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DIGHEMIII SURVEY OF THE N-N PROJECT COCHRANE AREA ONTARIO

FOR FSSO' MINERALS CANADA

BY DIGHEM SURVEYS & PROCESSING INC.

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SUMMARY AND RECOMMENDATIONS

A total of 249 km (155 miles) of survey was flown with DIGHEM^{III} system on May 12, 1986, approximately 70 km otheast of Cochrane, Ontario, for Esso Minerals Canada. ee Figure 1).

The survey detected many strong bedrock conductors ich are typical of graphite and/or sulphides. In dition, there are numerous anomalies which reflect orly-conductive sources of possible bedrock origin. Most

the conductors described in Section 1 of this report pear to warrant further investigation using appropriate rface exploration techniques. Areas of interest may be signed priorities for follow-up work on the basis of pporting geological and/or geochemical information. A mparison of the various geophysical parameters should be tremely valuable in mapping the geological units and ructural breaks within the survey area.

The entire survey area exhibits excellent potential as host for both strongly conductive massive sulphide posits and weakly conductive zones of auriferous neralization. Most of the interpreted bedrock conductors e considered to be of moderate to high priority as ploration targets.



FIGURE 1







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The Flight Records and Path Recovery Α.

EM Anomaly List В.

INTRODUCTION

A DIGHEM^{III} electromagnetic/resistivity/magnetic/VLF rvey totalling approximately 249 line-km was flown with a 0 m line-spacing for Esso Minerals Canada, on May 12, 86. Survey coverage consisted of a single survey grid with averse lines flown in an azimuthal direction of 5°/345°. The 249 km total includes two tie lines.

The survey block is located on N.T.S. map sheet 42H/8. a approximate centre of the survey area occurs at latitude 28'48"N/longitude 80°04'00"W. The survey results have an presented on separate map sheets for each geophysical rameter.

Aerospatiale Squirrel turbine An helicopter egistration C-GFHP) was provided by Frontier Helicopters The helicopter flew at an average airspeed of 130 km/h 1. th an EM bird height of approximately 33 m. Ancillary ipment consisted of a Sonotek PMH 5010 magnetometer with s bird at an average height of 48 m, a Sperry radio timeter, a Geocam sequence camera, an RMS GR33 digital aphics recorder, Sonotek SDS 1200 digital а data quisition system, а Herz Industries Totem-2A F-electromagnetometer with its sensor towed at an average ight of 55 m, and a DigiData 1640 9-track 800-bpi magnetic pe recorder. The analog equipment recorded four channels EM data at approximately 900 Hz, two channels of EM data approximately 7200 Hz, four channels of VLF-EM (total eld and quadrature components for two frequencies, (two bient EM noise channels (for the coaxial and coplanar ceivers), two channels of magnetics (coarse and fine unt), and a channel of radio altitude. The digital uipment recorded the EM data with a sensitivity of 20 ppm at 900 Hz and 0.40 ppm at 7200 Hz, the VLF field to 1%, and the magnetic field to one nT (i.e., one gamma). e VLF-EM receivers were tuned to 24.8 kHz (Seattle, shington-NLK) as the primary station and 21.4 kHz nnapolis, Md.-NSS) as an alternate signal source. The ta shown on the VLF map were derived from the transmitter Seattle. and the second second

In addition to the above equipment, a Del Norte Flying agman navigation system was employed to track the rcraft's progress across the ground. This information was corded in a range-range mode to an accuracy of 5 metres th a once-per-second update.

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Appendix A provides details on the data channels, their spective sensitivities, and the navigation/flight path overy procedure. Noise levels of less than 2 ppm are nerally maintained for wind speeds up to 35 km/h. Higher ids may cause the system to be grounded because excessive swinging produces difficulties in flying :d the The swinging results from the 5 m^2 of licopter. area ich is presented by the bird to broadside gusts. The **GHEM** system nevertheless be flown under wind can nditions that seriously degrade other AEM systems.

In areas where EM responses are evident primarily on e quadrature components, zones of poor conductivity are dicated. Where these responses are coincident with strong qnetic anomalies, it is possible that the inphase mponent amplitudes have been suppressed by the effects of qnetite. Most of these poorly-conductive magnetic atures give rise to resistivity anomalies which are only below background. If it is ightly expected that orly-conductive economic mineralization may be associated th magnetite-rich units, most of these weakly anomalous atures will be of interest. In areas where magnetite the inphase components become negative, the uses to

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parent conductance and depth of EM anomalies may be reliable.

Anomalies which occur near the ends of the survey nes, (i.e., outside the survey area), should be viewed th caution. Some of the weaker anomalies could be due to rodynamic noise, i.e., bird bending, which is created by normal stresses to which the bird is subjected during the imb and turn of the aircraft between lines. Such rodynamic noise is usually manifested by an anomaly on the axial inphase channel only, although severe stresses can fect the coplanar inphase channels as well.

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SECTION I: SURVEY RESULTS

eral Discussion

The survey covered a single grid with 249 km of flying, results of which are shown on separate map sheets for h parameter. Table I-1 summarizes the EM responses in survey area, with respect to conductance grade and erpretation.

The anomalies shown on the electromagnetic anomaly maps based on a near-vertical, half plane model. This model t reflects "discrete" bedrock conductors. Wide bedrock ductors or flat-lying conductive units, whether from ficial or bedrock sources, may give rise to very broad malous responses on the EM profiles. These may not ear on the electromagnetic anomaly maps if they have a ional character rather than locally а anomalous These broad conductors, which more closely racter. roximate a half space model, will be maximum coupled to horizontal (coplanar) coil-pair and should be more dent on the resistivity parameter. The resistivity maps, refore, may be more valuable than the electromagnetic

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TABLE I-1

EM ANOMALY STATISTICS - H-N PROJECT ,

CONDUCTOR GRADE	CONDUCTANCE RANGE	NUMBER OF RESPONSES
6	> 99 мноз	0
5	50-99 MHOS	0
4	20-49 MHOS	0
3	10-19 MHOS	2
2	5- 9 MHOS	10
1	< 5 MHOS	163
х	INDETERMINATE	98
TOTAL		273

CONDUCTOR		NUMBER OF
MODEL	MOST LIKELY SOURCE	RESPONSES
D	DISCRETE BEDROCK CONDUCTOR	25
В	DISCRETE BEDROCK CONDUCTOR	29
S	CONDUCTIVE COVER	211
E	EDGE OF WIDE CONDUCTOR	8
	the second se	
TOTAL		273

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(SEE EM MAP LEGEND FOR EXPLANATIONS)

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maly maps, in areas where broad or flat-lying conductors considered to be of importance. A contoured resistivity , based on the 7200 Hz coplanar data, is included with is report.

Excellent resolution and discrimination of conductors made possible by using a relatively fast sampling rate 0.1 sec and by employing a common frequency (900 Hz) on orthogonal coil-pairs (coaxial and coplanar). The sulting "difference channel" parameter permits Efferentiation of bedrock and surficial conductors, even bugh they exhibit extremely weak conductance in many Bes.

As previously mentioned in the introduction to this port, the effects of magnetite can reduce the positive plitude of the inphase responses and can yield negative phase responses in poorly conductive areas. It should be iterated that the effects of magnetite can yield higher verstated) apparent resistivities, lower (understated) EM nductance values, and erroneously shallow depth estimates. rthermore, the apparent dips of conductors may also be correct if they are flanking, or contained within, gnetite-rich units. There are some instances where the low frequency 00 Hz) inphase response is negative while the high equency (7200 Hz) inphase response is positive. Although effects of magnetite are frequency independent, and ould therefore yield equal negative excursions for both equencies, the higher frequency will yield a more positive sponse over zones of poor conductivity.

P-EM

The VLF map shows the contoured results of the filtered al field parameter (Seattle). As the VLF method is quite sitive to the angle of coupling between the conductor and propogated EM field, conductors which strike towards the station will usually yield a stronger response than ductors which are nearly orthogonal to it. The general ike in the survey area is approximately east-northeast/ t-southwest and therefore provides good coupling with the signal source.

In general, the VLF trends over most of the survey area w moderately good correlation with the interpreted EM ductors and resistivity trends. VLF anomalies which ear to transect the local geologic strike inferred from magnetic, resistivity and EM data, and those VLF trends

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ich appear to be truncated or offset, are likely due to Hults or shears. Lakes and swampy areas appear to have Hiluenced the VLF results.

The VLF parameter does not provide the same degree of solution available from the EM data. Closely-spaced inductors, conductors of short strike length or conductors ich are poorly coupled to the VLF field, may escape tection with this method. Erratic signals from the VLF ansmitters can also give rise to strong, isolated omalies which should be viewed with caution. Regardless these limitations, however, the VLF results have provided ditional structural information, particularly within the re resistive portions of the survey area.

gnetics

The total field magnetic map suggests the geology and ructure underlying the survey area is much more complex an is indicated on the O.G.S. Map 2410*.

At least two major magnetic units extend in an imuthal direction of 070° to 080° across the property.

Ontario Geological Survey, Map 2410, Twopeak Lake

ese are intersected by two parallel linear features which rike roughly 005° and 002° respectively from the extreme uthern ends of lines 10070 and 10220. The former trends e attributed to relatively non-conductive, magnetite-rich cks (possibly iron formation), while the latter are obably due to cross-cutting diabase dikes.

