

GEOLOGY AND GOLD DEPOSITS

MURGOLD RESOURCES

GOGAMA AREA

by

NORMINEX LIMITED

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INTRODUCTION

The property consists of 295 claims situated in Chester, Benneweiss and St. Louis townships about midway between Sudbury and Timmins. Convenient access to the area is provided by provincial highway 144. The claim group constitutes the eastern extension of the Swaze "greenstone belt".

Geological mapping at a scale of 1 inch to 400 feet was carried out from May to September, 1981. Vertical aerial photographs originally at 1 inch to 3 mile but enlarged to 1 inch to 400 feet provided the basic control for mapping. Surveyed township lines, highway 144, power lines, bush roads, and base lines established by Murgold Resources provided additional controls. Rock outcrops were examined along and on either side of lines spaced 300 to 600 feet apart. The entire property was covered in this manner even though large areas lacked rock exposure. Geological interpretation of such areas was greatly assisted by using filtered aeromagnetic data and air photo examination.

Chester and Benneweiss townships lie just north of the height of land marking the boundary between the Arctic and Atlantic watersheds. Although the area is topographically high and relief commonly exceeds 100 feet, large expanses of well-exposed bedrock are not common. Most of the high ground consists of glacial debris through which only small rock knolls protrude. Even these are usually covered by a thin veneer of dirt and moss. Much of the overburden consists of boulder till and glacial outwash, sand and gravel. The glacial deposits may range in thickness from a few feet to more than 100 feet in areas that were topographically low at the time the ice melted. A typical example of thick and extensive glacial deposits is the area bordering Mesomikenda Lake which occupies a fracture-controlled linear depression trending NNW. The shores of this lake are characterized by steep high bluffs of boulders, gravel and sand. The shores of the lake and the terrain for several hundred feet on either side of the lake are almost devoid of outcrop.

The only published geological report of the area encompassing the Murgold property is that by H. C. Laird, Geology of the Three Ducks Lake Area, being part III of the 41st annual report of the Ontario Department of Mines for 1932. Laird's mapping at 1 inch to 3/4 mile was done shortly after the original discovery of gold in Chester township.

The results of the current geological survey are presented on 3 sheets. Sheets 1 and 3 cover the northern parts of Chester and Benneweiss townships; sheet 2 covers that part of the property immediately south of the other two sheets.

Aeromagnetic map and aerial photos of the property provided much useful geological information. For example, a number of linear magnetic highs trending about 165° are clearly visible on the total field aeromagnetic maps. It has been determined that these are caused by magnetite-bearing diabase dykes whose magnetism permits them to be traced across areaslacking rock exposure. In addition to this common 165° trend a prominent and persistent magnetic feature (also caused by diabase) trends at 110° from the NW corner of the property to the western shore of Mesomikenda Lake (sheet 2). Across the lake, on the east shore the magnetic high (and the diabase) re-appear some 4500° southward. A narrower dyke lies north of and trends parallel to the larger diabase. This 110° trend is also a common direction for quartz veins in the area.

About 4000' south of Arethusa Lake (bottom of sheet 1) is an easterly trending unnamed lake (top of sheet 2) through which an easterly trending aeromagnetic feature passes. This apparently represents a structure along which shearing occurs in at least two places: 1) on the north shore of the above-mentioned lake and 2) near the west shore of Mesomikenda Lake (Southcamp Bay).

A computer filtering technique described in a separate report was applied to the aeromagnetic data for Chester township. This technique separates the total field magnetism into a residual portion representing near-surface magnetism and a regional portion representing deeper magnetism. The report on this procedure was completed prior to compilation of the geological mapping and the significance of the residual magnetism was not fully understood until the geology was being compiled. It is now apparent that residual magnetic values greater than 0 (and up to several hundred) gammas represent areas of gabbro-diorite, as contrasted with more granitic rocks which yield negative residual values. For example, the boundaries of the large irregular gabbro-diorite mass trending south from Mill Pond (see sheets 1 and 2) coincide remarkably well with the 0 contour of the residual magnetic map. This observation has permitted the delineation of

gabbro as contrasted with granodiorite in the largely overburn-covered area immediately west of Southcamp Bay (sheets 1 and 2) and of the gabbroic body in the southwestern part of the property (sheet 2).

Air photo interpretation was useful in revealing structural features even in areas where little bedrock is exposed. For example, fracture systems through Mill Pond and Arethusa Lake are well-defined. Both NNW and ESE shear and fracture zones are visible in the vicinity of the small lake 4000' south of Arethusa Lake. Less clearly defined but nonetheless visible is a fine "grain" in the vegetation trending at 110°. When first noted, this "grain" appeared to be a shadow effect. However, the shadows trend at about 255°. The 110° grain probably reflects small-scale pervasive foliation parallel to the larger 110° fractures and veins. Large 110° lineaments, the fine 110° "grain" and a series of NNW lineaments are visible on photo 45-191. Only the most prominent of the lineaments are shown on the geological map.

GEOLOGY OF THE PROPERTY

The property is underlain largely by intrusive igneous rocks ranging in composition from gabbroic to granitic and in texture from fine-grained to coarse-grained. The dominant feature in the western part of the property (sheets 1 and 2) where most of the known gold-bearing quartz veins occur, is a more or less rectangular but irregular northerly trending strip of gabbro to diorite sandwiched between a coarse-grained quartz-eye granodiorite on the east and a fine- to medium-grained alaskite granite on the west. Both of the granitoid bodies are intrusive into the gabbro-diorite. The outline of this gabbro-diorite coincides with the O contour on the residual magnetic field map. Similar mafic igneous rocks occur south and west of the west end of Benneweiss Lake and immediately west of Southcamp Bay on Mesomikenda Lake.

The diorite and gabbro and both types of granitoid rocks are intrusive into rhyolitic volcanic and greywacke-type sedimentary rocks. The rhyolites and greywacke, with associated conglomerates, occur along the northern edge of the property (Sheets 1 and 3) and in the area of Benneweiss Lake.

Irregular zones of hybrid material named hybrid felsite are found in a few places near the contacts of granitoid intrusions and older rocks.

The youngest and freshest rocks on the property are diabase dykes, occupying well-defined fracture zones trending approximately north and a little north of west. The various lithologies are briefly described below along with their most important petrographic features. More complete thinsection descriptions are included in the Appendix to this report.

RHYOLITE, RHYOLITE PORPHYRY

Fine-grained, in part porphyritic, light-coloured felsic rocks associated with and assumed to underlie the greywacke-conglomerate are either massive or foliated, in some places well banded. These rocks are classified as rhyolite but may more correctly be called rhyolitic crystal tuffs in cases where they are banded as along highway 144 near the top of Chester township (sheet 1). These rhyolitic rocks are hard, dense and weathered to a light cream colour. They commonly contain small rounded clear quartz grains and feldspar laths in a fine quartzo-feldspathic matrix. The rhyolites consist of 75% or greater quartz and feldspar accompanied by fine muscovite (sericite), chlorite, epidote, biotite and minor sphene.

GREYWACKE, CONGLOMERATE

Grey to greyish green well-bedded sedimentary rocks along the northern boundary of the property and around Benneweiss Lake are largely greywacke in composition. They are fine to medium grained and occur in beds an inch or so thick. Dips are generally near—vertical. Beds of conglomerate containing mafic to felsic pebbles are present in places.

The greywacke is poorly sorted with grain size ranging from ½ mm to 2 mm or more. They consist of angular to rounded clasts of quartz and feldspar in a matrix of finer quartz, chlorite, biotite, muscovite and sericite. The flaky minerals are generally aligned and impart a weak to strong schistosity to the rock.

GABBRO, DIORITE

The central part of Chester township (sheets 1 and 2) is underlain by a northerly trending complex of gabbro, diorite and quartz diorite intruded on the east by quartz-eye granodiorite and on the west by alaskite. A similar mafic complex occurs in the northeast part of Chester township and west and south of the west end of Benneweiss Lake. Various phases of this unit grade into one another without well-defined contacts. They range from medium to dark green in colour, from fine to coarse grained in texture and from massive to schistose in structure. Some of the fine-grained chloritic phases resemble basalts but no well-defined flow features were noted; these basaltic looking phases are distributed in an irregular manner throughout the gabbroic phases and it is assumed that they are a part of the gabbro-diorite complex.

The various phases of this complex, in part at least, are a result of alteration. The least altered material is typically gabbroic and consists dominantly of plagioclase tablets with coarse interstitial hornblende and a little quartz. The plagioclase is altered to fine epidote and the amphibole displays minor alteration to chlorite. In more altered phases the amphibole is almost completely altered to chlorite. In phases which appear chloritic the plagioclase becomes altered to epidote and chlorite and the matrix contains abundant biotite with a few remnants of chlorite from which it was derived. The more altered phases may also contain greater than 10% quartz occurring as introduced grains containing fluid (aqueous) inclusions.

Quartz-bearing phases are secondary quartz diorites. The alteration of these mafic rocks has been sequential and progressive in the sense that amphibole first alters to chlorite and plagioclase to epidote. This involves the addition of water. Chlorite then alters to biotite by the addition of further water and potash. This sequence is of some practical interest as it is the same sequence of alteration that occurs in wallrocks around quartz veins although in this case the alteration is even more intense, involving the production of chlorite and biotite schists and the introduction of numerous large opalescent quartz eyes into the mafic host rock.

The mafic rocks discussed above are important hosts to many of the gold-bearing quartz veins on the property, particularly near the contacts of the gabbro-diorite with the quartz-eye granodiorite and alaskite. Because of this it is important to outline the distribution of the gabbro-diorite, particularly in areas where rock outcrops are scarce. The filtered aeromagnetic data (residual magnetism) has been useful in this regard.

HYBRID FELSITE

Some granitized, feldspathized and silicified rocks of mixed character have been placed in this category, hybrid felsite. They occur mainly in two localities, one around the #20 vein system (sheet 1) and east of Three Ducks Lake (sheet 2). The hybrid rocks are of two types. One type consists of fine grained granitic material containing irregular inclusions of gabbrodiorite or "ghost-like" altered remnants of these rocks. The other type, which forms the bulk of this unit is a dense hard massive felsite, also containing "ghost-like" darker areas. The felsite consists of fine granular quartz and feldspar with variable amounts of biotite and muscovite. In places minute phenocrysts of plagioclase are visible on the weathered surface. The felsite represents silicified and feldspathized pre-existing rock or a fine border phase of the granitic rocks. Somewhat similar material, regarded as a border phase of the alaskite, occurs around the #3 vein system. Some of the rhyolite south of Benneweiss Lake is also similar to the felsite and it is possible that some or all of that rhyolite may indeed be felsite.

