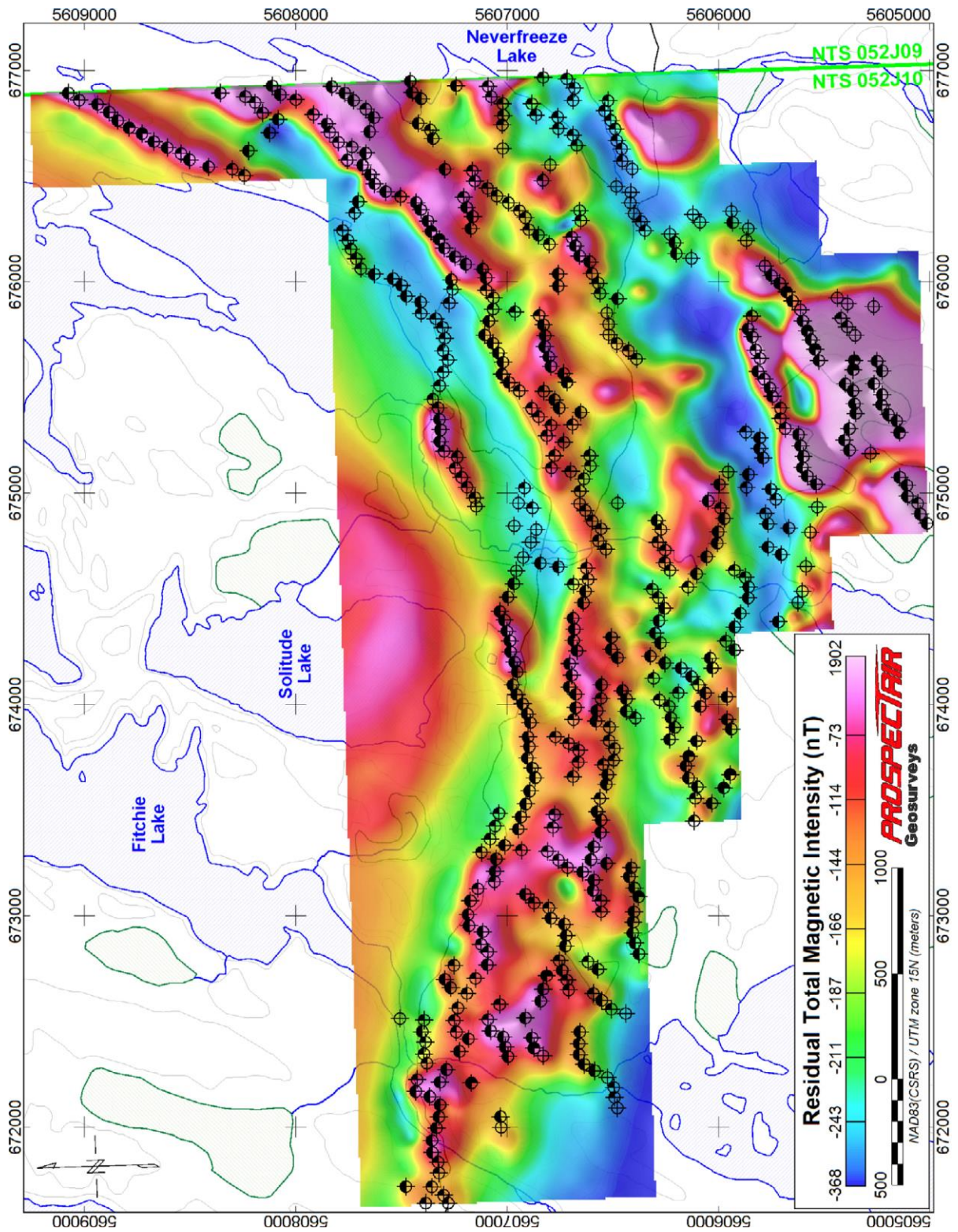


Figure 10: Total magnetic intensity with linear color distribution and TDEM anomalies