In addition to the linear features, there are several ag-like magnetic anomalies of limited extent. The rongest of these is centered at fiducial 2622 on line 190, where values of more than 4,700 nT above background evident. This interesting plug-like magnetic unit also sts one of the stronger bedrock conductors in the area, lich appears to have been drilled in previous years. Alcopyrite was reportedly intersected in one of the drill tes.

Magnetic relief varies from a low of 59,030 nT on line 580, to a high of about 63,700 nT at anomaly 10190F.**

The magnetic units shown on the total field contour map more clearly resolved and defined on the enhanced inetic map. Furthermore, there are several well-defined

This refers to EM anomaly "F" on survey line 10190.

hanced magnetic anomalies which are only weakly evident on total field map. One example is the feature which tends from fiducial 1097 on line 10631 to the south end of he 10770. The data enhancement technique is described in ction II of this report. It should be noted, however, at the algorithm used to enhance subtle positive magnetic pmalies does not highlight magnetic lows which may be due faulting, alteration or non-magnetic intrusions.

The magnetic results, in conjunction with the other ophysical parameters, should provide information which can used to effectively map the geology and structure in the cvey area.

sistivity

The resistivity map shows the conductive properties of survey area. Some of the resistivity lows (i.e., nductive areas) coincide with discrete bedrock conductors d others indicate weakly conductive overburden associated th water-covered areas. The resistivity patterns may aid plogic mapping and in extending the length of known zones.

The rocks and surficial cover underlying the survey ea are moderately resistive, usually yielding resistivity

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Lues of more than 1,000 ohm-m. Some swamps, however, Itain enough conductive material to lower the apparent sistivity values to less than 1,000 ohm-m. Zones of less in 800 ohm-m resistivity, which do not occur in low-lying water-covered areas, are generally considered to be due bedrock conductors. The resistivity maps provide a quick easy method of outlining all the highly conductive gets in the survey area.

Although the host rocks are generally quite resistive, y contain numerous, clearly-defined, strong EM anomalies ch exhibit responses typical of graphite and/or highly ductive sulphides. There are several conductive zones lined by the 1,000 ohm-m resistivity contour. It is eresting to note that these zones do not appear to be tricted to a single geologic unit, but occur within the metavolcanic mafic flows as well ped as the There are broad resistivity lows overlying asediments. granodiorite intrusive in the west central portion of survey grid. Most of the EM anomalies in this area, exhibit the characteristics ever, of conductive rburden, suggesting that these resistivity lows may lect surficial cover.

One of the more interesting resistivity lows is a -shaped zone which extends southeast from anomaly 10440C

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10480C, swinging west-southwest to anomaly 10390C and n south-southeast to 10430D. This feature may reflect a ded conductive unit or three intersecting conductive nds. The magnetic and VLF results suggest the latter othesis is more likely. In addition to the "Z"-shaped istivity low, there are at least six other lows which are ributed to bedrock conductors. These include the zones ociated with anomalies 10060C, 10120A, 10200B, 10430A, 00A and 10610B. These zones contain the most interesting physical targets in the survey area.

ctromagnetics

The electromagnetic profiles and the calculated istivity parameters for the 900 Ηz and 7200 Hz quencies, suggest that most of the survey area is covered a moderately thin layer of weakly conductive overburden. e e are some anomalies, however, which indicate overburden kness of up to 50 m. Surficial resistivities, based on 7200 Hz data, commonly yield values of more than 10 ohm-m.

Approximately 70% of the anomalous responses in the 'ey area consist of broad responses which exhibit the 'acteristics of a poorly-conductive half space, such as

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ht be expected from conductive overburden. The general k of anomalous responses on the difference channel ameters usually implies moderately broad or flat-lying rces. There are, however, some anomalies which yield erately narrow responses on the quadrature or high quency parameters and give rise to subtle responses on

Some of these coincide with difference channels. netic (bedrock) features and have therefore been given a or "S?" interpretive symbol. These classifications ote possible bedrock sources of weak conductance which be partially masked by the effects of magnetite and/or juctive overburden. Although the latter responses are sidered to be low priority targets on the basis of their physical signatures, they may reflect poorly-conductive izons (weakly mineralized faults or contacts?) which ld be of economic significance in the area. A more rough evaluation of these anomalies should therefore be ried out by one or more qualified professionals who have ess to, and can provide a meaningful compilation of, all ilable geophysical, geological and geochemical data for survey area.

The electromagnetic anomaly maps show the anomaly ations with the interpreted conductor type, dip, Juctance and depth being indicated by symbols. Direct

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inetic correlation is also shown if it exists. The strike ection and length of the conductors are indicated when malies can be correlated from line to line. When idying the map sheets for follow-up planning, consult the maly listings appended to this report to ensure that none the conductors are overlooked.

DUCTORS IN THE SURVEY AREA

It is beyond the scope of this report to describe all anomalous responses defined by the survey. The lowing text provides a brief description of the anomalies ch appear to be due to bedrock conductors. Although some the stronger anomalies may yield signatures which are racteristic of graphite and/or semi-massive to massive phides, they do not necessarily bear a direct ationship to economic mineralization. In areas where iferous mineralization "is considered to be the primary loration target, some of the very weak or magnetiteociated anomalies may be of greater importance than the onger, well defined conductors.

Where several conductors or conductive trends exhibit ilar characteristics, or appear to be related to a common logical unit or stratigraphic horizon, these have been uped into "Zones" for purposes of discussion. The zone lines shown on the EM map may approximate the limits of a ductive unit, a magnetic unit, or both.

maly: 10060C

This interesting anomaly gives rise to an isolated resistivity low of limited extent, located within a relative magnetic low. It reflects a strong, narrow bedrock conductor with a probable dip to the north. Although this non-magnetic conductor is not evident on adjacent lines, it is considered to be a high priority target. It is recommended that further investigation be carried out to check the causative source of this attractive anomaly.

malies: 10070A, 10110B-10120A

Anomaly 10070 is associated with a moderately strong magnetic anomaly, exhibiting а direct correlation of 430 nT. The effects of magnetite have suppressed the inphase responses and have probably yielded erroneously high resistivity values. The anomaly probably reflects conductive material associated with magnetite. It may be open to the west and could continue eastward through 10080A to 10120A.

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10110B-10120A reflects a Anomaly narrow, north-dipping conductor which is associated with the magnetic trend which hosts 10070A, same and may therefore be related to a similar causative source. The eastern portion, 10110B-10120A, does not appear to be as magnetic as the western portion, and probably reflects a higher concentration of conductive material within the magnetite-rich host rock. This conductor also correlates with a VLF anomaly and an isolated resistivity low which may enhance its significance. It is recommended that this area be subjected to further detailed investigation.

malies: 10100A-10130A, 10120B, 10120F

Anomalies in this group comprise broad, poorly-defined responses which are probably due to conductive overburden or a broad conductive rock unit. subtle inflections the There are on quadrature difference channels which may be indicative of weak, buried bedrock conductors. With the possible exception "edge effect", all 10120B, which denotes of an anomalies in this group are associated with a moderately strong resistivity low and a relative magnetic depression. This may reflect a felsic intrusion or a zone of alteration. These anomalies do not coincide with VLF anomalies, probably because of the attenuating effects of the conductive cover.

Most of the anomalies in this group are attributed primarily to conductive overburden and are therefore considered to be of moderately low priority.

malies: 10180xD-10220C, 10160xB, 10230A, 10230C

The anomalies in this group are contained in an area between two parallel north/south trending magnetic linears which are attributed to diabase dikes.

The conductor defined by 10180xD-10220C is the strongest feature in the survey area. It is associated with a very strong, plug-like magnetic anomaly, as evidenced by the 4,820 nT correlation with anomaly 10190F. This conductor reflects a narrow, northdipping, highly conductive and magnetic source, with a probable strike length of about 400 m. Pyrrhotite is considered to be a likely cause although chalcopyrite was reportedly intersected in one of the drill holes in this vicinity. This conductor appears to have been drilled in the past but additional work may be warranted. Anomalies 10160xB, 10230A and 10230C consist of isolated responses which may reflect bedrock conductors of short strike length. Poorly conductive material associated with magnetite is a likely cause of 10160xB. This anomaly, however, is located in close proximity to the northwestern flank of the strong magnetic plug, where it abuts a north/south trending diabase dike. This poorly-defined conductor is of interest because of the apparent structural deformation which may have influenced mineral deposition in the immediate area.

Anomaly 10230A is an interesting single-line response which occurs on the southern flank of a weak magnetic anomaly. The sharpness of this response suggests that a spheric noise spike may have been a contributing factor, although there is no supporting evidence on the spherics monitor. This anomaly, therefore, is considered to be due to a concentration of metallic sulphides of short strike extent. Its associated VLF anomaly may be due to a geological contact which occurs near the edge of a zone of conductive overburden. The coincident resistivity low, however, is more isolated in nature, and probably reflects the bedrock conductor. Further work is

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warranted to determine the causative source, the strike, and extent of this conductor.

Anomaly 10230C occurs at the northern edge of a well-defined magnetic anomaly and near the southern edge of a VLF anomaly. This weak anomaly is probably due to a partially masked, poorly-conductive bedrock feature which may be related to a (faulted?) contact.

malies: 10280A, 10300xA, 10320xB, 10390xD

Anomalies in this group consist of isolated responses which suggest "edge effects", i.e., resistivity contrasts at the edges of conductive units. It is difficult to ascertain whether the resistivity contrast occurs within the bedrock or is related to surficial cover.