ALASKITE

The granitic rocks exposed in the western part of the property are fine- to medium-grained grey massive alaskite generally containing less than 10% dark minerals. The alaskite intrudes the sedimentary rocks and the gabbro-diorite complex. The alaskite consists of about 45% each of plagio-clase and quartz, about 1 mm in size and about 10% fine sericite, biotite, chlorite, epidote and carbonate. In places fine disseminated pyrite is present even though the rock is massive.

QUARTZ-EYE GRANODIORITE

The granitoid rocks intruding the gabbro-diorite complex in the area of Mesomikenda Lake and eastward are coarse grained granodiorites containing prominent bluish opalescent quartz eyes. The rock varies from massive to faintly foliated and from light grey to orange-pink on the weathered surface.

The granodiorite consists of about equal amounts of plagioclase. orthoclase and quartz, totalling from 75% to 90% of the rock and in grains from 3 to 4 mm across. The remaining minerals are interstitial biotite, chlorite, muscovite, epidote and accessory sphene, apatite, and zircon. The most striking feature of the rock is the presence of prominent bluish quartz eyes consisting of either single large grains or round mosaics of smaller individual grains. Generally the "eyes" are unstrained and are crowded with abundant minute fluid (aqueous) inclusions. It is these fluid inclusions which impart the bluish opalescence to the quartz eyes. The fact that the quartz eyes are unstrained and retain their fluid inclusions even though the matrix in which they occur may be foliated, even schistose in places, indicates that they are late metasomatic and hydrothermal. This conclusion and the fact that many of the quartz veins in gabbro-diorite occur in proximity to the quartz-eye granodiorite and, furthermore, the observation that secondary opalescent quartz eyes occur in the gabbrodiorite wallrocks hosting veins suggests a direct genetic association between the granodiorite and gold mineralization. Certainly the abundance of aqueous inclusions in the quartz eyes attests to widespread hydrothermal activity.

DIABASE, TRAP

Medium-grained brown weathering diabase dykes 100 feet or more wide intrude all other rocks on the property. Narrower dykes, a foot or so wide, and the borders of the coarser diabase are fine-grained trap. Most of the diabase dykes trend in a northerly direction but a few trend slightly south of east.

The diabase is generally fresh undeformed and unaltered. In places, however, a foliation is developed. Fresh massive diabase consists of about 45% each of randomly oriented plagioclase laths and interstitial augite. Up to 10% magnetite is present, accounting for the pronounced magnetism of the diabase.

STRUCTURE

Most of the rocks are intrusive and fold structures, therefore, are generally obscure. Within the sedimentary sequence there is some suggestion that tops face north and that the greywacke-conglomerate stratigraphically overlie the rhyolitic rocks. The sedimentary rocks, both at Benneweiss Lake and in the northern part of the property, may occupy synclinal troughs.

The most prominent and significant structural elements are fracture and shear systems, some of which are occupied by diabase dykes. Although most of the dykes trend in a NNW or northerly direction, a second important trend is at 110°. Two parallel dykes trend in this direction across the property (sheet 1) and most of the gold-bearing quartz veins occur in fractures and shears striking at about 110°. There is clear evidence that repeated or prolonged deformation, pre-dating, contemporary with, and post-dating vein formation, has taken place along this direction. This temporally persistent fracturing, therefore, appears to have been important in "ground preparation" and in providing access for mineralizing fluids. Some of the evidence for repeated deformation is summarized below.

The earliest recognizable deformation involved development of schistosity in parts of the gabbro diorite and the conversion of amphibole to chlorite. This foliation trends at about 110° and may have preceded or have been contemporary with intrusion of granitoid rocks. Some phases of the quartz-eye granodiorite display a pervasive foliation at 110° that is characterized by parallel films of chlorite and biotite. This foliation must have developed after the granodiorite was solidified, but prior to complete consolidation, as the blue quartz eyes are unstrained and have not lost their fluid inclusions. Fracturing, with 110° trend, continued after consolidation of both the quartz-eye granodiorite and alaskite, and both these rock types contain gold-bearing quartz veins. In some veins the quartz is banded, indicative of re-opening and healing of fractures during vein formation. Finally the diabase striking at 110° north of the #10 vein, and which is younger than the granitoid rocks and the veins, is itself foliated in places.

Northerly trending fractures and diabases are also structurally important. Some of these clearly cut the quartz veins. However, whether or not they have displaced the veins is uncertain. The veins between Mill Pond and Arethusa Lake appear to pass uninterrupted across northerly trending diabase dykes.

On the other hand, the NNW fracture occupied by Mesomikenda Lake appears to have displaced the sedimentary rocks in a left hand manner (east side northward). Displacement of the gabbro-granodiorite contacts may be in a similar left-handed sense. The two diabase dykes at 110° north of #10 vein, however, appear to have been displaced in a right-handed manner. This apparent displacement of the 110° dykes may not be real. It is equally likely that the 110° dykes and the NNW fractures along Mesomikenda Lake are conjugate and formed at the same time.

Pesides the two shear-fracture directions discussed above some more easterly trending shears are present, for example through the lake at the top of sheet 2. Parts of the #20 vein system also trend about 090°. Other shear-fractures strike about 045°, for example north of the #10 vein.

GOLD DEPOSITS

Of the various fracture directions, at 045°, 090°, 110° and 165°, those trending at 110° are most commonly occupied by quartz veins. Although many of the veins first discovered in the early 1930's appeared to be individual isolated occurrences, it is now apparent that in many cases they occupy long persistent shear-fracture zones, the entire length of which have potential for further development. The shear-fractures tend to pinch and swell along strike and quartz veins occur in the wider portions of these zones. The rectangular area between Mill Pond, Arethusa Lake and the #10 vein is typical of this type of occurrence, where a series of parallel fractures at 110° contain numerous gold-bearing veins.

Although the importance of the 110° shear-fractures is obvious, those trending at 045°, 090° and 165° may also have potential for gold mineralization. The long east-west shear east of Southcamp Bay (top of sheet 2) contains quartz veins in at least two places. Similarly the #20 vein system consists of veins and shears at 090° as well as 110°.

Although the shear-fracture zones cut through the granodiorite and alaskite, as well as the gabbro-diorite, most of the known veins occur within the gabbro-diorite or in the peripheral parts of the major granodiorite and alaskite intrusions. Dyke-like protuberances or "noses" of granodiorite and alaskite invading gabbro-diorite are present at several vein locations.

The veins and fractures pinch and swell and may range from a few inches to 4 feet or more in width. Generally they are wider in the gabbro-diorite and, on passing into the granitoid rocks or hybrid felsite, they tend to narrow down abruptly or break up into numerous small separate shear-fractures an inch or so wide. The physical difference between gabbro-diorite, on the one hand, and the granitoid rocks, on the other, may be responsible for the different character of the shears and fractures in the two rock types. In addition, it is probable that deformation was initiated earlier and persisted longer in the gabbro-diorite than in the granodiorite and alaskite.

The types and degree of alteration associated with mineralization differ around veins in the gabbro-diorite and granitic rocks. In the granitic rocks the only notable alteration is silicification and development of muscovite or sericite adjacent to the veins. In gabbroic rocks the alteration is much more widespread, zoned and of different mineralogy.

Adjacent to the veins in gabbro and diorite, biotite is developed at the expense of chlorite and in some instances lenses of biotite schist ranging from a few inches to several feet wide are present. This biotite alteration resembles lamprophyre in places. Further outward from the vein both biotite and chlorite are present. In some cases, for example at #20 vein, prominent bluish opalescent quartz eyes have been introduced into the altered gabbro-diorite of the outer alteration zone. These are particularly abundant in altered gabbro adjacent to the quartz-eye granodiorite that intrudes the gabbro-diorite.

The alteration is clearly sequential, with amphibole first altered to chlorite and chlorite subsequently altered to biotite. The process involves addition of water and potash to the gabbro. This, along with the introduction of opalescent quartz eyes into the gabbro suggests a genetic link between the alteration and the blue quartz-eye granodiorite, which is a probable source of water (as evidenced by the abundant fluid inclusions in the quartz eyes) potash and silica. This suggestion is in accord with the presence of granodiorite in proximity to many veins. In any event, as documented below, there appears to be a positive correlation between the amount and intensity of alteration and the gold content of the vein zones.

The veins consist of white to grey glassy to sugary quartz containing various sulfides and metallic oxides. Chalcopyrite and pyrite are the dominant sulfides in all veins. Some also contain pyrrhotite, sphalerite and molybdenite. Rutile and ilmenite are generally present in minor amounts. A striking component of some veins is manganese, occurring as colloform masses of pyrolusite, psilomelane and manganite (?) along with probable alabandite (MmS). The manganese minerals are usually closely associated with chalcopyrite. These minerals weather readily and are undoubtedly the source of pronounced Mn anomalies detected in the geochemical test survey (separate report).

Sulfides are not restricted to the vein material but are also present in stringers and disseminations in the schist and altered wall-rocks around the veins. In altered rock containing introduced quartz eyes the sulfides commonly occur as rims around the quartz, another indication of a genetic link between mineralization and the quartz-eye granodiorite. Although there is a clear association between gold and sulfides, especially chalcopyrite, the gold does not occur directly with or bound up in the sulfide minerals. Instead, visible grains of free gold occur in the quartz gangue. Tellurides have also been reported from veins in this area but none were observed in the polished sections examined. Gold is not confined to the quartz but also occurs in the altered rock around the vein. Consequently where shearing and alteration are widespread the width of the gold-bearing zone may be much greater than the width of individual narrow quartz stringers.

A relationship between quartz veins and the quartz-eye granodiorite has been discussed. However, mineralization also seems to be associated with the alaskite that occurs west of the main gabbro-diorite bodies. Examples are the #3 vein, and parts of the west end of the #1 vein system (and the adjacent #17 zone). These occur in part or entirely in alaskite and the mineralization consists not only of quartz vein, but in large part as heavily disseminated pyrite, chalcopyrite and pyrrhotite in alaskite, much like a porphyry copper. Whether the alaskite was the source of the mineralization or whether it serves merely as its host is not clear. The problem is compounded by lack of any positive data indicating the age relationship between the alaskite and the quartz-eye granodiorite; it is assumed that the granodiorite is the younger, in which case it may also be the source of mineralization in the alaskite.

Some, but not all, of the veins are described below, mainly to illustrate the main features associated with the gold occurrences and which may serve as an aid in continued exploration of the property. Assay results quoted are from previous Murgold reports or from current files on material collected by Norminex Limited and by Mr. H. Kenty of Murgold Resources.

DESCRIPTION OF VEINS

The Strathmore (#10) vein illustrates many of the general features common to gold mineralization on the property. The vein occupies a shear trending 110° and dipping north in gabbro-diorite and a "nose" of quartz-eye granodiorite intruding gabbro-diorite. A diabase dyke trending NNW cuts the vein east of the shaft. A small dyke of fine granodiorite (too small to show on the map) cuts the gabbro-diorite northwest of the shaft. The level plan (facing south) at 1" = 10' illustrates the zoned nature of the alteration in diorite in the west drift. The alteration ranges from biotite schist adjacent to the main shear through chlorite-carbonate schist to less altered but silicified diorite.

