

REPORT NO. 275A

DIGHEMIII SURVEY

OF THE

BLAKELOCK TOWNSHIP AREA

ONTARIO

FOR

MINETA RESOURCES LTD.

BY

DIGHEM SURVEYS & PROCESSING INC.

RECEIVED

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MINING LANDS SECTION

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SUMMARY AND RECOMMENDATIONS

A total of 150 km (93 miles) of survey was flown with the DIGHEM^{III} system in December 1986, on behalf of Mineta Resources Ltd., over a property near Cochrane, Ontario.

The survey outlined several discrete bedrock conductors associated with areas of low resistivity. Most of these anomalies appear to warrant further investigation using appropriate surface exploration techniques. Areas of interest may be assigned priorities for follow-up work on the basis of supporting geological and/or geochemical information.

The area of interest contains several anomalous features, many of which are considered to be of moderate to high priority as exploration targets.

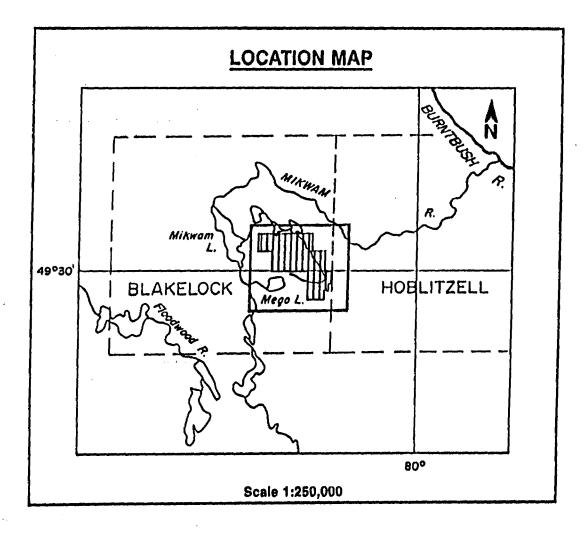


FIGURE 1

THE SURVEY AREA



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A. The Flight Record and Path Recovery

INTRODUCTION

A DIGHEM^{III} electromagnetic/resistivity/magnetic/VLF survey totalling 150 line-km (93 line-miles) was flown with a 100 m line-spacing for Mineta Resources Ltd., in December in the Cochrane area of Ontario (Figure 1).

The Aerospatiale 350B turbine helicopter flew at an average airspeed of 100 km/h with an EM bird height of approximately 30 m. Ancillary equipment consisted of a Sonotek PMH5010 magnetometer with its bird at an average height of 45 m, a Sperry radio altimeter, a Geocam sequence camera, an RMS GR33 digital graphics recorder, a Sonotek SDS1200 digital data acquisition system and a Digidata 1140 800-bpi magnetic tape recorder. 9-track The equipment recorded four channels of EM data at approximately 900 Hz, two channels of EM data at approximately 7200 Hz, four channels of VLF-EM information (total field quadrature components), two ambient EM noise channels (for the coaxial and coplanar receivers). two of magnetics (coarse and fine count), and a channel of radio The digital equipment recorded parameters, with the EM data to a sensitivity of 0.2 ppm at 900 Hz, 0.4 ppm at 7200 Hz, the VLF field to 0.1%, and the magnetic field to one nT (i.e., one gamma).

SECTION I: SURVEY RESULTS

CONDUCTORS IN THE SURVEY AREA

The survey covered a single grid with 150 km of flying, the results of which are shown on one map sheet for each parameter.

The electromagnetic anomaly map shows the anomaly locations with the interpreted conductor type, dip, conductance and depth being indicated by symbols. Direct magnetic correlation is also shown if it exists. The strike direction and length of the conductors are indicated when anomalies can be correlated from line to line. When studying the map sheets for follow-up planning, consult the anomaly listings appended to this report to ensure that none of the conductors are overlooked.

The resistivity map shows the conductive properties of the survey area. Some of the resistivity lows (i.e., conductive areas) coincide with discrete bedrock conductors and others indicate conductive overburden or broad conductive rock units. The resistivity patterns may aid geologic mapping and in extending the length of known zones.

The survey area is dominated by a highly magnetic feature striking east across the northern portion of the survey block. Another strong magnetic high is located near the northern edge of the survey boundary. There are also several less magnetic east-striking features and a weak north-striking high located between lines 10200 and 10240. This north-striking zone may be due to a diabase dike.

VLF responses in this area are very poor due to a thin but highly conductive near surface layer which appears to be channeling the horizontal VLF-EM field. The VLF map (Seattle - 24.8 kHz) provides very little information regarding bedrock conductors and structure.

This survey block contains several bedrock conductors of moderate conductivity thickness. It is strongly recommended that all ground geochemical, geophysical and geological information be used in order to try and differentiate between graphite and sulfide-type responses.

The strongest conductors are located on the south shore of Lake Mikwam. These responses are located on the edge of a magnetic high and may be due to mineralization along a contact. These anomalies should be investigated.

whereas the quadrature component tends to yield crossovers. Both appear as traces on the profile records. The total field data also are filtered digitally and displayed on a contour map, to facilitate the recognition of trends in the rock strata and the interpretation of geologic structure.

The response of the VLF-EM total field filter operator in the frequency domain (Figure II-3) is basically similar that used to produce the enhanced magnetic to (Figure II-2). The two filters are identical along the abscissa but different along the ordinant. The VLF-EM filter removes long wavelengths such as those which reflect regional and wave transmission variations. The filter sharpens short wavelength responses such as those which reflect local geological variations. The filtered total field VLF-EM contour map is produced with a contour interval of one percent.

Respectfully submitted, DIGHEM SURVEYS & PROCESSING INC.

S.J. Kilty Chief Geophysicist

SECTION II: BACKGROUND INFORMATION

Section II provides background information on products which are available from your survey data. Those products not obtained as part of the survey contract may be generated later from raw data which is available on your archive digital tape.

ELECTROMAGNETICS

classes, discrete and broad. The discrete class consists of sharp, well-defined anomalies from discrete conductors such as sulfide lenses and steeply dipping sheets of graphite and sulfides. The broad class consists of wide anomalies from conductors having a large horizontal surface such as flatly dipping graphite or sulfide sheets, saline water-saturated sedimentary formations, conductive overburden and rock, and geothermal zones. A vertical conductive slab with a width of 200 m would straddle these two classes.