Anomalies 10300xA and 10390xD both occur on the same east/west trending magnetic unit, yielding direct magnetic correlation of 180 nT and 310 nT respectively. Both may be due to weakly conductive, moderately broad bedrock units. The magnetic host appears to be an eastward continuation of the strong plug-dike unit which hosts conductor 10180xD-10220C, described previously. Response 10390xD is associated with negative inphase responses caused by magnetite and yields strong VLF correlation.

Anomaly 10280A is related to a diabase dike which is probably fault controlled. The anomaly occurs at a contact between resistive material to the west and more conductive material to the east. There is a possible sinistral offset of about 200 m associated with the inferred diabase.

Anomaly 10320xB probably reflects the northern edge of a zone of conductive overburden and is considered to be of very low priority.

maly: 10410B-10440C (Zone A)

This interesting zone comprises two intersecting conductors or a single folded conductive horizon which strikes east-northeast from 10410B to 10480D where it swings to the northwest to 10440C. The anomalies comprising this trend suggest a narrow, non-magnetic bedrock conductor which dips to the north. This conductor forms the north and central limbs of a "Z-shaped" resistivity low. All three limbs are associated with VLF anomalies. The central portion of the zone occurs on the north flank of a major east-northeast trending magnetic unit which may be loosely associated with a mapped iron formation. The southern limb also occurs on the north flank of a magnetic unit of similar intensity but shorter strike length. The magnetic contour patterns suggest that the north and south limbs may be separate from the central The VLF results show the north and south portion. limbs are associated with northwest/southeast trending linears, which may be due to faults or shears. The lack of any offsets in the central portion indicates that the major magnetic unit (and the conductive material?) either occurred post-faulting, or there was little or no lateral displacement during the faulting stage. It is recommended that this interesting zone be subjected to further investigation, if such work has not already been carried out.

Other anomalies in this area may also warrant follow-up, even if they have been given a "B?" or "S?" interpretive symbol. Such responses would include 10390xA, 10390C, 10400D, 10510xD and 10470D-10480E. Most of these anomalies are associated with resistivity lows and/or VLF anomalies.

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alies: 10430A-10530B, 10570A-10640xA

These two anomalous trends reflect separate segments of an attractive narrow, north-dipping conductor which strikes approximately east/west. Both segments of the conductor are associated with а continuous VLF anomaly, and give rise to well-defined resistivity lows. The western segment occurs in a relatively non-magnetic area, with the exception of anomaly 10430A which yields a direct magnetic correlation of 30 nT. The eastern segment is situated near the northern flank of a strong magnetic unit. Offsets in magnetic contour patterns suggest the two conductive segments are separated by a probable structural break which extends in a northerly direction from the south end of line 10410 to the north end of line 10610. This inferred break may be paralleled by two other breaks which strike through 10490xB-10590B and 10520xE-10650A.

Conductors 10430A-10520B and 10570A-10640xB are attractive targets which should be followed up on a high priority basis. Initial attention may be focused on the more conductive portions (such as 10430A, 10500B and 10620B), anomalies which yield magnetic correlation (10430A, 10590B and 10620B) and areas where the conductor appears to have been subjected to possible deformation.

to In addition the probable bedrock conductors scribed in the foregoing, there are several weaker malies which may be of interest. Most of these comprise orly-defined anomalies with an "S?" interpretive symbol, reflect weak bedrock conductors masked by .ch may erburden. Such anomalies may be upgraded if they relate with magnetic and/or VLF anomalies, or yield istivity values which are lower than those observed in er overburden-covered areas. Examples would include malies 10400C, 10250xA, 10350xA, 10500A, 10510xE, 10550A, 31xE, 10670B and 10730B, for example. Most of these are sociated with magnetic/VLF trends and/or resistivity dients which indicate possible geological contacts or erately broad conductors. A detailed follow-up program ill be required to investigate and properly assess the ative merits of the weak isolated responses as well as highly conductive trends.

Although at least one of the major conductors has been jected to drilling in previous years, there are numerous eresting conductors which apparently remain untested. lti-frequency horizontal loop system (MaxMin) or a domain equivalent (PEM) should suffice as an effective w-up tool although VLF may be adequate to locate the ger conductors in most areas.

The entire property is considered to be a potential of both massive sulphides and weakly conductive erous mineralization.

SECTION II: BACKGROUND INFORMATION

Section II provides background information on products ich are available from your survey data. Those products c obtained as part of the survey contract may be generated ter from raw data which is available on your archive gital tape.

ELECTROMAGNETICS

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DIGHEM electromagnetic responses fall into two general asses, discrete and broad. The discrete class consists of arp, well-defined anomalies from discrete conductors such sulfide lenses and steeply dipping sheets of graphite and lfides. The broad class consists of wide anomalies from nductors having a large horizontal surface such as flatly oping graphite or sulfide sheets, saline water-saturated limentary formations, conductive overburden and rock, and othermal zones. A vertical conductive slab with a width 200 m would straddle these two classes.

The vertical sheet (half plane) is the most common del used for the analysis of discrete conductors. All omalies plotted on the electromagnetic map are analyzed cording to this model. The following section entitled screte Conductor Analysis describes this model in detail,

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: luding the effect of using it on anomalies caused by ad conductors such as conductive overburden.

The conductive earth (half space) model is suitable for ad conductors. Resistivity contour maps result from the of this model. A later section entitled <u>Resistivity</u> <u>ping</u> describes the method further, including the effect using it on anomalies caused by discrete conductors such sulfide bodies.

metric interpretation

The geophysical interpreter attempts to determine the metric shape and dip of the conductor. Figure II-1 shows ical DIGHEM anomaly shapes which are used to guide the metric interpretation.

crete conductor analysis

The EM anomalies appearing on the electromagnetic map analyzed by computer to give the conductance (i.e., ductivity-thickness product) in mhos of a vertical sheet el. This is done regardless of the interpreted geometric pe of the conductor. This is not an unreasonable cedure, because the computed conductance increases as the

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•		\setminus		0			
line	vertical thin dike	dipping thin dike	vertical or dipping thick dike	sphere; horizontat disk; metal roof; small fenced yard	wide horizontai ribbon; large fenced area	 S = conductive overburden H = thick conductive cover or wide conductive rock unit E = edge effect from wide conductor 	flight line parallel to conductor
4/1	2/1	variabte	variable	1/4	variable	1/2	< 1/4

Typical DIGHEM anomaly shape

ectrical quality of the conductor increases, regardless of s true shape. DIGHEM anomalies are divided into six ades of conductance, as shown in Table II-1. The conducnce in mhos is the reciprocal of resistance in ohms.

Anomaly Grade	Mho Range
6	> 99
5	50 - 99
4	20 - 49
3	10 - 19
2	5 - 9
1	< 5

Table II-1. EM Anomaly Grades

The conductance value is a geological parameter because is a characteristic of the conductor alone. It generally independent of frequency, flying height or depth of rial, apart from the averaging over a greater portion of e conductor as height increases.¹ Small anomalies from eply buried strong conductors are not confused with small omalies from shallow weak conductors because the former 11 have larger conductance values.

Conductive overburden generally produces broad EM sponses which may not be shown as anomalies on the EM ps. However, patchy conductive overburden in otherwise

This statement is an approximation. DIGHEM, with its short coil separation, tends to yield larger and more accurate conductance values than airborne systems having a larger coil separation.

istive areas can yield discrete anomalies with a ductance grade (cf. Table II-1) of 1, or even of 2 for ducting clays which have resistivities as low as 50 -m. In areas where ground resistivities can be below 10 -m, anomalies caused by weathering variations and similar ses can have any conductance grade. The anomaly shapes n the multiple coils often allow such conductors to be ognized, and these are indicated by the letters S, H, G sometimes E on the map (see EM legend).

bedrock conductors, the higher anomaly grades For increasingly higher conductances. icate Examples: Insco copper discovery (Noranda, Canada) HEM's New grade 4 anomaly, as did the neighbouring lded а per-zinc Magusi River ore body; Mattabi (copper-zinc, Lake, Canada) Whistle (nickel, rgeon and Sudbury, ada) gave grade 5; and DIGHEM's Montcalm nickel-copper covery (Timmins, Canada) yielded a grade 6 anomaly. phite and sulfides can span all grades but, in any ticular survey area, field work may show that the ferent grades indicate different types of conductors.

Strong conductors (i.e., grades 5 and 6) are characteric of massive sulfides or graphite. Moderate conductors ades 3 and 4) typically reflect graphite or sulfides of a s massive character, while weak bedrock conductors ades 1 and 2) can signify poorly connected graphite or vily disseminated sulfides. Grade 1 conductors may not pond to ground EM equipment using frequencies less than 0 Hz.

The presence of sphalerite or gangue can result in deposits having weak to moderate conductances. As example, the three million ton lead-zinc deposit of tigouche Mining Corporation near Bathurst, Canada, lded a well defined grade 1 conductor. The 10 percent volume of sphalerite occurs as a coating around the fine ined massive pyrite, thereby inhibiting electrical duction.

Faults, fractures and shear zones may produce anomalies ch typically have low conductances (e.g., grades 1 2). Conductive rock formations can yield anomalies of conductance grade. The conductive materials in such k formations can be salt water, weathered products such clays, original depositional clays, and carbonaceous erial.