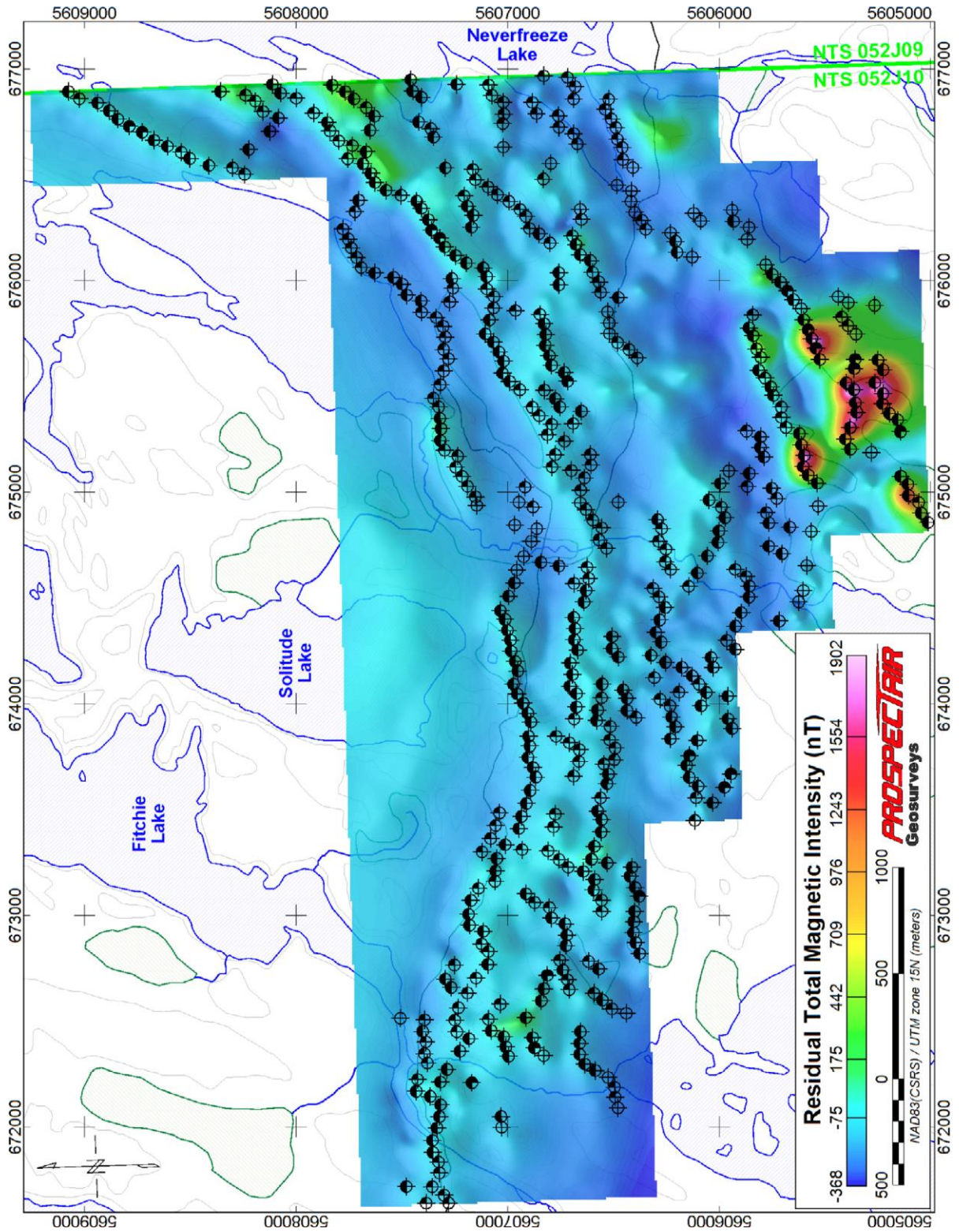


Figure 11: First vertical derivative of TMI and TDEM anomalies

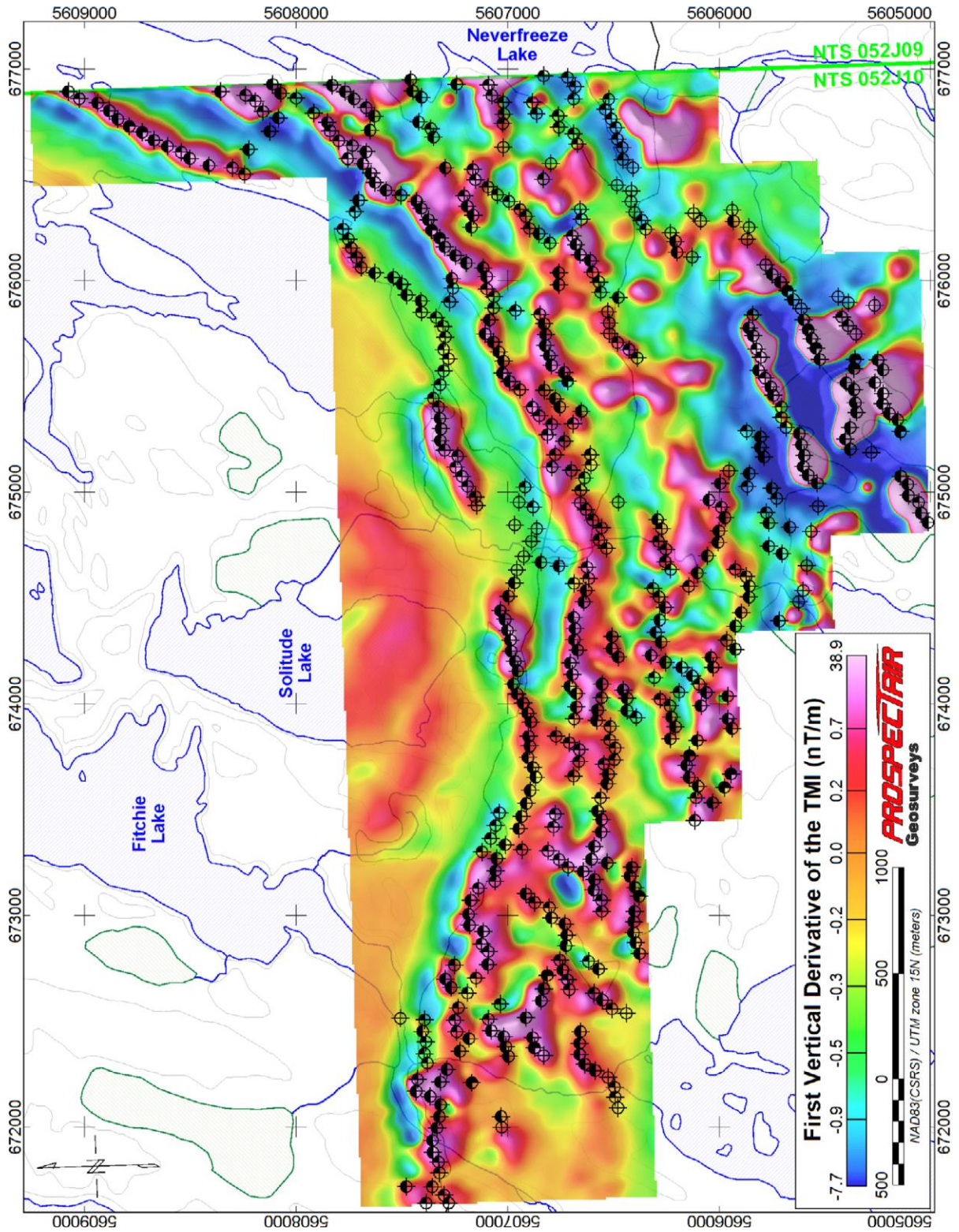


Figure 12: Magnetic tilt angle derivative and TDEM anomalies

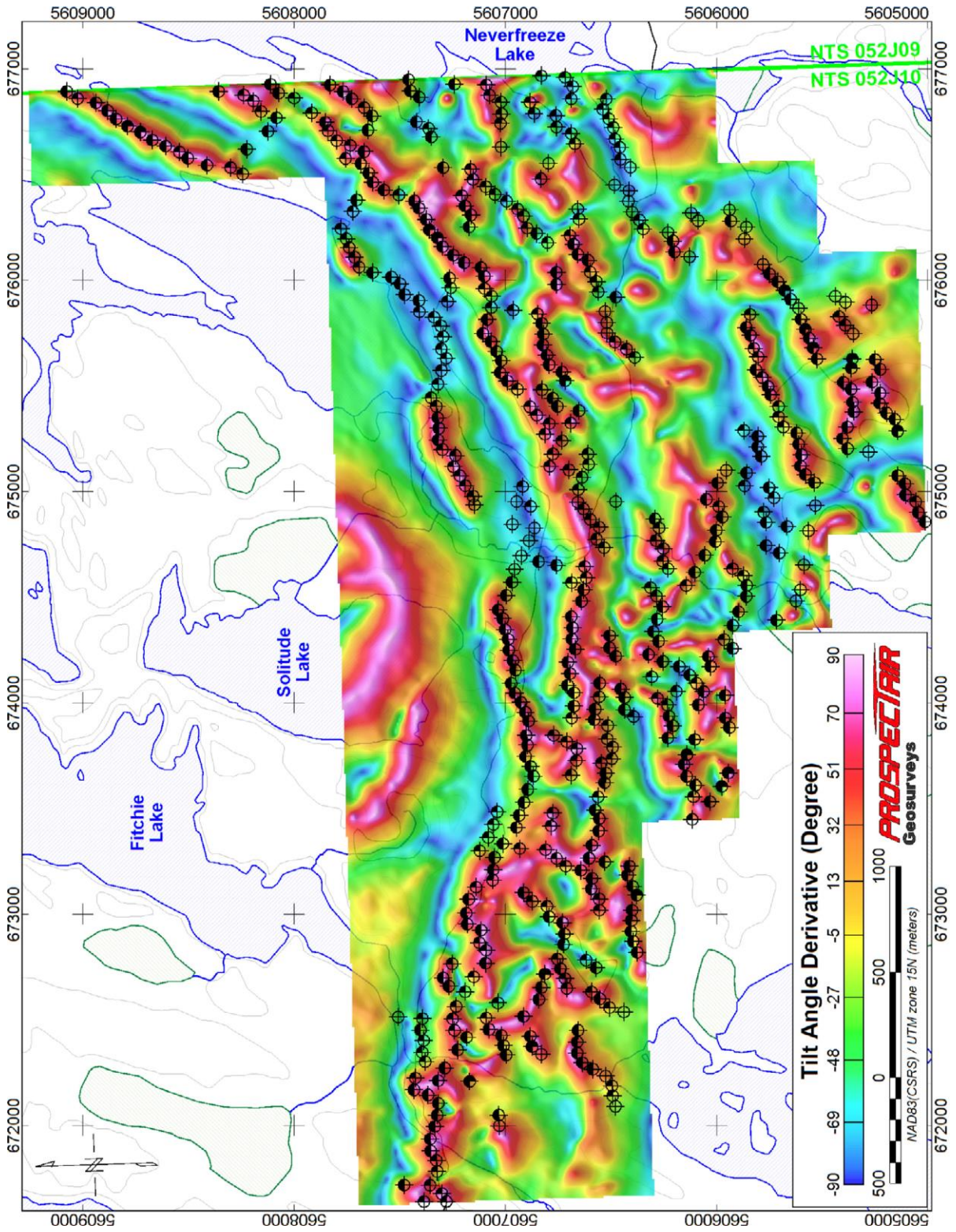
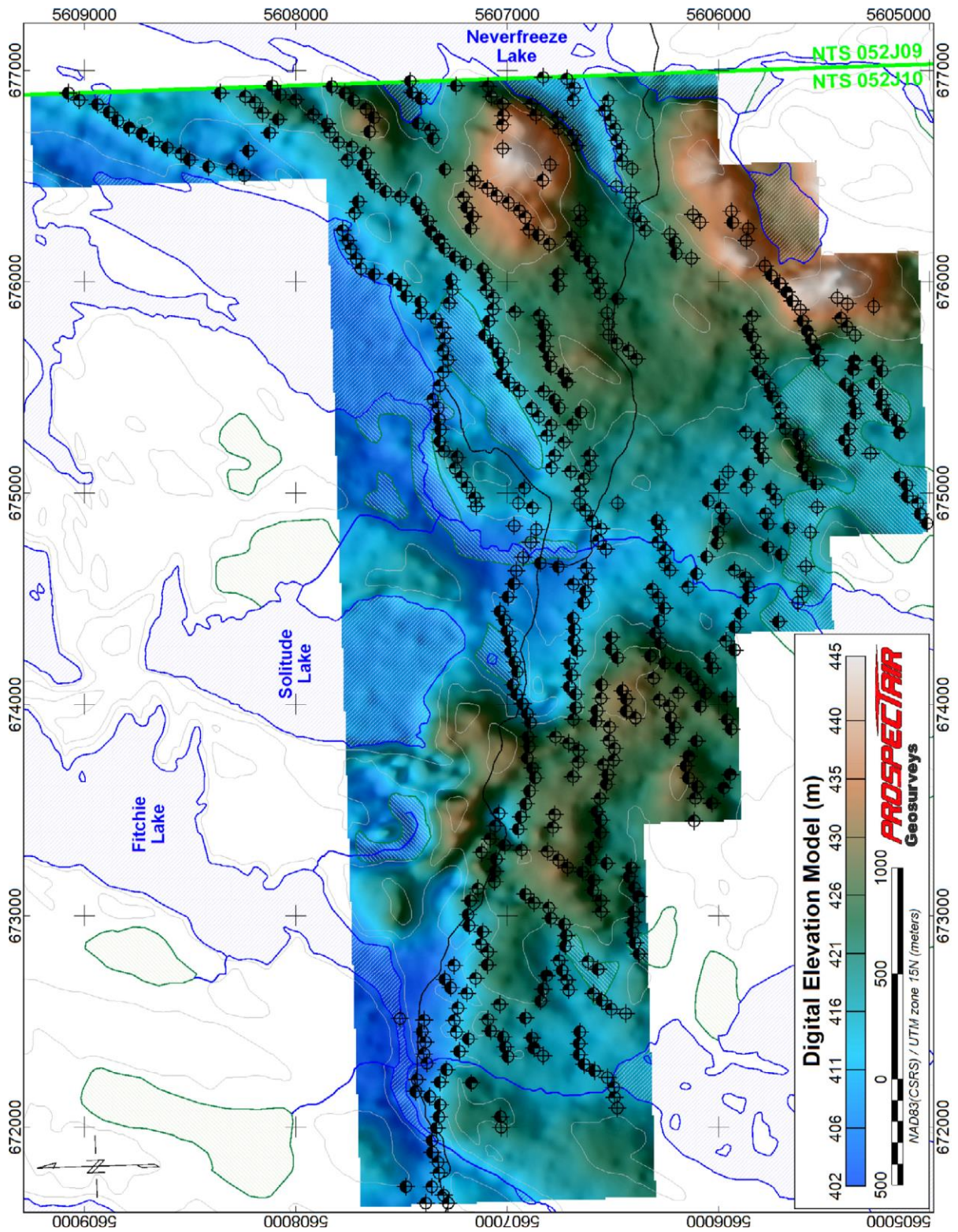