The vertical sheet (half plane) is the most common model used for the analysis of discrete conductors. All anomalies plotted on the electromagnetic map are analyzed according to this model. The following section entitled Discrete Conductor Analysis describes this model in detail,

including the effect of using it on anomalies caused by broad conductors such as conductive overburden.

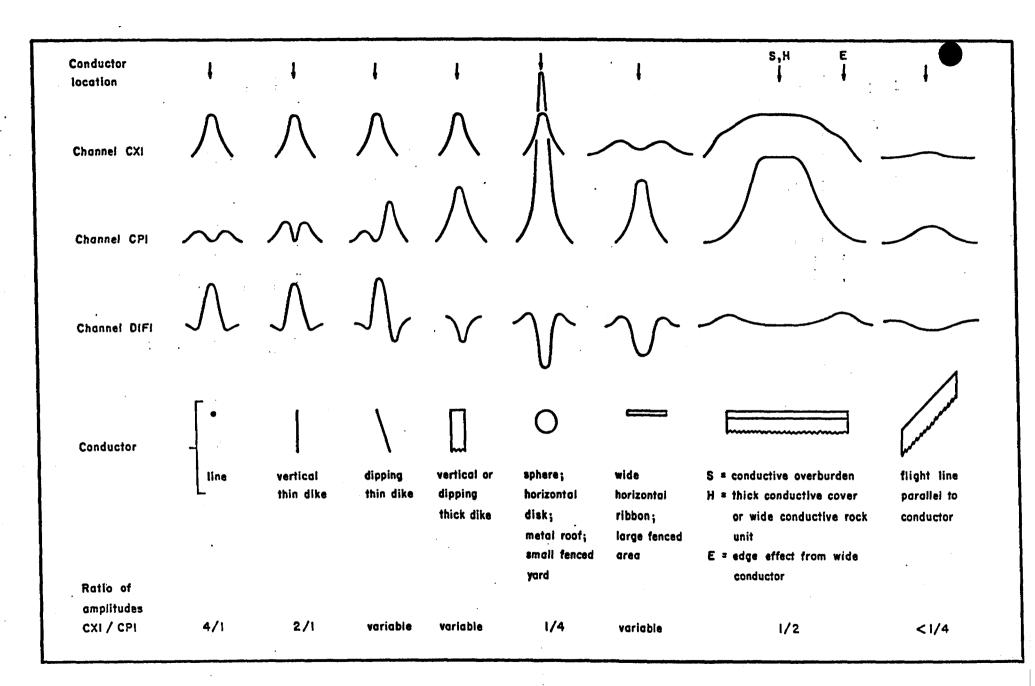
The conductive earth (half space) model is suitable for broad conductors. Resistivity contour maps result from the use of this model. A later section entitled Resistivity Mapping describes the method further, including the effect of using it on anomalies caused by discrete conductors such as sulfide bodies.

Geometric interpretation

The geophysical interpreter attempts to determine the geometric shape and dip of the conductor. Figure II-1 shows typical DIGHEM anomaly shapes which are used to guide the geometric interpretation.

Discrete conductor analysis

The EM anomalies appearing on the electromagnetic map are analyzed by computer to give the conductance (i.e., conductivity-thickness product) in mhos of a vertical sheet model. This is done regardless of the interpreted geometric shape of the conductor. This is not an unreasonable procedure, because the computed conductance increases as the



Typical DIGHEM anomaly shape

electrical quality of the conductor increases, regardless of its true shape. DIGHEM anomalies are divided into six grades of conductance, as shown in Table II-1. The conductance in mhos is the reciprocal of resistance in ohms.

Table II-1. EM Anomaly Grades

Mho Range
> 99
50 - 9 9
20 - 49
10 - 19
5 - 9
< 5

The conductance value is a geological parameter because it is a characteristic of the conductor alone. It generally is independent of frequency, flying height or depth of burial, apart from the averaging over a greater portion of the conductor as height increases. Small anomalies from deeply buried strong conductors are not confused with small anomalies from shallow weak conductors because the former will have larger conductance values.

Conductive overburden generally produces broad EM responses which may not be shown as anomalies on the EM maps. However, patchy conductive overburden in otherwise

This statement is an approximation. DIGHEM, with its short coil separation, tends to yield larger and more accurate conductance values than airborne systems having a larger coil separation.

resistive areas can yield discrete anomalies with a conductance grade (cf. Table II-1) of 1, or even of 2 for conducting clays which have resistivities as low as 50 ohm-m. In areas where ground resistivities can be below 10 ohm-m, anomalies caused by weathering variations and similar causes can have any conductance grade. The anomaly shapes from the multiple coils often allow such conductors to be recognized, and these are indicated by the letters S, H, G and sometimes E on the map (see EM legend).

For bedrock conductors, the higher anomaly grades indicate increasingly higher conductances. Examples: DIGHEM's New Insco copper discovery (Noranda, Canada) yielded a grade 4 anomaly, as did the neighbouring copper-zinc Magusi River ore body; Mattabi (copper-zinc, Sturgeon Lake, Canada) and Whistle (nickel, Sudbury, Canada) gave grade 5; and DIGHEM's Montcalm nickel-copper discovery (Timmins, Canada) yielded a grade 6 anomaly. Graphite and sulfides can span all grades but, in any particular survey area, field work may show that the different grades indicate different types of conductors.

Strong conductors (i.e., grades 5 and 6) are characteristic of massive sulfides or graphite. Moderate conductors (grades 3 and 4) typically reflect graphite or sulfides of a less massive character, while weak bedrock conductors

(grades 1 and 2) can signify poorly connected graphite or heavily disseminated sulfides. Grade 1 conductors may not respond to ground EM equipment using frequencies less than 2000 Hz.

The presence of sphalerite or gangue can result in ore deposits having weak to moderate conductances. As an example, the three million ton lead-zinc deposit of Restigouche Mining Corporation near Bathurst, Canada, yielded a well defined grade 1 conductor. The 10 percent by volume of sphalerite occurs as a coating around the fine grained massive pyrite, thereby inhibiting electrical conduction.

