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REPORT

ON

PHASE I

GEOPHYSICAL SURVEYS

SAVANT LAKE PROJECT

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FORM NO, LED HILP REPORT PARER - GRAND & TOV

June 8, 1977

W. G. Wahl Limited

Suite 1101, 302 Bay Street Toronto, Ontario M5H 2P3

June 8, 1977

Dr. E. L. Evans Director of Exploration Denison Mines Limited 4 King Street West Toronto, Ontario M5H 1C2

Dear Dr. Evans:

Submitted herewith is our report on:

PHASE I

GEOPHYSICAL SURVEYS

SAVANT LAKE PROJECT

Three major, apparently stratabound, conductive zones were located. The conductivity-thickness products are suggestive of sulphide mineralization. Depth to top estimates for these zones from electromagnetic and correlated magnetic data range from 45 to 80 metres. As this range is just near the threshold of detectability at 120m coil spacing for an horizontal-loop survey, it is recommended that selected areas be re-surveyed at 180m coil spacing utilizing all frequencies in order to screen some of the anomalies. With the sources at such depths, there may not be any surficial geologic expression and consequently a judicious diamond drilling programme is necessary.

INTRODUCTION

1

LOCATION AND ACCESS

The Savant Lake Claim Block is located on the eastern shore of Savant Lake, approximately 70 air miles northeast of Sioux Lookout, Ontario, and is most easily accessible by charter aircraft from Sioux Lookout. It lies within NTS Nos. 52 J/8 and 9.

GENERAL

A combined electromagnetic horizontal-loop and totalfield gradient magnetometer survey was conducted on the Savant Lake property in order to delineate the source of massive sulphide float found to the south of the property.

The interpreted results are subject to revision as more geological information is made available from field mapping in progress.

ELECTROMAGNETIC - HORIZONTAL-LOOP SURVEY

The electromagnetic data was observed employing an Apex MaxMin II horizontal-loop unit in the maximum-coupled mode. The transmitter frequencies were 444 Hz and 1777 Hz utilizing a 120 metre coil separation. Three distinctly anomalous zones were detected during the course of the survey. These anomalous areas have been designated as Anomaly A, ..nomaly B and Anomaly C for discussion purposes only, and no priority significance should be attached to their order of discussion.

ANOMALY A

Anomaly A is a northeast trending conductor having a strike length of about 500 metres. It is situated between lines 900 east and 1500 east and lies between stations 750 north and 795 north. The negative peak value (relative to background) ratios of the in-phase to quadrature (or out-of-phase) anomaly amplitudes average about 1.0 at 444 Hz and 0.55 at 1777 Hz. The decrease of this ratio with frequency, rather than the normally expected increase, indicates a complex-conductivity situation.

The low frequency depth to top estimate is 0.5 times the coil spacing or 60 metres, while at the high frequency it is 0.3 times the coil spacing or about 35 metres.

The width of the zone, as determined by the amount in excess of the coil spacing from the 0 cross-over points of the anomaly, varies from thin to about 30 metres. The variations in conductivity-thickness, along strike, only I artially correlate to the variations in the thickness, thereby implying a heterogeneity to the zone.

ANOMALY B

Anomaly B is a northeast trending conductor having a strike length of about 600 metres. It is situated from line 960 east to line 1560 east and lies between stations 420 north and 510 north (this is assuming a continuous conductor across a claim not at present held by W. G. Wahl Limited). The discontinuous magnetic high associated with this zone suggests that it may be a pyrrhotite containing conductor. The estimated depth to top is 0.35 times the coil separation or 42 metres at 444 Hz and 0.15 times the coil separation or 18 metres at 1777 Hz. The zone differs in thickness from 15 metres to 45 metres. The maximum and minimum conductivity-thickness products are not always coincident with the maximum and minimum thicknesses of the conductor, thereby implying a heterogeneity to the zone.

ANOMALY C

Anomaly C consists of a north-south trending conductor that extends from line 420 east to line 1720 east. The estimated conductivity-thickness products decrease with an increase in frequency. The conductivity-thickness values are 36 mhos to 120 mhos at 444 Hz and 20 mhos to 40 mhos at 1777 Hz. The depth to top estimates at both frequencies average about 0.55 times the coil separation or 66 metres. On average, the zone appears to be less than 30 metres thick. As this conductor has a high conductivity-thickness product, appears to lie within the magnetically mapped felsic intrusive, and is found near the edge of swamp, it presents an enigma.

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MAGNETOMETER SURVEY

The magnetometer survey was conducted using a geoMetrics G-816 total-field magnetometer in the gradient configuration. Stations were occupied every 15 metres and two readings per station were obtained.

Based on the contoured magnetic expression, the survey area can easily be divided into two major units. The southeast portion has a low, broad uniform magnetic expression and correlates to the regionally mapped felsic intrusive. The northwest portion consists of a series of narrow, parallel, continuous to lensoid, northeast-trending magnetic highs, 2,000 gammas and greater, above a background of 59,000 gammas. This zone reflects the regionally mapped mafic to intermediate volcan.cs.

CORM FIG. 112 ATT P. REPORT PAPER - GRANDA TO-

The sequence of the mafic to intermediate volcanics consists predominantly of pillowed lavas and subsidiary fine- to medium-grained massive flows of the same composition with local pillow breccias, flow-top breccias, amygdaloidal flows, porphyritic flows, volcanic conglomerate, lapilli-tuff, tuff-breccia, crystal tuff, variolitic flows and metasedimentary iron formation present. The different rock types imply quite a range in the magnetic expressions.

A very strong magnetic anomaly (up to 10,000 gammas above background) trends north to northwest between lines 1260 east and 1380 east. This anomaly is also coincident with a strong positive in-phase electromagnetic anomaly and consequently must be due to magnetite.

Magnetic lenses in excess of 3,000 gammas above

background are correlated to conductors in quite a few cases. These are probably pyrrhotite lenses and/or more iron-rich parts of some horizons. Zones in excess of 3,000 gammas are thought ' to reflect flows and pillowed lavas. Zones in the 1,500 to 1,900 gamma range are probably tuffs or more felsic components.

Over the lake, the estimated depth to top varies from 15 metres to 30 metres or so. The station spacing of every 15 metres, however, precludes any more precise determinations. The depth to top estimates associated with the conductors is at least 30 metres in most cases and as such, is comparable to the electromagnetic interpretations, within limits of resolution.

Off-sets and abrupt terminations of some of the magnetic trends suggest at least two east-west trending faults. The sense of the off-sets in both cases is left-lateral.

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CONCLUSIONS

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The causative body for Anonaly A may be a relatively narrow, stratabound zone of sulphide mineralization. As the estimated conductivity-thickness product is moderate, the composition of the zone could be of a more disseminated nature and/or sphalerite and galena rich.

The causative body of Anomaly B may be a zone of sulphide mineralization containing pyrrhotite as evidenced by sporadic high magnetic lenses. The lensoid magnetic expression and variable conductivity-thickness product also reflect the heterogeneity of the zone.

Anomaly C is physically correlatable to an edge of a swamp, but it does give a distinct in-phase anomaly rather than a quadrature anomaly resulting in conductivity-thickness of up to 120 mhos. As it is parallel to the trend of the other conductors and is at the stratigraphic bottom and physically at depth, it warrants further investigation.

RECOMMENDATIONS

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As all the anomalies appear to be at depths of 45 metres or more, and the present geophysical expression is weak, selected areas should be re-surveyed at 180 metre coil separation utilizing all four transmitter frequencies in order to better screen or define the nature of these anomalies. Since these zones are deep, and may not have a surficial expression, diamond drilling is a must.

All of which is respectfully submitted.

Yours very truly,

W. G. WAHL LIMITED

A. Gubins, B.A.Sc.

June 8, 1977



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REPORT ON

PHASE II

GEOLOGICAL SURVEY

SAVANT LAKE PROJECT

November 9, 1977

W. G. Wahl Limited

FORM MOLLETALL P. REPORT PAMER - GRAND& TOV

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Suite 1101, 302 Bay Street Toronto, Ontario M5H 2P3

November 9, 1977

Dr. E. L. Evans Director of Exploration Denison Mines Limited P. O. Box 40 Royal Bank Plaza Toronto, Ontario M5J 2K2

Dear Dr. Evans:

Submitted herewith is our report entitled:

PHASE II

GEOLOGICAL SURVEY

SAVANT LAKE PROJECT

The geologic mapping determined that both an epigenetic environment and a syngenetic depositional environment occurred on the Savant Lake property.

Epigenetic mineralization in the form of pyrite, chalcopyrite, sphalerite and galena was noted in quartz veins. Although no surficial expression of syngenetic sulphide mineralization was found, the presence of iron formation suggests an environment favourable to the deposition of massive sulphide type mineralization.



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INTRODUCTION

LOCATION AND ACCESS

The Savant Lake Claim Block is located on the eastern shore of Savant Lake, approximately 70 air miles northeast of Sioux Lookout, Ontario, and is most easily accessible by charter aircraft from Sioux Lookout. It lies within NTS Nos. 52 J/8 and 9.

GENERAL

The geological mapping programme discussed in this report and completed by W. G. Wahl Limited in June, 1977, was part of an integrated geophysical, geochemical and geological programme initiated to evaluate the economic potential of the Savant Lake area.

The geophysical portion of this programme was completed during March, 1977. At that time, conductive zones were delineated on the basis of the MaxMin II horizontal loop electromagnetic data. The highly variable magnetic responses reported for the study area, and the lack of geological data, precluded a detailed assessment of the geophysical data until the geologic mapping had been completed. Both the electromagnetic and magnetic data were clarified by the geologic mapping.

The results of the geochemical work will be reported at a later date.



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GENERAL GEOLOGY

Archean mafic and intermediate rocks of basalticandesitic composition are present as massive fine-grained pillowed lavas, chlorite schists and medium-grained tuffs. These rock types represent the majority of the volcanic rocks in the study area. Discontinuous magnetite-rich iron formations are present and represent the only reliable marker horizon within the otherwise massive mafic volcanics. Numerous quartz veins of varying width and mineralization also occur. The other major rock type, also Archean in age, is a large felsic (granodiorite) intrusive, restricted to the east and southeast portion of the study area.

Mineralization, in some cases significant, is predominantly associated with the quartz veins, and minor concentrations of disseminated pyrite were noted in the mafic metavolcanic rocks.



GEOLOGY

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GENERAL

On a regional scope the rocks of the Savant Lake property are representative of the Mafic to Intermediate Metavolcanic and Foliated Felsic Intrusive rocks, which are shown in their proper sequence in Table 1. The remainder of the rock types outlined in the table were not encountered on the property although minor felsic metavolcanics and metasedimentary rock units were reported from diamond drilling to the north and northwest of the property.