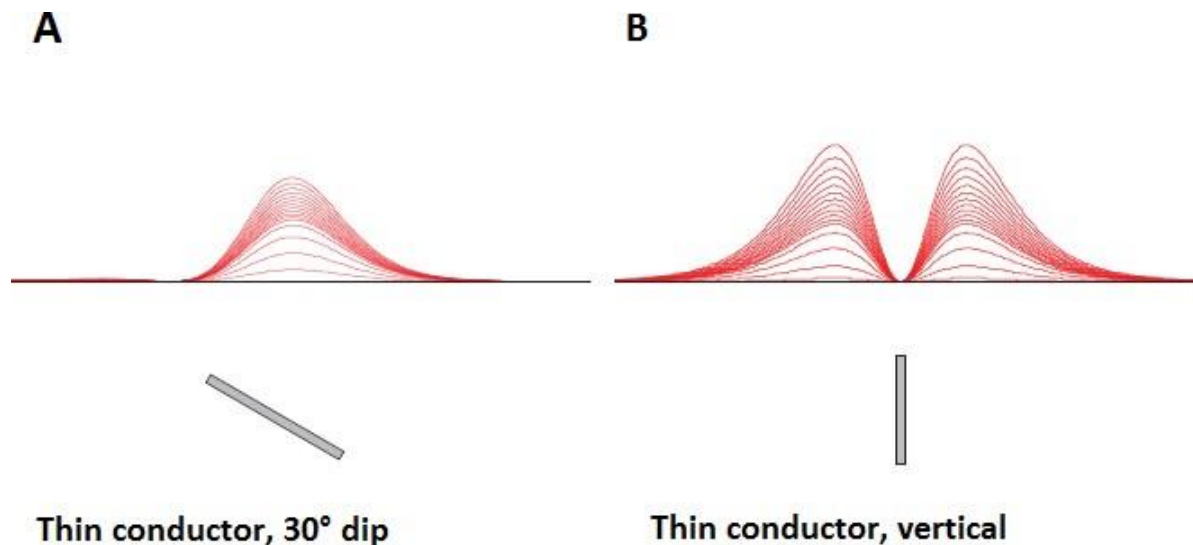
Figure 13: Digital elevation model and TDEM anomalies



Time-Domain Electromagnetic data

There is no automatic picking program involved in the interpretation procedures of the ProspecTEM system. Identification of the EM anomalies is made from the EM profiles. Most of the time, the location of anomalies is based on the assumption that the causative source is a somewhat thick or flat lying conductor, which would generate an anomaly mostly centered over the conductor (Figure 12, A). It is important to understand that some other conductive bodies could generate a strong EM response that is offset from the mass centre of the source. For instance, a thin conductor with a steep dip would generate an “M” shape anomaly (Figure 12, B), with the stronger shoulder on the dip side. Therefore, caution must be taken when planning work at the location of an anomaly. It is recommended to combine other available geoscientific information and to review the EM anomaly location before to investigate an anomaly of interest.