Faults, fractures and shear zones may produce anomalies which typically have low conductances (e.g., grades 1 and 2). Conductive rock formations can yield anomalies of any conductance grade. The conductive materials in such rock formations can be salt water, weathered products such as clays, original depositional clays, and carbonaceous material.

On the electromagnetic map, a letter identifier and an interpretive symbol are plotted beside the EM grade symbol. The horizontal rows of dots, under the interpretive symbol, indicate the anomaly amplitude on the flight record. The

vertical column of dots, under the anomaly letter, gives the estimated depth. In areas where anomalies are crowded, the letter identifiers, interpretive symbols and dots may be obliterated. The EM grade symbols, however, will always be discernible, and the obliterated information can be obtained from the anomaly listing appended to this report.

The purpose of indicating the anomaly amplitude by dots is to provide an estimate of the reliability of the conductance calculation. Thus, a conductance value obtained from a large ppm anomaly (3 or 4 dots) will tend to be accurate whereas one obtained from a small ppm anomaly (no dots) could be quite inaccurate. The absence of amplitude dots indicates that the anomaly from the coaxial coil-pair is 5 ppm or less on both the inphase and quadrature channels. Such small anomalies could reflect a weak conductor at the surface or a stronger conductor at depth. The conductance grade and depth estimate illustrates which of these possibilities fits the recorded data best.

Flight line deviations occasionally yield cases where two anomalies, having similar conductance values but dramatically different depth estimates, occur close together on the same conductor. Such examples illustrate the reliability of the conductance measurement while showing that the depth estimate can be unreliable. There are a

number of factors which can produce an error in the depth estimate, including the averaging of topographic variations by the altimeter, overlying conductive overburden, and the location and attitude of the conductor relative to the flight line. Conductor location and attitude can provide an erroneous depth estimate because the stronger part of the conductor may be deeper or to one side of the flight line, or because it has a shallow dip. A heavy tree cover can also produce errors in depth estimates. This is because the depth estimate is computed as the distance of bird from conductor, minus the altimeter reading. The altimeter can lock onto the top of a dense forest canopy. This situation yields an erroneously large depth estimate but does not affect the conductance estimate.

Dip symbols are used to indicate the direction of dip of conductors. These symbols are used only when the anomaly shapes are unambiguous, which usually requires a fairly resistive environment.

A further interpretation is presented on the EM map by means of the line-to-line correlation of anomalies, which is based on a comparison of anomaly shapes on adjacent lines. This provides conductor axes which may define the geological structure over portions of the survey area. The absence of

conductor axes in an area implies that anomalies could not be correlated from line to line with reasonable confidence.

DIGHEM electromagnetic maps are designed to provide a correct impression of conductor quality by means of the conductance grade symbols. The symbols can stand alone with geology when planning a follow-up program. The actual conductance values are printed in the attached anomaly list for those who wish quantitative data. The anomaly ppm and depth are indicated by inconspicuous dots which should not distract from the conductor patterns, while being helpful to those who wish this information. The map provides an interpretation of conductors in terms of length, strike and dip, geometric shape, conductance, depth, and thickness (see below). The accuracy is comparable to an interpretation from a high quality ground EM survey having the same line spacing.

The attached EM anomaly list provides a tabulation of anomalies in ppm, conductance, and depth for the vertical sheet model. The EM anomaly list also shows the conductance and depth for a thin horizontal sheet (whole plane) model, but only the vertical sheet parameters appear on the EM map. The horizontal sheet model is suitable for a flatly dipping thin bedrock conductor such as a sulfide sheet having a thickness less than 10 m. The list also shows the

resistivity and depth for a conductive earth (half space) model, which is suitable for thicker slabs such as thick conductive overburden. In the EM anomaly list, a depth value of zero for the conductive earth model, in an area of thick cover, warns that the anomaly may be caused by conductive overburden.

Since discrete bodies normally are the targets of EM surveys, local base (or zero) levels are used to compute local anomaly amplitudes. This contrasts with the use of true zero levels which are used to compute true EM amplitudes. Local anomaly amplitudes are shown in the EM anomaly list and these are used to compute the vertical sheet parameters of conductance and depth. Not shown in the EM anomaly list are the true amplitudes which are used to compute the horizontal sheet and conductive earth parameters.

X-type electromagnetic responses

DIGHEM maps contain x-type EM responses in addition to EM anomalies. An x-type response is below the noise threshold of 3 ppm, and reflects one of the following: a weak conductor near the surface, a strong conductor at depth (e.g., 100 to 120 m below surface) or to one side of the flight line, or aerodynamic noise. Those responses that

have the appearance of valid bedrock anomalies on the flight profiles are indicated by appropriate interpretive symbols (see EM map legend). The others probably do not warrant further investigation unless their locations are of considerable geological interest.

The thickness parameter

DIGHEM can provide an indication of the thickness of a steeply dipping conductor. The amplitude of the coplanar anomaly (e.g., CPI channel on the digital profile) increases relative to the coaxial anomaly (e.g., CXI) as the apparent thickness increases, i.e., the thickness in the horizontal (The thickness is equal to the conductor width if the conductor dips at 90 degrees and strikes at right angles to the flight line.) This report refers to a conductor as thin when the thickness is likely to be less than 3 m, and thick when in excess of 10 m. Thick conductors are For base metal indicated on the EM map by crescents. exploration in steeply dipping geology, thick conductors can be high priority targets because many massive sulfide ore bodies are thick, whereas non-economic bedrock conductors are often thin. The system cannot sense the thickness when the strike of the conductor is subparallel to the flight line, when the conductor has a shallow dip, when the anomaly amplitudes are small, or when the resistivity of the environment is below 100 ohm-m.

Resistivity mapping

of widespread conductivity Areas are commonly encountered during surveys. In such areas, anomalies can be generated by decreases of only 5 m in survey altitude as well as by increases in conductivity. The typical flight record in conductive areas is characterized by inphase and quadrature channels which are continuously active. EM peaks reflect either increases in conductivity of the earth or decreases in survey altitude. For such conductive areas, apparent resistivity profiles and contour maps are necessary for the correct interpretation of the airborne The advantage of the resistivity parameter is that anomalies caused by altitude changes are virtually eliminated, so the resistivity data reflect only those anomalies caused by conductivity changes. The resistivity analysis also helps the interpreter to differentiate between conductive trends in the bedrock and those patterns typical of conductive overburden. For example, discrete conductors will generally appear as narrow lows on the contour map and broad conductors (e.g., overburden) will appear as wide lows.