Pleistocene and Recent features dominate sections of the study area and are represented by till, ribbed moraines and muskeg swamps.

The geology of the area as mapped by the Ontario Department of Natural Resources has been published on or in the following maps and reports:

Moore, E. S., Lake Savant Area, District of Thunder Bay; Ont. Dept. Mines, Vol. 37, Pt. 4, p. 53 - 82, 1928 (Pub. 1929). Accomp. by Map 37j, 1 in. = 2 mi.

Skinner, R., Geology of the Sioux Lookout Map Area, Ontario, on Part of the Superior Province of the Precambrian Shield (52J); Can. Geol. Surv. Pap. 68-45, 10 p. Accomp. by Map 14-1968, 1 in. = 4 mi.

Prelim. Geol. Map. P.962, Pashkokogan-Caribou Lakes Sheet, Dist. of Thunder Bay, 1 in. = 2 mi.

In order to relate the geology of the Savant Lake property to the regional geologic setting, reference was made to Ontario Ministry of Natural Resources Geoscience Report 160, "Geology of McCubbin, Poisson, and McGillis Townships",

TABLE 1 TABLE OF LITHOLOGIC UNITS FOR McCUBBIN, POISSON AND McGILLIS TOWNSHIPS, (from ODM Geosc. Rpt. 160, 1977)

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PHANEROZOIC CENOZOIC

OUATERNARY

Swamp and stream deposits (unconsolidated) PLEISTOCENE

Silt, sand, gravel and boulders (unconsolidated)

UNCONFORMITY

PRECAMBRIAN LATE PRECAMBRIAN MAFIC INTRUSIVE ROCKS Diabamb

INTRUSIVE CONTACT

EARLY PRECAMBRIAN (ARCHEAN) FELSIC INTRUSIVE ROCKS MASSIVE FELSIC INTRUSIVE ROCKS

Porphyritic granodiorite; granodiorite, quartz monzonite; aplite

dikes

INTRUSIVE CONTACT

FOLIATED FELSIC INTRUSIVE ROCKS Granodiorite to quartz monzonite, granodiorite to quartz monzonite with biotite-rich xenoliths, hybrid granitic rocks

INTRUSIVE CONTACT

METAMORPHOSED FELSIC INTRUSIVE ROCKS HERON LAKE STOCK Trondhjemite; quartz-rich_t.ondhjemite; quartz-poor trondhjemite; quartz, quartz-feldspar porphyry^b

INTRUSIVE CONTACT

MAFIC INTRUSIVE ROCKS

Metagabbro^a; metadiorite^a, porphyritic metagabbro^{ab}

INTRUSIVE CONTACT

METASEDIMENTS

ARENACEOUS METASEDIMENTS

Greywacke, siltstone; tuffaceous metasediments; siliceous sandstone, in part tuffaceous metasediments, subgreywacke sandstone

FERRUGINOUS METASEDIMENTS

Greywacke, subgreywacke and sandstone; siltstone; quartz-magnetite iron formation; tuffaceous metasediments; jasper chertmagnetite iron formation

CONGLOMERATIC METASEDIMENTS Polymictic orthoconglomerate with metavolcanic clasts; polymictic gravitud boulder-bearing orthoconglomerate

FAULT CONTACT

METAVOLCANICS

FELSIC TO INTERMEDIATE METAVOLCANICS

Rhyodacitic to rhyolitic metavolcanics; dacitic to andesitic metavolcanics; tuff; lapilli-tuff; tuff-braccia to agglomerate; fine-to medium-grained flows, porphyritic flows; quartz-carbonate schist; reworked tuff

MAFIC TO INTERMEDIATE METAVOLCANICS

Chloritic metavolcanics; actinolitic metavolcanics, pillowed lavas; pillow breccia; flow-top breccia; fine-to medium-grained flows^C; crystal tuff; iron formation; volcanic conglomerate; variolitic flows

- a occurs as small dikes or sills
- b occurs as small dikes only
- c may be intrusive in part

d - may be extrusive in part

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(Savant Lake Area), District of Thunder Bay, by W. D. Bond, 1977. The three townships covered in this report lie immediately to the south and southwest of Savant Township.

ROCK TYPES

Mafic - Intermediate Metavolcanics

The mafic to intermediate metavolcanic rocks of the study area are restricted to the northwest portion of the area. Basaltic to andesitic pillowed lavas and subsidiary fine- to medium-grained flows occur, the basalts being most common. Bond (1977) reports local occurrences of pillow breccia, flow top breccia, amygdaloidal flows, porphyritic flows, tuff breccia, crystal tuff and variolitic flows present in the mafic metavolcanic sequences to the south of the Savant Lake property. Bring of a restricted distribution, these features, although not identified on the property, may be present but hidden by glacial material.

In addition to the massive flows, a chlorite schist and a "tuffaceous" unit were mapped. The chlorite schist is fine-grained and highly schistose, reflecting the high chlorite content. The "tuffaceous" unit is medium-grained and schistosity is only poorly developed. The massive pillowed lavas and flows are fine-grained and massive. Schistosity is only developed along flow boundaries. Both the chlorite schist and "tuffaceous" unit are present as discontinuous lenses of varying thickness, thereby preventing surface delineation of these units over any distance.

The chemical data of the mafic metavolcanic samples demonstrate that both basaltic and andesitic members predominate in the survey area (Fig. 1). Figures 2 to 4 illustrate that both of these rock types occur as flows, tuffs and schists, with the basaltic varieties accounting for the greater









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proportion of samples collected.

In the southeast portion of the property, these metavolcanics are intruded by the intrusive felsic unit.

Iron Formations

Interbedded with the mafic and intermediate metavolcanic sequence are magnetite-rich iron formations. These units are typically highly magnetic, fine-grained and well laminated, and range from less than 5 cm up to 1 metre in thickness. The iron formations, although numerous, were found to have a limited spatial distribution in the stratigraphic sequence on the property. Although surface exposure of the iron formation was similarly intermittent due to glacial cover, the magnetic data proved a useful tool in extending the surface mapping of the distribution of the iron formations.

Chemically, the iron formations sampled compare well with other Precambrian iron formations in that the Al, Na, K, Ti, P, Ca, Mg and Sr concentrations are low. Iron (as Fe) content of the samples ranges between 32.9% and 41.7%.

Foliated Felsic Intrusive Rocks

The felsic intrusion is located in the eastern and southeastern portion of the study area. Based on mineralogical studies (Bond, 1977), the felsic unit is granodioritic to quartz monzonitic in composition. Chemical data of samples collected on the property, illustrated in Fig. 1 , further support this rock type classification. The unit is massive and weathers light pink to white. It is generally mediumgrained with coarser grained sections.

The contact of the granodiorite with the host mafic metavolcanics is marked by an altered sheared zone of granodiorite. This zone is characterized by alteration of the plagioclase feldspars, fractured quartz phenocrysts and a predominance of sericite with minor chlorite. Chemical data of samples from this contact zone are in close agreement with the data of the unaltered granodiorite (Fig.5). Mineralogically, the decrease in Na and K occurs due to the alteration of the plagioclase and some of the potassic feldspars. Sericite is the dominant alteration product. The sheared zone is . foliated parallel to the contact with quartz grains being rounded and flattened parallel to foliation, rendering a quartz eye appearance. Minor sulphide mineralization in the form of pyrite and arsenopyrite is present. Alteration of the mafic metavolcanic host rock is marked by an increased zone of chloritization. The width of the altered zone, including both sheared granodiorite and altered host rocks, is variable but is generally about 30 metres. Stringers of the granodiorite are noted in the mafic metavolcanic units near the granodiorite contact.

Quartz Veins

The quartz veins sampled on the property vary from unaltered, highly siliceous and unmineralized to intensely altered, mineralized types. Figure 5 illustrates the position of the quartz veins relative to both the altered and unaltered granodiorite. No distinction between the altered and unaltered quartz veins is possible on this figure. This is due to the high silica content of both these groups, implying that any





variations of the remaining major elements are of such a minor nature that their overall effect is small. Separation is, however, possible using the trace elements, as outlined in Table 2. On Map 1 the quartz veins have been separated into groups Q_1 and Q_2 , based on degree of visible alteration and mineralization. Massive blebs, stringers and disseminated forms of mineralization are present and consist of pyrite, pyrrhotite, chalcopyrite, galena and minor sphalerite. Significant quantities of silver are also reported. Alteration in the form of sericite and chlorite in the contact zone between the quartz and host rock is variable, being most prominent about the Q_1 mineralized quartz veins to minor about the Q_2 mineralized quartz veins. Dissemination of sulphides into the wall lock is minor.

One of the mineralized quartz veins contained 7.06% copper, while another contained zinc and lead values of 4.95% and 4.03% respectively. Barren quartz veins contained copper, lead and zinc values of 6 ppm, 17 ppm and ND respectively.

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Both Q_1 and Q_2 quartz veins are extensively clustered about fault traces and are most numerous where the structure is most complex, as in the area bounded by lines 1080 east and 1140 east between the base line and 200 south.

Unlike the Q_1 veins, the Q_2 veins appear to be more stratigraphically controlled and there is no distinguishable mineralogical alteration halo within the host metavolcanics.

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TRACE ELEMENTS, QUARTZ VEINS

TABLE 2.

STRUCTURE

Schistosity

Schistosity (S_0) is poorly developed for the majority of the mafic sequences observed. The pillowed lavas and massive flows are for the most part featureless. Only at the contact between flows is S_0 developed. The chlorite schist, and to a lesser extent the "tuffaceous" units, exhibit well develoted schistosity. The majority of the schistosity measured is a set of the schistosity measured for the otherwise massive units, especially where marker horizons such as the iron formation units were not present. Dips of the measured foliations were steep, varying + 20^o from vertical.

Folding

Micro-folds of carbonate and quartz lenses within the mafic metavolcanics, contorted configuration of the mapped iron formations and absence of a predominant S₁ cleavage indicate that the area of study has been subjected to one major period of folding, with possible minor secondary periods of folding. These features support the observations of Bond (1977), who reports only one major period of folding, having a northeastscathwest axis. Pe also reports that the tops of pillowed lavas to the southeast of the property indicate that the mafic sequences are overturned and are now facing northwest. This observation was confirmed on the property, where pillow tops indicate a "stratigraphic up" direction to the northwest. The "tuffaceous" units and schistosity supports a similar trend of folds observed by Bond (1977).

Faulting

At least two periods of faulting can be identified on the property. The first period, confined to the mafic sequences, represents a major stage of uplift faulting, while the second, post intrusive stage, is typified by limited displacement strike-slip and uplift faulting.