In the east drift the alteration is less intense and less extensive; the east drift lacks the intensely altered biotite schist but contains only chloritic schist. The main shear passes eastward into quartz-eye granodiorite which is little altered. This granodiorite is much coarser-grained than the narrow dyke on surface.

There is a distinct relationship between the extent and intensity of alteration and the gold content of the different parts of the vein zone. Average gold values in the west drift where alteration in diorite is intense (biotite schist) are 1.03 oz./ton (channel samples) or 0.62 oz./ton (bulk samples). By comparison average gold values in the east drift, where alteration is less intense (chloritic) and where granodiorite is the dominant host rock, are 0.19 oz./ton (channel) or 0.18 oz./ton (bulk). In detail the correlation between alteration is even more striking. In the west drift a 70-foot section of vein with biotite alteration averages 1.38 oz./ton (channel) compared to a 30-foot section with only chlorite alteration which averages 0.16 oz./ton. Similarly in the east drift a 55-foot section with chloritic alteration averages 0.28 oz./ton compared to a 30-foot section in granodiorite with little alteration which averages 0.01 oz./ton.

The #20 vein zone, discovered during the summer of 1981 consists of a number of branching shears trending in a general easterly direction and dipping both north and south along a length of about 520 feet and across a width of 100 to 250 feet (for the whole zone, not all of which is sheared or mineralized). The shears occur in altered diorite, in diorite-granodiorite breccia, and in silicified and feldspathized rock, hybrid felsite. The

diorite is intruded from the east by a dyke-like body of quartz-eye grano-diorite and from the west by alaskite. The diorite country rock hosting the shears is altered to chlorite and biotite and impregnated with prominent bluish quartz eyes. The shears range from a few inches to several feet in width and consist of biotite schist containing about 40% biotite, 25% muscovite and 15% quartz crowded with fluid inclusions. Individual quartz veins, 1" to 5" or 6" wide occupy the shears. The veins contain pyrite, chalcopyrite, pyrrhotite and manganese oxides. Assays from well-mineralized grab samples contain in excess of 2 oz./ton gold. Eastward the mineralized zone passes from altered diorite, through diorite-granodiorite breccia into fine dense hybrid felsite. The shears, alteration and mineralization appear to weaken eastwards.

The #3 vein north of Arethusa Lake occurs in a very fine-grained border phase of the alaskite that somewhat resembles the hybrid felsite at the eastern part of the #20 zone. The shear-fracture at the #3 occurrence varies from a few inches to about 4 feet in width and is occupied by a relatively narrow quartz vein containing pyrite and chalcopyrite along a length of about 300 feet. Alteration is not widespread nor intense and consists mainly of silica and muscovite that is confined to the shear itself. Although the eastern end of the vein and the fracture pinch out next to a diabase dyke, the strike extension of the #3 zone aligns with the #10 vein further east. Gold assays from the #3 vein range from 0.04 to more than 3 oz./ton.

North of and parallel to the #3 vein is the #2 vein occurring in altered gabbro-diorite. It has been exposed for about 600 feet east from the road leading to Arethusa Lake. The same zone appears to extend westward, perhaps not as a continuous zone, for 2400 feet to the south shore of Mill Pond.

The #1 vein (original Gomac vein discovered in the early 1930's) has been exposed intermittently by trenching and stripping for about 2500 feet in an ESE direction. The western part of this zone occurs in diorite intruded by alaskite. Parts of the diorite are altered to chloritic and biotitic schist. Grab samples from the western end of the vein range from trace to 2.4 oz./ton in gold. Also near the western part of this zone are lenses of alaskite containing abundant disseminated sulfides, both chalcopyrite and pyrite with pyrrhotite in places. One such occurrence, just south of the #1 vein and designated the #17 zone contains mainly pyrite where exposed.

The other, just north of the #1 vein carries abundant disseminated chalcopyrite.

The central part of the #1 system is located at the Gomac shaft and averages about 3 feet wide with gold values ranging from about 0.01 to 1.2 oz./ton. The eastward extension was uncovered by stripping in 1980. The mineralized zone there varies from a little more than a foot to 10 feet or more in width, although not entirely mineralized across this width. Sulfides present include pyrite, chalcopyrite, molybdenite, sphalerite, pyrrhotite along with visible gold. Assays run into the ounces per ton in selected grab samples.

The #15 "vein" consists of disseminated pyrite and chalcopyrite in quartz-eye granodiorite. The mineralized zone, unlike many of the veins on the property trends NE. Assays up to 0.4 oz./ton Au have been obtained from this zone.

The #16 vein lies close to and parallel to the major diabase dyke which trends at 110 across the property (sheet 1). It has been exposed for over 700 feet and consists of a shear and quartz stringers in silicified and feldspathized diorite. The mineralized zone contains pyrite and chalcopyrite. Gold assays up to 2.9 oz./ton have been obtained from grab samples. A large barren quartz-carbonate vein occurs just north of and parallel to the #16 zone. The occurrence, in close proximity to one another, of the #16 vein proper, the adjacent quartz carbonate vein and the large fracture occupied by the diabase a few hundred feet to the south suggest the possibility that this general area may be one of extensive fracturing suitable for the occurrence of further veins.

DISCUSSION OF CONTROLS ON MINERALIZATION

The vein systems discussed above all occur on that part of the property, covered by sheet 1 of the geological map, where most of the exploratory work has been done to date. The general conditions apparently favourable for mineralization in that part of the property are:

- 1. Suitable shear-fracture zones. Most of these trend about 110° but others, at 090° and 045° are also favourable.
- 2. Presence of host gabbro-diorite. Veins and mineralization also occur in the granitoid rocks but these generally do not support structures as strong as those in gabbro-diorite; nor have the granitoid rocks the alteration that appears to accompany better gold mineralization.
- 3. Presence of chlorite and, in particular, biotite alteration. The presence of secondary opalescent quartz eyes in diorite-gabbro may indicate proximity to mineralization as may areas of fine dense silicified and feldspathized rock, hybrid felsite.
- 4. Proximity to dyke-like masses, "noses" and protuberances of granodiorite or alaskite invading the gabbro-diorite.

According to the above guidelines the entire block of gabbro-diorite around Mill Pond, Arethusa Lake and the #10 vein merits intensive detailed exploration. It is apparent that at least 4 parallel fracture systems are present there and gold-bearing veins probably occur elsewhere along these structures or between them. Similar favourable conditions appear to be present in the southward continuation of the gabbro-diorite (sheet 2). NNW, ESE and E-W fractures are present in gabbro-diorite in that area. Some quartz veins and mineralization are also present.

The gabbro diorite occurring immediately west of Southcamp Bay (Mesomikenda Lake, Sheets 1 and 2) is not well exposed but in two places the long easterly fracture (top of sheet 2) contains quartz veins, one in the central gabbro-diorite and one in the gabbro-diorite near Southcamp Bay.

Quartz veins and shears are present in several places in gabbro-diorite east of Mesomikenda Lake. Along highway 144 (Southeast corner of sheet 1) two small shears occur just south of where the diabase crosses the highway.

The most northerly shear yielded trace in gold and the southerly one gave 0.01 oz./ton gold. Some 1200 feet further south along the highway a quartz vein occurring in altered diorite yielded 0.20 oz./ton gold. Still further south (sheet 2) a number of veins are exposed in road cuts along the highway east of Southcamp Bay. Grab samples contained trace amounts of gold. East of this locality and west of Robitaille Lake, at the eastern boundary of the property, a poorly exposed shear occurs in gabbro-diorite. Two grab samples yielded trace and two others yielded 0.01 oz./ton gold.

The area west and south of the west end of Benneweiss Lake is underlain in part by gabbroic rocks. Although no veins were encountered during mapping, chlorite—carbonate schists are developed in places and this area is considered favourable for the occurrence of quartz veins.

Little is known about the mineral potential of the greywacke suite sedimentary rocks on the property. These should be structurally favourable for vein occurrence. In places, as along the east shore of Mesomikenda Lake (top of sheet 1) they appear heavily sheared. The former Jerome gold mine in Osway township to the west occurred in similar sedimentary rocks. The greywacke-like sedimentary rocks along the north shore of Benneweiss Lake are heavily sheared in places and sulfides with quartz are present in several places. Although there is little positive indication of their potential the greywacke-conglomerate should be thoroughly prospected.

Large parts of the property are underlain by monotonous expanses of quartz-eye granodiorite and alaskite. Veins occur in the border phases of these rocks and in offshoots from the larger bodies intruding the gabbro-diorite. Within the interiors of the large alaskite and granodiorite intrusions there is little indication of gold mineralization. In places, foliated phases of both these granitoid types contain finely disseminated pyrite but little else.

In terms of lithology rock types may be rated from apparently most favourable to apparently least favourable as follows:

- 1. Gabbro-diorite
- 2. Border phases and offshoot of the granodiorite and alaskite invading gabbro-diorite
- 3. Greywacke-conglomerate
- 4. Interiors of alaskite and granodiorite batholiths.

Accordingly, there remain large areas of structurally and lithologically favourable rocks on the property which warrant further exploration.

September 24, 1981

Dag F. DAVIES, Ph.D



NORMINEX LIMITED

MINERAL EXPLORATION & RESEARCH

R. A. Cameron, Ph.D., P.Eng. J. F. Davies, Ph.D., P.Eng. R. E. Whitehead, Ph.D.

I, J. J. Pavies, vice-president, Norminex Limited of the City of Sudlury in the Province of Ontario, do herely contify:

- J. That I am a graduate in goology of the University of Lamitoin (P.Sc., M.Sc.) and the University of Toronto (Ph.D.).
- 2. What I have spent 35 years in the practise of goology.
- 3. That I am a Jellow of the Geological Association of Canada, a member of the Society of Mconomic Geologists, a member of the Canadian Institute of Timing and Metallurgy and a Jellow of the Timeralogical Association of Canada.
- h. That I did personally map and supervise geological investigations of the Eurgold Property and did propers the the geological maps and write the report dated September 1981.
- 5. That I have no interest, direct or indirect, in Eurgold Resources Inc.

Pated at Sudlamy Objections of Parch, 1983

J. l. Pavice, Ph.D.



THIN SECTION .

D-5

Fine light cream weathering quartzo-feldspathic with small quartz "eyes"

Quartz 60% Plagioclase 20% Muscovite-sericite 20% Carbonate

Fine to medium quartz and larger quartz eyes (1-2mm) Plagioclase remnants, now largely altered to epidote and sericite.

Name: Rhyolite

D-7

Medium-grained quartzo-feldspathic light cream weathering.

Quartz Fedlspar 70%

Muscovite 30%

Mosaic of fresh interlocking quartz and plagioclase with randomly oriented micas. Common large quartz "eyes" 1 to 2 mm.