The resistivity profile (see table in Appendix A) and the resistivity contour map present the apparent resistivity using the so-called pseudo-layer (or buried) half space model defined by Fraser (1978)². This model consists of a resistive layer overlying a conductive half space. depth channel (see Appendix A) gives the apparent depth below surface of the conductive material. The apparent depth is simply the apparent thickness of the overlying apparent depth (or thickness) layer. The resistive parameter will be positive when the upper layer is more resistive than the underlying material, in which case the apparent depth may be quite close to the true depth.

The apparent depth will be negative when the upper layer is more conductive than the underlying material, and will be zero when a homogeneous half space exists. The apparent depth parameter must be interpreted cautiously because it will contain any errors which may exist in the measured altitude of the EM bird (e.g., as caused by a dense tree cover). The inputs to the resistivity algorithm are the inphase and quadrature components of the coplanar coil-pair. The outputs are the apparent resistivity of the

Resistivity mapping with an airborne multicoil electromagnetic system: Geophysics, v. 43, p. 144-172.

conductive half space (the source) and the sensor-source distance. The flying height is not an input variable, and the output resistivity and sensor-source distance are independent of the flying height. The apparent depth, discussed above, is simply the sensor-source distance minus the measured altitude or flying height. Consequently, errors in the measured altitude will affect the apparent depth parameter but not the apparent resistivity parameter.

The apparent depth parameter is a useful indicator of simple layering in areas lacking a heavy tree cover. The DIGHEM system has been flown for purposes of permafrost mapping, where positive apparent depths were used as a measure of permafrost thickness. However, little quantitative use has been made of negative apparent depths because the absolute value of the negative depth is not a measure of the thickness of the conductive upper layer and, therefore, is not meaningful physically. Qualitatively, a negative apparent depth estimate usually shows that the EM anomaly is caused by conductive overburden. Consequently, the apparent depth channel can be of significant help in distinguishing between overburden and bedrock conductors.

The resistivity map often yields more useful information on conductivity distributions than the EM map. In

comparing the EM and resistivity maps, keep in mind the following:

- (a) The resistivity map portrays the absolute value
 of the earth's resistivity.
 (Resistivity = 1/conductivity.)
- (b) The EM map portrays anomalies in the earth's resistivity. An anomaly by definition is a change from the norm and so the EM map displays anomalies, (i) over narrow, conductive bodies and (ii) over the boundary zone between two wide formations of differing conductivity.

The resistivity map might be likened to a total field map and the EM map to a horizontal gradient in the direction of flight³. Because gradient maps are usually more sensitive than total field maps, the EM map therefore is to be preferred in resistive areas. However, in conductive areas, the absolute character of the resistivity map usually causes it to be more useful than the EM map.

³ The gradient analogy is only valid with regard to the identification of anomalous locations.

Interpretation in conductive environments

Environments having background resistivities below 30 ohm-m cause all airborne EM systems to yield very large responses from the conductive ground. This usually prohibits the recognition of discrete bedrock conductors. The processing of DIGHEM data, however, produces six channels which contribute significantly to the recognition of bedrock conductors. These are the inphase and quadrature difference channels (DIFI and DIFQ), and the resistivity and depth channels (RES and DP) for each coplanar frequency; see table in Appendix A.

The EM difference channels (DIFI and DIFQ) eliminate up to 99% of the response of conductive ground, leaving responses from bedrock conductors, cultural features (e.g., telephone lines, fences, etc.) and edge effects. An edge effect arises when the conductivity of the ground suddenly changes, and this is a source of geologic noise. While edge effects yield anomalies on the EM difference channels, they do not produce resistivity anomalies. Consequently, the resistivity channel aids in eliminating anomalies due to edge effects. On the other hand, resistivity anomalies will coincide with the most highly conductive sections of conductive ground, and this is another source of geologic

noise. The recognition of a bedrock conductor in a conductive environment therefore is based on the anomalous responses of the two difference channels (DIFI and DIFQ) and the two resistivity channels (RES). The most favourable situation is where anomalies coincide on all four channels.

The DP channels, which give the apparent depth to the conductive material, also help to determine whether a conductive response arises from surficial material or from a conductive zone in the bedrock. When these channels ride above the zero level on the digital profiles (i.e., depth is negative), it implies that the EM and resistivity profiles are responding primarily to a conductive upper layer, i.e., conductive overburden. If both DP channels are below the zero level, it indicates that a resistive upper layer exists, and this usually implies the existence of a bedrock conductor. If the low frequency DP channel is below the zero level and the high frequency DP is above, this suggests that a bedrock conductor occurs beneath conductive cover.

The conductance channel CDT identifies discrete conductors which have been selected by computer for appraisal by the geophysicist. Some of these automatically

selected anomalies on channel CDT are discarded by the geophysicist. The automatic selection algorithm is intentionally oversensitive to assure that no meaningful responses are missed. The interpreter then classifies the anomalies according to their source and eliminates those that are not substantiated by the data, such as those arising from geologic or aerodynamic noise.

Reduction of geologic noise

geophysical Geologic noise refers unwanted to responses. For purposes of airborne EM surveying, geologic noise refers to EM responses caused by conductive overburden and magnetic permeability. It was mentioned above that the EM difference channels (i.e., channel DIFI for inphase and DIFQ for quadrature) tend to eliminate the response of conductive overburden. This marked a unique development in airborne EM technology, as DIGHEM is the only EM system which yields channels having an exceptionally high degree of immunity to conductive overburden.

Magnetite produces a form of geological noise on the inphase channels of all EM systems. Rocks containing less than 1% magnetite can yield negative inphase anomalies caused by magnetic permeability. When magnetite is widely

distributed throughout a survey area, the inphase EM channels may continuously rise and fall reflecting variations in the magnetite percentage, flying height, and overburden thickness. This can lead to difficulties in recognizing deeply buried bedrock conductors, particularly if conductive overburden also exists. However, the response of broadly distributed magnetite generally vanishes on the inphase difference channel DIFI. This feature can be a significant aid in the recognition of conductors which occur in rocks containing accessory magnetite.