On the electromagnetic map, a letter identifier and an erpretive symbol are plotted beside the EM grade symbol. horizontal rows of dots, under the interpretive symbol, icate the anomaly amplitude on the flight record. The

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rtical column of dots, under the anomaly letter, gives the timated depth. In areas where anomalies are crowded, the tter identifiers, interpretive symbols and dots may be literated. The EM grade symbols, however, will always be scernible, and the obliterated information can be obtained on the anomaly listing appended to this report.

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The purpose of indicating the anomaly amplitude by dots to provide an estimate of the reliability of the conducnce calculation. Thus, a conductance value obtained from a large ppm anomaly (3 or 4 dots) will tend to be accurate ereas one obtained from a small ppm anomaly (no dots) uld be quite inaccurate. The absence of amplitude dots dicates that the anomaly from the coaxial coil-pair is 5 ppm or less on both the inphase and quadrature channels. 5 ich small anomalies could reflect a weak conductor at the surface or a stronger conductor at depth. The conductance g ade and depth estimate illustrates which of these sosibilities fits the recorded data best.

Flight line deviations occasionally yield cases where two anomalies, having similar conductance values but characterized and the same conductor. Such examples illustrate the reliability of the conductance measurement while showing that the depth estimate can be unreliable. There are a

- II-7 - .

mber of factors which can produce an error in the depth timate, including the averaging of topographic variations the altimeter, overlying conductive overburden, and the cation and attitude of the conductor relative to the ight line. Conductor location and attitude can provide an roneous depth estimate because the stronger part of the nductor may be deeper or to one side of the flight line, because it has a shallow dip. A heavy tree cover can also produce errors in depth estimates. This is because the impth estimate is computed as the distance of bird from nductor, minus the altimeter reading. The altimeter can lock onto the top of a dense forest canopy. This situation relds an erroneously large depth estimate but does not a fect the conductance estimate.

Dip symbols are used to indicate the direction of dip conductors. These symbols are used only when the anomaly shapes are unambiguous, which usually requires a fairly cosistive environment.

A further interpretation is presented on the EM map by reans of the line-to-line correlation of anomalies, which is tased on a comparison of anomaly shapes on adjacent lines. It is provides conductor axes which may define the geological E:ructure over portions of the survey area. The absence of

- II-8 -

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E ductor axes in an area implies that anomalies could not E correlated from line to line with reasonable confidence.

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DIGHEM electromagnetic maps are designed to provide correct impression of conductor quality by means of the inductance grade symbols. The symbols can stand alone is ductance grade symbols. The symbols can stand alone is ductance values are printed in the attached anomaly list is those who wish quantitative data. The anomaly ppm and eith are indicated by inconspicuous dots which should not is stract from the conductor patterns, while being helpful to those who wish this information. The map provides an interpretation of conductors in terms of length, strike and is, geometric shape, conductance, depth, and thickness (see elow). The accuracy is comparable to an interpretation is m a high quality ground EM survey having the same line is acing.

The attached EM anomaly list provides a tabulation of pomalies in ppm, conductance, and depth for the vertical bet model. The EM anomaly list also shows the conductance and depth for a thin horizontal sheet (whole plane) model, ut only the vertical sheet parameters appear on the map. The horizontal sheet model is suitable for a flatly pping thin bedrock conductor such as a sulfide sheet eving a thickness less than 10 m. The list also shows the

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istivity and depth for a conductive earth (half space) el, which is suitable for thicker slabs such as thick ductive overburden. In the EM anomaly list, a depth ue of zero for the conductive earth model, in an area of the cover, warns that the anomaly may be caused by beductive overburden.

Since discrete bodies normally are the targets of i surveys, local base (or zero) levels are used to compute al anomaly amplitudes. This contrasts with the use true zero levels which are used to compute true EM litudes. Local anomaly amplitudes are shown in the anomaly list and these are used to compute the vertical let parameters of conductance and depth. Not shown in the anomaly list are the true amplitudes which are used to pute the horizontal sheet and conductive earth mameters.

+ ype electromagnetic responses

DIGHEM maps contain x-type EM responses in addition > EM anomalies. An x-type response is below the noise > eshold of 3 ppm, and reflects one of the following: a > k conductor near the surface, a strong conductor at depth > g., 100 to 120 m below surface) or to one side of the l ght line, or aerodynamic noise. Those responses that we the appearance of valid bedrock anomalies on the flight ofiles are indicated by appropriate interpretive symbols ee EM map legend). The others probably do not warrant wrther investigation unless their locations are of considerable geological interest.

l : thickness parameter

DIGHEM can provide an indication of the thickness of steeply dipping conductor. The amplitude of the coplanar romaly (e.g., CPI channel on the digital profile) increases Elative to the coaxial anomaly (e.g., CXI) as the apparent ckness increases, i.e., the thickness in the horizontal line. (The thickness is equal to the conductor width if in conductor dips at 90 degrees and strikes at right angles > the flight line.) This report refers to a conductor as in when the thickness is likely to be less than 3 m, and ck when in excess of 10 m. Thick conductors are icated on the EM map by crescents. For base metal i loration in steeply dipping geology, thick conductors can high priority targets because many massive sulfide ore mies are thick, whereas non-economic bedrock conductors often thin. The system cannot sense the thickness when is strike of the conductor is subparallel to the flight e, when the conductor has a shallow dip, when the anomaly a litudes are small, or when the resistivity of the a ironment is below 100 ohm-m.

istivity mapping

widespread conductivity are commonly Areas of countered during surveys. In such areas, anomalies can E generated by decreases of only 5 m in survey altitude as E.1 as by increases in conductivity. The typical flight E ord in conductive areas is characterized by inphase and drature channels which are continuously active. Local v peaks reflect either increases in conductivity of the : th or decreases in survey altitude. For such conductive reas, apparent resistivity profiles and contour maps are Elessary for the correct interpretation of the airborne The advantage of the resistivity parameter is ella. hat anomalies caused by altitude changes are virtually liminated, so the resistivity data reflect only those romalies caused by conductivity changes. The resistivity ralysis also helps the interpreter to differentiate between onductive trends in the bedrock and those patterns typical f conductive overburden. For example, discrete conductors ill generally appear as narrow lows on the contour map nd broad conductors (e.g., overburden) will appear as de lows.

The resistivity profile (see table in Appendix A) and resistivity contour map present the apparent resistivity the so-called pseudo-layer (or buried) half space ig 5 defined by Fraser (1978)². This model consists of sesistive layer overlying a conductive half space. The p:h channel (see Appendix A) gives the apparent depth) w surface of the conductive material. The apparent r:h is simply the apparent thickness of the overlying The apparent depth (or thickness) Eistive layer. rameter will be positive when the upper layer is more Eistive than the underlying material, in which case the Firent depth may be quite close to the true depth.

The apparent depth will be negative when the upper yer is more conductive than the underlying material, and 11 be zero when a homogeneous half space exists. The parent depth parameter must be interpreted cautiously cause it will contain any errors which may exist in the sured altitude of the EM bird (e.g., as caused by a dense (e cover). The inputs to the resistivity algorithm are inphase and quadrature components of the coplanar 1-pair. The outputs are the apparent resistivity of the

esistivity mapping with an airborne multicoil electroagnetic system: Geophysics, v. 43, p. 144-172. Figure 1 and the sensor-source of the source) and the sensor-source shance. The flying height is not an input variable, is the output resistivity and sensor-source distance are inpendent of the flying height. The apparent depth, is simply the sensor-source distance minus is measured altitude or flying height. Consequently, is in the measured altitude will affect the apparent is parameter but not the apparent resistivity parameter.

The apparent depth parameter is a useful indicator simple layering in areas lacking a heavy tree cover. DIGHEM system has been flown for purposes of permafrost ping, where positive apparent depths were used as a sure of permafrost thickness. However, little quantitae use has been made of negative apparent depths because absolute value of the negative depth is not a measure of thickness of the conductive upper layer and, therefore, not meaningful physically. Qualitatively, a negative arent depth estimate usually shows that the EM anomaly is sed by conductive overburden. Consequently, the apparent th channel can be of significant help in distinguishing : ween overburden and bedrock conductors.

The resistivity map often yields more useful informaion on conductivity distributions than the EM map. In

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paring the EM and resistivity maps, keep in mind the owing:

- (a) The resistivity map portrays the absolute value of the earth's resistivity.
 (Resistivity = 1/conductivity.)
- (b) The EM map portrays anomalies in the earth's resistivity. An anomaly by definition is a change from the norm and so the EM map displays anomalies, (i) over narrow, conductive bodies and (ii) over the boundary zone between two wide formations of differing conductivity.

The resistivity map might be likened to a total and map and the EM map to a horizontal gradient in the restriction of flight³. Because gradient maps are usually resensitive than total field maps, the EM map therefore to be preferred in resistive areas. However, in conducte areas, the absolute character of the resistivity map tally causes it to be more useful than the EM map.