The first period of faulting, having a general northwest-southeast trend, and being perpendicular to the major northeast-southwest fold axis, may therefor be related to, but later than, the period of folding. Although strike-slip movement may have occurred, the primary movement is believed to be that of uplift. The abrupt termination of rock types, and greater representation of varied rock types, on opposite sides of the fault, would tend to support this. The fault traces within the mafic sequences are also readily apparent on the magnetic data map, Map 2. The fault locations further explain the abrupt termination of the electromagnetic conductors.

The second period of faulting, having a northeastsouthwest trend, is parallel to both the major Savant Lake fault and the major fold axis. The second set of faults is best illustrated in the felsic intrusion. On Map 1 only major traceable lineations within the granodiorite intrusion are marked. The northeast-southwest lineament of the granodiorite outcrop, expressed as ridges in the field, in places exceeding 10 metres in elevation, indicates an extensive period of faulting. These faults that are so pronounced in the granodiorite, may be present in the mafic sequence but are not as readily apparent. Minor shearing and silicification were noted along the fault traces. The lack of significant remobilization or introduction of fluids along the fault traces is supported by the continuity of magnetic data across the traces. Remobilization along the first period's traces may have occurred, but the relative movement in either the northwestsoutheast or northeast-southwest direction is believed minor because the contact zone between mafic and granodiorite units is generally regular.

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The predominance of chlorite, and to a lesser degree, sericite, indicates a lower greenschist regional metamorphic grade in the area of study.

Bond (1977) reports local development of amphibolitegrade contact metamorphism along the boundary of the larger intrusive bodies to the south. Contact metamorphism of this grade was not observed in the study area.



GEOLOGIC HISTORY

The geologic history of the Savant Lake area, and more precisely the area within the claim block studied, is complex, involving a variety of geologic events. The inferred history of the property is outlined below, from youngest to oldest, relative to the present.



The following discussion reverses the sequence given in the above table.

The predominance of massive flows and pillowed lavas implies a period of "quiet" volcanism intermittently interrupted by more explosive periods, indicated by the "tuffaceous" units.

Periods of quiescence are known to have occurred
during the deposition of the mafic sequence, evidenced by the presence of iron formations. The rapid interchange of stratigraphic horizons in the vicinity of the iron formations, however, indicated that this quiescence was interrupted by periods of volcanism, in some cases explosive in nature.

The next major geologic event to affect the area of study was a period of folding occurring about a northeastsouthwest trending axis that resulted in the development of very steep to isoclinal to overturned folds. This folding in turn resulted in the development of schistosity essentially parallel to bedding. Contemporaneous with the folding, regional metamorphism left primary structures of the sedimentary iron formations intact, permitting bedding - schistosity correlations.

Faulting accompanied the folding of the mafic sequence on the property and is marked by major northwest-southeast traces. Introduction of mineralized quartz veining accompanied the initial faulting stage, resulting in the hydrothermal alteration of the surrounding host rocks. The degree and extent of alteration is variable and is most extensive where clusters of veins occur. The period of faulting culminated with a major uplift, thrusting taking place along earlier formed fault fractures.

The next major geologic event to affect the study area was the intrusion of the granodiorite stock. The ensuing contact metamorphism resulted in intensified chloritization and sericitization of the wall rock at the contact. The intrusive contact is best illustrated by the abrupt termination of the iron formation units of the central grid area against the granodiorite. Subsequent faulting of the granodiorite stock may have resulted from either cooling or later periods of deformation. Some movement along the fault traces may have occurred although displacement along the granodiorite-related faults is not believed to be substantial.

Following the second period of faulting, a second set of unmineralized quartz veins (Q_2) were emplaced. A post granodiorite intrusion emplacement of these veins is suggested by their identification in the granodiorite mass (sample 123). These veins are similar to those described by Bond (1977), who believed them to be metamorphic in origin.

The effect of the granodiorite intrusion, faulting and introduction of the Q_2 veins on the Q_1 veins was extensive. The combination of these three events resulted in the intense shearing of the original mineralized Q_1 veins, giving them a false "pyroclastic" texture. Alteration is reflected in intense sericitization with minor chlorite.

Glaciation was the next event that affected the prospect area, resulting in the deposition of a coarse boulder till of varying thickness. Ribbed (washboard) moraines, located immediately to the east of the property, and glacial striations, indicate a south-southwest direction of ice movement. Numerous pieces of mineralized float material were located and subjected to chemical analysis. The chemical data indicates that at least two depositional environments are represented by the float, a hydrothermal environment similar to that present for the quartz veins, and a chemical precipitate type environment typical of the volcanogenic massive sulphide class. A more in-depth study of the mineralized float will be discussed in the forthcoming report on the geochemistry of the area. Recent deposits of bogs and swamps are restricted to the north, east and south of the property and it is only along these property margins that the recent deposits become the major geographic feature.

All of which is respectfully submitted.

Yours very truly,

W. G. WAHL LIMITED

John L. Wahl, Ph.D. Consulting Geologist

JLW/mhc

November 9, 1977



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PHASE III

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ROCK GEOCHEMICAL PROGRAMME

SAVANT LAKE PROJECT

November 23, 1977

FORM NO. 1474(1.P. HEPORT PAPER - GRAPO 4 TOV

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W. G. Wahl Limited

Suite 1101, 302 Bay Street Toronto, Ontario M5H 2P3

November 23, 1977

Dr. E. L. Evans Director of Exploration Denison Mines Limited P. O. Box 40 Royal Bank Plaza Toronto, Ontario M5J 2K2

Dear Dr. Evans:

Submitted herewith is our report entitled

PHASE III

ROCK GEOCHEMICAL PROGRAMME

SAVANT LAKE PROJECT

The rock geochemical programme identified the granodiorite intrusive and the Q_1 quartz veins as sources of hydrothermal alteration, and the existence of patterns characteristically associated with massive sulphide mineralization of the sedimentary volcanogenic type. The hydrothermal alteration was identifed on the basis of major element concentration variations.

Two prominent massive sulphide-related alteration patterns were identified on the property. Pepresentations of areas of alteration that occupy both proximal and distal positions relative to the primary vent area, the most favourable location for sulphide deposition, were recognized. The absence of any pronounced trace element alteration within the major element alteration zone, especially in the proximal location, indicates that the primary vent is located at depth. This supports the geophysical data where deep plunging sources are indicated. From the rock geochemical data a depth to the source of alteration is believed to be not in excess of 120 metres.

The mineralized lead-silver float was found to be similar to the Q_1 quartz veins distributed throughout the

claim block in both degree of alteration and mineralization.

It is recommended that the geophysical conductors be re-surveyed with the MaxMin II electromagnetic system, employing various frequencies and cable lengths to arrive at a more quantitative estimation of the source geometry.

It is further recommended that the Q, quartz vein material already collected be subjected to additional analytical determinations, for gold. If significant amounts of gold were to be found, further field evaluation of the quartz vein systems located on the property would be warranted.

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LIST OF MAPS

(in back pocket)

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MAP 2. ROCK GEOCHEMISTRY AND GEOLOGY, 1" = 400'



INTRODUCTION

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LOCATION AND ACCESS

The Savant Lake Claim Block is located on the eastern shore of Savant Lake, approximately 70 air miles northeast of Sioux Lookout, Ontario, and is most easily accessible by charter aircraft from Sioux Lookout. It lies within NTS Nos. 52 J/8 and 9.

GENERAL

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The rock geochemical programme discussed in this report, completed by W. G. Wahl Limited in 1977, was part of an integrated geophysical, geological and geochemical programme initiated to evaluate the economic potential of the Savant Lake area.

The rock geochemical survey carried out over the Savant Lake property and surrounding area involved the collection of 174 rock samples, the locations of which are given on Map 1 (refer to Appendix I for Sample Collection and Preparation). Each sample was crushed and ground to -200 mesh, and subjected to a complete HF-HClO₄-HNO₃ acid digestion (refer to Appendix II for Sample Digestion and Instrument Parameters). The resulting sample solutions were analysed on a Plasma Torch Emission System from which element determinations for Al, Ti, Fe, Na, K, Ca, Mg, Mn, P, Cu, Pb, Zn, Co, Cr, Zr, V, Sr, Cd and Be were made. Sample variation and reproducibility during the analytical period was monitored by the ingestion of standard



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samples at the beginning, throughout and at the end of the run (refer to Appendix III for Sample Variation and Reproducibility). A total of 3,298 units of analytical data were acquired and are tabulated as to rock type in Appendix IV.

GENERAL GEOLOGY

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Archean mafic and intermediate volcanic rocks of basaltic-andesitic composition are present as massive finegrained pillowed lavas, chlorite schists and medium-grained tuffs. These rock types represent the majority of the volcanic rocks in the study area. Discontinuous magnetite-rich iron formations are present and represent the only reliable marker horizon within the otherwise massive mafic volcanics. Numerous quartz veins of varying width and mineralization also occur. The other major rock type, also Archean in age, is a large felsic (granodiorite) intrusive, restricted to the east and southeast portion of the study area.

Mineralization, in some cases significant, is predominantly associated with the quartz veins, and minor concentrations of disseminated pyrite were noted in the mafic volcanic rocks.



ROCK GEOCHEMISTRY

GENERAL

The employment of rock geochemistry for exploration is based on the ability of this technique to detect specific element distribution patterns within rocks that can be directly to the presence of ore forming processes. Maximum chemical variations occur from processes that result in the formation of new minerals and/or the destruction of pre-existing minerals. Of the ore forming processes, the movement of hot hydrothermal fluids are known to result in extensive mineralogic changes, a fact that is reflected in the chemistry of the rocks through which the fluids pass. Formation or destruction of minerals requires the movement either into or out of the system of the major rock forming elements Ca, Na, K, Mg, Fe, Al and Ti. Geochemical studies of hydrothermal alteration have demonstrated that of these elements, Al and Ti remain virtually unchanged while Ca and Na characteristically decrease in concentration and Mg, Fe and K increase in concentration. Mineralogically the decrease in Na and Ca is reflected in the destruction of feldspars, predominantly the plagioclase, while the increase in K, Mg and Fe is reflected predominantly in the development of chlorite and sericite. The recognition of these general characteristic distribution patterns provides a means for a rapid testing of chemical data derived from rock samples. Plotting of unknown data on a triangular diagram with apices of Na + Ca, Mg + Fe, and K, would show those samples that have been intensely altered by hydrothermal activity to lie in the Mg + Fe field

while those samples exhibiting unaltered or regional character would be shown to lie outside this range.

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The distribution of the Savant Lake basaltic and andesitic samples on such a diagram are given on Figures 2 and 3. It is apparent from the distribution that the majority of both the basaltic and andesitic rock units, within and about the immediate vicinity of the property, fall in a tightly grouped area of the plots equivalent to unaltered or regional samples. The inherently low concentration of K in mafic volcanic rocks is reflected in the clustering of the samples along the Mg + Fe and Na + Ca side of the plot. Despite the concentration of samples within the unaltered range there is also a pronounced scattering of samples trending from the unaltered range toward and terminating in the intensely altered field. Such a scattered distribution is strongly suggestive of an increasing degree of hydrothermal alteration.