Name: Rhyolite

D-6

Light cream weathering, grey fresh surface, fine quartzo-feldspathic. Power line, South Benneweiss

Quartz 60% Sericite 20% Biotite 10%

Epidote 5% Chlorite 5%

Fine mosaic of quartz grain and aligned flakes of Mica. No feldspar identified but may be present mixed with fine quartz. Several larger fresh unstrained quartz grains.

Name: Rhyolite

JF-66

Fine light cream quartzo-feldspathic. South of power line, south of Benneweiss.

Quartz 32% Sericite 32% Chlorite 32% Sphene 3%

Fine-grained mixture of quartz and aligned flakes of mica and chlorite. Some larger euhedral unstrained quartz.

Name: Rhyolite

RW-7

"Porphyritic" light cream weathering fine-grained

Quartz Plagioclase 90% Chlorite Nuscovite 10%

Fine interlocking mosaic of fresh clear quartz and feldspar. Many phenocrysts of feldspar up to 3 mm.

Name: Rhyolite porphyry

B-3

Fine-medium grey granular sedimentary West of Weeduck Lake

Quartz Feldspar Epidote Chlorite Biotite Sphene

Rounded clasts of altered feldspar < 1 mm in matrix of fine quartz, chlorite, biotite and epidote. Feldspars are altered to epidote. Some round areas consisting of mosaics of quartz and coarse epidote.

Name: Greywacke

D-13

Sedimentary rock East of Weeduck Lake

Feldspar Quartz Biotite Epidote Chlorite Carbonate

Irregular clasts of quartz and feldspar up to 1 mm in fine quartzo-feldspathic matrix with abundant aligned flakes of biotite and chlorite and irregular grains of epidote. Unsorted.

Name: Greywacke

Y-1

Grey-green foliated sedimentary, Hwy 144

Feldspar Quartz Chlorite Sericite Carbonate

Very irregular grain size, unsorted, from ‡ mm to 1 - 2 mm. Larger grains are angular to sub-rounded feldspar and quartz. Chlorite in moderately large flakes; large masses of extremely fine sericitic material.

Name: Greywacke

A-10

Gabbro, South of B.L. 150S

Plagioclase 50% Chlorite Amphibole 30 Biotite Quartz 10 Epidote Sericite

Coarse (>3mm) partly zoned plagioclase with fresh interstitial amphibole. Rock is massive.

Name: Gabbro

M-12

Gabbro, south of Strathmore

Plagioclase 45% Quartz 10
Amphibole 30 Chlorite 5
Epidote 10

Coarse highly epidotized plagioclase and interstitial amphibole and quartz. Some chlorite alteration of amphibole.

Name: Gabbro

gar Pa

F-105

Diorite south of Arethusa Lake

Plagioclase 50% Chlorite
Biotite 20% Epidote
Quartz 10% Sphene
Opaque

Coarse (>2 mm) highly epidotized and chloritized plagioclase. Interstitial biotite with some remnant chlorite. Interstitial quartz. Rock is massive but highly altered.

Name: Gabbro-Diorite

JF-100

"Diorite" between Three Ducks and Mesomikenda, South of $B_{\bullet}L_{\bullet}$ 150 S_{\bullet}

Plagioclase Horneblende Epidote Quartz Biotite Opaque

Large euhedral but altered (epidote) tablets of plagioclase with interstitial hornblende. Granular epidote in matrix. Quartz corrodes hornblende. Biotite closely associated with quartz and biotite contains remnants of hornblende. Quartz crowded with fluid inclusions.

Name: Diorite-gabbro

X-1

Quartz-eye granodiorite, Hwy 144

Coarse feldspars and quartz with irregular areas and stringers of fine granular quartz mixed with schistose chlorite and/or sericite and carbonate. Large quartz eyes are slightly strained but contain fluid inclusions.

Name: Quartz-eye granodiorite

JF-132

Massive quartz-eye granodiorite from Benneweiss Twp.

Plagioclase	30%	Epidote)
Orthoclase	30%	Biotite > 10%
Quartz	30%	Muscovite /

Coarse plagioclase and orthoclase (to 4 mm) and rounded quartz grains that are slightly strained but contain abundant fluid inclusions. Interstitial mica and granular epidote

Name: Quartz-eye granodiorite

D-46

Massive blue quartz-eye granodiorite

Quartz	40%	Chlorite)
Plagioclase orthoclase	40%	Sericite (8% Sphene (
Biotite	12%	Apatite \int

Interlocking quartz and feldspars which are partly sericitized. Quartz crowded with fluid inclusions. Biotite interstitial.

Name: Granodiorite

J-43

Rusty weathering quartz-eye granodiorite, east of Hwy 144

Quartz Plagioclase	35% 30%	Biotite Carbonate	3% 5%
Orthoclase	20%	Zircon	210
Muscovite	5%	Pyrite	

Coarse (to 4-5 mm) interlocking quartz and feldspar Quartz slightly strained and packed with fluid inclusion. Biotite and muscovite in interstitial foliated masses.

Name: Quartz-eye granodiorite

D-11

Fine - med. granular granitic, east of Weeduck Lake

Quartz 45% Sericite Plagioclase (albite-oligoclase) 40% Biotite Carbonate 15% Epidote Chlorite

Round interlocking quartz grains and plagioclase laths average 1 mm. Feldspars are sericitized. Minor interstitial biotite chlorite, sericite and carbonate. Massive granular

Name: Alaskite granite

D-12

Fine-grained light grey granitic rock North of Weeduck

Quartz 45% Sphene Plagioclase 45% Sericite Chlorite Biotite

Mass of interlocking quartz and slightly sericitized plagioclase. Plagioclase is part corroded by granular quartz. Grain size generally ≤ 1 mm. Massive.

Name: Alaskite

A-34

Fine-medium granitic

Plagioclase 47% Quartz 45%

Biotite
Sericite 8%
Pyrite

Fine (<1 mm) interlocking quartz and feldspar; minor interstitial biotite. Granular massive

Name: Alaskite

JF-12

Fresh medium-grained diabase

Plagioclase 45% Augite 45%

Magnetite 8% Quartz 2%

Randomly oriented fresh laths of plagioclase with interstitial augite. Prominent magnetite grains. Minerals fresh and unaltered.

Name: Diabase.

THIN SECTIONS FOR #10 VEIN STRATHMORE DEPOSIT

Four sections from underground characterize the rock types and alteration associated with the gold-bearing quartz vein.

J.F.-1

Altered diorite-gabbro

Plagioclase	40%	Carbonate 10%
Quartz	25%	Sericite
Chlorite	20%	Opaque (sulfide)

Highly altered plagioclase tablets replaced by chlorite and lesser carbonate and sericite. Quartz occurs in two modes: 1) as large grains with rows of secondary fluid inclusions and 2) as mosaics of polygonal or lamellar grains, commonly intergrown with chlorite. The coarse quartz replaces plagioclase in part. Carbonate is late and in places forms veinlets cutting all other minerals. Quartz and carbonate are metasomatic and introduced.

Name: Altered Gabbro-diorite

J.F.-4

Silicified gabbroic

Plagioclase	40%	Epidote 5%
Quartz	30%	Carbonate
Chlorite	15%	Zircon
Amphibole	5%	Sulfide

Plagioclase corroded by later quartz. Plagioclase also alters to fine granular epidote (saussurite)
Chlorite is remnant after amphibole

Name: Silicified Gabbro-diorite

J.F.-2

Biotite schist

Green biotite
White muscovite
Green chlorite

65%

Quartz 25% Rutile-leucoxene Sulfide

Set in a fine highly schistose matrix of micas and chlorite are fresh undeformed quartz grains which abut against the micas. The micas are alteration products of chlorite. The quartz contains abundant small fluid inclusions. Biotite schist containing introduced quartz.

Name: Biotite schist

J.F.-3

Blue quartz-eye granodiorite

Plagioclase 50% Orthoclase 50% Quartz 35% Sericite

Chlorite Carbonate Zircon Pyrite

Coarse-grained (3-4 mm) feldspars and quartz.

Feldspars are sericitized and difficult to determine composition. Untwinned feldspars probably orthoclase. Quartz occurs in mosaics of coarse grains exhibiting only slight deformation extinction. Quartz contains abundant fluid inclusions. Some feldspars are broken and cut by vein-like masses of chlorite and/or carbonate. Quartz "eyes" replace plagioclase.

Name: Granodiorite

THIN AND POLISHED SECTIONS #20 VEIN

These sections reveal some important relationships between alteration and mineralization.

Section #20-T-l is a dark grey coarse dioritic rock similar to the surrounding diorite-gabbro mass, but containing prominent bluish white quartz eyes.

Sections #20-T-2a to #20-T-2c characterize a biotite alteration envelope around a quartz-sulfide stringer. #20-T-5 is biotite schist, an intense alteration product along shears.

#20-T-1

Dark grey coarse diorite with bluish white quartz eyes

Plagioclase	40%	Carbonate	5%
Quartz	30%	Epidote -	minor
Chlorite	12%	Sulfides	2%
Biotite	1 <i>2</i> %		

Subhedral plagioclase 1-2 mm in various stages of replacement by chlorite and quartz. Quartz, > 2 mm in rounded "eyes" crowded with fluid inclusions. Carbonate medium-grained interstitial. Chlorite randomly oriented and distributed throughout. Biotite almost invariably occurs as alteration of chlorite. The sulfides are intimately associated with biotite and chlorite.

 K^{\dagger} metasomatism of chlorite and metasomatic quartz eyes.

20-T-2a

Fine dark schistose adjacent to quartz-sulfide stringer

Quartz	50%	Carbonate	5%
Plagioclase		Epidote	8%
Biotite Chlorite	20% 10%	Sulfides	8%

Aligned foliated flakes of biotite and chlorite in a fine - medium matrix of quartz and plagioclase grains, most of which are elongated. Some larger quartz eyes are slightly strained but generally younger than the biotite and chlorite against which they abut. Biotite exceeds chlorite and is a product of chlorite alteration. Sulfides intimately associated with biotite.

Name: Biotite schist

20-T-2B

Section is 1.5" from veinlet and marks the outer edge of the biotite envelope which is in contact with a coarse grained lighter material.

Fine dark biotitic		Coerser light	
Plagioclase Quartz	20% 30%	Plagioclase Quartz	50% 1 2%
Chlorite	20%	Chlorite	12%
Biotite	20%	B iotite	1.2%
Carbonate	5%	Carbonate	10%
S ulfide	7%	S ulfide	5%
Epidote	Tr	Ep idote	Tr
Sphene	Tr	Sphene	${ t Tr}$

Biotite envelope is similar to #20-T-2a. Coarser light is similar to #20-T-2c. Sharp contrast in grain size and fabric. The schist is foliated whereas the lighter phase (diorite) has randomly oriented larger plagioclase and quartz.