The gradient analogy is only valid with regard to the identification of anomalous locations.

r pretation in conductive environments

Environments having background resistivities below tim-m cause all airborne EM systems to yield very e responses from the conductive ground. This usually bits the recognition of discrete bedrock conductors. processing of DIGHEM data, however, produces six rels which contribute significantly to the recognition drock conductors. These are the inphase and quadrature rence channels (DIFI and DIFQ), and the resistivity and channels (RES and DP) for each coplanar frequency; see in Appendix A. なると、ためたちないないないなるが、たちにたいいうない

The EM difference channels (DIFI and DIFQ) eliminate o 99% of the response of conductive ground, leaving ponses from bedrock conductors, cultural features (e.g., ponne lines, fences, etc.) and edge effects. An edge (ot arises`when the conductivity of the ground suddenly piges, and this is a source of geologic noise. While edge (ot yield anomalies on the EM difference channels, they not produce resistivity anomalies. Consequently, the stivity channel aids in eliminating anomalies due to effects. On the other hand, resistivity anomalies coincide with the most highly conductive sections of puctive ground, and this is another source of geologic

- II-16 -

The recognition of a bedrock conductor in a tive environment therefore is based on the anomalous bases of the two difference channels (DIFI and DIFQ) the two resistivity channels (RES). The most favourable the ion is where anomalies coincide on all four channels.

The DP channels, which give the apparent depth to the stive material, also help to determine whether a stive response arises from surficial material or from a stive zone in the bedrock. When these channels ride the zero level on the digital profiles (i.e., depth is rive), it implies that the EM and resistivity profiles esponding primarily to a conductive upper layer, i.e., totive overburden. If both DP channels are below the level, it indicates that a resistive upper layer s, and this usually implies the existence of a bedrock ictor. If the low frequency DP channel is below the level and the high frequency DP is above, this suggests t a bedrock conductor occurs beneath conductive cover.

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The conductance channel CDT identifies discrete ductors which have been selected by computer for maisal by the geophysicist. Some of these automatically

- II-17 -

cted anomalies on channel CDT are discarded by the hysicist. The automatic selection algorithm is intionally oversensitive to assure that no meaningful i onses are missed. The interpreter then classifies the r alies according to their source and eliminates those if are not substantiated by the data, such as those , ing from geologic or aerodynamic noise.

1 ction of geologic noise

Geologic noise refers to unwanted geophysical plonses. For purposes of airborne EM surveying, geologic like refers to EM responses caused by conductive overburden magnetic permeability. It was mentioned above that EM difference channels (i.e., channel DIFI for inphase DIFQ for quadrature) tend to eliminate the response of bluctive overburden. This marked a unique development airborne EM technology, as DIGHEM is the only EM system i :h yields channels having an exceptionally high degree ...munity to conductive overburden.

Magnetite produces a form of geological noise on the pase channels of all EM systems. Rocks containing less in 1% magnetite can yield negative inphase anomalies ised by magnetic permeability. When magnetite is widely > ibuted throughout a survey area, the inphase EM chanmay continuously rise and fall reflecting variations : e magnetite percentage, flying height, and overburden ness. This can lead to difficulties in recognizing > y buried bedrock conductors, particularly if conductive : urden also exists. However, the response of broadly to ibuted magnetite generally vanishes on the inphase forence channel DIFI. This feature can be a significant in the recognition of conductors which occur in rocks taining accessory magnetite.

ragnetite mapping

The information content of DIGHEM data consists of a ination of conductive eddy current response and magnetic reability response. The secondary field resulting from luctive eddy current flow is frequency-dependent and isists of both inphase and quadrature components, which positive in sign. On the other hand, the secondary d resulting from magnetic permeability is independent frequency and consists of only an inphase component which negative in sign. When magnetic permeability manifests elf by decreasing the measured amount of positive whase, its presence may be difficult to recognize. se anomaly (e.g., in the absence of eddy current flow), resence is assured. In this latter case, the negative > nent can be used to estimate the percent magnetite : nt.

A magnetite mapping technique was developed for the lanar coil-pair of DIGHEM. The technique yields channel (see Appendix A) which displays apparent weight percent netite according to a homogeneous half space model.⁴ The d can be complementary to magnetometer mapping in Compared to magnetometry, it is far less : ain cases. itive but is more able to resolve closely spaced quetite zones, as well as providing an estimate of the t int of magnetite in the rock. The method is sensitive to 43 magnetite by weight when the EM sensor is at a height 30 m above a magnetitic half space. It can individually solve steeply dipping narrow magnetite-rich bands which separated by 60 m. Unlike magnetometry, the EΜ inetite method is unaffected by remanent magnetism or 1 netic latitude.

The EM magnetite mapping technique provides estimates f magnetite content which are usually correct within a

Refer to Fraser, 1981, Magnetite mapping with a multicoil airborne electromagnetic system: Geophysics, v. 46, p. 1579-1594.

for of 2 when the magnetite is fairly uniformly fributed. EM magnetite maps can be generated when fetic permeability is evident as indicated by anomalies the magnetite channel FEO.

Like magnetometry, the EM magnetite method maps bedrock features, provided that the overburden is macterized by a general lack of magnetite. This masts with resistivity mapping which portrays the thined effect of bedrock and overburden.

gnition of culture

Cultural responses include all EM anomalies caused by r-made metallic objects. Such anomalies may be caused by rictive coupling or current gathering. The concern of the serpreter is to recognize when an EM response is due to ture. Points of consideration used by the interpreter, on coaxial and coplanar coil-pairs are operated at a mon frequency, are as follows:

Channels CXS and CPS (see Appendix A) measure 50 and 60 Hz radiation. An anomaly on these channels shows that the conductor is radiating cultural power. Such an indication is normally a guarantee that the conducfor is cultural. However, care must be taken to ensure hat the conductor is not a geologic body which strikes cross a power line, carrying leakage currents.

In flight which crosses a "line" (e.g., fence, telephone ine, etc.) yields a center-peaked coaxial anomaly and an m-shaped coplanar anomaly.⁵ When the flight crosses the cultural line at a high angle of intersection, the amplitude ratio of coaxial/coplanar (e.g., CXI/CPI) is 4. Such an EM anomaly can only be caused by a line. The geologic body which yields anomalies most closely resembling a line is the vertically dipping thin dike. Such a body, however, yields an amplitude ratio of 2 rather than 4. Consequently, an m-shaped coplanar anomaly with a CXI/CPI amplitude ratio of 4 is virtually a guarantee that the source is a cultural line.

A flight which crosses a sphere or horizontal disk yields center-peaked coaxial and coplanar anomalies with a CXI/CPI amplitude ratio (i.e., coaxial/coplanar) of 1/4. In the absence of geologic bodies of this geometry, the most likely conductor is a metal roof or

See Figure II-1 presented earlier.

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small fenced yard.⁶ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.

A flight which crosses a horizontal rectangular body or wide ribbon yields an m-shaped coaxial anomaly and a center-peaked coplanar anomaly. In the absence of geologic bodies of this geometry, the most likely conductor is a large fenced area.⁶ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.

EM anomalies which coincide with culture, as seen on the camera film, are usually caused by culture. However, care is taken with such coincidences because a geologic conductor could occur beneath a fence, for example. In this example, the fence would be expected to yield an m-shaped coplanar anomaly as in case #2 above. If, instead, a center-peaked coplanar anomaly occurred, there would be concern that a thick geologic conductor coincided with the cultural line.

t is a characteristic of EM that geometrically dentical anomalies are obtained from: (1) a planar onductor, and (2) a wire which forms a loop having imensions identical to the perimeter of the equivaent planar conductor.

above description of anomaly shapes is valid The when the culture is not conductively coupled to the environment. In this case, the anomalies arise from inductive coupling to the EM transmitter. However, when the environment is quite conductive (e.g., less than 100 ohm-m at 900 Hz), the cultural conductor may be conductively coupled to the environment. In this latter case, the anomaly shapes tend to be governed by current gathering. Current gathering can completely distort the anomaly shapes, thereby complicating the identification of cultural anomalies. In such circumstances, the interpreter can only rely on the radiation channels CXS and CPS, and on the camera film.

TOTAL FIELD MAGNETICS

The existence of a magnetic correlation with an EM maly is indicated directly on the EM map. An EM anomaly h magnetic correlation has a greater likelihood of ng produced by sulfides than one that is non-magnetic. ever, sulfide ore bodies may be non-magnetic (e.g., the d Creek deposit near Timmins, Canada) as well as magnetic g., the Mattabi deposit near Sturgeon Lake, Canada).

'he magnetometer data are digitally recorded in circraft to an accuracy of one nT (i.e., one gamma). igital tape is processed by computer to yield a l field magnetic contour map. When warranted, the Elic data also may be treated mathematically to enhance ragnetic response of the near-surface geology, and an thed magnetic contour map is then produced. The sise of the enhancement operator in the frequency domain lustrated in Figure II-2. This figure shows that the t and components of the airborne data are amplified : mes by the enhancement operator. This means, for tle, that a 100 nT anomaly on the enhanced map reflects T anomaly for the passband components of the airborne

The enhanced map, which bears a resemblance to a hard continuation map, is produced by the digital hass filtering of the total field data. The enhancement huivalent to continuing the field downward to a level he the source) which is 1/20th of the actual sensorhe distance.

Because the enhanced magnetic map bears a resemblance a ground magnetic map, it simplifies the recognition trends in the rock strata and the interpretation of

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2

Frequency response of magnetic operator.

-II-26-

gical structure. It defines the near-surface local gy while de-emphasizing deep-seated regional features. imarily has application when the magnetic rock units teeply dipping and the earth's field dips in excess degrees.