Figure 14: Example of EM response over thin conductors

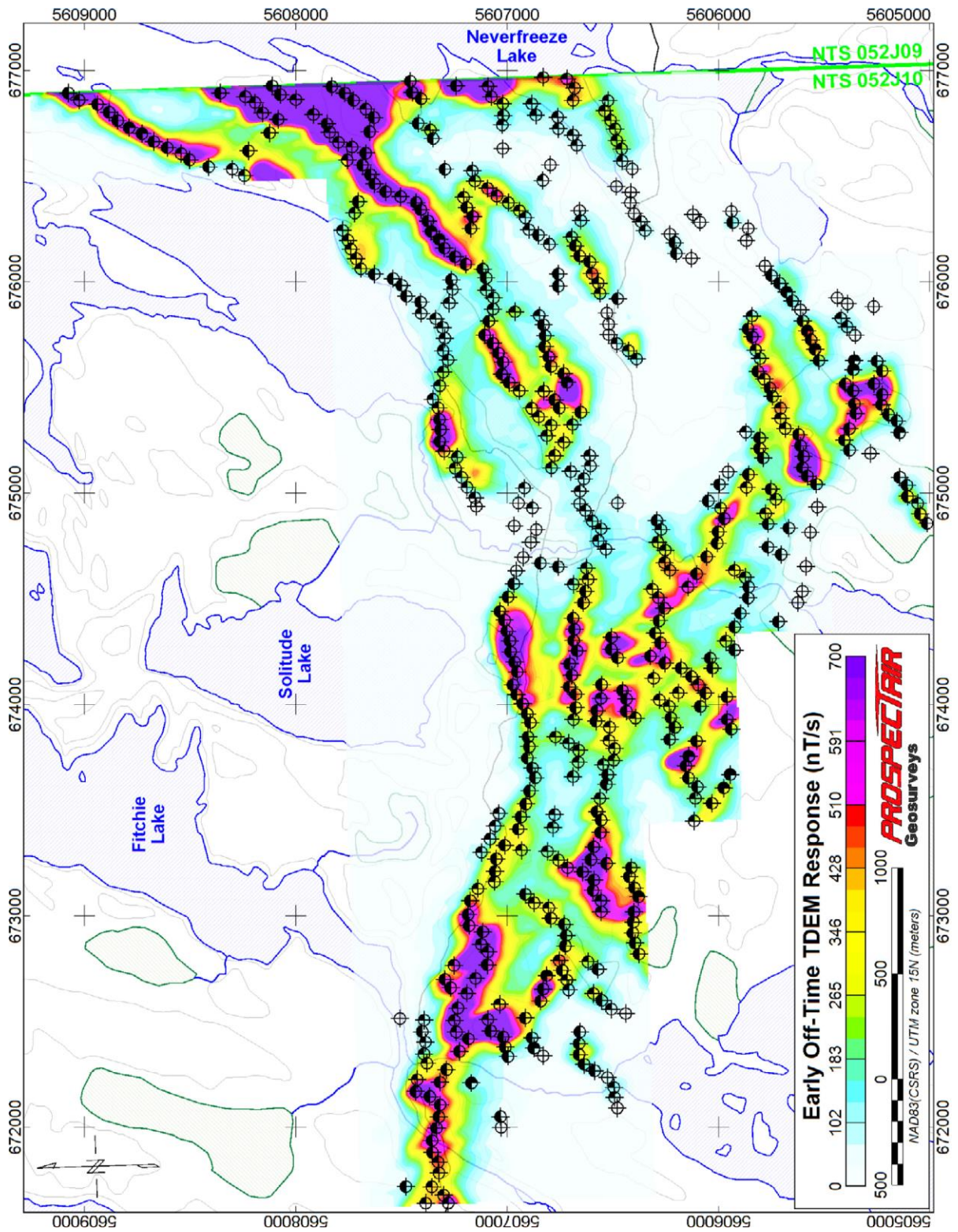


The classification of anomalies is based on the calculated time constant (TAU). The EM time constant is a general measure of the speed of decay of the electromagnetic response and reflects the “conductance quality” of a source. The decay rate of the secondary EM field recorded by the TDEM system is a function of the conductivity and geometry of conductors detected. A weak conductor, such as shallow conductive overburden, will show rapid response decay, thus a small value of the time constant. Conversely, a good conductor, such as a graphite or sulphide orebody, will have a response decaying slowly, relating to a large TAU value. The TAU is calculated using proprietary software and is derived from the best exponential least squares fit for channels Z13 to Z27. Calculating TAU for low amplitude anomalies that have their first off-time channel (channel 13) amplitude smaller than 75 nT/s can yield unreliable results given the weak response. As well, in some rare cases, despite stronger response of the first off-time channel, noise in the mid to late channels can cause the TAU estimation to be unreliable. No best fit were tried on these noisy or low signal anomalies and an arbitrary minimal time constant of 0.10 msec was attributed. Moreover, the resulting exponential best fit of the decay curve is extrapolated to the zero delay time, which can be used to compare the amplitude of anomalies.

On the Solitude Lake block, 539 EM anomalies are identified, classified and listed (Appendix C). All marginal/weak anomalies with TAU lower than 0.25 msec are included in a group represented by an empty circle on the anomaly map. In total, 46 anomalies are reported in this class. The remaining anomalies are classified in 4 other groups, with time-constant considered small (0.25 to 0.50 msec, 205 anomalies), intermediate (0.50 to 0.75 msec, 270 anomalies), strong (0.75 to 1.00 msec, 18 anomalies) and very strong (over 1.00 msec, 0 anomalies). These anomalies are reported on all the figures of this section, and the symbols used are similar to the legend on the maps. The early off-time map (Figure 13) provides a good overview of the TDEM response amplitude distribution.

EM anomalies seen in the area are characteristic of graphite and sulphide conductors. They are most often narrow, of variable amplitude, and usually with larger TAU values compared to typical conductive overburden EM responses, which is denoting better quality conductors. The orientation of conductive lineaments is clearly aligned with magnetic trends, which suggests that most conductive sources are indeed embedded in the bedrock. Furthermore, many conductive lineaments actually show a positive correlation to magnetic lineaments (Figures 7 to 10), which indicates that sulphides (including pyrrhotite), or possibly magnetite, are likely to compose at least part of the conductive sources. The most magnetic conductors are likely related to magnetite rich horizons, with possibly some contribution from sulphides occurrences, while the less magnetic ones could pertain to mineralized veins or smaller size sulphides occurrences with limited amounts of pyrrhotite. These conductors detected by the survey can be considered of interest for exploration in the Solitude Lake exploration context. This being said, some of the isolated and marginal anomalies, especially those found in topographic depressions, could also relate to local occurrences of conductive overburden.

Figure 15: Early off-time TDEM response and anomalies



XI. WORK RECOMMENDATION

The discussion on the geological implication of the survey data is minimal in this report. A more general study including information regarding the local geology and all other geoscience data available in the area would be necessary to extract the full potential of the geophysical data and help to confirm and prioritize exploration targets.

EM anomalies detected by this survey could be investigated with basic ground prospecting methods at first. If interesting results are obtained, or if overburden proves too thick for prospecting, it is recommended to use ground resistivity/IP or EM techniques, depending on the nature of the sources, to accurately define targets for stripping and/or drilling. The implementation of a geochemical soil sampling program or of a till sampling program could also help further prioritize outlined anomalies.

In addition, given the geological context that may be considered prospective for gold mineralization, it is also recommended to use the newly acquired magnetic data, together with known local geological information, to carry out a comprehensive structural interpretation. In the case of gold lode deposits, the geophysical signature is often very subtle given the absence of marked physical properties contrast. The best approach is rather indirect, and consists in looking for geophysical signatures typical of faults and deformation structures, where gold bearing dilation zones can develop. The recommended structural interpretation work could help identifying structures that could then be investigated further.

XII. FINAL PRODUCTS

Digital line data

The Geosoft database is provided with the channels detailed in Table 5.

Table 6: MAG-TDEM line data channels

No.	Name	Description	Units
1	UTM_X	UTM Easting, NAD-83, Zone 15N	m
2	UTM_Y	UTM Northing, NAD-83, Zone 15N	m
3	Lat_deg	Latitude in decimal degrees (WGS-84)	Deg
4	Long_deg	Longitude in decimal degrees (WGS-84)	Deg
5	GPS_Z	Helicopter altitude (w.r.t. MSL)	m
6	Gtm_sec	Second since midnight GMT	Sec
7	Radar	Ground clearance given by the radar altimeter	m
8	DEM	CDED Digital Elevation Model (w.r.t. MSL)	m
9	Terrain	Digital Elevation Model calculated from GPS and Radar	m
10	Mag_Raw	Raw magnetic data	nT
11	Mag_Lag	Lagged magnetic data	nT
12	Gnd_mag	Base station magnetic data	nT
13	Mag_Cor	Magnetic data corrected for diurnal variation	nT
14	TMI	Fully levelled Total Magnetic Intensity	nT
15	TMIres	Residual TMI (IGRF removed)	nT
16	OFF_TIME	Amplitude of Off-time channels (13 to 36)	nT/s

Maps

All maps are referred to NAD-83 in the UTM projection Zone 15 North, with coordinates in metres. Maps are at a 1:10,000 scale. They are provided in PDF, PNG, Geotiff and Geosoft MAP formats for the products detailed in Table 6.

Table 7: Maps delivered

No.	Name	Description
1	DEM+FlightPath_Claims	Digital Elevation Model with flight path and properties claims
2	TMI	Residual Total Magnetic Intensity
3	FVD	First Vertical Derivative of the TMI
4	TILT	Tilt Angle Derivative of the TMI
5	Early_OffTime	Early_Off-Time TDEM response (Channel 13)
6	TDEM_Profiles+Anomalies	TDEM profiles with anomalies
7	TILT +TDEM_Anomalies	Tilt Angle Derivative of the TMI with TDEM anomalies

Grids

All grids are referred to NAD-83 in the UTM projection Zone 15 North, with coordinates in metres. Grids are provided in Geosoft GRD format, with a 10 m grid cell size, as well as in the Geotiff format for the products listed in Table 7.

Table 8: Grids delivered

No.	Name	Description	Units
1	DEM	CDED Digital Elevation Model	m
2	TERRAIN	Digital Elevation Model measured by helicopter	m
3	TMI	Total Magnetic Intensity	nT
4	FVD	First Vertical Derivative of TMI	nT/m
5	SVD	Second Vertical Derivative of TMI	nT/m ²
6	TMIres	Residual TMI (IGRF removed)	nT
7	TILT	Tilt Angle Derivative of the TMI	Degree
8	Early_Off-Time	Early Off-Time TDEM response (Channel 13)	nT/s
9	Mid_Off-Time	Mid Off-Time TDEM response (Channel 20)	nT/s
10	Late_Off-Time	Late Off-Time TDEM response (Channel 27)	nT/s

Project report

The report is submitted in PDF format. The anomaly table presented in annex is also provided as a separate Excel spreadsheet.