20-T-2c

Medium-grained grey massive diorite, 4" from quartz-sulfide stringer and 1" from outer edge of biotite alteration envelope.

Plagioclase	50%	Biotite	5%
Quartz	20%	Carbonate	5%
Chlorite	15%	Sulfides	5%

Randomly oriented plagioclase 1 - 2 mm with interstitial quartz, carbonate and chlorite. Fine chlorite also replaces some plagioclase. Quartz grains contain fluid inclusions and are metasomatic. Some chlorite beginning to alter to biotite. Also some discrete flakes of biotite. Note that compared to the biotite schist envelope (#20-T-2a) this rock contains less total chlorite plus biotite and of these two minerals chlorite far exceeds biotite. The total plagioclase plus quartz is also somewhat greater in #20-T-2c (diorite).

Name: Slightly altered diorite

20-T- 5

Biotite schist, #20 vein, occurs as narrow zones along shear in altered diorite- abbro.

Biotite	40%	Epidote	8%
Quartz	15%	Sulfides	8%
Muscovite	25%	Chlorite	4%

Fine foliated mass of biotite and muscovite with a few small flakes of chlorite. Dotted throughout are small anhedral unstrained quartz grains and subhedral epidote. Both the micas and quartz are crowded with fluid inclusions.

Name: Biotite schist (alteration)

POLISHED SECTIONS

G1 -1

Gomac vein #1 - east of Mill Pond

Megascopic: Chalcopyrite and pyrite in fine sugary

quartz with small brownish metallic

grains, possibly pyrrhotite.

Chalcopyrite dominant.

Microscopic:

Pyrite

Chalcopyrite
Pyrrhotite
Molybdenite
Sphalerite

Ilmenite & rutile

Sulfides occur as separate and isolated grains or clusters of grains interstitial to gangue. Some of the chalcopyrite and molybdenite are intergrown, pyrrhotite and chalcopyrite also occur as intergrowths. Ilmenite and rutile are invariably intergrown with one another. No native gold or tellurides were observed.

S-1

Strathmore underground, west drift, at face

Chalcopyrite

Pyrite

Pyrrhotite - minor

Sphalerite? - minor small grain

Gold

- 1 small grain

Chalcopyrite occurs as large irregular grains and as stringers of small grains cutting the gangue. Pyrite is cubic and in part replaced by chalcopyrite.

Gold occurs "free" in gangue and not with sulphides.

Strathmore underground, west drift at face.

Pyrite
Chalcopyrite
Pyrrhotite
Sphalerite - minor
Rutile

Pyrite cubic and skeletal. Chalcopyrite in irregular grains and with associated pyrrhotite. Minor sphalerite and rutile. No gold visible in section.

20-P-1

#20 vein, west pit

Pyrite
Chalcopyrite
Pyrrhotite
Alabandite
Pyrolucite and/or psilomelane

Section dominantly irregular grains of chalcopyrite with lesser cubic and skeletal pyrite and minor pyrrhotite.

Common alabandite (MnS) and colloform alteration products, dominantly psilomelane but some pyrolusite in oriented cleavage-like masses. Some replacement of Mn minerals by fine rims of pyrite.

20-P-2

#20 vein, west pit
"Mineralized" altered diorite containing opalescent
quartz "eyes"

Chalcopyrite
Pyrrhotite
Pyrite
Rutile
Ilmenite

Disseminated grains of chalcopyrite, some containing cores of pyrrhotite. Minor very fine pyrite grains. The sulfides commonly occur as rims around the opalescent "blue" quartz eyes.

20-P-3

#20 vein, west pit

Chalcopyrite
Alabandite
Pyrolucite
Psilomelane
Pyrrhotite
Rutile
Pyrite

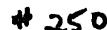
Section dominantly large grains of chalcopyrite intimately associated with large irregular and colloform masses of manganese minerals. Minor pyrrhotite and pyrite as discrete grains, but also considerable fine pyrite replacing alabandite.



Ministry of Natural Resources

Report of Work

(Geophysical, Geological, Geochemical and Expenditures)



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 Note: Only days credits calculated in the "Expenditures" section may be entered in the "Expend. Days Cr." columns.

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Geotechnical Report Approval

File 2. 4889

lining Lands Co	omments		
		· · · · · · · · · · · · · · · · · · ·	
6. Geophysics	•		
omments			
Approved o: Geology - E	Wish to see again with corrections Expenditures What Kusa	ha.	Signature
o: Geology - E			Signature
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Ontario

Ministry of Natural Resources

GEOPHYSICAL – GEOLOGICAL – GEOCHEMICAL TECHNICAL DATA STATEMENT

TO BE ATTACHED AS AN APPENDIX TO TECHNICAL REPORT FACTS SHOWN HERE NEED NOT BE REPEATED IN REPORT TECHNICAL REPORT MUST CONTAIN INTERPRETATION, CONCLUSIONS ETC.

Type of Survey(s)		
Township or Area Chief	he Jacon ships	MANANC OF A MAC TO AMEDICAN
Claim Holder(s)	de Jacon shegar	MINING CLAIMS TRAVERSED List numerically
	Harios en'Il Carrieros. Leg Suddency Will. Minecutting to office)	(prefix) (number)
Total Miles of Line Cut	41.6	
SPECIAL PROVISIONS CREDITS REQUESTED	DAYS per claim	
ENTER 40 days (includes line cutting) for first survey.	ElectromagneticMagnetometerRadiometric	
ENTER 20 days for each additional survey using same grid.	Other Geological Geochemical	
AIRBORNE CREDITS (Special provise MagnetometerElectromagn (enter d		
DATE: THE Y DE SE SIGNA	Author of Report or Agent	
Res. GeolQualif	ications	
Previous Surveys File No. Type Date	Claim Holder	
	RECEIVED	
	્રાંN·~-3-1982·····	
	MINING LANDS SECTION	
	A Sure of the second se	TOTAL CLAIMS

GEOPHYSICAL TECHNICAL DATA

GROUND SURVEYS - If more than one survey, specify data for each type of survey

Number of Stations		Number o	f Readings	
Station interval		Line spaci	ng	
Profile scale			·····	
Contour interval				
Instrument				
	onstant			
•	method			
	in interval (hours)			
	on and value			
	m and value			
Instrument				
Coil configuration _				
Coil separation				
Accuracy				
Method:	☐ Fixed transmitter	☐ Shoot back	☐ In line	☐ Parallel line
Frequency		(specify V.L.F. station)		
Parameters measure	·d	· ·		
		•		
	A STATE OF THE STA			
	4.4.111			
Corrections made_			AND THE STREET S	
	nd location			
Instrument				
Method	Domain	☐ Fr	equency Domain	
Parameters - On ti	me	Fr	equency	
Off ti	me	Ra	inge	
- Delay	time			
Integ	ration time	, .		
Power		. 1		
Electrode array				
Electrode spacing _				
Type of electrode				

INDUCED POLARIZATION

SELF POTENTIAL	
Instrument	Range
Survey Method	
RADIOMETRIC	
Instrument	
Values measured	
Energy windows (levels)	
Height of instrument	Background Count
Size of detector	
Overburden	
	(type, depth — include outcrop map)
OTHERS (SEISMIC, DRILL WELL	LOGGING ETC.)
•	,
Instrument	
Accuracy	
Parameters measured	
Additional information (for understa	nding results)
AIRBORNE SURVEYS	
• 11	
Instrument(s)	
, .	(specify for each type of survey)
Accuracy	(specify for each type of survey)
Aircraft used	
	method
	I in a Specime
	Line Spacing
Miles flown over total area.	Over claims only

GEOCHEMICAL SURVEY - PROCEDURE RECORD

Numbers of claims from which samples taken					
-					
Total Number of Samples	ANALYTICA	AL METHODS	<u>S</u>		
Type of Sample(Nature of Material)	Values expressed in:	per cent			
Average Sample Weight		p. p. m.			
Method of Collection		p. p. b.			
	Cu, Pb, Zn, Ni, Co	, Ag, Mo,	As,-(circle)		
Soil Horizon Sampled	Others				
Horizon Development	Field Analysis (tests)		
Sample Depth	Extraction Method				
Terrain	Analytical Method		**************************************		
	Reagents Used				
Drainage Development	Field Laboratory Analysis	3			
Estimated Range of Overburden Thickness	No. (test				
	Extraction Method				
	Analytical Method				
	Reagents Used				
SAMPLE PREPARATION					
(Includes drying, screening, crushing, ashing)	Commercial Laboratory (•		
Mesh size of fraction used for analysis					
	Extraction Method				
	Analytical Method				
	Reagents Used				
General	General				
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1982 11 04

- Mr. William L. Good Mining Recorder, Timmins

File No. 2.4889

As you requested, enclosed are the geological survey maps (in duplicate) covering 295 mining claims in Chester, Benneweis and St. Louis Townships Nos P 471952 et al.

Before these plans may be examined for assessment work credits they must have the following information added:

- The location of all geological traverse lines must be indicated.
- 2. The Hining claim numbers and boundaries must be shown.
- 3. Each map should show a north direction and a key map.
- 4. Each map should be signed and dated by the author of the report Mr. R.A. Cameron.

As we have no record of Mr. Cameron having previously submitted a geological report it will be necessary for him to provide a statement of his qualifications including his education background and work history.

In the event that Mr. Cameron is no longer available, another qualified geologist may sign the maps provided he furnishes a declaration or affidavit to the effect that he accepts responsibility for the contents of the maps and report.

I understand that you will be contacting a representative of the holder of the mining claims with respect to clearing up this matter.

E.F. Anderson Director Land Management

Whitney Block, Room #6450 Queen's Park, Toronto Telephone: 416/965-1380

CHARLES L. McALPINE, C.A.

Suite 2308, 120 Adelaide Street West, Toronto, Canada M5H 1T1 Telephone Bus. (416) 363-8046 Res. (416) 483-9390 Telex 06523873

DELIVERED

March 28, 1983

Mining Land Section, Lands Management Branch, Ministry of Natural Resources, Room 6450, Whitney Block, Parliament Buildings, Toronto, Ontario M7A 1W3

Attention: Mr. Fred W. Matthews

Re: Murgold Resources Inc.

Dear Sir:

CLM/jac: Encls.

It is my understanding that the original Geological Report, prepared by Norminex Limited for Murgold's assessment work purposes, had certain deficiencies and was rejected.

I have now had Norminex revise their report and make the necessary corrections. Two copies of this revised report are enclosed herewith.

I trust that the revised report will be acceptable, and I should appreciate it if you could advise me when you have made your review.

Yours truly,

C. L. McAlpine

C. L. M. alp-o McAlpine Per you

RECEIVED

MAR 28 1983

MINING LANDS SECTION

10

#250

1983 07 28

2.4889

Mr. William L. Good Mining Recorder Ministry of Natural Resources 60 Wilson Avenue Timmins, Ontario P4N 2S7

Dear Sir:

RE: Geological Survey on Mining Claims P 19992 et al in the Townships of Chester, Benneweis and St. Louis

The Geological Survey assessment work credits as listed with my Notice of Intent dated July 7, 1983 have been approved as of the above date.