VLF-EM

VLF-EM anomalies are not EM anomalies in the entional sense. EM anomalies primarily reflect eddy ints flowing in conductors which have been energized ctively by the primary field. In contrast, VLF-EM elies primarily reflect current gathering, which is a : nductive phenomenon. The primary field sets up ents which flow weakly in rock and overburden, and these co collect in low resistivity zones. Such zones may be to massive sulfides, shears, river valleys and even n formities.

The Herz Industries Ltd Totem VLF-electromagnetometer trees the total field and vertical quadrature trents. Both these components are digitally recorded in dircraft with a sensitivity of 0.1 percent. The total trees yields peaks over VLF-EM current concentrations



nre 3

Frequency response of VLF-EM operator.

s: the quadrature component tends to yield crossovers. ppear as traces on the profile records. The total data also are filtered digitally and displayed on a u: map, to facilitate the recognition of trends in the s:rata and the interpretation of geologic structure.

the response of the VLF-EM total field filter operator frequency domain (Figure II-3) is basically similar used to produce : at the enhanced magnetic map II-2). The two filters are identical along the ∴ e sa but different along the ordinant. The VLF-EM e removes long wavelengths such as those which reflect nal and wave transmission variations. The filter) ns short wavelength responses such as those which y t local geological variations. The filtered total | VLF-EM contour map is produced with a contour interval 1 percent.

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MAPS ACCOMPANYING THIS REPORT

> sheets accompany this report:

🗉 magnetic Anomalies	1 map sheet
vity (7200 Hz)	1 map sheet
ield Magnetics	1 map sheet
a Magnetics	1 map sheet
Total Field (Seattle)	1 map sheet

Respectfully submitted, DIGHEM_SURVEYS & PROCESSING INC.

av

Paul A. Smith Geophysicist

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APPENDIX A

THE FLIGHT RECORDS

analog and digital flight records were produced. by profiles were recorded on chart paper in the t during the survey. The digital profiles were t later by computer and plotted on electrostatic eper at a scale of 1:10,000. The analog and digital t are listed in Tables A-1 and A-2 respectively.

Table A-2, the log resistivity scale of 0.06 In means that the resistivity changes by an order I tude in 16.5 mm. The resistivities at 0, 33 and I) from the bottom of the digital flight record are vely 1, 100 and 10,000 ohm-m.

FLIGHT PATH RECOVERY

i craft positioning and post-survey recovery of
f position was accomplished through the use of a Del
lying Flagman positioning system. This electronic
t on system operates in the 8 gHz band and is
t ce range limited by hills and by the curvature of the

1 31	Parameter	Sensitivity	Designation on
H :		per mm	digital profile
)	<pre>coaxial inphase (900 Hz) coaxial quad (900 Hz) coplanar inphase (900 Hz) coplanar quad (900 Hz) coplanar quad (900 Hz) coplanar inphase (7200 Hz) coplanar quad (7200 Hz) coaxial monitor altimeter magnetics, coarse magnetics, fine VLF-total: Seattle VLF-total: Seattle VLF-quad: Seattle VLF-quad: Annapolis</pre>	2.5 ppm 2.5 ppm 2.5 ppm 2.5 ppm 5.0 ppm 5.0 ppm 3 m 10 nT 2 nT 2% 2% 2% 2%	CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPQ (900 Hz) CPI (7200 Hz) CPQ (7200 Hz) CXS ALT MAG

Table A-1. The Analog Profiles

Table A-2. The Digital Profiles

1		Observed parameters	Scale units/mm
)))))))	Hz) Hz) Hz) Hz) Hz) Hz) Hz)	magnetics bird height vertical coaxial coil-pair inphase vertical coaxial coil-pair quadrature ambient noise monitor (coaxial receiver) horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature	20 nT 6 m 2 ppm 2 ppm 2 ppm 2 ppm 2 ppm 2 ppm 2 ppm 2 ppm
		Computed Parameters	
	Hz) Hz) Hz) Hz) Hz) Hz)	difference function inphase from CXI and CPI difference function quadrature from CXQ and CPQ conductance log resistivity log resistivity apparent depth apparent depth	2 ppm 2 ppm 1 grade .06 decade .06 decade 6 m 6 m

Flying Flagman uses two ground based transponder which transmit distance information back to the r. The onboard Central Processing Unit then takes distances and determines the helicopter position to the two ground stations. This is accomplished y second. The ground stations are set up well away survey area and are positioned such that the ross the survey blocks at an angle between 30° and ifter site selection, a baseline is flown at right to a line drawn through the transmitter sites to an arbitrary coordinate system for the survey The distance from each ground transmitter site

range-range data is transposed during data i g into an arbitrary x-y coordinate system based on etion of the two transmitter sites. This x-y grid eterred to the base map by correlating a number of topographical features to the navigational data The use of numerous visual tie-in points serves two to correct for distortions in the photomosaic (if d to accurately relate the navigation data to the map

Εi

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APPENDIX B

EM ANOMALY LIST

	AOC 90	XIAL 00 HZ	COPI 90	LANAR DO HZ	COPI 72(LANAR 00 Hz	. VER	FICAL IKE	. HORI:	zontal Eet	CONDUC EAR	CTIVE Fh
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e S	2M	PPM	PPM	PPM	PPM	PPM	. MHOS	M DEPTH-	. MHOS	M	OHM-M	M
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	•	•	•				•	•	•			
	(F	LIGHT	r 5)).			•		•			
	0	4	0	7	8	49	. 1	0	. 1	40	771	0
							•		•			
	1)	12 12	Ľ 5, 1) 10	20	132	•	n	•	22	563	0
	-1	9	0	15	29	96	. 1	1	. 1	26	591	0
	•	-	•				•		•			·
	(F	LIGHT	r 5))			•		•			
	0	7	1	11	17	70	. 1	0	. 1	39	550	0
	1	5	1	8	20	51	. 1	0	. 1	45	576	0
	(16	T.TCH	r 5	۱			•		• '			
	0	4	2	, 8	15	55	. 1	0	• • 1	63	435	5
	0	4	1	4	3	26	. 1	0	. 1	10	817	Ō
	4	4	5	5	10	23	. 8	32	. 1	107	126	60
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	(F	LIGH	r 5)	•	47	•		•		1005	•
	-4	4 3	-9	4 3	-2	17	. 3	23	. 1	. 111	1035	0
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	(F	LIGHT	r 5)			•		•			•
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	-2	9	-2	14	20	114	. 1	1	. 1	26	584	0
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	1.5	אין איז	r 5	\			•	• .	•			
	1	6	1	, 12	22	79	. 1	0	. 1	31	687	0
	0	3	Ó	. 7	15	22		6	. 1	43	751	0
				• • •			•		•			
	(F	LIGH	r 5)			•		•			
	1	6	2	12	29	64	. 1	0	. 1	21	440	0
	0	4	1	9	18	58	. 1	0	. 1	34	562	0
	(1	71.1681	r s	۱			•		•			
	0	2	0	, 3	10	24	. 1	0	. 1	9	748	0
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											•	
	MAI	TED DI	EPTH	MAY B	E UNR	ELIABL	E BECA	USE THE	STRON	GER PA	RT.	
	HE.	CONDI	JCTOR	MAY	HE DE	EPER O	клоо	NE SIDE	OF TH	E FLIG	нт .	

1. OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

•

	:0A 90	XIAL)0 HZ	COPI 90	LANAR 00 Hz	COP1 720	LANAR DO HZ	•	VER1 D	FICAL IKE	•	HORI: SHI	zontal Eet	CONDUC EAR	CTIVE FH
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	• •	4 7	1	12	01 A	50	•	1	י ח	٠	1	30	675	0
	U	'	1	13	4	20	:	1	U	•	1	21	041	v
	(F	LIGH	r 5)			•			•				
	2	5	1	9	21	12		2	12	•	1	64	843	0
	-3	9	-3	16	15	118	•	1	1	•	1	30	606	0
	1	4	2	8	36	47		-1	17		1	19	479	0
	0	4	-1	2	21	21	•	1	25	•	1	28	614	0
	0	8	0	13	19	31	•	1	2	•	1	29	622	Ö
	/ 5	אד דרשי	n 6	、			٠			٠				
	۰۱ ۱	10	L 0	, 16	23	119	•	1	1	•	1	27	614	0
	n	.0	0	12	12	20	•	1	1		1	40	713	Ő
	Ŭ	5		÷		20	•	•	•		•	40		Ŭ
	(F	LIGH	r 5)			•							
	0	4	0	6	10	29	•	1	0	•	1	61	819	0
	-2	5	-4	7	2	26	•	1	0	٠	1	58	806	0
	<i>.</i>			、			•			٠				
	(ピ - つ	LIGH.	L' 4.) c	ć	F 2	•	1	2	•	1	E A	702	0
	-2	ວ າ	-1	10 20	0	52	•	1	14	•	1	24 02	703 054	0
	- <u>2</u>	2	1	2	17 17	13	•	1	14	٠	1	00 40	702	0
	1	* *	0	57	14	44	٠	י ר	12	٠		442	103	0
		2	U	'	1.4	. 20	•	2	15		•	52	000	Ŭ
	(F	LIGH	r 5)										
	0	4	-1	6	6	23	•	1	9		1	61	806	0
	0	1	0	3	5	20	•	1	2		1	78	863	1
	2	3	1	3	13	24	•	1	0	•	1	15	497	0
	1 -		n 4	,			•			٠				
	- 2	N DIGU	L . 41 _ 1	, E	11	12	•	1	0	•		51	001	٥
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	(F	LIGH	r 5)										
	1	5	1	6	16	. 39	•	1	0	•	1	45	662	0
	1.	ייסד דכטי	т <i>А</i>	۱			٠			•				
	0		L' *1 1	، د	5	16	•	1	0	•	. 1	60	755	٥
	1	4	1 2	10	. 1E	01	•	1	0	•	1	00	/55	U 0
	י ה	0 7	2	12	13	02 20	•	1	0	•	1	21	830	U 0
	2	21	-0	14	14 2/	24	•	11	17	•	1	24	543	· 0
	J	21	- 5	4	64	64	•		17	•	•	64		v
r 1	MAI	TED D	EPTH	MAY B	E UNR	ELIAB	LE	BECA	USE TH	E	STRON	GER PA	RT .	
: F	ΗE		JCTOR	MAY	BE DE	EPER ()R	70 0	NE SID	E	OF TH	E FLIG	HT .	