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Alteration is further indicated by the respective major element frequency percent histograms given on Figures 4 through 7 inclusive. For those major elements, Fe, Mg and K, that typically increase with proximity to an alteration zone, the distributions are marked by pronounced negative skews for Fe and K, for both rock types. The Mg population is also negatively skewed, but the skew is less pronounced than that of the Fe and K. These distributions are indicative of an addition in element concentrations due to secondary processes. For the major elements that decrease with proximity to an alteration zone, Ca and Na, a positive skewed distribution is apparent. The degree of skewness is significantly greater for Ca and Na.



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FIGURE 4

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FIGURE 5

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FIGURE 6

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To determine the areal distribution of those samples that reflect weak to intense alteration, the Mg + Fe contribution to the triangular plot for each sample was plotted on the surface geologic map in an attempt to delineate any distinguishable patterns and to possibly explain the cause of the apparent alteration (Map 2).

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QUARTZ VEIN RELATED ALTERATION

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The surface geologic mapping carried out over the study area revealed that two distinct quartz vein systems were present and that each system represented a separate mode of formation. Segregation into groups was made on degree of visible mineralogic alteration and amount of mineralization. The earliest emplaced group of quartz veins, designated Q_1 , were highly mineralized and exhibited a high degree of mineralogic alteration. The second or later group of quartz veins, designated Q_2 , were generally unaltered and carried only minor amounts of ore elements. The mafic wall rocks enclosing the Q_1 veins exhibited pronounced mineralogic alteration whereas the mafic wall rock enclosing the Q2 veins exhibited no detectable mineralogic alteration features. On the basis of the noted differences between the two quartz vein systems a hydrothermal origin of the Q_1 veins was suggested, while the mode of formation of the Q_{2} veins was suggested to be a result of regional metamorphic processes.

The rock geochemistry of the mafic wall rocks reaffirms these hypotheses. Two supportive examples are given by the geochemical data derived from the mafic wall rock enclosing the quartz veins represented by sample 27 and sample 8. These two examples were chosen on the basis of their location, that is, remote from any other possible source of alteration. Sample 27 is representative of a Q_2 unmineralized vein and the geochemical data derived from the mafic wall rock (sample 28) exhibits no major element alteration characteristics. Sample 8, however, represents a Q_1 mineralized vein within an enclosing mafic wall rock (sample 9) which exhibits an intense degree of major element variations characteristically associated with hydrothermal alteration. The area of influence of the hydrothermal activity associated with the Q_1 veins is not believed to exceed 30 metres on either side of the vein. The recognition of altered mafic rock associated with the quartz veins explains several of the samples classified as weakly and intensely altered based on major element concentrations identified on the property.

GRANODIORITE RELATED ALTERATION

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The granodiorite intrusive results in further hydrothermal alteration of the mafic volcanic rocks identified on the property. Samples 121 and 125 were collected in close proximity to the granodiorite contact and both exhibit an intense degree of alteration based on major element concentrations. With increasing distance from the contact, for example sample 118, degree of alteration decreases. The lack of any detectable degree of alteration in samples 25, 30 and 31, located in the vicinity of the suspected northern intrusive contact implies that hydrothermal activity associated with the granodiorite intrusive does not extend any significant distance into the wall rocks. This conclusion supports the geologic findings where only minimal host rock mineralogic alteration attributable to the granodiorite intrusion was noted.

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The minimal apparent hydrothermal alteration associated with the granodiorite intrusion indicates that the large anomalous mineralogic and geochemical zone about the Q_1 and Q_2 quartz vein concentration located between Line 1020 East and Line 1080 East and the Base Line and Line 200 South is not due to any significant degree to the granodiorite intrusive but is a result of hydrothermal activity associated with the formation of the Q_1 quartz veins.

MASSIVE SULPHIDE RELATED ALTERATION

PULLE PART

The data discussed in the preceeding sections enabled the source determination of several of the weakly and intensely altered samples identified on the property. However, a significant number of altered samples remain to be explained. Although sporadic anomalous values are expected in any normal population, the concentration in a given area of a significant number of samples which were classified as weakly or intensely altered warrants furt' study. Several such sample concentrations present on the property can not be adequately explained by the presence of either the quartz veins or the granodiorite. It is these anomalous grouping of weakly and intensely altered samples for which a primary sulphide-bearing fumerolic source is suggested. Two areas in particular form the basis of the following discussion.

One of these areas is located in the vicinity of the

main concentration of iron formations centered about Lines 1320 East and 1380 East and north of the Base Line to the lake. The iron formations, by their very nature of formation, mark a horizon along which conditions compatible with sulphide deposition are known to occur. The presence of the iron formation also indicates that significant amounts of iron were present at a time of deposition, the source of which may be volcanic. Further, iron formations provide stratigraphic control in that they are always located stratigraphically above and/or laterally from massive sulphide concentrations. On the basis of geologic data, a stratigraphic younging direction for the map area was determined to be to the northwest. To the south of the iron formation, which would be the footwall rocks by this geologic interpretation, the geochemistry of the mafic volcanic rocks outlines a broad zone of samples that exhibit weak to intense alteration characteristics. In this zone the degree of alteration increases with proximity to the iron formations. To the north of the iron formations the mafic volcanic rocks exhibit regional chemical characteristics. Knowing that primary alteration is present in all cases in the footwall, the geologic interpretation with respect to stratigraphy is confirmed by the geochemistry. The width of the alteration zone delineated does not extend more than 100 metres into the footwall suggesting a position for this area that is distal to the fumerolic vent. Similarly the iron formation located about Line 600 East and the Base Line indicates a younging direction to the southwest by virtue of the alteration being present on the north side of the iron formation.

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The second area of anomalous concentration of altered samples is that area bounded by the two faults running essentially parallel to Lines 720 East and 960 East and by the granodiorite in the southeast and by Savant Lake to the northwest. Dividing this area is a narrow stretch of water that forms the inlet of the small bay located to the south of the property. As will be demonstrated, distinct geochemical, geological and geophysical differences exist between the mafic volcanic rocks located to the northwest and southeast of this stretch of water, suggesting that a stratigraphic break occurs in this area. For these reasons and for ease of discussion, this area of the inferred stratigraphic break will be referred to as "The Narrows Break". The most apparent chemical differences between the north and south blocks is the greater percentage (80%) of the mafic volcanic samples identified to the south of The Narrows Break that exhibit weak to intense alteration characteristics. To the north of The Narrows Break the majority of the samples reflect unaltered or regional characteristics. The presence of a geophysically inferred iron formation located immediately to the north of The Narrows Break also suggests that a change in the depositional history occurs in this area. Further geophysical evidence for a stratigraphic break at The Narrows is implied by the identification of an electromagnetic conductive source following The Narrows Break. The anomalous samples located in the large area to the southeast of The Narrows Break suggest a stratigraphic position that is more proximal to the fumerolic vent than is present under the iron formation. It is therefore suggested that the mafic volcanic sequences located to the

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northwest of The Narrows Break represent hanging wall rocks while those to the southeast of The Narrows Break represent footwall rocks within close proximity to the primary vent. Although the varying degrees of alteration suggest proximity to a zone of alteration the geophysical data indicates the source of the identified conductors to be at a depth of at least 60 metres. The geochemistry supports this interpretation in that the primary vent and the area most likely to be associated with massive sulphides is also at depth. It is known that detectable major element alteration limits exceed those of the associated trace element alteration zone. Characteristically within the trace element alteration zone Cu, Pb, Zn, Co and Cr increase while Ni and Si corease. Erratic high values for those trace elements hat increase characterize the trace element alteration zone resulting in a pronounced negative skew of the respective element frequency distribution. From the frequency percent histograms for these trace elements for both basaltic and andesitic rock types given on Figures 8 through 11 inclusive, no distribution patterns indicative of any significant element introduction is apparent for any of the samples collected. Conversely, for those elements which decrease, Ni and Sr, the decrease in concentration would result in a positive skew distribution. From the histograms only Sr exhibits a distinguishable positive skew which may be explained by Sr substitution for Ca in rock forming processes, in which case Sr would reflect similar alteration characteristics previously noted for Ca. Comparison of the respective Ca and Sr distributions supports this contention.

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FIGURE 8

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FIGURE 10

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FIGURE II

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The presence of massive sulphide mineralization in the vicinity of the study area was confirmed with the identification of numerous pieces of highly mineralized material along the northwest shore of The Narrows Break. The sulphides, predominantly pyrite, are well-bedded and although exhibiting a massive texture were found to possess low conductivity. Fourteen separate pieces of mineralized material were identified in the space of only 60 metres along the shore line. These mineralized samples were not in situ but are not believed to be glacially transported because of their location outside the primary glacial depositional area, absence of other glacial debris, angularity and limited distribution. Similar material was not identified either up or down ice from the area of sulphide concentration. Fragments of massive sulphide-rich material similar in appearance to that found on the beach were identified incorporated within nearby mafic volcanic rocks. The incorporation of sulphide rich material in a stratigraphically younger rock indicates a source of the sulphide material to be from a position immediately lower in the stratigraphic section. All indications, geological, geochemical and geophysical, would suggest that The Narrows Break is the most likely source area. Collection of this previously unknown sulphide material was facilitated by the unusually low water levels experienced in Savant Lake during the period of investigation. The low water levels also enabled sampling of a glacially transported boulder, sample 137, that exhibited prominent sedimentary bedding features. Sample 137 is located approximately 800 metres down ice from the area designated as The Narrows Break. Chemical

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analysis of the boulder proved to be siderite. With the identification of the sedimentary iron formation, the identification of laminated siderite-rich material can be easily reconciled with the massive sulphide genetic model. Deposition of the siderite and the iron formations would follow actual sulphide deposition.

MINERALIZED QUARTZ FLOAT AND QUARTZ VEINS

The final aspect of the geochemical study to be considered in this report deals with the mineralized float (sample 131) located to the south of the property. There has been extensive exploration for the source of this lead-rich material since its discovery in the early 1950's. In the area of study glacial striae and ribbed moraines indicate a southerly direction of ice movement. Utilizing chemical data of this sample and that of the other quartz-rich material collected on the property and plotting it on a triangular plot with apices Al, Mg + Ca and Na + K, Figure 12 illustrates that the float sample 131 falls in the distribution field of the other quartz material. This in turn indicates that the float is quartz vein material similar to that located on the property. In addition, the highly altered texture and significant mineralization suggest that the float boulder is representative of the Q₁ quartz vein variety. The highly variable trace element concentration experienced for the Q_1 veins identified on the property precluded precise designation of the source of the float. The trace element concentration within the float sample does however most approximate that of sample 64.