Please inform the recorded holder of these mining claims and so indicate on your records.

Yours very truly,

E.F. Anderson Director Land Management Branch

Whitney Block, Room 6450 Queen's Park Toronto, Ontario M7A 1W3 Phone: (416)965-1380

D. Kinvig:mc

cc: Murgold Resources Inc.
Suite 2308
120 Adelaide Street West
Toronto, Ontario
M5H 1T1

cc: Resident Geologist Timmins, Ontario Property Committee

to year



Technical Assessment Work Credits

File		
	2.	4889

1983 07 07

Recorded Holder MURGOLD RESOURCES INC.					
Township or Area	ND ST. LOUIS TOWNSHIPS				
CHESTER, DEMNEWETS	ND 31. E0013 TOWNSHIFS				
Type of survey and number of Assessment days credit per claim	Mining Claims Assessed				
Geophysical	P 19992 P 539404 to 14 inclusive 20009 539416 to 21 inclusive				
Electromagnetic da					
Magnetometer da	471953 to 58 inclusive 543827 473667 to 69 inclusive 543993 473672 to 95 inclusive 543995 - 96				
Radiometric da	s 473698 - 99 546981 to 7000 inclusive				
induced polarization da	475770 00 10 110103170				
Section 86 (\$8) da	507667 to 69 inclusive 549016 to 19 inclusive 515048 to 58 inclusive 549108 to 17 inclusive				
Geologicald0	515328 to 30 inclusive 549294 515335 - 36				
Geochemical da	528055 to 59 inclusive				
Man days [] Airborne [000002				
Special provision X Ground L	538523 to 25 inclusive 539105 to 29 inclusive				
	539136 to 55 inclusive 539181 to 83 inclusive				
Credits have been reduced because of pa coverage of claims.	539279 to 98 inclusive 539308				
Credits have been reduced because of correct to work dates and figures of applicant.	539310 to 12 inclusive 539314 to 17 inclusive 539319 to 22 inclusive 539324 to 28 inclusive				
77(16) Special credits under section 88 RNS Kfor the follow	ng mining claims				
20 days Geological P 473696 -97 473700 to 03 inclusive 473716 - 17 473743 to 46 inclusive 538087 538089	P 539309 539313 539323 543818 - 19 543994				
No credits have been allowed for the following min					
P 471952 P 538088 473092 539318 473102 539415 473670 - 71 546980 473709 549003 515059 549018	Insufficient technical data filed				



1983 07 07

Our file: 2.4889

Mr. William L. Good Mining Recorder Ministry of Natural Resources 60 Wilson Avenue Timmins, Ontario P4N 2S7

Dear Sir:

Enclosed are two copies of a Notice of Intent with statements listing a reduced rate of assessment work credits to be allowed for a technical survey. Please forward one copy to the recorded holder of the claims and retain the other. In approximately fifteen days from the above date, a final letter of approval of these credits will be sent to you. On receipt of the approval letter, you may then change the work entries on the claim record sheets.

For further information, if required, please contact Mr. F.W. Matthews at 416/965-1380.

Yours very truly,

Anderson

Director

Land Management Branch

Whitney Block, Room 6450 Queen's Park Toronto, Ontario M7A 1W3

Phone: 416/965-1316

D. Kinvig:mc

Murgold Resources Inc cc: Gogama, Ontario

Mr. G.H. Ferguson Mining & Lands Commissioner Toronto, Ontario



Notice of Intent for Technical Reports

1983 07 07

2.4889

An examination of your survey report indicates that the requirements of The Ontario Mining Act have not been fully met to warrant maximum assessment work credits. This notice is merely a warning that you will not be allowed the number of assessment work days credits that you expected and also that in approximately 15 days from the above date, the mining recorder will be authorized to change the entries on his record sheets to agree with the enclosed statement. Please note that until such time as the recorder actually changes the entry on the record sheet, the status of the claim remains unchanged.

If you are of the opinion that these changes by the mining recorder will jeopardize your claims, you may during the next fifteen days apply to the Mining and Lands Commissioner for an extension of time. Abstracts should be sent with your application.

If the reduced rate of credits does not jeopardize the status of the claims then you need not seek relief from the Mining and Lands Commissioner and this Notice of Intent may be disregarded.

If your survey was submitted and assessed under the "Special Provision-Performance and Coverage" method and you are of the opinion that a re-appraisal under the "Man-days" method would result in the approval of a greater number of days credit per claim, you may, within the said fifteen day period, submit assessment work breakdowns listing the employees names, addresses and the dates and hours they worked. The new work breakdowns should be submitted direct to the Lands Management Branch, Toronto. The report will be re-assessed and a new statement of credits based on actual days worked will be issued.

1982 07 05 2.4889

Mining Recorder
Ministry of Natural Resources
199 Larch Street
Sudbury, Ontario
P3E 5P9

Dear Sir:

We have received reports and maps for a Geological survey submitted under Special Provisions (credit for Performance and Coverage) on mining claims S 471952 et al in the Townships of Chester/Benneweis.

This material will be examined and assessed and a statement of assessment work credits will be issued.

Yours very truly,

E.F. Anderson Director Land Management Branch

Whitney Block, Room 6450 Queen's Park Toronto, Ontario M7A 1W3 Phone: 416/965-1316

- J. Skura/sc
- c.c. Norminex Limited Sudbury, Ontario Attn: R.A. Cameron
- c.c. Murgold Resources Inc Gogama, Ontario

Mongrey Veransen - Dice Marie ; June 1/22. It will a diction of Brown the 100 ching hands decles all. The Theed Totalhens. Men: Let the grant in grather across mount would your clos de Gener office last week (Geolege). and thoughty mint to youth) and und you Arthur, golden de lete die in remodern program Lind 195100 from the claim last of Cack Espart av there Claimen de met. Julang to the Wange of Company. The August . Gane, Truly. RECEIVED Haracof A Kersly Land Management Branch RECEIVED JUN 1 U 1982 de la Microton. MINING LANDS SECTION



NORMINEX LIMITED

MINERAL EXPLORATION & RESEARC

R. A. Cameron, Ph.D., P.Eng. J. F. Davies, Ph.D., P.Eng. R. E. Whitehead, Ph.D. had son it

1274 Hastings Crescent Sudbury, Ontario, Canada P3A 2R5

October 28, 1981

Murgold Resources Inc., C/o Mr. H. Kenty, Gogama, Ontario

Dear Harold:

The following list will provide you with the amount of time spent on the Murgold Gogama property.

John Fonji May 19 to Aug 31, 1981 /05
Bruce Miller May 19 to Aug 28, 1981 /02
Peter Jurenovskis May 19 to Aug 30, 1981 /04

R. Whitehead, R. Cameron and J. F. Davies: 174 58 days in field 3x58

J. F. Davies: 19 days map preparation and report writing

19 504×1= 3528 days

I hope this will be satisfactory.

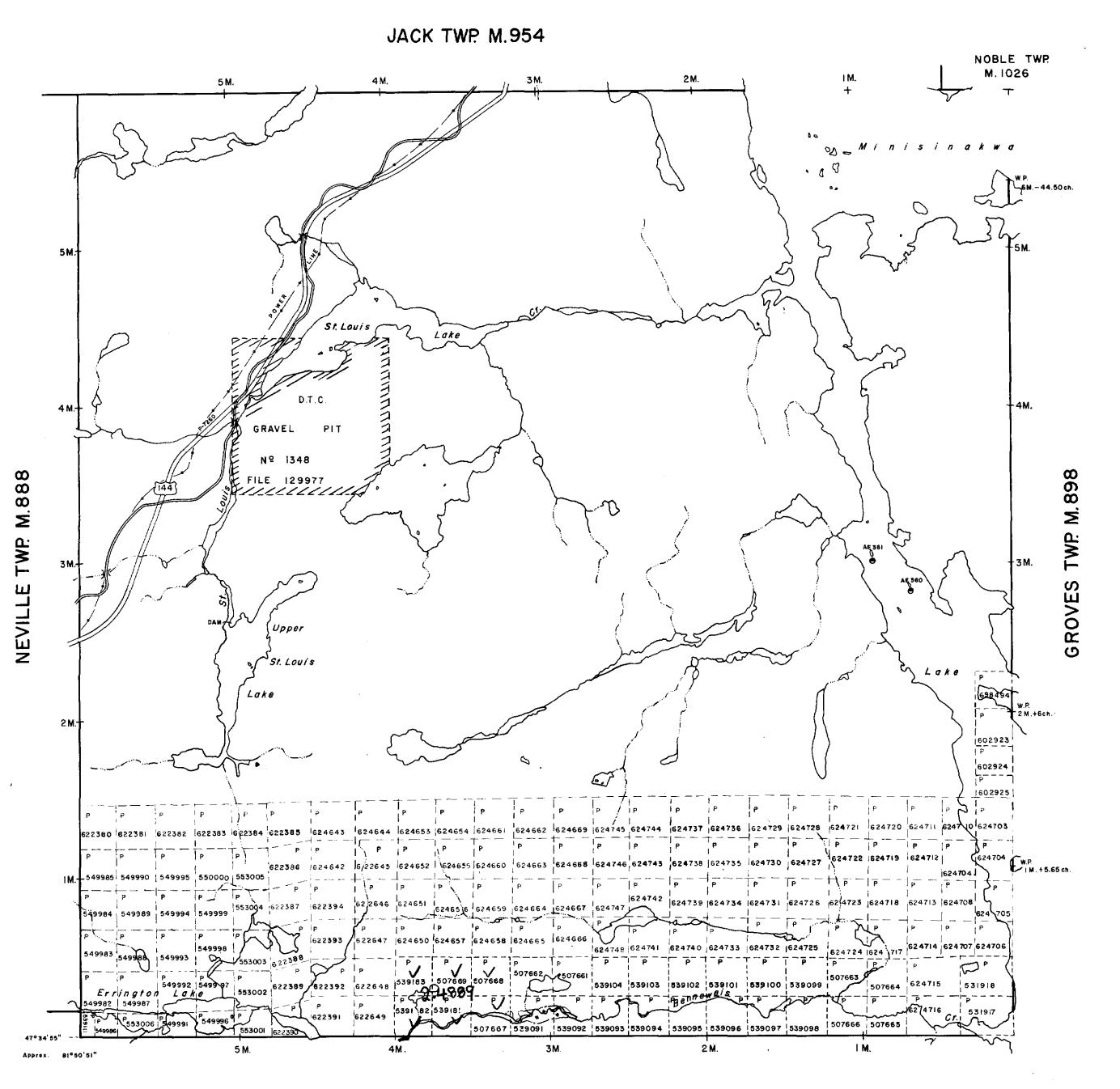
Yours truly,

R. Whitehead

RW/mw

The Trelaming list of Confliques. Jose man append an Suphypical Survey him tertlering Total days Robert Sent - may 18. - June ich-Richard Weight- Wing 19. June 30 de-Jac Marie - May 19. July 22 mil Leanard Cenit - June 7. July 22. 20 43 64 Total days . 172 days Touten movent work & theafting. sand From leason - May 19 - aug 4. 63. Minnia Palas. May 19 - augst. 66 Ken Rilley. 10my 22 June 25-3.5 Rong Tibles May 22 June 25. 35 July 9. 120 7. July 9. 120 15. July 1 July 15. 30 U. Thorwas 333 1. hatarest 10 24 301 days Talal Jatal- 172. 30/x7 = 2/07 days.