 HE CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIGHT .

 I, OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

0A) 90	XIAL 0 HZ	COPL 90	ANAR 0 HZ	VAR COPLANAR . VERTICAL . HORIZONTAL CON HZ 7200 HZ . DIKE . SHEET E		CONDUC	CTIVE CH					
					•	-		•				
E L	QUAD	REAL	QUAD	REAL	QUAD .	COND	DEPTH*	. COND	Depth	RESIS	DEPTH	
$\mathbf{E} \in \mathbf{M}$	PPM	PPM	PPM	PPM	PPM .	MHOS	М	. MHOS	М	онм-м	М	
15	тант	5)			•			•				
2	19	3	31	77	232	1	0	. 1	15	365	0	
1	42	22	22	60	66.	18	5	. 2	64	54	33	
			:		•			•				
(F)	LIGHT	4)			•			•			•	
0	5	1	5	8	44.	1	0	. 1	59	633	0	
1	15	3	27	/2	194 .	i A	U	. 1	10	300	17	
9	10	5		19	88 .		9	. 1	28	516	0	
Ŭ	•	•	Ŭ		••••	•	•	•			•	
(F	LIGHT	5)			-			•				
1	10	3	17	54	125 .	1	0	. 1	22	372	0	
4	4	2	2	4	15 .	6	36	. 1	86	564	9	
1	5	2	9	20	• 73 •	1	0	• 1	35	457	U	
(F	LIGHT	4)			•			•				
1	6	1	8	5	61.	2	16	. 1	68	819	0	
•1	5	0	10	14	46.	1	1	. 1	32	647	0	
0	6	-1	7	6	60 .	. 1	10	. 1	45	721	0	
2	10	2	21	43	153 .	1	0	. 1	18	519	0	
2	6	1	10	25	37.	1	2	. 1	28	630	0	
1	5	0	10	20	77.	1	0	. 1	33	707	U O	
J	3	U	4	Ø			U	• •	13	912	U	
(F	LIGHT	5)	•••					•				
-1	13	1	21	35	161 .	1	2	. 1	21	539	0	
0	7	-1	10	6	46.	. 1	2	. 1	31	693	0	
0	5	-1	7	12	60 .	1	0	. 1	41	741	0	
1	3	1	4	7	30.	1	0	• 1	11	841	0	
(ፑ	т.тснт	<u>ل</u> ه ا			•			•				
1	4	2	6	16	42	່ 1	13	. 1	57	554	0	
1	3	1	- 5	12	30 .	. 1	0	. 1	18	744	0	
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(F	T.TGHT	י מו			•	,		•				
0	3		6	13	43	. 1	0	. 1	55	794	0	
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(F	LIGHT	2 5))				•	•			•	
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HE	CONDU	CTOR	MAY	BE DE	EPER OF	10 0	NE SIDE	OF TH	E FLIG	HT .	•	
1., 0	R BEC	CAUSE	OF A	SHAL	LOW DI	PORO	VERBURI	EN EFF	ECTS.	•		

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02	AXIAL	COPI	LANAR	COPI	LANAR	•	VER!	TICAL	•	HORIZ	CONTAL	CONDUC	CTIVE
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E L	QUAD	REAL	QUAD	REAL	QUAD	. (COND	DEPTH	۰.	COND	DEPTH	RESIS	DEPTH
E M	PPM	PPM	PPM	PPM	PPM	. 1	HOS	М	•	MHOS	M	ОНМ-М	М
						•			٠		•		
(1	FLIGHT	5)	1	25	62	٠		0	٠		20		0
1	8	1	10	20	63 76	•	1	0	٠	1	30	649	0 0
1	o	4	10	11	70	•	•	U	•	I	40	049	U
()	FLIGHT	4)											
1	3	1	5	6	39	•	1	0		1	7	1101	0
		•.	•••			•			•				
()	FLIGHT	5))			•							
1	5	0	9	9	59	•	1	0	•	1	37	709	0
0	3	1	7	9	• 17	٠	1	0	٠	1	46	705	0
0	2	1	4	7	34	•	1	0	•	1	10	1071	0
0						٠			٠				
2	8 E DIGUI	- 1	12	16	91	•	1	0	•	1	28	640	٥
£ +	U	•	<u>,</u> , , , ,			•	•	. 0	•	•	20	040	v
()	FLIGHT	5))										
-1	· 4	0	7	5	59		1	0	•	1	71	863	0
0	4	2	6	14	23	•	1	0	•	1	32	500	0
0	5	1	7	8	59	•	1	4	•	1	41	582	0
						•			•				
()	FLIGHT	5)) _			٠		•	٠				•
0	4	1	5	11	26	•	1	0	٠	1	76	837	0
0	RT TOUM					•			•				
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()	FLIGHT	5))										
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()	FLIGHT	4)			•			٠				
0	6	0	8	9	63	•	1	0	٠	1	46	761	0
2	2	1	4	11	29	٠	1	0	٠	1	11	945	0
2	7	1	8	15	37	٠	1	8	٠	1	32	649	0
0	9	0	14	19	9	•		0	٠	1	33	685	0
1	FT.TCH	1 5'	\			•			•				
1	۲ <u>۱۱</u> ۹۱۱ ۸		/ 8	15	29	•	1	3	•	1	50	378	4
2	5	2	5	16	30	:	1	0		1	17	574	0
2	4	3	2	16	17		4	38		1	61	318	13
0	6	1	3	16	71		1	0		1	7	800	0
Ő	8	-1	12	5	54		1	2		1	55	763	ů 0
Ţ	-	,		-		•	-	_		-			
()	FLIGHT	4)							,			
0	6	1	10	16	34	•	1	0	•	1	46	633	0
												•	
1 A	TED DE	EPTH 1	MAY B	E UNR	ELIABI	LE .	BECA	USE THI	E	STRON	GER PA	RT .	
' IE	CONDU	ICTOR	MAY	BE DE	EPER (DR '	TO O	NE SID	E D ~	OF TH	E FLIG	HT .	
1 (OR BEC	AUSE.	OF A	SHAL	มเมพ กา	1 1 1	OR O	VERBUR	DE	N EFF	ECTES . "		

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í	ОА 90	XIAL 0 HZ	COPI 90	ANAR)0 HZ	COP1 72	LANAR 00 Hz	•	VER! D	FICAL IKE	•	HORI: SH	zontal Eet	CONDUC EAR	CTIVE FH
.	Ŧ	סמוזס		01130	זגיזמ	OUND	•	COND		•	CONT		DECTO	משמים מ
- L -	ы м	DAUQ	DDM	עמע	DDM	DDW DDW	•	MHOS	M DEFIN	•	MHOS	M	OHW-W	M
Ŧ	14	EEM	E E PI	E E FI		E E E E	•	PAIOD	1.1	•	14100	11		14
	(ም	T.TGHT	י א	•			•			•				
	2	5	, 0	4	13	23	•	2	6	•	1	73	922	0
	•	5	v	•				-	v	•	•		,	v
	(F	TTGHT	۰ 5											
	0	7	. 3, 1	11	15	80		1	0		1	45	607	0
	2	5	2	6	18	33	Ī	2	8	1	1	64	340	13
	0	3	1	5	11	40		1	0		1	67	671	0
	-	-	-	-			•							
	(F	LIGHI	r 4))						•				
	4	6	5	7	17	21		5	23	•	1	99	152	51
	4	7	3	5	19	17	•	4	20	•	· 1	, 71	306	20
	3	7	-4	10	4	81	•	1	3		1	33	647	0
	1	5	-1	10	9	79		1	0		1	41	730	0
							•			•				
	(F	LIGHI	r <u>5</u>)) .			٠			٠				
	1	4	1	6	6	49	٠	1	5	٠	1	57	685	0
	3	8	1	4	6	24	•	2	9	•	1	106	571	19
	6	5	3	4	13	9	•	7	25	٠	1	86	487	13
	<i>.</i>						٠			٠				
	(F	LIGHT	° 4)) 	~	20	٠		•	٠		01	1012	•
	0	4	1	2	5	30	٠	1	0	•	1	21	1013	0
	1	3	4	4	14	5	•	1	0	٠	1	59	165	43
	4	5	5	5	16	19	٠	5	19	٠	1	85	188	35
	170	* ****					٠			•				
	2 (F	V V	· 5/	11	30	25	٠		18	•	1	77	90	37
	5	10	5	12	34	23	•		10	•	1	50	307	57
	1	. 10 	1	9	27	72	•	1	10	•	1	52	700	n n
	•	т		,	Ŭ		•	•			•	52	/00	
	(F	LIGH	r 4)											
	4	7	1	4	9	28		3	20		1	91	563	16
	3	6	5	8	5	24		3	19	•	1	87	118	44
	6	12	7	17	50	58		4	7		· 1	46	193	. 7
	0	3	2	3	10	21		1	19	•	1.	53	506	0
	(F	LIGHT	r 5))						•				
	2	5	2	2	11	12	•	3	19	•	1	103	269	45
	3	3	4	2	7	8	•	8	53	•	1	137	140	85
	2	7	4	10	15	77		2	13	٠	1	50	380	6
	-1.	6	2	8	15	62	•	1	0	•	1	. 34	516	0
	0	7	1	13	12	. 93	•	1	.6	•	1	29	586	0
							٠			•				· ·
	(F	LIGH	r 4))			٠			•				
	8	17	4	17	37	73	•	3	4	٠	1	42	225	4
~	1	ים חסו			-			0000			CUDAN		•	
ل ۲	1A1	נע עבט זיקונרסס	erth l	MAI B	יים מפ מים מפ	EDED (പല്പ വല	BECA	VOE THE	2 7	OP IIII OP IIII	GER PA	um KT •	
,	<u>رتان</u>	עווועט יפפ פו	JUTOK	1941 VD 2	סם טב דג נים	LOW D	JK		NEDDUDI NE DINI	יי קר	or Tri N EPP	r ruigi Rome	GT •	
3	- u	까 모다	CUUSE	Or A	עאמס	υ υη Π.		UR U	A DUDOUL	ער	us DEE	- C L J V		