EM magnetite mapping

The information content of DIGHEM data consists of a combination of conductive eddy current response and magnetic permeability response. The secondary field resulting from conductive eddy current flow is frequency-dependent and consists of both inphase and quadrature components, which are positive in sign. On the other hand, the secondary field resulting from magnetic permeability is independent of frequency and consists of only an inphase component which is negative in sign. When magnetic permeability manifests itself by decreasing the measured amount of positive inphase, its presence may be difficult to recognize. However, when it manifests itself by yielding a negative

inphase anomaly (e.g., in the absence of eddy current flow), its presence is assured. In this latter case, the negative component can be used to estimate the percent magnetite content.

A magnetite mapping technique was developed for the coplanar coil-pair of DIGHEM. The technique yields channel FEO (see Appendix A) which displays apparent weight percent magnetite according to a homogeneous half space model.4 method can be complementary to magnetometer mapping in Compared to magnetometry, it is far less certain cases. sensitive but is more able to resolve closely spaced magnetite zones, as well as providing an estimate of the amount of magnetite in the rock. The method is sensitive to 1/4% magnetite by weight when the EM sensor is at a height of 30 m above a magnetitic half space. It can individually resolve steeply dipping narrow magnetite-rich bands which are separated by 60 m. Unlike magnetometry, the magnetite method is unaffected by remanent magnetism or magnetic latitude.

The EM magnetite mapping technique provides estimates of magnetite content which are usually correct within a

Refer to Fraser, 1981, Magnetite mapping with a multicoil airborne electromagnetic system: Geophysics, v. 46, p. 1579-1594.

factor of 2 when the magnetite is fairly uniformly distributed. EM magnetite maps can be generated when magnetic permeability is evident as indicated by anomalies in the magnetite channel FEO.

Like magnetometry, the EM magnetite method maps only bedrock features, provided that the overburden is characterized by a general lack of magnetite. This contrasts with resistivity mapping which portrays the combined effect of bedrock and overburden.

Recognition of culture

Cultural responses include all EM anomalies caused by man-made metallic objects. Such anomalies may be caused by inductive coupling or current gathering. The concern of the interpreter is to recognize when an EM response is due to culture. Points of consideration used by the interpreter, when coaxial and coplanar coil-pairs are operated at a common frequency, are as follows:

1. Channels CXS and CPS (see Appendix A) measure 50 and 60 Hz radiation. An anomaly on these channels shows that the conductor is radiating cultural power. Such an indication is normally a guarantee that the conduc-

tor is cultural. However, care must be taken to ensure that the conductor is not a geologic body which strikes across a power line, carrying leakage currents.

- 2. A flight which crosses a "line" (e.g., fence, telephone line, etc.) yields a center-peaked coaxial anomaly and an m-shaped coplanar anomaly. When the flight crosses the cultural line at a high angle of intersection, the amplitude ratio of coaxial/coplanar (e.g., CXI/CPI) is 4. Such an EM anomaly can only be caused by a line. The geologic body which yields anomalies most closely resembling a line is the vertically dipping thin dike. Such a body, however, yields an amplitude ratio of 2 rather than 4. Consequently, an m-shaped coplanar anomaly with a CXI/CPI amplitude ratio of 4 is virtually a guarantee that the source is a cultural line.
- 3. A flight which crosses a sphere or horizontal disk yields center-peaked coaxial and coplanar anomalies with a CXI/CPI amplitude ratio (i.e., coaxial/coplanar) of 1/4. In the absence of geologic bodies of this geometry, the most likely conductor is a metal roof or

⁵ See Figure II-1 presented earlier.

small fenced yard.⁶ Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.

- 4. A flight which crosses a horizontal rectangular body or wide ribbon yields an m-shaped coaxial anomaly and a center-peaked coplanar anomaly. In the absence of geologic bodies of this geometry, the most likely conductor is a large fenced area. Anomalies of this type are virtually certain to be cultural if they occur in an area of culture.
- 5. EM anomalies which coincide with culture, as seen on the camera film, are usually caused by culture. However, care is taken with such coincidences because a geologic conductor could occur beneath a fence, for example. In this example, the fence would be expected to yield an m-shaped coplanar anomaly as in case #2 above. If, instead, a center-peaked coplanar anomaly occurred, there would be concern that a thick geologic conductor coincided with the cultural line.

⁶ It is a characteristic of EM that geometrically identical anomalies are obtained from: (1) a planar conductor, and (2) a wire which forms a loop having dimensions identical to the perimeter of the equivalent planar conductor.

above description of anomaly shapes is valid 6. when the culture is not conductively coupled to the environment. In this case, the anomalies arise from inductive coupling to the EM transmitter. However. when the environment is quite conductive (e.g., less than 100 ohm-m at 900 Hz), the cultural conductor may be conductively coupled to the environment. latter case, the anomaly shapes tend to be governed by Current gathering can completely current gathering. distort the anomaly shapes, thereby complicating the identification of cultural anomalies. In such circumstances, the interpreter can only rely on the radiation channels CXS and CPS, and on the camera film.

TOTAL FIELD MAGNETICS

The existence of a magnetic correlation with an EM anomaly is indicated directly on the EM map. An EM anomaly with magnetic correlation has a greater likelihood of being produced by sulfides than one that is non-magnetic. However, sulfide ore bodies may be non-magnetic (e.g., the Kidd Creek deposit near Timmins, Canada) as well as magnetic (e.g., the Mattabi deposit near Sturgeon Lake, Canada).

The magnetometer data are digitally recorded the aircraft to an accuracy of one nT (i.e., one gamma). The digital tape is processed by computer to yield a total field magnetic contour map. When warranted, the magnetic data also may be treated mathematically to enhance the magnetic response of the near-surface geology, and an enhanced magnetic contour map is then produced. The response of the enhancement operator in the frequency domain is illustrated in Figure II-2. This figure shows that the passband components of the airborne data are amplified 20 times by the enhancement operator. This means, for example, that a 100 nT anomaly on the enhanced map reflects a 5 nT anomaly for the passband components of the airborne data.