Respectfully submitted,



Joël Dubé, P.Eng.

July 14th 2021

XIII. REFERENCES

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Thomson, J; 1968: Diamond drill logs. Canadian Nickel Co. AFRI document 52J10SE0007

XIV. Statement of Qualifications

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E-mail: jdube@ddgeoscience.ca

I, Joël Dubé, P.Eng., do hereby certify that:

1. I am a Professional Engineer specialized in geophysics, President of Dynamic Discovery Geoscience Ltd., registered in Canada.
2. I earned a Bachelor of Engineering in Geological Engineering in 1999 from the École Polytechnique de Montréal.
3. I am an Engineer registered with the Ordre des Ingénieurs du Québec, No. 122937, and a Professional Engineer with Professional Engineers Ontario, No. 100194954 (CofA No. 100219617), with the Association of Professional Engineers and Geoscientists of New Brunswick, No. L5202 (CofA No. F1853), with the Association of Professional Engineers of Nova Scotia, No. 11915 (CofC No. 51099), with Engineers Geoscientists Manitoba, No. 43414. (CofA No. 6897), with Professional Engineers & Geoscientists Newfoundland & Labrador, No. 10012 (PtoP No. N1134) and with the Northwest Territories Association of Professional Engineers & Geoscientists, No. L4447 (PtoP No. P1414).
4. I have practised my profession for 22 years in exploration geophysics.
5. I have not received and do not expect to receive a direct or indirect interest in the properties covered by this report.

Dated this 14th day of July, 2021

Joël Dubé, P.Eng. #100194954

XV. Appendix A – Survey block outline

Solitude Lake Block

Easting	Northing
676143	5605054
674811	5605010
674796	5605473
674353	5605458
674338	5605922
673453	5605892
673437	5606356
671668	5606297
671622	5607692
676487	5607854
676441	5609243
676888	5609258
676997	5606011
676555	5605996
676570	5605533
676128	5605518

XVI. Appendix B – Property claims numbers covered by the survey

Tenure number
653237
653250
657731
657732
657733
657734
657735

XVII. Appendix C – Solitude Lake block TDEM anomaly table

Line	UTM_X (m)	UTM_Y (m)	ID	Time Constant (msec)	Amplitude at zero delay (nT/s)
70	671639	5607280	70.01	0.49	1225
80	671676	5607301	80.01	0.45	732
80	671639	5607386	80.02	0.44	701
90	671716	5607357	90.01	0.44	705
100	672089	5606479	100.01	0.10	0
100	671782	5607323	100.02	0.41	749
100	671717	5607480	100.03	0.51	148
110	672140	5606493	110.01	0.39	180
110	671833	5607327	110.02	0.41	852
120	672186	5606496	120.01	0.38	255
120	671997	5607029	120.02	0.10	0
120	671878	5607354	120.03	0.51	727
130	672233	5606527	130.01	0.10	0
130	672048	5607032	130.02	0.57	107
130	671935	5607356	130.03	0.56	943
140	672265	5606587	140.01	0.51	188
140	671994	5607335	140.02	0.51	1036
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150	672047	5607323	150.02	0.51	716
160	672342	5606661	160.01	0.65	459
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170	672394	5606660	170.01	0.56	316
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190	672421	5606883	190.02	0.58	213
190	672377	5607010	190.03	0.55	536
190	672270	5607288	190.04	0.65	577
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200	672425	5607019	200.02	0.60	1214
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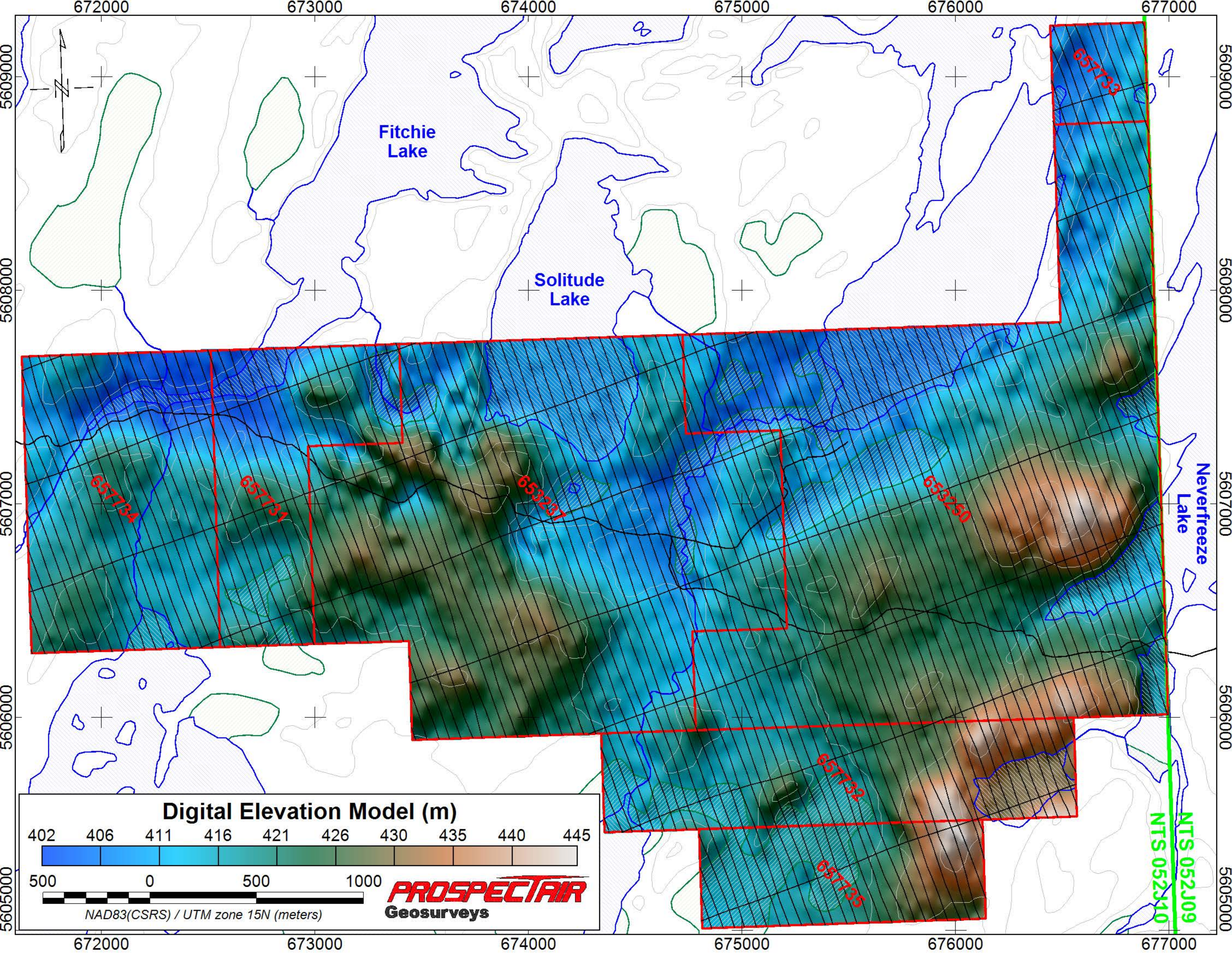
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850	676184	5606208	850.02	0.63	155
850	676037	5606599	850.03	0.48	806
850	675978	5606759	850.04	0.60	101
850	675870	5607067	850.05	0.50	265
850	675782	5607306	850.06	0.35	205
860	676331	5605940	860.01	0.10	0
860	676280	5606088	860.02	0.10	0
860	676223	5606230	860.03	0.10	0
860	676093	5606609	860.04	0.60	373
860	676036	5606759	860.05	0.70	101
860	675922	5607065	860.06	0.44	276
860	675824	5607338	860.07	0.33	193
870	676316	5606121	870.01	0.10	0
870	676241	5606348	870.02	0.40	211
870	676121	5606658	870.03	0.69	478
870	675959	5607100	870.04	0.50	245
870	675898	5607274	870.05	0.38	205
870	675849	5607408	870.06	0.31	186
880	676283	5606366	880.01	0.41	146
880	676169	5606678	880.02	0.58	357
880	676017	5607100	880.03	0.51	284
880	675963	5607259	880.04	0.47	173