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423699	57505/12	539/38/11	539323
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BENNEWEIS TWP. M.658

NOTES

400' surface rights reservation along the shores of all lakes and rivers

DATE OF ISSUE

APR 19 1983

Ministry of Natural Resources
TORONTO

LEGEND

PATENTED LAND
PATENTED FOR SURFACE RIGHTS ONLY

LEASE

LICENSE OF OCCUPATION

CROWN LAND SALES

LOCATED LAND

CANCELLED

MINING RIGHTS ONLY

SURFACE RIGHTS ONLY

HIGHWAY & ROUTE NO.

ROADS

TRAILS

RAILWAYS

POWER LINES

MARSH OR MUSKEG

MINES

ST. LOUIS

TOWNSHIP OF

*used only with summer resort locations or when space is limited

DISTRICT OF SUDBURY

PORCUPINE MINING DIVISION

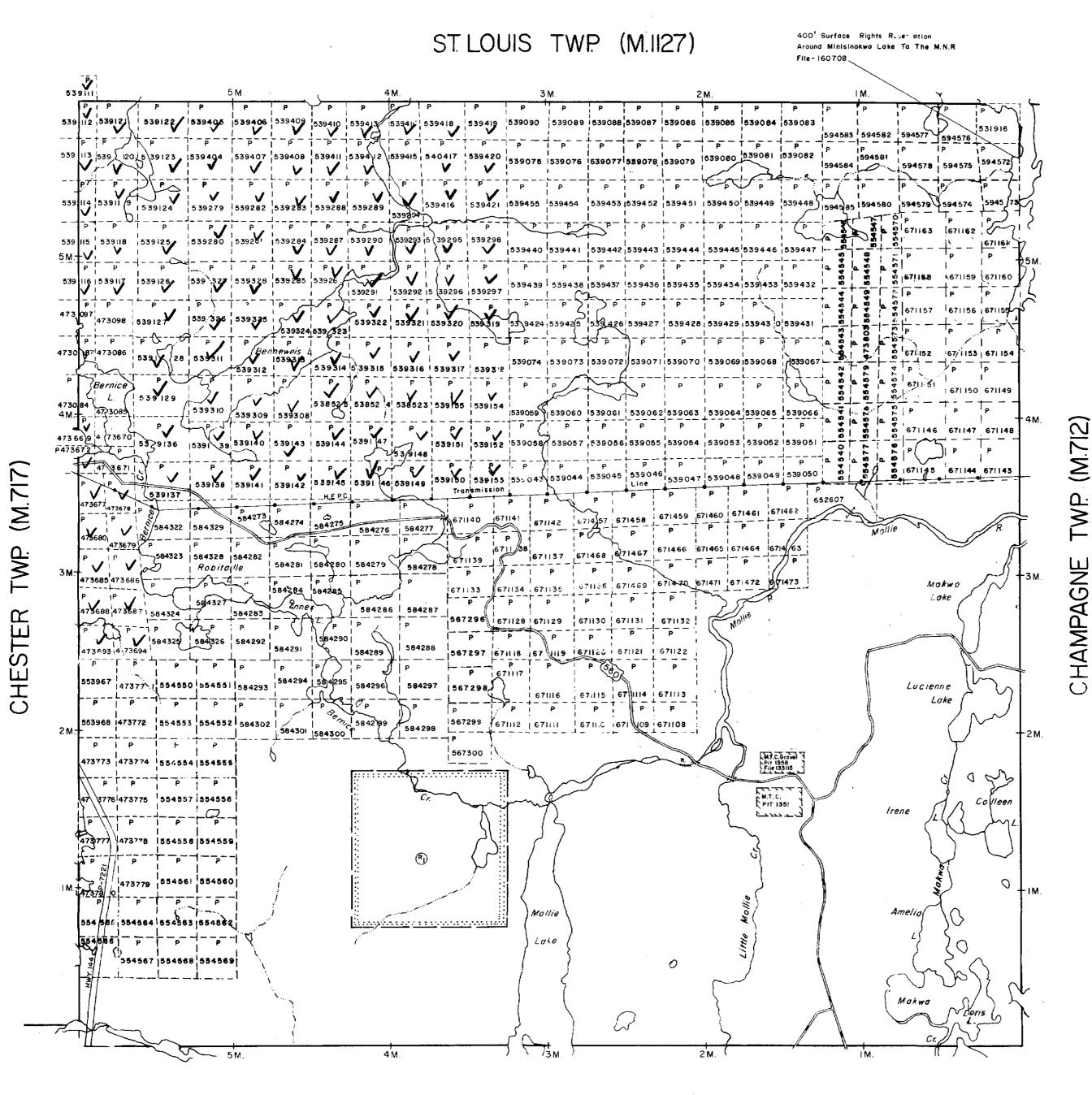
DATE JAN. '72.

SCALE: 1 INCH = 40 CHAINS (1/2 MILE)

ONTARIO

MINISTRY OF NATURAL RESOURCES

SURVEYS AND MAPPING BRANCH



THE TOWNSHIP OF

BENNEWEIS

DISTRICT OF SUDBURY

PORCUPINE

MINING DIVISION

SCALE: 1-INCH 40 CHAINS

LEGEND

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SURFACE RIGHTS ONLY	S.R.O.
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KING'S HIGHWAYS	—
RAILWAYS	
POWER LINES	**************************************
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MARSH OR MUSKEG	(* * * * * * * * * * * * * * * * * * *
MARSH OR MUSKEG MINES	*****

NOTES

400' Surface rights Reservation around all Lakes

Areas withdrawn from staking under Section 43 of the Mining Act (R.S.O. 1970) Order Nº File Disposition

(R)- NR.W.2/78 133115

19/10/78

DATE OF ISSUE

APR 19 1983

TORONTO

PLAN NO. M. 658

ONTARIO

MINISTRY OF NATURAL RESOURCES

SURVEYS AND MAPPING BRANCH

VROOMAN TWP (M.II73)

INVERGARRY TP. M.948

THE TOWNSHIP OF

CHESTER

DISTRICT OF SUDBURY

PORCUPINE MINING DIVISION

SCALE: 1-INCH=40 CHAINS

DISPOSITION OF CROWN LANDS
DISPUSITION OF CROWN LANDS
PATENT, SURFACE AND MINING RIGHTS
" , SURFACE RIGHTS ONLY 🖯
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LEASE, SURFACE AND MINING RIGHTS
" , SURFACE RIGHTS ONLY =
" , MINING RIGHTS ONLY =
LICENCE OF OCCUPATION _
ROADS
IMPROVED ROADS
KING'S HIGHWAYS -
RAILWAYS
POWER LINES
MARSH OR MUSKEG (+ +)
MAINEC

NOTES

400 Surface Rights Reservation along the shores of all lakes and rivers.

Flooding rights to 1200 contour reserved to Ont. Hydro ▼7543, Loc HY 34 file 10621