The enhanced map, which bears a resemblance to a downward continuation map, is produced by the digital bandpass filtering of the total field data. The enhancement is equivalent to continuing the field downward to a level (above the source) which is 1/20th of the actual sensor-source distance.

Because the enhanced magnetic map bears a resemblance to a ground magnetic map, it simplifies the recognition of trends in the rock strata and the interpretation of

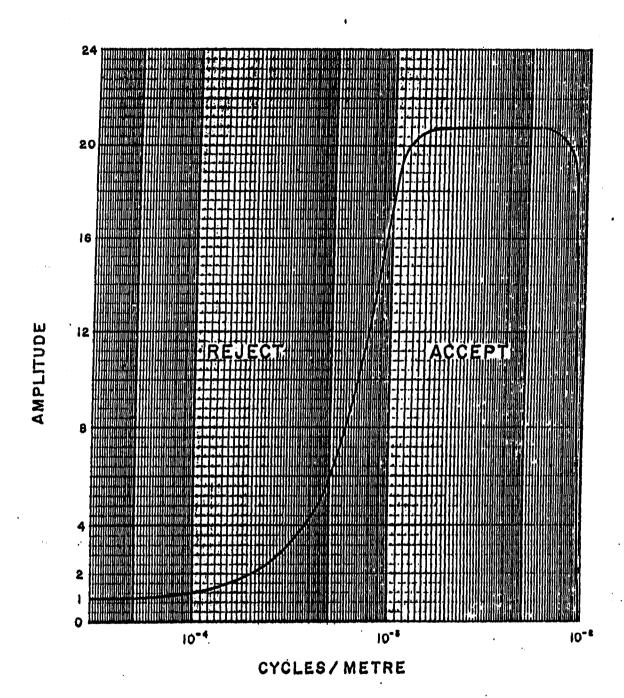


Figure 2 Frequency response of magnetic operator.

geological structure. It defines the near-surface local geology while de-emphasizing deep-seated regional features. It primarily has application when the magnetic rock units are steeply dipping and the earth's field dips in excess of 60 degrees.

VLF-EM

VLF-EM anomalies are not EM anomalies in the conventional sense. EM anomalies primarily reflect eddy currents flowing in conductors which have been energized inductively by the primary field. In contrast, VLF-EM anomalies primarily reflect current gathering, which is a primary field sets up non-inductive phenomenon. The currents which flow weakly in rock and overburden, and these tend to collect in low resistivity zones. Such zones may be due to massive sulfides, shears, river valleys and even unconformities.

The Herz Industries Ltd Totem VLF-electromagnetometer measures the total field and vertical quadrature components. Both these components are digitally recorded in the aircraft with a sensitivity of 0.1 percent. The total field yields peaks over VLF-EM current concentrations

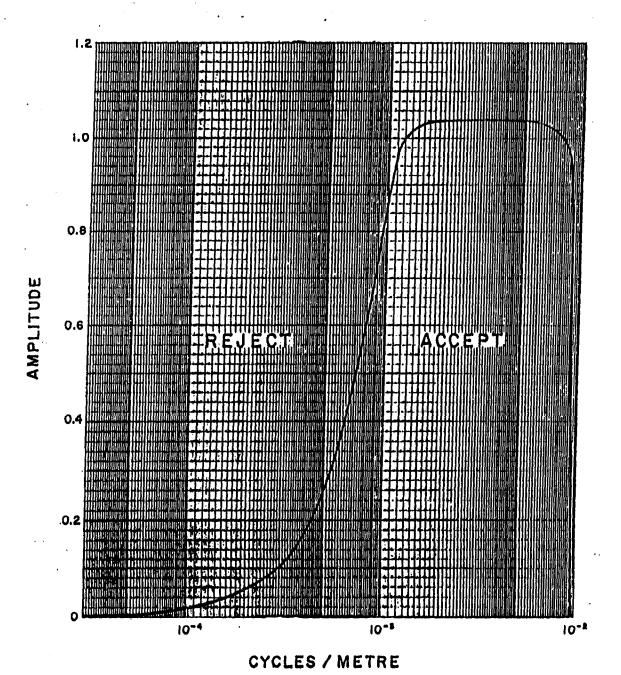


Figure 3 Frequency response of VLF-EM operator.

whereas the quadrature component tends to yield crossovers. Both appear as traces on the profile records. The total field data also are filtered digitally and displayed on a contour map, to facilitate the recognition of trends in the rock strata and the interpretation of geologic structure.

The response of the VLF-EM total field filter operator in the frequency domain (Figure II-3) is basically similar that used to produce the enhanced magnetic (Figure II-2). The two filters are identical along the abscissa but different along the ordinant. The VLF-EM filter removes long wavelengths such as those which reflect regional and wave transmission variations. The filter sharpens short wavelength responses such as those which reflect local geological variations. The filtered total field VLF-EM contour map is produced with a contour interval of one percent.

APPENDIX A

THE FLIGHT RECORDS

Both analog and digital flight records were produced. The analog profiles were recorded on chart paper in the aircraft during the survey. The digital profiles were generated later by computer and plotted on electrostatic chart paper at a scale of 1:15,000. The analog and digital profiles are listed in Tables A-1 and A-2 respectively.

In Table A-2, the log resistivity scale of 0.06 decade/mm means that the resistivity changes by an order of magnitude in 16.5 mm. The resistivities at 0, 33 and 67 mm up from the bottom of the digital flight record are respectively 1, 100 and 10,000 ohm-m.

NAVIGATION EQUIPMENT

Aircraft positioning and post-survey recovery of aircraft position was accomplished through the use of a Del Norte positioning system. This electronic navigation system operates in the UHF band and is therefore not as range limited by hills and by the curvature of the earth, as are the more common radar positioning systems.