OR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS. •

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	:0A 90	XIAL 0 HZ	COPI 90	ANAR	COPI 720	LANAR . DO HZ .	VER? D	FICAL IKE	. HORIS	ZONTAL EET	CONDUC EAR	CTIVE TH
3	Ъ	OUAD	REAL	OUAD	REAL	OUAD .	COND	DEPTH*	. COND	DEPTH	RESIS	DEPTH
]	'n	PPM	PPM	PPM	PPM	PPM /.	MHOS	М	. MHOS	М	OHM-M	м
						•			•			
	(F 1	FLIGHT 7	4)	1.4	10	• •	1	0	•	20	563	0
	I	'	2	1 41	10		I	0	• •	20	502	U
	(1	LIGHT	5))		•			•			
	2	1	3	3	б	11 .	1	0	. 1	37	579	7
	15	21	17	21	52	57.	8	7	• 1	68	68	34
	(F	LIGHT	4))		•			•			
	8	12	4	10	28	14.	4	7	. 1	61	302	13
	1	2	2	3	6	8.	2	33	. 1	128	524	33
	(F	LIGHT	51)		•			•			
	6	6	5	10	27	8.	5	10	. 1	64	233	16
						•			• '			
	(F	LIGHT	4)) _	4 -	•- •			• '			
	3	5	3	5	15	27.	3	16	. 1	115	205	59
	2	, D	I	12	11	91.	I	U	• •	49	530	U
	(F	LIGHT	5)			•			•	×.		
	1	2	3	4	11	5.	1	0	. 1	32	737	0
	•1	8	2	17	25	134 .	1	0	. 1	28	503	0
	(1	T.TGHT	4		- ···	•		•	•			
	1	10	1	20	20	157	1	0	•	25	602	0
						•	•		•			
	(1	FLIGHT	5)		•		_	•			
	0	7	1	12	6	94.	1	2	. 1	54	749	0
	(1	LIGHT	4)				,	•			
	2	. 4	2	4	21	21 .	1	0	• 1	59	280	32
	/-			:		•			•.			
	3	5	· 5, 4	, 5	20	13	4	23	•	109	162	58
	Ū	Ū	-			· · · · · · · · · · · ·			•	105	102	
	(1	LIGHT	4))		•			•			
	3	8	1	4	14	10 .	2	16	. 1	142	1035	0
	(1	at toum	. 5	· ·		•			•			
	1	2 STICUL	່ ວຸ ວ	י ז	6	10	1	0	• ·	54	763	10
	3	6	2	3	15	8.	4	16	. 1	110	420	30
	-					•	-		•			
	(1	FLIGHT	4)		•			•			
	0	4	1	7	. 9	50 .	1	0	. 1	56	695	0
tj	'1A'	red de	PTH 1	MAY B	E UNR	ELIABLE	BECA	USE THE	STRON	GER PA	RT	
ę	ΗE	CONDU	CTOR	MAY	BE DE	EPER OR	TO O	NE SIDE	OF TH	E FLIG	HT .	
Ŀ	. (OR BEC	AUSE	OF A	SHAL	LOW DIF	ORO	VERBURD	EN EFF	ECTS.	•	

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900	HZ	COPL 90	ANAR 0 HZ	COPI 720	LANAR DO HZ	•	VER! D	rical Ike	. HORI . SH	zontal Eet	CONDUC EAR	CTIVE FH
	מאוז	DFAT.	חגוזה	DENT.	חגוזס	•	റസ്ഥ	рарын Работна	COND	DRPTH	RESTS	ההסעה
M I	PPM	PPM	PPM	PPM	PPM	•	MHOS	M	. MHOS	M	OHM-M	M N
						•			•			•
(FL	IGHT	4)				•			•			
6	8	9	14	38	25	•	6	0	. 1	64	88	26
•						•			•			
(FL	IGHT	5)		10	400	٠		•	•	25		
2	/	U	14	18	103	•		0	• •	35	5/1	
/	8	6	/	25	14	•	1	4	• •	90	105	5
U	5	I	9	9	37	•	1	U	• 1	00	695	(
(FL)	IGHT	4)				•			•			
0	7	2	7	13	1		1	0	. 1	42	411	· (
3	5	1	3	11	9		3	21	. 1	107	803	
•1	6	Ó	9	12	92		1	0	. 1	41	732	(
0	7	Ő	11	4	87	•	1	0	. 1	50	759	(
						•			•			
(FL	IGHT	5)		•		•			•		•	
1	5	1	7	8	11	•	1	0	• 1	54	732	(
2	1	2	2	15	3	•	1	0	. 1	48	534	11
0	9	~2	: 16	11	127	•	1	0	• 1	33	678	(
(D T)	TOUM					•			•			
<u>т</u> р. 1	10U1	· · · · · · · · · · · · · · · · · · ·	2	10	10	•	1	0	• •	21	765	
י כ	2 0	- 1	11	10	19	•	1	0	• •	50	700	
°∠ .1	12	-4	21	15	167	•	1	0	• •	25	501	
•		L	~ '	, ,	107	•		Ŭ	• •	25	551	•
(FL	IGHT	5)	•									
0	8	-2	16	12	127	•	1	0	. 1	33	699	(
									•			
(FL	IGHT	· 4)			•				•			
2	7	-2	11	6	91	•	1	0	1	47	752	(
·	_					•			•			
(FL.	IGHT 7	5)	4.4	2		•		,	• •	40	746	
	· ' .	~ 	11	3	90	•	1	4	• 1	49	/40	(
(FL	тсят	5)					·		•			
2	4	1	б	15	39	•	2	9	. 1	57	682	1
-	•	·	•				-	•			002	
(FL	IGHT	4)							•			
0	3	0	6	11	16	•	1	11	1	40	756	(
1 -	10	2	19	39	136	•	1	0	. 1	19	487	(
		-				•			•			
(FL	IGHT	5)		_		•		-	•			
1	3	1	4	. 7	. 12	•	1	0	. 1	27	489	(
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T E CONDUCTOR MAY BE DEEPER OR TO ONE SIDE OF THE FLIG TOR BECAUSE OF A SHALLOW DIP OR OVERBURDEN EFFECTS.

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	IOA	XIAL	COPI	LANAR	COPI	LANAR	•	VER	FICAL	•	HORI	HORIZONTAL		CONDUCTIVE	
	90	JU HZ	90	JO HZ	120	OO HZ	•	D.	LKE	٠	SH	EET.	EAR.	гн	
F	۲	OUAD	REAL	OUAD	REAL	OUAD	•	COND	DEPTH		COND	DEPTH	RESIS	DEPTH	
ŗ	M	PPM	PPM	PPM	PPM	PPM	•	MHOS	М	•	MHOS	М	ОНМ-М	М	
							•			•					
	(F	LIGHT	r 5))			•								
	• 1	12	2	20	40	145	•	1	0	•	1	20	538	0	
							٠			٠					
	()	LIGH	Ľ 4.	, _		~ 7	٠	~	10	•		~~~	0.00	~ /	
	2	4	2	5	11	27	٠	2	18	٠	1	88	266	34	
	1	4	2	/	17	4/	•	1	0	٠	1	54	529	0	
	(7	TITGH	r 5')			•			•					
	·1	4	1	5	13	33		1	0		1	88	801	1	
	Ó	4	0	6	13	46		1	2		1	61	718	Ó	
							•			•					
	(E	LIGHT	r 4))			•								
	2	5	1	8	14	60	•	1	2	•	1	44	754	0	
	_						•			٠					
	(F	LIGHI	r 5))			•		_	٠					
	0	4	1	5	1	31	•	1	0	٠	1	16	534	0	
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