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CONCLUSIONS

The completion of the rock geochemical survey marks the completion of the first stage of exploration for sulphide mineralization on the Savant Lake Claim Block. During this exploration a geophysical programme, geologic mapping and a rock geochemical survey were completed. The primary conclusions are given below.

GEOPHYSICAL SURVEYS

The electromagnetic surveys identified three major conductive zones, the causative bodies of which are believed to be massive sulphides that are at a depth of greater than 60 metres. The conductors are believed to be plunging to the south ...though precise determinations were not possible from the data due to the depth to the conductors being equal to or greater than the coil separation used during the survey. Reliable interpretation of the gradient magnetic data was not possible until completion of the geologic mapping, at which time re-evaluation of the data was possible.

The near shore location of the conductor following The Narrows Break (Conductor B) permitted both geological and rock geochemical interpretations to be related to this conductor. The locations of the remaining conductors, however, lying beneath Savant Lake and beyond the limits of interpretation of both the detailed geologic and rock geochemical surveys precluded a meaningful discussion at this time.

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GEOLOGICAL MAPPING

The geological mapping determined that both an epigenetic environment and a syngenetic depositional environment occurred on the Savant Lake property.

Epigenetic mineralization in the form of pyrite, chalcopyrite, sphalerite and galena was noted in quartz veins. Although no surficial expression of syngenetic sulphide mineralization was found, the presence of iron formation suggests an environoment favourable to the deposition of massive sulphide type mineralization.

GEOCHEMICAL SURVEY

The rock geochemical programme identified the granodiorite intrusive and the Q_1 quartz veins as sources of hydrothermal alteration and the existence of alteration patterns characteristically associated with massive sulphide mineralization of the sedimentary volcanogenic type.

The alteration patterns on the property that are similar to those characteristically associated with massive sulphides were suggested as representing both distal and proximal alteration patterns. The distal alteration pattern was found to lie stratigraphically below the iron formations of the north-central portion of the claim block. The alteration exhibiting more proximal conditions was found to be located in the footwall rocks stratigraphically below a geologically anomalous stratigraphic break. The absence of any pronounced trace element alteration within the major element alteration zone indicates that the primary vent, and similarly the most likely location for sulphide deposition, is located at depth. This supports the geophysical data where a deep plunging source is indicated. From the rock geochemical data a depth to the source of the alteration is not believed to be in excess of 120 metres.

The mineralized lead-silver float was found to be similar to the Q_1 quartz veins distributed throughout the claim block.

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RECOMMENDATIONS

The encouraging geophysical, geological and rock geochemical results obtained from this phase of exploration warrant continued work on the Savant Lake Claim Block.

It is recommended that the geophysical conductors be re-surveyed with the MaxMin II electromagnetic system, employing various frequencies and cable lengths to arrive at a more quantitative estimation of the source geometry.

Before further field work is undertaken with respect to the Q_1 quartz veins, it is recommended that additional analytical determinations for gold be made on the Q_1 quartz vein material already collected. This data would determine whether or not the primary Q_1 quartz vein solutions contain any gold with the copper, lead, silver and zinc already determined to be present. If significant amounts of gold were found, further field evaluation of the quartz vein systems located on the property would be recommended.

All of which is respectfully submitted.

Yours very truly,

W. G. WAHL LIMITED

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ohn L. Wahl, Ph.D.

Toronto, Ontario November 23, 1977

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APPENDIX I SAMPLE COLLECTION AND PREPARATION CONTRACTOR OF A or strict train CONTRACTOR OF

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SAMPLE COLLECTION AND PREPARATION

- Approximately 1 to 1.5 kg of rock were collected from each outcrop site. An attempt was made in all cases to collect material from below the weathered surface.
- From each sample one piece was retained as a hand specimen for later reference.
- 3. Each sample was then crushed to 1/2 size and split several times to obtain a 50 gm homogenious sample.
- The 50 gms were then ground to -200 mesh in a "puck and ring" grinder.

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5. From this ground sample a 0.025 gm aliquot was used for digestion.



PREPARED FOR W. G. WAHL LTD.

This report describes the analytical procedures used for the analysis of rock powders sent to Barringer Research Limited, by Dr. John Wahl. It covers sample preparation and instrumental analysis by MRFPES.

I. SAMPLE PREPARATION

- (a) 0 250 grams of sample were weighed out into a 30 ml. teflon beaker.
- (b) Seven mls. of HF and 2.5 mls. of HNO_3 -HClO_A (3:2) were added.
- (c) The mixture was refluxed for 2 hours then evaporated to dryness over medium heat.
- (d) Another 2.5 mls. of HNO_3-HCIO_4 were added, and the mixture evaporated to dryness until $HCIO_4$ fumes disappeared.
- (e) The residue was taken up in 4 percent HNO₃ and made up to 25 mls. in a volumetric flask.
- (f) The resulting solution was used for multielement analysis.

II. INSTRUMENTAL PARAMETERS

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An inductively coupled argon plasma was used in the spectroscopic excitation source, the emissions for which were quantified by a QA-317 optical emission spectrometer (Applied Research Laboratories). A radio frequency current, at a frequency of 27.12 MHz and power of 1,600 watts is supplied to a two turn induction coil that surrounds the quartz plasma torch.

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The plasma gas flow is maintained at 1 1/min. and an outer shield of argon coolant at 10 1/min. isolates the plasma from the walls of the torch and keeps it from melting. Only high purity 99.999 percent argon (Liquid Carbonic) was used.

A cross flow nebuliser which requires argon at a pressure of 12 psi was used to aspirate the sample at 2 - 3 mls./min. The solution goes into a scaled-down Scott nebuliser chamber where the large droplets are removed by impaction, the very fine aerosols, (approximately 3 percent of the total amount nebulised) remaining is injected into the centre of the plasma by an argon carrier gas flow at 1.0 $\frac{+}{-}$ 0.1 1/min.

Emission measurements were made at a height of 16.5 mm above the induction coil. Optimisation of this observation position in the plasma is achieved by measuring the vertical and the horizontal profile of Mg emission.

The optical system consists of a 1-metre spectrometer, Paschen-Runge mounting, with an entrance slit of 12 microns and exit slits of 50 microns. It has a Bausch and Lomb grating with 1,960 grooves/mm. (blazed at 3,000 Å) giving a reciprocal linear dispersion of 5.1 Å/mm. and a spectral bandpass of 0.26 Å.

Emission signals were measured by multiplier photo tubes and their electrical outputs integrated over a 10-second period. Charge amp?ifiers for each element hold the signal for sequential transmission to an A/Dconverter and a digital data dump in millivolts (mv) is obtained from an HP-Thermoprinter, Model 9866.

The entire system is kept in a controlled environment room at 21° C $\stackrel{+}{-}$ 1° C and 50 percent relative humidity to minimise wavelength variations in the multi-channel optical system.

III. CALIBRATION AND SAMPLE ANALYSIS

Stock solutions for the preparation of multielement standards are prepared from Specpure or Analar grade reagents. The range of concentrations per

- 2 -

element (typically 3 - 4 orders of magnitude above the detection limit) is designed to cover that normally expected in the samples being analysed. Samples that exceed the calibration range are diluted 1:10 or 1:100, using a Repipet automatic dilutor.

At the start of the day, the instrument is calibrated using a series of multielement standard solutions (to provide at least 3 calibration points, excluding the zero standard]. A least-squares linear/parabolic fit is applied by a programmable Hewlett-Packard calculator, Model 9830, coupled to an HP Thermo-printer, Model 9866. The calibration coefficients are regularly updated throughout the day by aspiration of the standards.

Approximately one minute is required for the complete analysis of one sample. Uptake rate of the solution is 2 - 3 mls./min. A sample is first aspirated for 20 seconds before the emission signal is integrated for 10 seconds. At the end of the integration period, the integration signal is digitally transmitted to the calculator. The thermo-printer outputs in ppm of the analyte in solution, with the number of significant digits internally controlled by the calculator, based on the detection limits for that particular run.

Software developed by Barringer Research is used for further data reduction and data listing.

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IV. DETECTION LIMITS

	(ppm)		nd/d (bbw)		hð\d (bbw)
λq	10	Eu	10	Se	150
AÌ	1	Fe	1	Si	l
As	500	x	1000	Sn	500
Au	50	Mg	1	Sr	.1
В	15	Mn	5	Te	100
Be	.1	Mo	20	Ti	1
Ca	1	Na	1000	υ	90
Cđ	5	Ni	1	v	1
Co	2	P	1000	W	200
Cr	1	Pb	5	Zn	l
Cu	0.5	Ba	5		

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The above detection limits in the solid samples were obtained by the method.

V. LITERATURE REFERENCES

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- Velmer A, Fassel and Richard N. Kniseley, "ICP-Optical Emission Spectroscopy", <u>Analytical Chemistry</u>, 46 (1974): 1110A - 1120A, 1155A - 1164A.
- J. D. Winefordner, J. J. Fitzgerald, and N. Omenetto, "Review of Multielement Spectroscopic Methods", <u>Applied Spectroscopy</u>, 29 (1975]: 369 - 383.



SAMPLE VARIATION AND REPRODUCIBILITY

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Element reproducibility and sample variation were monitored throughout the analytical process by the inclusion of replicate samples at the beginning of the analytical run, after every 35 samples during the run and again at the end of the run. For this monitoring process sample SL 1 was used for which six separately digested portions were required. The analytical results, for the replicate samples, are given on the accompanying table.