880	675902	5607410	880.05	0.51	141
890	676321	5606401	890.01	0.10	0
890	676210	5606695	890.02	0.40	333
890	676176	5606801	890.03	0.50	131
890	676059	5607116	890.04	0.60	269
890	676008	5607265	890.05	0.54	160
890	675929	5607477	890.06	0.65	100
900	676373	5606418	900.01	0.10	0
900	676288	5606652	900.02	0.32	289
900	676220	5606852	900.03	0.60	144
900	676081	5607196	900.04	0.53	871
900	675981	5607499	900.05	0.63	109
910	676421	5606416	910.01	0.10	0
910	676335	5606658	910.02	0.10	0
910	676245	5606898	910.03	0.50	163
910	676117	5607247	910.04	0.56	1297
910	676013	5607540	910.05	0.71	107
920	676450	5606484	920.01	0.10	0
920	676302	5606914	920.02	0.39	243
920	676158	5607298	920.03	0.62	1470
920	676034	5607629	920.04	0.58	157
930	676533	5606409	930.01	0.10	0
930	676335	5606943	930.02	0.41	509
930	676251	5607172	930.03	0.52	614
930	676199	5607322	930.04	0.80	1286
930	676060	5607693	930.05	0.40	534
940	676570	5606456	940.01	0.31	249
940	676372	5606995	940.02	0.32	540
940	676308	5607164	940.03	0.50	925
940	676237	5607361	940.04	0.78	1151
940	676108	5607715	940.05	0.36	742
950	676627	5606459	950.01	0.29	512
950	676478	5606832	950.02	0.29	171
950	676405	5607050	950.03	0.37	1026
950	676351	5607192	950.04	0.53	438
950	676286	5607376	950.05	0.68	1164
950	676154	5607740	950.06	0.28	757
960	676667	5606482	960.01	0.28	564
960	676552	5606796	960.02	0.10	0
960	676441	5607094	960.03	0.37	966
960	676398	5607207	960.04	0.42	479
960	676340	5607409	960.05	0.55	994
960	676195	5607754	960.06	0.30	599
970	676721	5606490	970.01	0.33	452
970	676645	5606671	970.02	0.36	176
970	676475	5607154	970.03	0.45	386
970	676373	5607431	970.04	0.54	1376
970	676242	5607779	970.05	0.32	335
980	676402	5607504	980.01	0.47	1694
980	676762	5606511	980.02	0.26	630
980	676693	5606689	980.03	0.43	169
980	676523	5607165	980.04	0.43	182
980	676323	5607722	980.05	0.32	443
990	676808	5606540	990.01	0.28	615
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990	676425	5607575	990.05	0.60	2117
990	676378	5607705	990.06	0.51	364
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1070	676746	5607853	1070.03	0.51	3052
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1070	676543	5608413	1070.05	0.60	112
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1080	676703	5608124	1080.04	0.69	405
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1090	676605	5608544	1090.03	0.47	986
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1100	676861	5608000	1100.02	0.48	2149
1100	676797	5608157	1100.03	0.42	3146
1100	676633	5608609	1100.04	0.65	1057
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1110	676662	5608672	1110.03	0.69	1000
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1130	676891	5608358	1130.01	0.50	1397
1130	676727	5608789	1130.02	0.80	954
1140	676761	5608843	1140.01	0.61	1021
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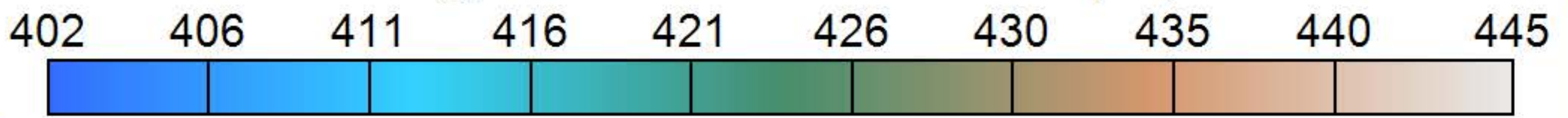


Fitchie Lake

Solitude Lake

Neverfreeze Lake

Digital Elevation Model (m)



NAD83(CSRS) / UTM zone 15N (meters)