Table A-1. The Analog Profiles

Channel Number	Parameter	Sensitivity per mm	Designation on computer profile
CXI CXQ CPI1 CPQ1 CPI2 CPQ2 ALT MAGC MAGF VLFT VLFQ	coaxial inphase (900 Hz) coaxial quad (900 Hz) coplanar inphase (900 Hz) coplanar quad (900 Hz) coplanar inphase (7200 Hz) coplanar quad (7200 Hz) altimeter magnetics, coarse magnetics, fine VLF-total: Seattle VLF-quad: Seattle	2.5 ppm 2.5 ppm 5.0 ppm	CXI (900 Hz) CXQ (900 Hz) CPI (900 Hz) CPQ (900 Hz) CPI (7200 Hz) CPQ (7200 Hz) AL/T MAG

Table A-2. The Digital Profiles

Channel Name (Freq)		Observed parameters	Scale units/mm
	(900 Hz) (900 Hz) (900 Hz) (7200 Hz)	magnetics bird height vertical coaxial coil-pair inphase vertical coaxial coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature horizontal coplanar coil-pair inphase horizontal coplanar coil-pair quadrature	20 nT 6 m 2 ppm
		Computed Parameters	
SIGT	(900 Hz) (900 Hz) (7200 Hz) (900 Hz)	difference function inphase from CXI and CPI difference function quadrature from CXQ and CPQ conductance log resistivity log resistivity apparent depth apparent depth	2 ppm 2 ppm 1 grade .06 decade .06 decade 6 m

AB-SK-464

uses two ground based transponder The Del Norte stations continuously interrogated by the helicopter mounted unit and which transmit distance information back to the helicopter. The onboard Central Processing Unit then takes the two distances and determines the helicopter position relative to the two ground stations. This is accomplished once every second. The ground stations were set up well away from the survey area and were positioned such that the signals crossed the survey blocks at an angle between 30° and 150°. After site selection, the aircraft then flew a baseline at right angles to a line drawn through the transmitter sites. The minimum distance recorded when flying this baseline established the arbitrary coordinate system used to fly the survey area. The final step was to establish the location of the first flight line on the map or photomosaic. This line was then flown while pressing a "start of line" and "end of line" switch, thereby establishing both survey boundaries and line direction. distance from each ground transmitter site (range-range) was continuously recorded digitally.

The range-range data was transposed during data processing into an arbitrary x-y coordinate system based on the location of the two transmitter sites. This x-y grid data was then transferred to the base map by correlating a number of prominent topographical features to the

navigational data points. The use of numerous visual tie-in points served two purposes: to correct for distortions in the photomosaic (if any) and to accurately relate the navigational data to the map sheet.

900

7	12 -1V	f'	WBL	18 .	55
ype c	Surve.				
		ΑT	RBORNE	SUI	RVEY

BLAKELOCK TWP

M-19643 M - 20603

86

Claim Holder(s

Ernest Sicard / E. Mord

Val Gagne, Survey Company

Ont. Connaught, Ont.

DIGHEM SURVEYS & PROCESSORS INC.

Day Mo. Yr.

Total Miles of line Cut

Name and Address of Author (of Geo-Technical report)

Steve Kilty, 228 Matheson Blvd Mississauna

Steve	Kilty, 228 Mat	neson B	31vd. E.,	_Mississauga	a, Ont.	L4Z 1X1		
Credits Requested per Each Claim in Columns at right Mining Claims Traversed (List in numerical sequence)								
Special Provisions	Geophysical	Days per		lining Claim	Expend.	М	ining Claim	Expend.
	Cappinysical	Claim	Prefix .	Number	Days Cr.	Prefix	Number	Days Cr.
For first survey: Enter 40 days, (This	- Electromagnetic		L	860883		L	878341	
includes line cutting)	- Magnetometer			860884			878342	
For each additional survey: using the same grid:	- Radiometric			860885			878343	
Enter 20 days (for each)	- Other			860903			878344	
, <u> </u>	Geological			860904			878345	
	Geochemical			860905			878346	
Man Days	Geophysical	Days per Claim		860906			878347	
Complete reverse side and enter total(s) here	- Electromagnetic			860907			878357	
URDER LAKE MINING DIVSK	netometer			860908			878358	
MEGELLY	S Radiometric			860909			878359	·
	Other			860910			878360	
11 11 DEC 22 19	Geological		Control of the second	860911			878361	
1. wind	Geochemic			860912			878362	
Airborne Credits		Days per Claim		860913			878363	
Note: Special provisions credits do not apply	Electromagnetic	40		860914			878364	
to Airborne Surveys.	Magnetometer	40		860915			878365	
	Radiometric			860916			878366	
Expenditures (excludes power				860917			878367	
Type of Work Performed 4	La G 3 / 3 /			860918			878368	
Performed on Claim(s)				860919			878369	
	1010			860920			878370	
•	MG LANCE JET	(ا		860921			878465	
Calculation of Expenditure Days Total Expenditures	7	otal Credits		860922			878466	
			12 Tan (2) 0. 1	000722	<u> </u>	Section of the	070400	لحسسا
\$ Instructions	continued on page two claims covered by this 73					73		
Total Days Credits may be ap	portioned at the claim h	older's				-	<u></u>	
choice. Enter number of days credits per claim selected For Office Use Only Total Days Cripate Recorded								

Date

in columns at right.

Recorded Holder or Agent (Signature) December 17/86

Certification Verifying Report of Work

Total Days Cr. Date Recorded DEC 2 2

Date Certified

I hereby certify that I have a personal and intimate knowledge of the facts sey forth in the Report of Work annexed hereto, having performed the work or witnessed same during and/or after its completion and the annexed report/is true.

Name and Postal Address of Person Certifying

Maurice Hibbard (agent)

Cedar Hill, Connaught, Ontario PON 1AO

December 17/86

1362 (85/9)

i Kara Barrangaya

1362 (85/9)

837 (85/12)



Ministry of Northern Development and Mines

Geophysical-Geological-Geochemical Technical Data Statement

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KNA	
1110	

TO BE ATTACHED AS AN APPENDIX TO TECHNICAL REPORT FACTS SHOWN HERE NEED NOT BE REPEATED IN REPORT TECHNICAL REPORT MUST CONTAIN INTERPRETATION, CONCLUSIONS ETC.