SAND & GRAVEL

(i) M. T. C. Pit No. 1349 (ii) "Gravel Pit No. 1649 (iii) "No. 1385

@ QUARRY PERMIT

CANCELLED

BENNEWEISS

AREAS WITHDRAWN FROM STAKING

S.R. - SURFACE RIGHTS

DESCRIPTION QRDER No. DATE DISPOSITION FILE

(A) SEC. 45 W.77/80 19/2/80 S.R. 171509

(B) 25/8/77 S.R. 183096

DATE OF ISSUE

APR 18 1983

Ministry of Natural Resources TORONTO

PLAN NO.- M 717

ONTARIO

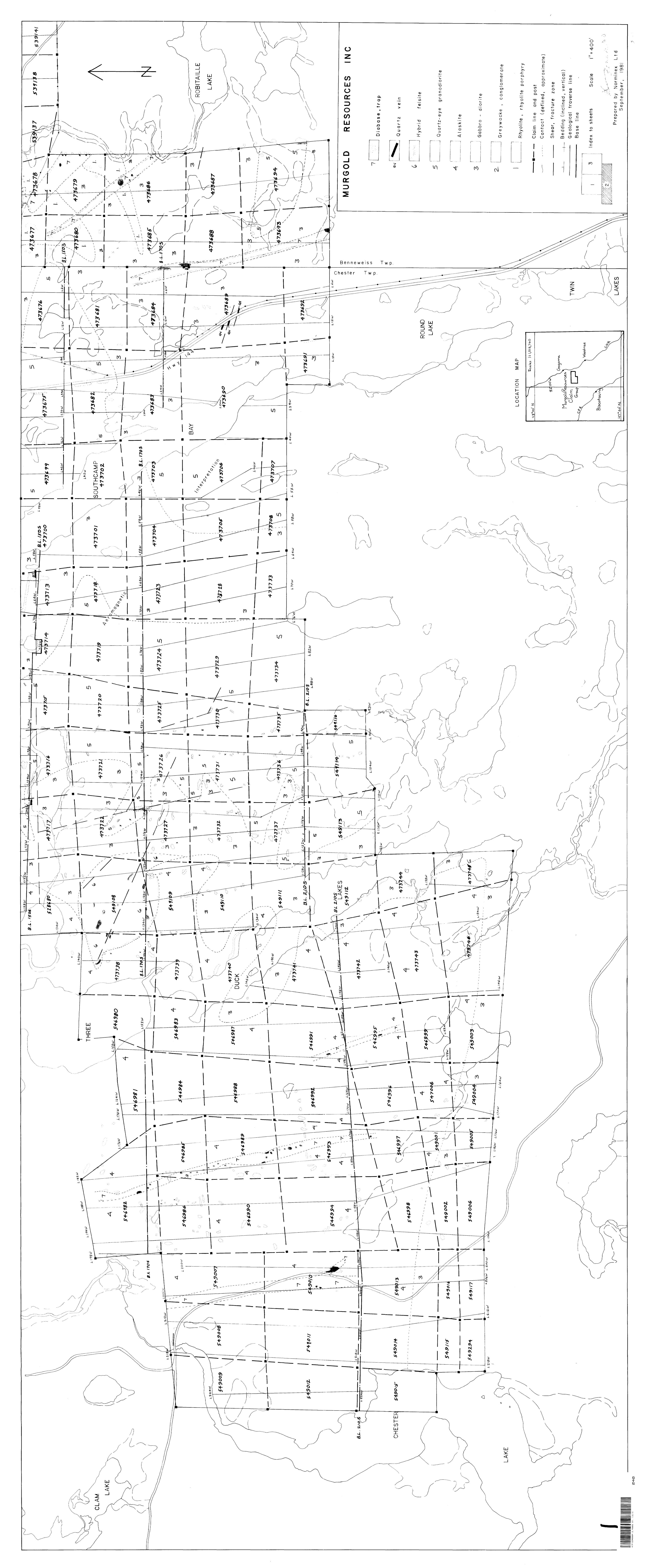
MINISTRY OF NATURAL RESOURCES

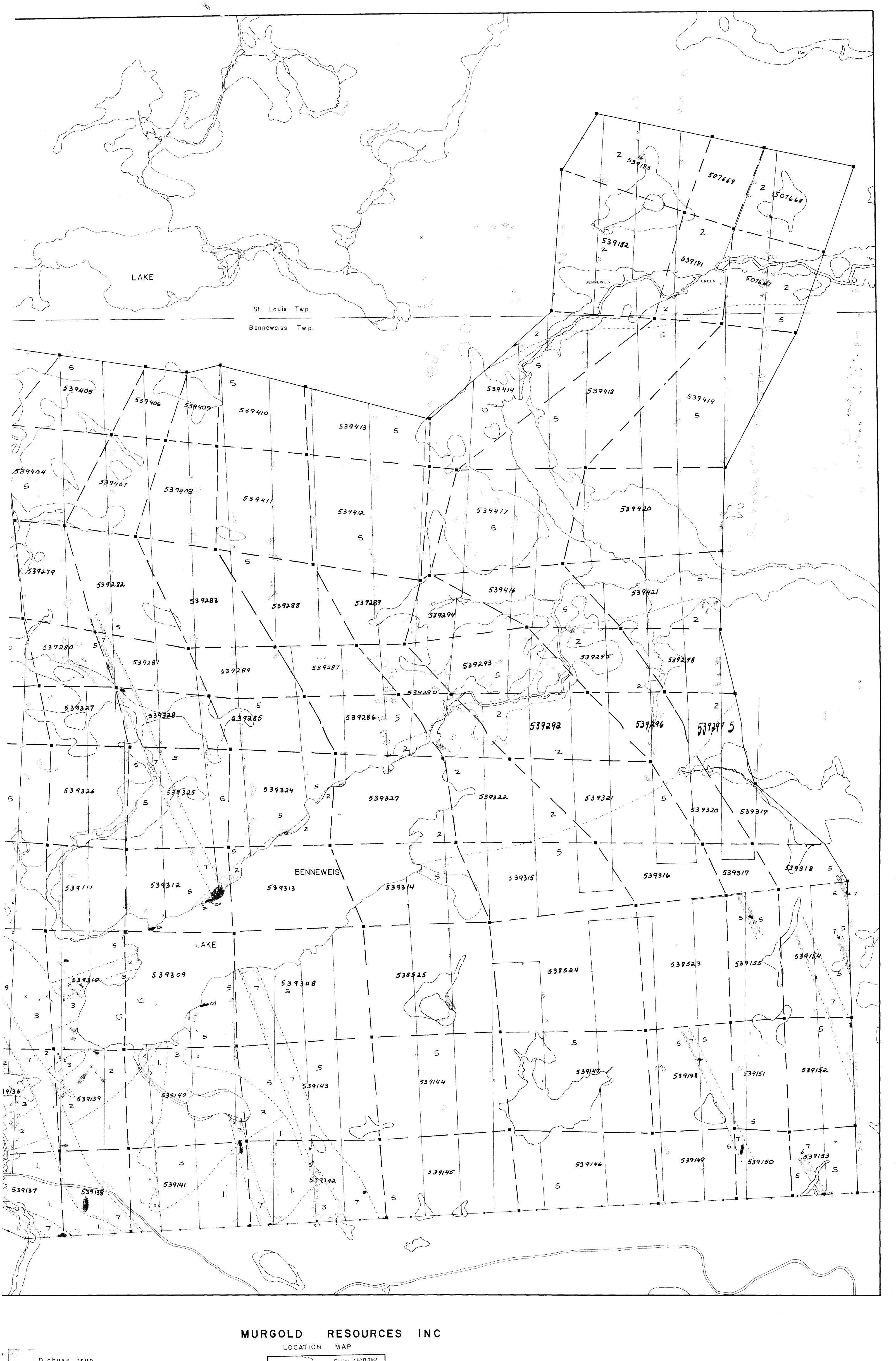
SURVEYS AND MAPPING BRANCH

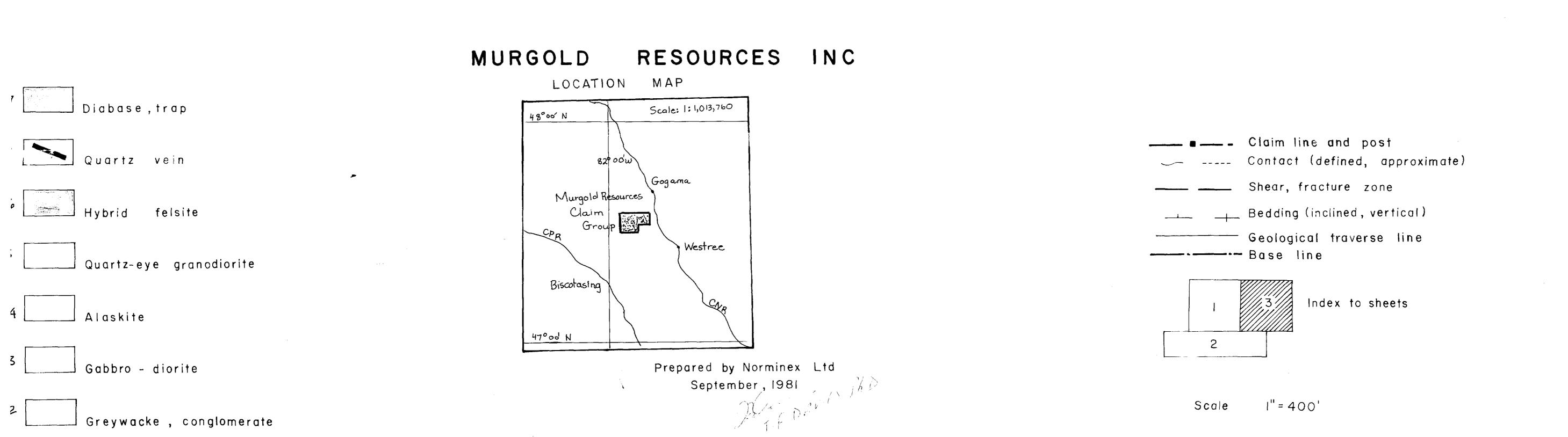


M.1188

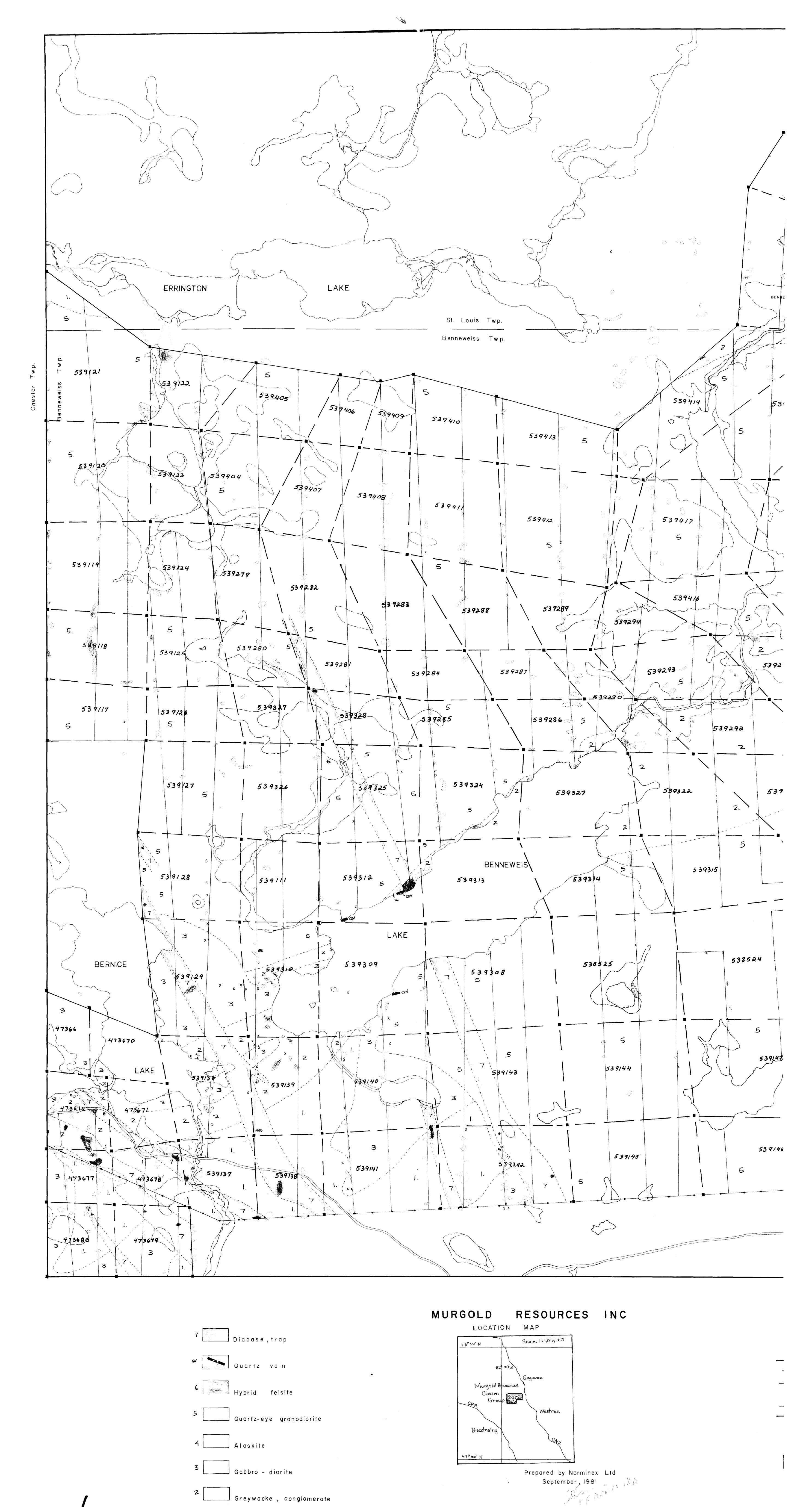








Rhyolite, rhyolite porphyry



41P:25W0004 2.4889 ST. LOUIS

Rhyolite, rhyolite porphyry