PROSPECTAIR
Geosurveys

657733

657734

657731

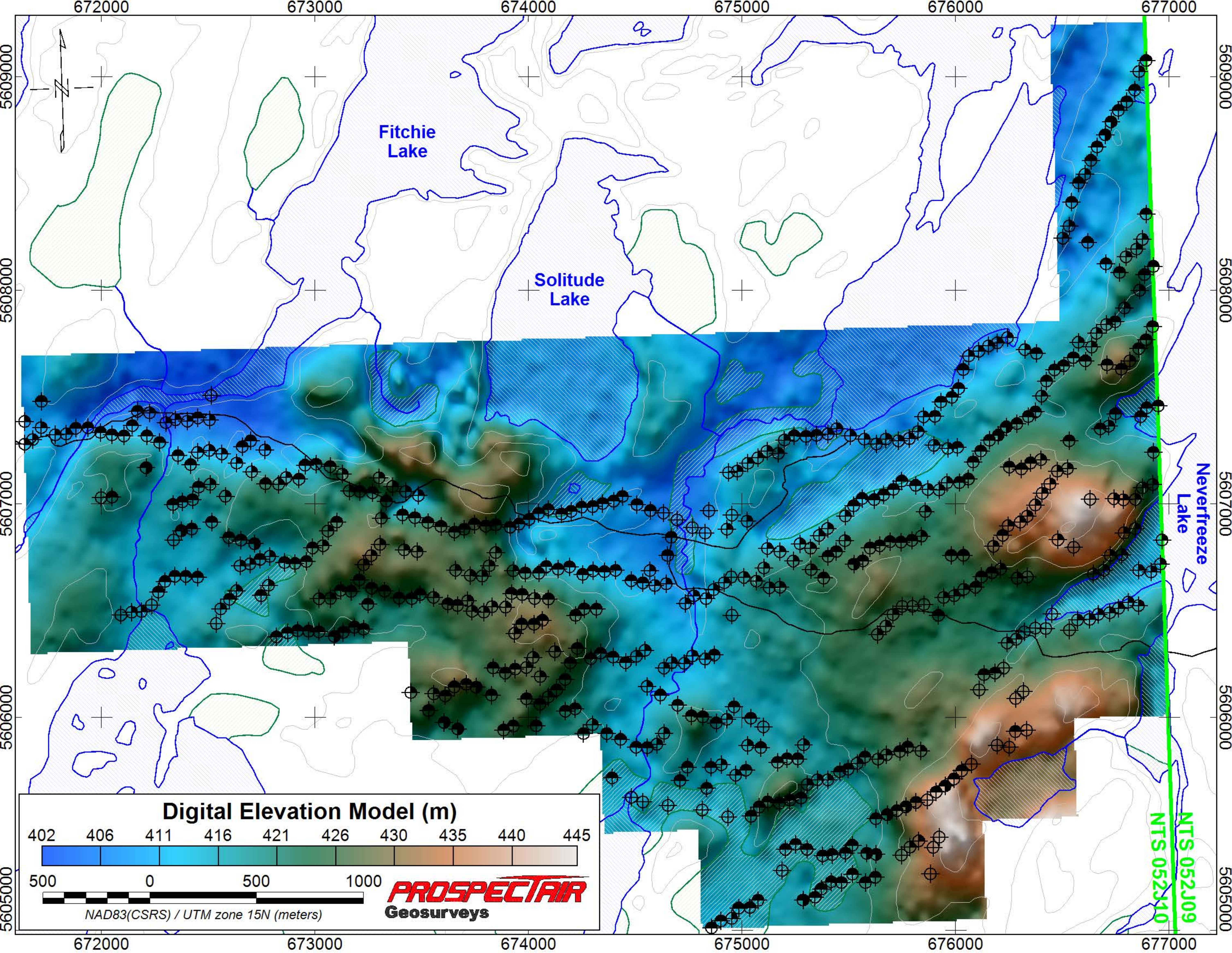
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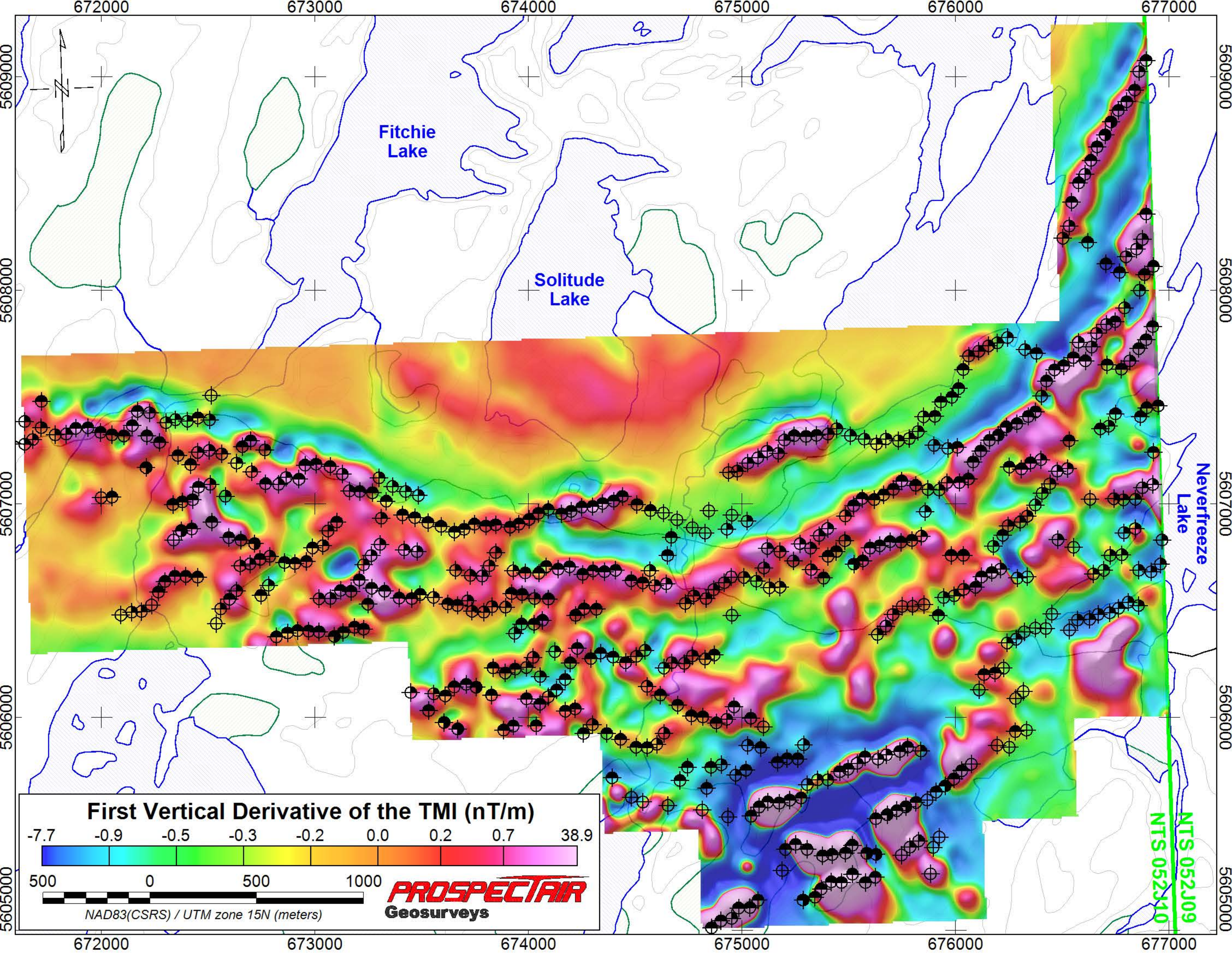
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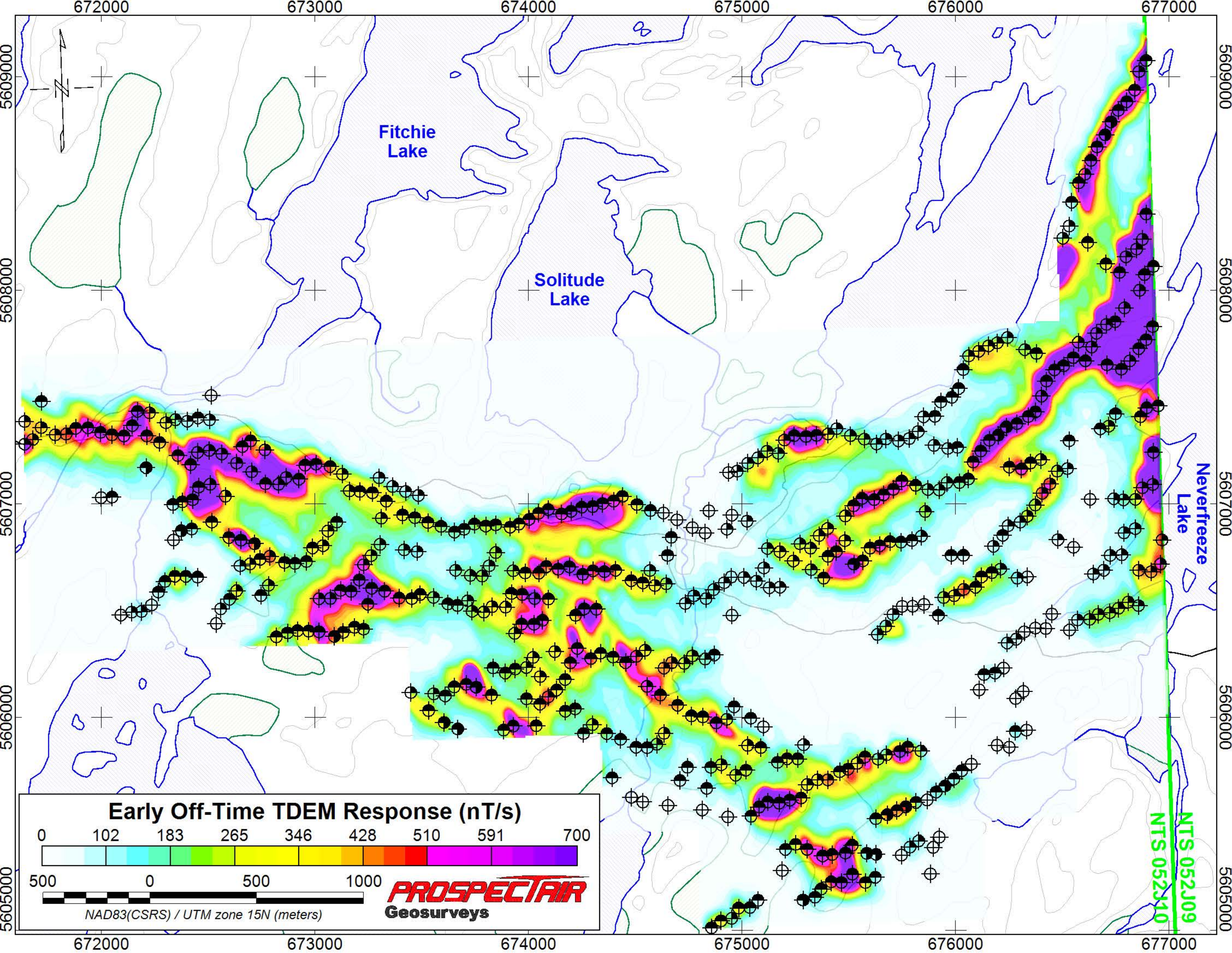
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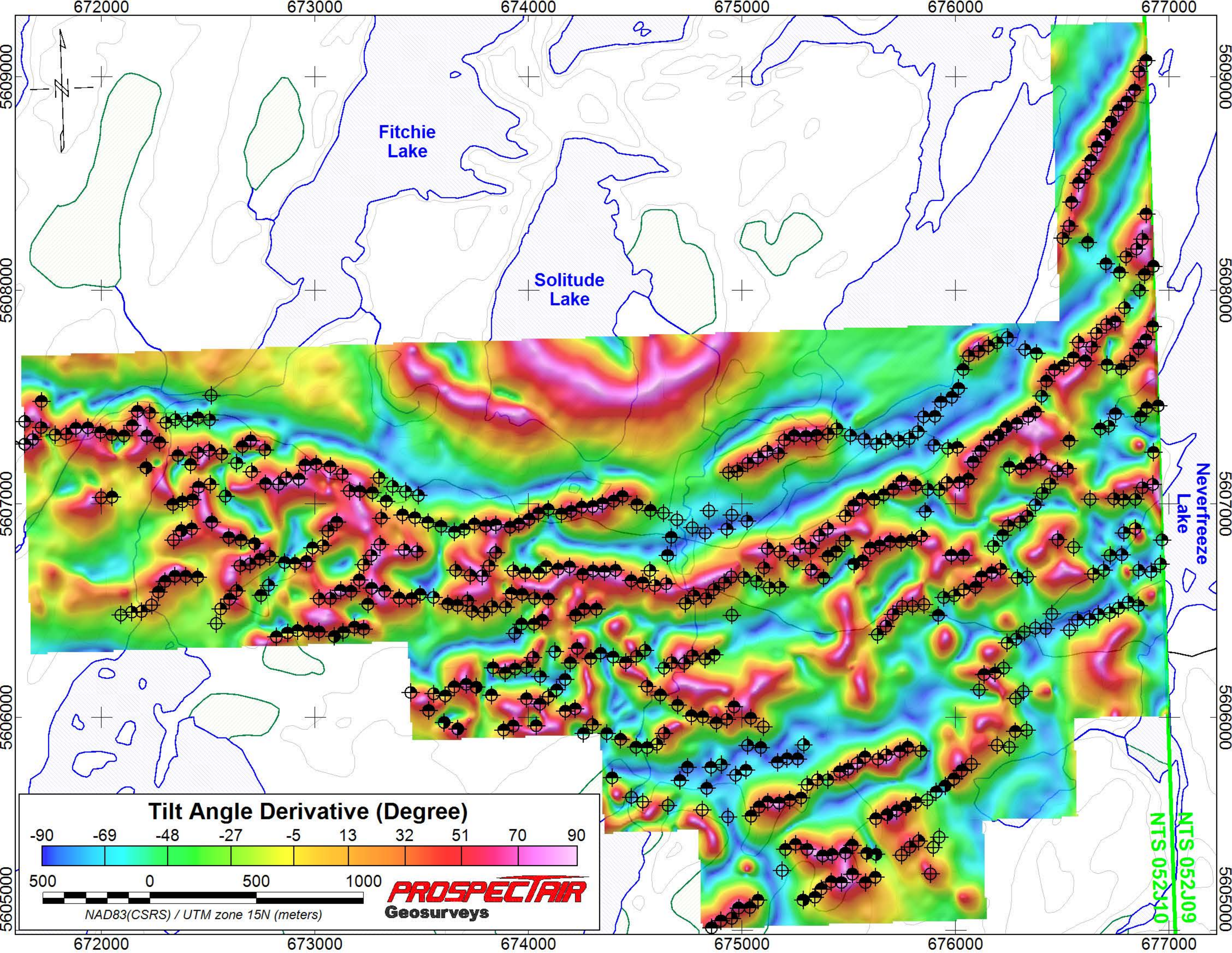