A high degree of reproducibility for the major elements, Al through P on the table, is indicated by the low standard deviation values calculated for these elements. For the trace elements, Cu through V on the table, the reproducibility is not as good as for the major elements and is reflected in the. higher reported standard deviation values. These higher standard deviation values however, were not unexpected because the trace elements are generally present as or incorporated within sulphide complexes. In any particular sample the sulphides represent a heterogenic distribution resulting in inhomogeneities in the ground portion. It is these inhomogeneities that cause the greater variation of the trace element data. FORMING, 147-431-P. REPORT PAPER - GRANDA 105

STANDARD SL 1

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		<u>A1</u>	Fe	Ca	Mg	Ti	Na	<u> </u>	Mn	P
S	xg	8.17	8,17	7.60	3.41	0.6223	2.30	0.2323	0.1393	0.0865
Lion	SD	0.016	0.013	0.019	0.066	0.021	0.024	0.073	0.045	0.020
cent	Max	8.71	8.57	8.15	4.08	0.6592	2.45	0.2766	0.1581	0.0900
cent	Min	7.87	7.94	7.24	2.86	0.5848	2.14	0.1799	0.1219	0.0800
l Ouc	Range	0.84	0.63	0.91	1.22	0.0744	0.31	0.0967	0.0362	0.0100

	Cu	Zn	<u>Ni</u>	Co	Cr	2r	Sr	<u>v</u>
xg	65.0	102.0	102.0	84.0	350.0	59.0	316.0	192.0
SD	0.117	0.020	0.025	0.397	0.397	0.500	1.149	0.418
Max	87.0	105.0	109.0	540.0	600.0	60.0		302.0
Min	39.0	93.0	92.0	55.0	228.0	30.0	103.0	27.0
Range	48.0	12.0	17.0	485.0	372.0	30.0		275.0
	xg SD Max Min Range	<u>Cu</u> xg 65.0 SD 0.117 Max 87.0 Min 39.0 Range 48.0	Cu Zn xg 65.0 102.0 SD 0.117 0.020 Max 87.0 105.0 Min 39.0 93.0 Range 48.0 12.0	Cu Zn Ni xg 65.0 102.0 102.0 SD 0.117 0.020 0.025 Max 87.0 105.0 109.0 Min 39.0 93.0 92.0 Range 48.0 12.0 17.0	Cu Zn Ni Co xg 65.0 102.0 102.0 84.0 SD 0.117 0.020 0.025 0.397 Max 87.0 105.0 109.0 540.0 Min 39.0 93.0 92.0 55.0 Range 48.0 12.0 17.0 485.0	CuZnNiCoCrxg65.0102.0102.084.0350.0SD0.1170.0200.0250.3970.397Max87.0105.0109.0540.0600.0Min39.093.092.055.0228.0Range48.012.017.0485.0372.0	Cu Zn Ni Co Cr Zr xg 65.0 102.0 102.0 84.0 350.0 59.0 SD 0.117 0.020 0.025 0.397 0.397 0.500 Max 87.0 105.0 109.0 540.0 600.0 60.0 Min 39.0 93.0 92.0 55.0 228.0 30.0 Range 48.0 12.0 17.0 485.0 372.0 30.0	Cu Zn Ni Co Cr Zr Sr xg 65.0 102.0 102.0 84.0 350.0 59.0 316.0 SD 0.117 0.020 0.025 0.397 0.397 0.500 1.149 Max 87.0 105.0 109.0 540.0 600.0 60.0 Min 39.0 93.0 92.0 55.0 228.0 30.0 103.0 Range 48.0 12.0 17.0 485.0 372.0 30.0

xg = geometric mean

SD = standard deviation

Max = maximum

Min = minimum

	BRL	MgO Recommended	BRL	TiO ₂ Recommended	BRL	MnO Recommended
AGV-1 RSD % *	1.39 1.6 -6.7	1.49	1.06 4.4 -1.9	1.08	0.095 1.3 -3.1	0.098
BRC-1 RSD ♥ △ ♥	2.70 0.5 -18	3.28	2.18 2.5 -2.2	2.23	0.174 1.2 -1.1	0.176
G-2 RSD ₪ ∆ ₪	0.753 1.4 -3.7	0.782	0.503 1.2 -5.3	0.531	0.033 0.5 -11	0.037
GSP-1 RSD ♥ △ ♥	0.965 0.9 +0.8	0.957	0.691 2.7 -1.1	0.699	0.040 1.5 -9.1	0.044
SY-2 RSD \ A \	2.10 1.1 -21	2.66	0.132 1.2 -12	0.15	0.292 1.4 -8.8	0.32

COMPARATIVE DATA FOR U.S.G.S. AND C.R.M.P. ROCK STANDARDS (MAJOR ELEMENTS, PERCENT)

* n = 4

- 43 -

	BRL	SiO ₂ Recommended	BRL	Al ₂ 0 ₃ Recommended	BRL	Fe ₂ 0 ₃ Recommended	BRL	CaO Recommended
AGV-1 RSD \$ * Δ \$	60.0 0.8 +1.6	59.00	17.7 1.9 +4.1	17.01	6.96 2.5 +2.4	6.80	5.04 1.2 +1:2	4.98
BCR-1 RSD % A %	55.6 0.3 +2.1	54.48	13.6 0.7 -0.4	13.66	14.6 1.0 +8.1	13.51	6.82 0.5 -1.9	6.95
G-2 RSD ♥ △ ♥	69.9 2.0 +1.0	69.19	16.2 1.5 +5.6	15.34	2.67 1.1 -3.6	2.77	1,96 1.4 -1.5	1.99
GSP-1 RSD ₪ ∆ ₪	67.7 0.5 +0.6	67.28	15.9 0 +5.2	15.12	4.49 1.4 +3.7	4.33	2.05 0.6 +1.0	2.03
SY-2 RSD & · A &	63.9 0.9 +6.4	60.07	11.8 1.1 -2.9	12.15	6.08 1.2 -4.1	6.34	7.74 1.1 -3.6	8.03

COMPARATIVE DATA FOR U.S.G.S. AND C.R.M.P. ROCK STANDARDS (MAJOR ELEMENTS, PERCENT)

n = 4

- 44 -

	BRL	Ba Recommended	BRL	2n Recommended	BRL	Ni Recommended
AGV-1 RSD \$* A \$	1290 1.6 -9.5	1410	106 3.0 -5.4	112	15 22 -16	17.8
BCR-1 RSD % A %	695 5.7 -12	790	137 0.5 +3.8	132	16 5.5 +6.7	15.0
G-2 RSD \$ ∆ \$	1940 0.3 -0.5	1950	89.6 2.8 +20	74.9	3.1 8.0 -52	6.4
GSP-1 RSD & A &	1390 5.4 +2.2	1360	109 3.4 -24	143	7 16 -35	10.7
SY-2 RSD & A &	508 7.6 +13	450	205 1.1 -18	250	12 3.3 +20	10

1 45; 1

COMPARATIVE DATA FOR U.S.G.S. AND C.R.M.P. ROCK STANDARDS (TRACE ELEMENTS, PARTS PER MILLION)

*n = 4

	0		1		 			
	BRL	Cu Recommended	BRL	Sr Recommended	BRL	V Recommended	BRL	Be Recommended
AGV-1 RSD %* A %	61.7 3.0 -3.1	63 . 7	654 3.0 -0.5	657	116 3.0 -4.1	121	2.2 2.6 +22	1.8
BRC-1 RSD % Å %	22.1 9.2 -1.3	22.4	315 0.4 -8.7	345	362 1.2 -5.7	384	2.9 6.4 +11	2.6
G-2 RSD ♥ ∆ ♥	11.9 4.2 +11	10.7	488 1.9 +5.4	463	38 1.2 +2.7	37.0	1.9 1.8 -21	2.4
GSP-1 RSD % & %	36.6 1.2 +4.0	35.2	238 0.9 -3.6	247	56 0.8 +7.7	52.0	1.3 3.1 · +62	0.8
SY-2 RSD \$ A \$	5.4 13 +8.0	5	242 2.9 -3.2	250	50 1.5 0	50	14.6 1.3 -27	20

- 46 -

COMPARATIVE DATA FOR U.S.G.S. AND C.R.M.P. ROCK STANDARDS (TRACE ELEMENTS, PARTS PER MILLION)

* n = 4



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	DDH	SAMPLE C FUDIG	LASSIFIC SAMPLE NUMBER	RT	LOCATION	AND RE FE	SULIS FOR CA	TOTAL MG	PLASMA DE	TERM INAT I	ONS K	MN	p	4 F
··•••••	•	· · · · · · · · · · · · · · · · · · ·		с	81500	-197966	83301	30503	6400 -		° 2500	147:		
	f	ι,	4	o	85500	81200	71974	33405	5120	20100	21.30	124.	870	
• •		i)	క	c	81900	826û©	85100	29000	5.40	5500C	50.00	1520	920	
	1	ð	6	G	83800	83604	56601·	71000	5340	22800	3020	1396	960	
	T	· · · · · · · · · · · · · · · · · · ·				79000-	56600	68307	5877	26700	\$300	1 \$80		
	C	••	16	0	82300	85400	69100	33300	66 90	21200	2000	1336	950	
	4	ζı.	17	0	83100	83800	72000	33340	576r.	2330r	55) ()	1320	8 <i>ò</i> Ĵ	
	ť	0	19	6	78100	84400	7⇒40î-	361 00	6200	21200	2	1220	840	
			27	5	80300	84000	74.857	32400	6610	-19300	1.301	1150	881-	
	* 1	.:	21	õ	81800	87306	65800	36100	5840	17800	25 <i>i</i> i)	1350	850	י בג
		¢	24	U	86300	77766	62500	44804	5310	24000	28.30	1260	880	9
	(()	25	U	80830	73200	72700	38160	5530	24600	1830	1120	840	I
	••••••				82666	75100	64300	41300	6530	22700-	55.5.4	111:	875-	
			29	Q	81400	86 . 5 6	755 PT	35107	7 1.	24601	2	1470	950	
	۰.	•	32	5 · · ·	777(;0	841 C C	71500	33300	657.	21906	1510	1370	830	
	٢	2	33	9	77260	76200	83864	33300	6195	20800	22.1	153	860	
		· · · · · ·	34			72700	494707	41301	4280	19707	2250	1130		
	,	'n	38	¢	74330	76400	65000	39100	592	18460	22 15	128.	835	
		÷	41	ť	81260	2.466	6666	43605	5830	18300	24 < 6	1300	940	
		L.	48	Ō	86540	79746	696.00	4360.	4934	17860	2330	1160	830	

DDH	SAMPLE C FUOTG	LASSIF1 SAMPLE NUMBER	CATION RT	• LUCATIC	N, AND PE FE	SULTS FOR	P TOTAL MG	PLASMA DI TT PPM	ETERMINATI NA	ONS K	MN	р
	5	45-	0	83200	49400	54300	~6770¢	4140-	17500	22010		790
C,	e	52	o	81920	84610	73000	43600	6360	18500	4660	1410	920
•	, n	53	0	80056	81000	709C0	35100	6636	24200	1354	1630	930
6	Q	- 54	Ċ	81906	92600	656 <i>0</i> ù	34200	6720	19890	22.20	1590	941
с. С			0	-733350	67200	57900 1	3819 0'''	4986~		2201	1.76	700
1,	G	58	0	81800	76000	67800	43600	5070	17900	3495	1200	710
¢	C	59	0	93500	77206	76100	41300	5130	20800	1900	1180	860
L	•2	61)	C C	84900	79200	76766	41300	4930	16980	1500	1370	900
····- ,	<u>ې</u>	61	6	98590	77600	69200	39105	4730	22800	2 80 t	1310	96
	6	62	o	84900	81150	68460	44800	5370	21406	1856	160.0	88
17	à	63	o 1	85400	83730	82800	34200	3391	237.00	2640	1310	101
• .	Q.	.78	υ	132000	75800	53800	42401	4760	5500H	51.79	1230	109
	مانچ است. ح	82	0	69600	77706	55800	33300	5740	19300	3650	1490	76
••	:)	84	e	75500	79.066	74960	36100	5710	23500	1815	1530	83
ť	e	87	C .	77800	85700	26007	79900	4826	70 90	6 9.00	1286	95
	•	88	c	85100	77400	65901	46230	4720	20700	2560	1410	81
• • • •	່			51602	···1116¢¢···	48300	26500	10400	19200	2700	1719	 118
	4	96	~	744.30	107600	46360	26500	11300	23501	2500	1510	116
	. :	97	C	96000	10033	66600	38100	5370	20600	28.0	1690	95
. •	t.	99	(98700	81400	62000	38109	6414	22200	3295	1630	95