Type of Survey(s) AIRBGANE	INPUT EN AND MAGNETIC	
Township or Area BUNKELOCK	MINING CLAIMS TRAVERSED	
Claim Holder(s) ERNGT SIG	List numerically	
Survey Company <u>PIGHERI</u> SU Author of Report <u>STEVE KIL</u>		—
Address of Author $M/SS/SS$	_	- L860884 L878341 L87846C
Covering Dates of Survey Nov	-	
Total Miles of Line Cut	(unecutting to ottice)	L860903 L878943 L878472
		1860904 1878344 1878473
SPECIAL PROVISIONS CREDITS REQUESTED	DAYS Geophysical per claim	L860905 L878345 L878474
	-Electromagnetic	L860906 1878346 L878475
ENTER 40 days (includes line cutting) for first	-Magnetometer	L860907 L878347 L878477
survey.	-Radiometric	<u> </u>
ENTER 20 days for each additional survey using	-Other	L860909 L878 358 L878479
same grid.	Geological	
8	Geochemical	L860910 L878359 L878480
AIRBORNE CREDITS (Special provisi		L860911 L878360 L880219
Magnetometer Electromagn (enter da	etic <u>40 </u>	- <u>L860912 L 878361 4880270</u>
DATE: FEB 4 /87 SIGNA	TURE: Moda Co-c Author of Report or Agent	L860913 L87836Z L880221
		- L860914 L878363 L880722
Par Cool Ougliti	(analama	L860915 L978364 L880223
Res. Geol. Qualifi Previous Surveys	cations	- 1860916 1878365 1880226
File No. Type Date	Claim Holder	L860917 L878366 L880227
		L860918 L878367 L880228
		L860919 L878368 L880229
		L860920 L878369 L883 678
		1860921 L878370 L883679
		TOTAL CLAIMS 73

GEOPHYSICAL TECHNICAL DATA

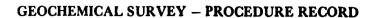
GROUND SURVEYS - If more than one survey, specify data for each type of survey

N	Number of Stations	Number	of Readings	· · · · · · · · · · · · · · · · · · ·
	station interval			
	rofile scale	_	_	
	ontour interval	77.44		
	_			
9	Instrument			
MAGNETIC	Accuracy - Scale constant			
	Diurnal correction method			
	Base Station check-in interval (hours)			
	Base Station location and value			
ان ان	Instrument			
ELECTROMAGNETIC	Coil configuration			
	Coil separation			
WA	Accuracy			
R	Method: Fixed transmitter	Shoot back	☐ In line	☐ Parallel line
CC	Frequency			
国	Parameters measured	(specify V.L.F. station)		
	Talameters measured			
	Instrument			
	Scale constant			
IX	Corrections made			
RAVITY				
SR	Base station value and location			
	Elevation accuracy			
	Instrument			
	Method		Frequency Domain	
	Parameters - On time		Frequency	
H	- Off time		Range	
XII	- Delay time			
STI	— Integration time			
RESISTIVITY	Power			
~	Electrode array			
	Electrode spacing			
•	Type of electrode			

INDUCED POLARIZATION

SELF POTENTIAL	
Instrument	Range
Survey Method	
Corrections made	
RADIOMETRIC	
Instrument	
Values measured	
Energy windows (levels)	
Height of instrument	Background Count
Size of detector	
Overburden(type, depth - include outcrop	
(type, depth — include outcrop	map)
OTHERS (SEISMIC, DRILL WELL LOGGING ETC.)	
Type of survey	
Instrument	
Accuracy	
Parameters measured	
Additional information (for understanding results)	
AIRBORNE SURVEYS	
Type of survey(s) INPUT MKNT ELECTROMAGNETIC MAGNET	TIC SURVEY
Instrument(s) Sonote & PM+15010 Magnetometer an RMS GRS	33 digited graphics recorder a Sonotek SDS170
	y of 0,2 ppm at 900 Hz, 0,4ppm at 7200 Hz
Aircraft used Aster 350Df whine Italicapter	
Sensor altitude 30 metres	
Navigation and flight path recovery method Del Norte N	avigation
Aircraft altitude 40 metres	Line Spacing 100 model
Miles flown over total area 150 line 1cm	Over claims only all clasms

) icd





Total Number of Samples	ANALYTICAL METHODS
Type of Sample(Nature of Material) Average Sample Weight	p. p. m. 🖳
Method of Collection	
	Cu, Pb, Zn, Ni, Co, Ag, Mo, As, (circle)
Soil Horizon Sampled	Others
Horizon Development	
Sample Depth	Extraction Method
Terrain	·
	Reagents Used
Drainage Development.	Field Laboratory Analysis
Estimated Range of Overburden Thickness	No. (tests)
	Extraction Method
	Analytical Method
	Reagents Used
SAMPLE PREPARATION (Includes drying, screening, crushing, ashing)	Commercial Laboratory (tests)
Mesh size of fraction used for analysis	Name of Laboratory
	Extraction Method
	Analytical Method
	Reagents Used
General	General



Ministry of Northern Development and Mines

Geophysical-Geological-Geochemical Technical Data Statement

File	
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		1.000 0.00		
• •			INPUT EM AM HAGNE	310
Township or Area BLAICELOCK TOLNSIMP				MINING CLAIMS TRAVERSED
Claim Hold	der(s)		· · · · · · · · · · · · · · · · · · ·	List numerically
	<u> </u>	· · · · · · · · · · · · · · · · · · ·		Pg 2 Commontaging application of
Survey Con	mpany			
Author of	Report			(prefix) (number) (
	-			
Covering D	ates of Surv	/ey	(linecutting to office)	[88368]
	s of Line Cu		(linecutting to office)	L883 682
Total Willes	s of Line Cu			L883683
SPECIAL	L PROVISIO	ONS	DAYS	L883684
	S REQUES		Geophysical per claim	
			-Electromagnetic	L883 685
	40 days (inc		Magnetometer	1883686
line cutti survey.	ing) for first		-Radiometric	
•	20 days for	an ah	-Other	
	20 uays 101 al survey usi		Geological	
same grid	•	· ·	Geochemical	-
AIDDODAG	E CDEDITE			
		- Electromag	ision credits do not apply to airborne surve neticRadiometric days per claim)	***************************************
DATE:		SIGNA	ATURE:	
			Author of Report or Ager	ent
Res. Geol		Quali	fications	
Previous Su		~		
File No.	Туре	Date	Claim Holder	
••••••••		,	***************************************	

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•••••••			•••••••••••••••••••••••••••••••	77
				TOTAL CLAIMS_73