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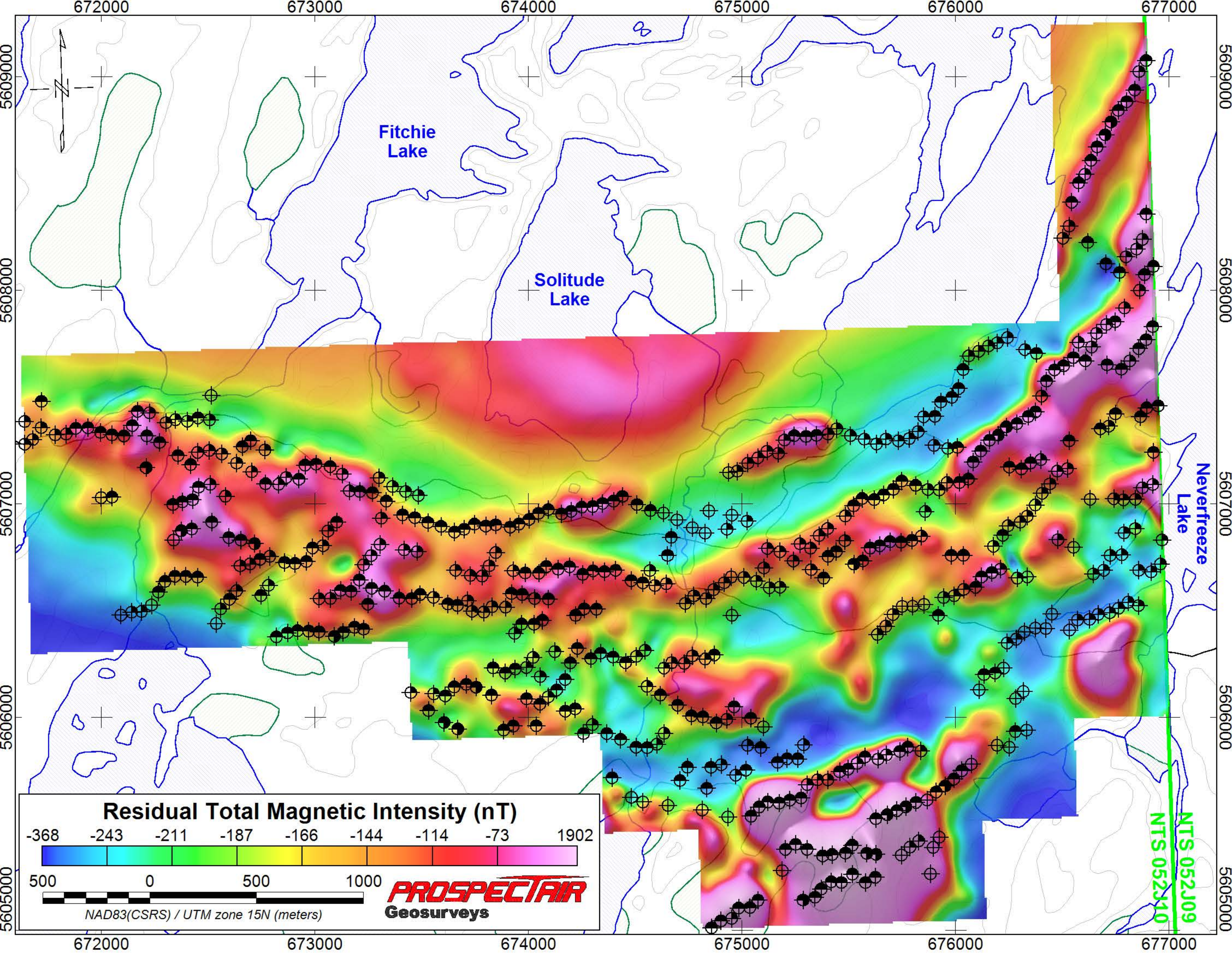
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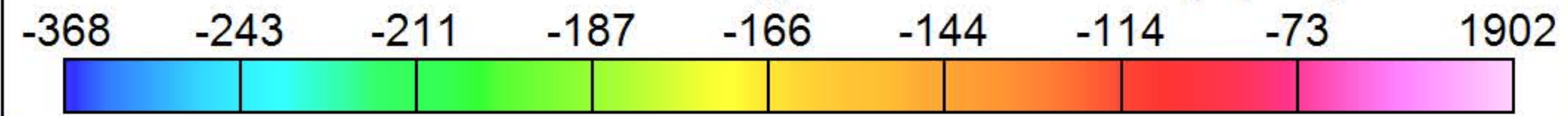


Fitchie
Lake

Solitude
Lake

Neverfreeze
Lake

Residual Total Magnetic Intensity (nT)



NAD83(CSRS) / UTM zone 15N (meters)

PROSPECTAIR
Geosurveys

NTS 052J10
NTS 052J10