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•				4 (1 1.) 2		с. Сам								* .)
	ррн	SAMPLF C FOOTG	LASSIFI SAMPLE NUMBER	CATION	LDCATIC	DN. AND RE	SULTS FOR CA	TUTAL P	LASMA DE TI	TERNINAT	ICNS K	. MN	ρ	
			101	<u>.</u>	75200			-26500		-\$6900	2500-		1220	، ۲ بر میں میں م
	6	c	107	ō	77500	87200	44600	3710	561	21370	1.8/92	1720	870	
- .		- •,	110	0	79600	S1400	64000	39100	5770	14000	5500	1680	870	•
		, o	117	0	77000	103000	45500	39100	7270	24300	3600	1550	1110	
		····· 0	125		84000	97900	1920	85607		4120	4100	10:17	910-	
	- Ĺ	-	132	C	80600	81000	51200	42400	6820	29000	2506	1380	1050	
	· (,	ti.	133	6	71900	71306	7660A	54600	3380	16400	2200	1490	750	
	4.	ů.	136	c	81000	118400	39400	38100	6820	20000	2400	2050	1170	
				0	80400-	\$9300	73000			17300	1720	1 65 7	930-	
	e,	3	139	e	84800	85700	76700	39101	6160	20230	17.56	1530	870	
	ć.	э	140	o	52900	67700	27600	38100	3886	7150	1200	989	600	
	ť	Ċ.	142	0	85500	84206	71700	381 0,0	6(50	16700	1400	1560	960	
			143		83866	90206	69600 -	- 41300	5820		18.0	1740		
	\.	6	149	G	85500	84200	88600	38100	5920	19900	17:00	1670	930	
•	,		151	r · ·	74506	7300C	83000	39107	4850	12400	1708	1470	850	
	,		152	C'	75800	8060 0	65000	43600	4990	19660	19:00	1720	880	
			153"	<u> </u>	влеас	75.00	68700	40200	5330 -	12307	1600	1430		
		ņ	154	6	78606	75360	74983	39100	5170	14200	1700	1350	850	
			163	0	84500	85700	2496.	65600	3650	9680	1150	1630	770	
	•		164	n	00 26.0	83504	8. 964	43665	549:	13400	17.04	1430	840	

 $\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{n}\sum_{j=1}^{n}\sum_{i=1}^{n}\sum_{j=1}^{$

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DDH	SAMPLE (FODTG	CLASSIFIC SAMPLE NUMBER	CATION: RT	LOCATION AL	AND RE	SULTS FOR	TOTAL MG	PLASMA DE TI	TERMINAT	IONS K	Mi)	p
		170		78900		37600	43605	5400	17200	1400	-1:52:0	852
0	û	172	0	77500	9020C	72800	36100	6600	17700	1906	1520	94 0
4,	\$	174	0	43300	53200	37300	21200	5160	1350	500	877	570
e	Ģ.	11	0	83350	72800	31800	58000	525¢	20800	221	992	920
6	· · · · · · · · · · · · · · · · · · ·	23				**************************************	° 39100 °	54.30m	- 21800	2300	-1420	870-
ú	2	28	0	79600	83100	76800	35100	6790	20700	2000	1370	890
6	Ú.	. 3 0	ð · · · ·	79400	7910C	68200	39100	5300	20000	3000	1290	820
¢.	ц,	31	0	73200	86200	79999	41300	6690	17600	2801	1170	900
•	<u> </u>			80200	77500	35800	47.300	5810			1170-	880-
r.	J	76	Ċ	91500	81100	66900	40200	5180	21600	55 96	1180	930
6	9	94	Q ·	85600	79650	65200	41300	4680	22200	3600	1390	790
C	(,	95	Q	78530	113000	65830	27300	16836	19400	3600	1880	1190
		IC 8	0	77800			39100	5170		2434	1670	900-
	r,	169	o	69504	84900	57493	37100	5086	10800	6200	1589	870
• •)	112	0	74600	164000	28500	597NJ	7530	19400	24).	1290	1010
ſ	•)	113	ز	769.10	84406	69000	41300	537a	19800	2200	1570	880
	· · · · · · · · · · · · · · · · · · ·	123		71500	84855	"966 - "	39100	5540	-1667g	1894	1530	
t		146	3	70100	167390	344 4	31500	12700	24000	1844	1390	1090
		148	U .	88600	161106	- 33601	46101	8520	26610	5836	1580	1030
• •		• • •		72000	CA7.1C	4440	4870.	3770	180.20	4666	1835	- 880

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	Barred 15 - pr- 440 - 1	- a- ar a raar ababbat builded at	Ma untitus Aunderstation as yo ditatas	14	mantine anni al airsonile an airson an an an		یند و در ایرونه ایم زیادیون در ایو اینو برو ا	1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (1.40 (م <mark>س</mark> وم به رابید. ا	entrolausjatis sa nasas caasis oo oo		a sanan ar sa sa ar a a an T	under ungegenzenspesieller die der einen einen einen	
	ррн	SAMPLE C FUUTG	LASSIFI SAMPLE NUMBER	ICATION. E RT R	LOCATION	AND FE	RESULTS FOR CA	TUTAL MG	PLASMA DE	TERMINATI NA	ONS K	Mfi	p p	

1894 - 1994 - 1997 - 1997 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 - 1994 -	c	·	155	<u>, </u>	72900	78300	64300	-36100		- 20300	5530		850-	
	•	0	152	0	93700	163000	27700	67700	5030	19500	1900	1700	950	
•·· •• • • • • •	1975 - 1 9	• • • • •	162	່ ປີ 🛄	71500	75400	17860	40200	5560	26800	1406	1150	850	
	2	Ŭ	165	Ŭ,	83500	87600	57400	46100	4880	12800	1620	1590	940	
			36	0			43000		3940	-17800	···· 1900 ···	1410	820	
	e	0	47	e	80000	79900	59208	42400	5220	22000	5500	1189	850	
	ů Tri v	о ^с	51	<u>c</u>	81500	54100	64600	42400	6470	17400	2204	1240	930	- 12 ⁻ -
	Ç	a	55	C	83200	83300	57900	48700	5230	16700	2300	1160	730	÷
<u> </u>	<u></u>	6	80		88500		66500	-44800-	5250	<u></u>	1670	1.540	820	
	u		83	C	90700	88100	59000	39100	5110	18100	2300	1520	850	្រុ ហ
	۲,	, n .	92	0	83500	8236.0	69600	40200	4410	13700	46 213	1580	870	ũ
	ς,	, U	129	G	80900	88500	68706	34200	6280	13500	2500	1650	880	ł
	· •			0		105700	53400	-43600	4970		5000	1910-	1170-	·
	ů	; ex	161	S	69600	89466	37300	448-30	4445	18890	1700	1645	920	
	ن. م)	166	o	80300	86906	56430	44800	5466	14500	1630	1450	880	
	ζ,		107	n	83100	83200	64300	46100	5846	15100	1900	1519	870	
		, , , , , , , , , , , , , , , , , , ,	168		65300	36693	122000	34200	476r	5360	··· 1255*		780-	
	U.	. 43	171	G	83800	81200	. E93 00	41300	5440	22406	1600	1659	880	
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	DDH	SAMPLE C FUUTG	LASSIFI SAMPLE NUMBER	CATIC	DN. LDCA	TICN, AND ZN	RESULTS F	UR TOTA	AL PLAS	MA DETER SR	MINATION AU	5 AS	3N	мо
									A	1				
•		от то по то с		00	112	120	1.5	· D · · · ·	286	105		13 - 1		
	÷)	U O	4	Дарана Дарана	120	168	88 2.1	(, 2	264	101	1, 1)	د. م	с 	U O
		ц,	ະ • 6	С	77 10 A	114 Q.)	DI RA	n n	217	197 16 A	c a	. • •	0	0
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1	Ministry of Natural Resources	File A A A A	
	GEOPHYSICAL – GEOLOGICAL – GEOCH TECHNICAL DATA STATEMEN	IEMICAL T	800 800
	TO BE ATTACHED AS AN APPENDIX TO TECHNIC FACTS SHOWN HERE NEED NOT BE REPEATED I TECHNICAL REPORT MUST CONTAIN INTERPRETATION,	AL REPORT IN REPORT CONCLUSIONS ETC.	
	Type of Survey(s) Ceophysical -Geological -Geochemical Township or Area Javant Claim Holder(s) WiG. WAHL LTD, 1101, 302 BAY ST TORONTO ONT MSH 2P3 Survey Company W. G. WAHL LTD Author of Report J.L. WAHL AND A. Gualass Address of Author 101 - 302 BAY ST TORONTO ONT MSH 2P3 Covering Dates of Survey Jule '77 Harcone, 100 '77 (linecutting to office) Total Miles of Line Cut 1910 miles SPECIAL PROVISIONS Geophysical DAYS Per claim Geophysical DAYS	PROJECTS UNIT MINING CLAIMS TRAVERSED List numerically May	
	ENTER 40 days (includes line cutting) for first survey. Electromagnetic 40 ENTER 20 days for each additional survey using same grid. Radiometric 20 Geological 20 10	$\frac{7}{34} PH = 376609$ $\frac{3}{4} PH = 376610$ $\frac{1}{2} C PH = 376611$ $\frac{1}{2} C PH = 376612$	lf space unufficient, at
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Ministry of Natural Resources GEOPHYSICAL – GEOLOGICAL – GEOCHE TECHNICAL DATA STATEMENT	File
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Type of Survey(s) Scophysical-Scalegical-Scale	MINING CLAIMS TRAVERSED List numerically PA - 375985 (prefix) (number) PA - 375986 PA - 375988 PA - 375988 PA - 376607 PA - 376608 PA - 376608 PA - 376608 PA - 376610 PA - 376612
	TOTAL CLAIMS

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GROUND SURVEYS - If more than one survey, specify data for each type of survey

2	Sumber of Stations 1520 May, 730 EMNumber of Readings 4920 May, 1460EM									
5	tation interval 30 meters EM, 15 meters Mag Line spacing 60 meters									
ł	rofile scale $1^{\prime\prime} = 20\%$									
(ontour intervalasinducetted									
MAGNETIC	Instrument <u>Servetric 6-816 Magnetonck</u> Accuracy - Scale constant <u>I gemma</u> , +59000 gemmes Diurnal correction method <u>time atrupilated</u> Base Station check-in interval (hours) <u>1/2 - 1 hour</u> Base Station location and value <u>al base line statung</u> standardized as included by =									
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ETIC	Coil configuration have a tol									
N	Coil separation 400 feet									
AMO	$\frac{1}{\sqrt{3}} \frac{1}{\sqrt{2}} \frac{1}{\sqrt{2}$									
IRC	Method: Fixed transmitter Shoot back I In line Parallel line									
EC	Frequency 444 Hz AND 1777 HZ									
E	(specify V.L.F. station)									
	Instrument									
	Scale constant									
λIJ	Corrections made									
AV										
S	Base station value and location									
	Elevation accuracy									
	Instrument									
	Method Time Domain									
	Parameters On time Frequency									
A	- Off time Range									
IIV	- Delay time									
SIS	Integration time									
RE	Power									
	Electrode array									
	Electrode spacing									
	Type of clectrode									

NYDOGAD FOLTARIZATION

N	umber of Stations_1520_Mag730_FM	Number of Readings 4	520 Mag 1460 EM								
St	ation interval 30mors EM, 15mors Mag	Line spacingOnen	erg []								
Pr	ofile scale 1"=20%		•								
C	ontour interval Az Indicated										
J.	Instrument <u>Geometric G-BIG Magnetone</u>	ter									
11	Accuracy - Scale constant - 1 gemma , +59,0	008									
	Diurnal correction method time extrapolated										
MM	Base Station check-in interval (hours) 1/2 to 1 hour		· · · · · · · · · · · · · · · · · · ·								
	Base Station location and value <u>all base line</u> sh	ations standardized as in	dicated by								
الا											
112	Coil configuration horisontal										
3	Coil separation 400 feet										
C N	Accuracy = 4.9 TE = 10										
	Method:	Church had									
1	Frequency 444 HZ ALD 1377 HZ										
1	(specify V.L.F. station)										
	Parameters measured <u>preseter and guade</u>	ATURE									
	Instrument										
H	Scale constant										
	Corrections made										
5	Base station value and location										
	Elevation accuracy										
	Instrument		·								
	Method 🔲 Time Domain	E Frequency Don	nain								
	Parameters - On time	Frequency									
	Off time	Range									
2	- Delay time										
2	Integration time										
Y FY	Power										
-1	Electrode array										
	Electrode spacing										
	Type of electrode										
1.											
		•									

WALLAND CONTRACTOR

GEOCHEMICAL SURVEY **PROCEDURE RECORD**

Numbers of claims from which samples taken_33 	1207, 376608, 376609, 37660, 376610, 376612, 375985
Total Number of Samples175 Type of Sample(Nature of Material) Average Sample Weight1-11/2_kg Method of CollectionQutr 598	ANALYTICAL METHODS Values expressed in: p. p. m. p. p. b.
Soil Horizon Sampled	Cu, Pb (n Ni) (Ng, 10) As-(circle) Others Fe, Al Carle, Na, K, P. Ylm, Cr, Cd, Se, Fiel J Analysis (

Drainage Development
Estimated Range of Overburden Thickness

Sample Depth_____

Terrain_____

SAMPLE PREPARATION

(Includes drying, screening, crushing, ashing)

Mesh size of fraction used for analysis____ Grushing & Grinding to -200 mesh.

General

Reagents Used _____ Field Laboratory Analysis No. (______tests) Extraction Method Analytical Method _____ Reagents Used _____ Commercial Laboratory (_____175_____tests) Name of Laboratory Barringer Resurch. Extraction Method Correcte Analytical Method Plasma

Extraction Method HE-HCIOy-HNO3-

Analytical Method_ Plasma

Reagents Used HF-HCLOy - HND2

General _____

GEOCHEMICAL SURVEY - PROCEDURE RECORD

otal Number of Samples175	ANALYTICAL METHODS Values expressed in: per cent p. p. m. p. p. b. Cu Pb (n. Ni Co (k) Mr, As, circle Others Al, Fr, Mn, Ca, Ma, K, P, Cl, Ca, Be Field Analysis (
bil Horizon Sampled	ANALYTICAL METHODS Values expressed in: per cent p. p. m. p. p. b. Cu Pb (n. Ni Co (k) Mo, As, circle Others Al, Fe, Mn, Ca, Ma, K, P, Cl, Ca, Be. Field Analysis (
verage Sample <u>Qut to p</u> chod of Collection <u>Qut to p</u> bil Horizon Sampled orizon Development ample Depth	Values expressed in: p. p. m. p. p. m. p. p. b. Cu Pb (n. Ni Co (g) (Mr, (As, circle Others Al, Fe, (In, Ca, Mg, Ma, K, F, Cd, G, Be. Field Analysis (
verage Sample Weight <u>1-1/2 kg</u> ethod of Collection <u>Outers p</u> oil Horizon Sampled orizon Development ample Depth	p. p. m. p. p. b. p. p. b. Cu Pb (n. Ni) Co (k) Mo, (As, circle Others Al, Fe, Ma, Ca, Ma, K, P, Ca, Be. Field Analysis (
ethod of Collection Out trop	Cu Pb (n. Ni Co (Mo, As, circle Others Al, Fr, Mn, Ca, Mq, Ma, K, P, Cd, Ch, Be Field Analysis (
bil Horizon Sampled orizon Development ample Depth errain	Cu) (Pb) (n, (Ni,) Co, (k) (My, (As, icircle Others A) Fr, (In, Cu, Mg, Mg, K, F, Cl, S, Be Field Analysis (
oil Horizon Sampled orizon Development ample Depth errain	Others Al, Fr. An, G., Ma, K., P. Cl.G., Be. Field Analysis (test Extraction Method Analytical Method
orizon Development ample Depth	Field Analysis (test Extraction Method Analytical Method
ample Depth	Extraction Method Analytical Method
crrain	Analytical Method
	Reagents Used
rainage Development	Field Laboratory Analysis
stimuted Range of Overhurden Thickness	No (
simated Range of Overbuilden Threeness	Extraction Method
	Analytical Method
	Analytical Method
	Reagents Usca
SAMPLE PREPARATION	Commercial Laboratory (175 tes
(Includes drying, screening, crushing, ashing)	Name of Laboratory Parting T
esh size of fraction used for analysis	Extraction Method
Irushed + Ground To -200 mesh	- Anthrical Mathed Discond
	- Reagents Used <u>Hr - mkvs - 11 CUU</u>
eneral	General
	-
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	-
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Dotario

Ministry of Natural Resources

Your file:

Our file: 2.2544

1978 10 06

Mrs. Doris Cosco Acting Mining Recorder Ministry of Natural Resources P.O. Box 669 Court House Sioux Lookout, Ontario POV 2TO

Dear Mrs. Cosco:

Re: Mining Claims Pa. 375985 et al. Savant Township File 2.2544

The Geophysical (Electromagnetic & Magnetometer), Geochemical and Geological assessment work credits as listed with my Notice of Intent dated April 21, 1978 have been <u>approved</u> as of the above date.

Please inform the recorded holder of these mining claims and so indicate on your records.

Yours very truly,

Ciciacol

J.R. McGinn Director Lands Administration Branch

Whitney Block, Room 6404 Queen's Park Toronto, Ontario M7A 1W3 Phone: 416/965-6918

DN:ie

cc: W.G. Wahl Ltd. Toronto, Ontario

> Resident Geologist ' Sioux Lookout, Ontario



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		KEVISED
 	 < -	

(Supersedes	Notice	of	Intent	statement	dated	April	21st.	1973)	
Recorded Holder									

W.M. Wath Londted

Savart Townshi;

TONTSHI THAT A

Type of survey and number of Assessment days credit per claim	Mining Clzims Assessed
Geophysical	
Electromagneticdays	Pa. 375985 to 88 inclusive
Magnetometer <u>20</u> days	376607 to 12 "
Radiometric days	
Induced polarization days	
Section 86 (18) days	
Geological days	
Geochemical days	
Man days 🗌 🛛 Airborne 🗌	
Special provision 🔀 Ground 🕱	
Credits have been reduced because of partial coverage of claims.	
Credits have been reduced because of corrections to work dates and figures of applicant.	
Special credits under section 86 (15a) for the following	mining claims
	· ·
No credits have been allowed for the following mining claims	
not sufficiently covered by the survey	Insufficient technical data filed
1	~
,	
The Mining Recorder may reduce the above credits if necessary in order that the total number of approved assessment days recorded on each claim does not exceed the maximum allowed as follows; Geophysical — 80; Geological — 40; Geochemical — 40; Section 86(18) 60;	

Ministry of Natural Resources

Ontario

Your file:

1979 01 12



58 3195W

Mr. Albert Hanson Mining Recorder Ministry of Natural Resources P.O. Box 669, Court House Sioux Lookout, Ontario POV 2TO

Dear Sir:

Re: Geophysical (Electromagnetic & Magnetometer) Geological and Geochemical surveys Mining Claims Pa. 375985 et 21. Savant Twp. File 2.2544

I have re-examined and assessed the Magnetometer survey on the above mining claims in regards to additional data (new plan) being submitted.

35) 275

Attached herewith is a revised statement which supersedes the list sent with my Notice of Intent dated April 21st, 1978 and is finalized as of the previous 'date of approval' October 6th, 1978.

Yours very truly, Morton J.B Acting Director Lands Administration Branch Whitney Block, Room 6450 Queen's Park Toronto, Ontario M7A 1W3 Phone: 416/965-6918 OJ:ie مانزر cc: W.G. Wahl Limited Toronto, Ontario Attn: Mr. John L. Wahl Resident Geologist

Sioux Lookout, Ontario



Gillis Twp.

FOR ADDITIONAL

INFORMATION

SEE MAPS:

527/095W-0024 #1-#7







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. <u>527/0950-0024-#3</u> DENISON MINES LTD.

WH HY PEL SAVANT LAKE PROJECT THATTE BY HEY MAGNETOMETER SURVEY SPENCES C 444 Corrected Total Field Values in Gammas 52-J 9 *ы*, , BACKGROUND: 59,000 gammas 9#5 N2 ANILLEU W.G.WAHL Limited 1cm = 30m





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52J095W2354 52J095W0024 ENDOGOKI LAKE

٢ 527/095W-0024-#4 DENISON MINES LTD. D PROFEESIO DRAWN BY REV. SAVANT LAKE PROJECT TRACED BY REV. D. G. WAHL CONTOURED MAGNETIC DATA Carlos - ----APPROVED REV. Corrected Total Field Values in Gammas N.7. S. 52-J-9 REV. aà Background — 59,000 gammas DWG. NO. W.G.WAHL Limited 1cm = 30m My I Wah

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